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BULLETIN OF THE GEOLOGICAL SURVEY OF JAPAN

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Special Issue:
Scientific results from InterRad XV in Niigata 2017 (Proceedings)



表紙の写真

InterRad XV の集合写真

InterRad XV (国際放散虫研究者協会第 15 回会議) の開会式で撮影された全参加者の集合写真 (2017 年 10 月 23 日, 新潟大学中央図書館ライブラリーホール)。世界 16 カ国 (オーストラリア, 中国, フランス, ドイツ, インドネシア, イタリア, 日本, 韓国, モンゴル, フィリピン, ロシア, スロベニア, スペイン, スイス, トルコ, アメリカ) から総勢 187 名の参加があった。1980 年代から, 産総研地質調査総合センターはナショナルセンターとして日本国内の放散虫研究を支援してきており, 2017 年 10 月 20 日~11 月 1 日に開催された InterRad XV を共催した。

(写真: InterRad XV in Niigata 2017, 文: 中江 訓)

Cover Photo

Group photograph of InterRad XV

Group photograph of the all participants in the opening ceremony of InterRad XV (15th meeting of the International Association of Radiolarists) taken at the Library Hall in the Central Library of Niigata University on 23 October, 2017. A total of 187 participants from 16 countries (Australia, China, France, Germany, Indonesia, Italy, Japan, Korea, Mongolia, the Philippines, Russia, Slovenia, Spain, Switzerland, Turkey and the United States of America) attended the conference. Since 1980s, the Geological Survey of Japan has been supporting radiolarian studies in Japan as a national center of geological research and thus co-hosted InterRad XV which was organized from 20 October to 1 November 2017.

(Photograph by InterRad XV in Niigata 2017, Caption by NAKAE Satoshi)

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GSJ Bulletin Special Issue: Scientific results from InterRad XV in Niigata 2017 (Proceedings)

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Keywords: radiolaria, biostratigraphy, paleoceanology, geology, tectonics, InterRad.

InterRad is an international association of radiolarists, and is a non-profit organization that promotes research on all aspects of radiolarian biology, ecology, taxonomy, evolution, paleobiology, paleoecology, paleobiogeography and biostratigraphy. Since 1978, with the aim of providing an opportunity to exchange ideas for understanding all aspects of radiolarian-related topics beyond the radiolarian society, InterRad has convened conferences regularly every three years in which a wide range of research papers have been presented. The 15th meeting of the International Association of Radiolarists (InterRad XV) was held mainly in Niigata, Japan from 20 October to 1 November 2017, co-hosted by the Geological Society of Japan, the Palaeontological Society of Japan, the Society of Science on Form, Japan, and the Geological Survey of Japan, AIST (National Institute of Advanced Industrial Science and Technology). A total of 187 participants from 16 countries (Australia, China, France, Germany, Indonesia, Italy, Japan, Korea, Mongolia, the Philippines, Russia, Slovenia, Spain, Switzerland, Turkey and the United States of America) attended the conference. The next meeting (InterRad XVI), once decided during the business meeting to be held in Ljubljana, Slovenia in September of this year, is postponed until September 2021 due to the COVID-19 pandemic.

The scientific sessions, held at Niigata University on 23–27 October, focused on the five thematic topics (1–5) devoted as a special symposium reflecting the prevailing research directions and provided other eight general themes (6–13) to wide range of radiolarian studies. They are (1) Paleooceanography of Tethys and Panthalassa (chairs: S. Takahashi and P. O. Baumgartner), (2) Cenozoic paleoceanography in marginal seas (chairs: T. Itaki, Y. Okazaki and R. W. Jordan), (3) Biology and paleobiology of shelled Protista (chairs: K. Kimoto and F. Not), (4) An interface between function and evolution (chairs: Y. Tokuda and Y. Shiino), (5) Jurassic–Cretaceous boundary (chairs: A. Matsuoka and G. Li), (6) Insightful studies for radiolarians (chairs: Y. Aita and J. Rogers), (7) Biosiliceous records (chairs: J. Rogers and Y. Aita), (8) Modern oceanography (chairs: S. R. Hori and K. Kuwahara), (9) Paleobiogeography (chairs: K. Kuwahara and S. R. Hori), (10) Evolution and diversity (chairs: W. H. He and M. Chiari), (11) Biostratigraphy (chairs: M. Chiari and W. H. He), (12) Tibetan tectonics (chairs: T. Danelian and H. Luo) and (13) European tectonics (Chairs: H. Luo and T. Danelian). These sessions attracted 128 papers including oral and poster presentations and the abstracts were published as the Volume 40 of *Radiolaria*, the formal newsletter of InterRad.

The proceedings of the InterRad XV have been separately published as the special issues of *Island Arc* (Sashida *et al.*, 2019), *Paleontological Research* (Matsuoka *et al.*, 2019) and *Revue de Micropaléontologie*. This time the special issue of the *Bulletin of the Geological Survey of Japan* (this issue) is newly released to give representative research topics discussed at the InterRad XV as well as some later invited articles (five research articles, one report paper and one note and comment, together with Frontispiece, are included).

Radiolarian biochronological study:

Suzuki and Gawlick (2020) describe a well-preserved radiolarian fauna from bedded radiolarites of the Fludergraben section in the Northern Calcareous Alps, Austria. These radiolarites deposited just above the Klaus Formation dated by ammonites at latest Callovian or the Callovian–Oxfordian boundary, thus this fauna is undoubtedly assigned to the early Oxfordian age. New index species including *Kilinora spiralis*, *Fultacapsa sphaerica*, *Protunuma japonicus* and *Pseudoeucyrtis reticularis*, which were first appeared in the early Oxfordian, can be distinguished from long-lasting radiolarian species coming from the Callovian. The authors discuss these results and redefined the *Williriedellum dierschei* Zone (lower–middle Oxfordian), which was previously ranked as a subzone in the *Zhamoidellum ovum* Zone, on the basis of the new index species. These new findings fill a gap

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in the definition of the Oxfordian by radiolarians and result in a better resolution of the radiolarian biostratigraphy.

Nishizono and Yonemitsu (2020) report the first discovery of radiolarian fauna from seven localities of the uppermost Toyora Group, which is one of the representative Lower to Middle Jurassic strata in Japan and is famous for containing abundant ammonoid. The radiolarian faunas consist of representative species of the *Transhsuum hisuikyoense* and *Striatojaponocapsa plicarum* zones, which have formerly been correlated to the Aalenian–Bajocian and the lower Bathonian respectively. The authors point out that this age-assignment is slightly older than the age determined by previously reported ammonoids and inoceramids.

Ito (2020) first compiled the classification of intra-formational structures of striped chert observed in the Jurassic accretionary complexes in the Inner Zone of Southwest Japan, and their radiolarian ages. Furthermore, the author examined the relation between the type of striped chert and the age at four sections in Ashikaga and Sano cities of Tochigi Prefecture, Japan, and consequently summarizes the striped chert might be useful as an alternative age index for the Triassic at least for present four sections.

Radiolarian fauna related to Jurassic accretionary tectonics in Japan:

Uchino and Suzuki (2020) demonstrate the geological map, lithology and radiolarian age of accretionary complexes in the North Kitakami Belt in the northeastern Shimokita Peninsula, Tohoku, Japan. The authors extracted radiolarian fossils such as *Eucyrtidiellum cf. pyramis* from mudstone near the previously U–Pb dated sandstone and indicate that it is the Kimmeridgian as well as the depositional age of the sandstone. This radiolarian age proves that the accretionary complex in the peninsula is tectono-stratigraphically divided into the Late Jurassic and the Early Cretaceous units. They also illustrate the revised schematic compilation diagram of the ages and lithostratigraphic columns among the North Kitakami Belt.

Isotopic analyses for paleoceanic environmental study:

Bôle *et al.* (2020a, b) conducted isotopic analyses by respectively measuring $\delta^{30}\text{Si}$ and $\delta^{18}\text{O}$ of radiolarian tests for understanding the global silica cycle and paleoceanographic environment. Bôle *et al.* (2020a) measured $\delta^{30}\text{Si}$ of the Mesozoic radiolarian molds in the Inuyama section, central Japan by SIMS, indicating that the range of $\delta^{30}\text{Si}$ (-0.3 to 2 ‰) is consistent with that of modern and the Cenozoic ones, and that the 10-Myr scale trend of $\delta^{30}\text{Si}$ of the Mesozoic radiolarian molds from 250 Ma to 180 Ma is overall out-of-phase relation with biogenic silica (BSi) burial flux. This relation contradicts with the interpretation of $\delta^{30}\text{Si}$ as a productivity proxy. Bôle *et al.* (2020b) also measured $\delta^{18}\text{O}$ of Mesozoic radiolarian molds from Japan, Italy, Switzerland and Romania by SIMS. The result shows that the range from 19.8 to 35.8 ‰ is consistent with that of modern and the Cenozoic radiolarian tests from deep-sea cores of the equatorial Pacific. A slightly positive excursion during the Early–Middle Triassic, a high plateau in the Late Triassic, a negative excursion in the Early Jurassic, a slightly positive excursion in the Middle Jurassic and a few low values for the Cretaceous are recognized, although the Early Jurassic negative excursion is not consistent with $\delta^{18}\text{O}$ trend of less-diagenetic low–Mg calcite shells in shallow marine Tethys. This phenomenon implies a potential preservation of an environmental component even after the diagenesis of biogenic silica.

Bibliographic lists related to radiolarian studies by GSJ:

Ito *et al.* (2020) made a large effort to compile previous radiolarian-related publications by the Geological Survey of Japan, including geological maps, bulletins, cruise reports and newsletters with bibliographic lists from 1950 to 2019. The compilation effort aims to provide bibliographic lists related to radiolarians for future reference.

Since 1980s, the Geological Survey of Japan has been supporting and leading the radiolarian studies in Japan as a national center of geological research (see Ito *et al.*, 2020), and the consequent editing and publishing of this special issue under co-hosting InerRad XV are another form of the supporting. We believe that the articles in the issue will contribute to and profoundly affect future radiolarian researches not only by themselves but also in collaborating with other fields of geological sciences.

Acknowledgements: We express our appreciation to all the authors for their contribution to the current advances in radiolarian studies on biostratigraphy, paleoceanology and tectonic interpretation, and also acknowledge the great efforts contributed by all of the reviewers in providing valuable comments and suggestions, which have improved the quality of the articles in this special issue. We are grateful to the Editor-in-Chief, Dr. SUZUKI Atsushi and the editorial board members of the Bulletin, who support us during the editing process. In addition to the above, as both a member of the organizing committee and the guest editors, we specially thank the Geological Survey of Japan for co-hosting InterRad XV.

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Guest Editors:
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Radiolarian-inspired art design: Simplification and identification

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The makeup of a certain organism has been applied to human design activities, including architecture and art. Radiolaria, a type of holoplanktonic protozoa, contain siliceous shells that develop into various forms. Several artists have become interested in geometrically complex structures of radiolarians, and have created works of the art based on radiolarians (e.g. Hart, 1998, 2000; Morgante, 2017; Vones, 2018). These artificial expressions are diverse, ranging from the realistic to the abstract, and have been applied to several materials, such as the simplified image (Fig. 1A), silver model (Fig. 1B) and bead model (Figs. 1C, D). De Wever (2016) and Jungck *et al.* (2019) introduced several reproductions and architectural designs that were inspired by radiolarians. Nagai and Shiraki (2017a, b) reported on hand-sized realistic radiolarian models as an educational tool, which were either made in Europe or the United States from the late 19th century.

When an abstraction inspired by real organisms is created, the original forms are often simplified. Some scientific information is therefore lost. However, the simplified images of radiolarians illustrated by Moria (Fig. 1A)

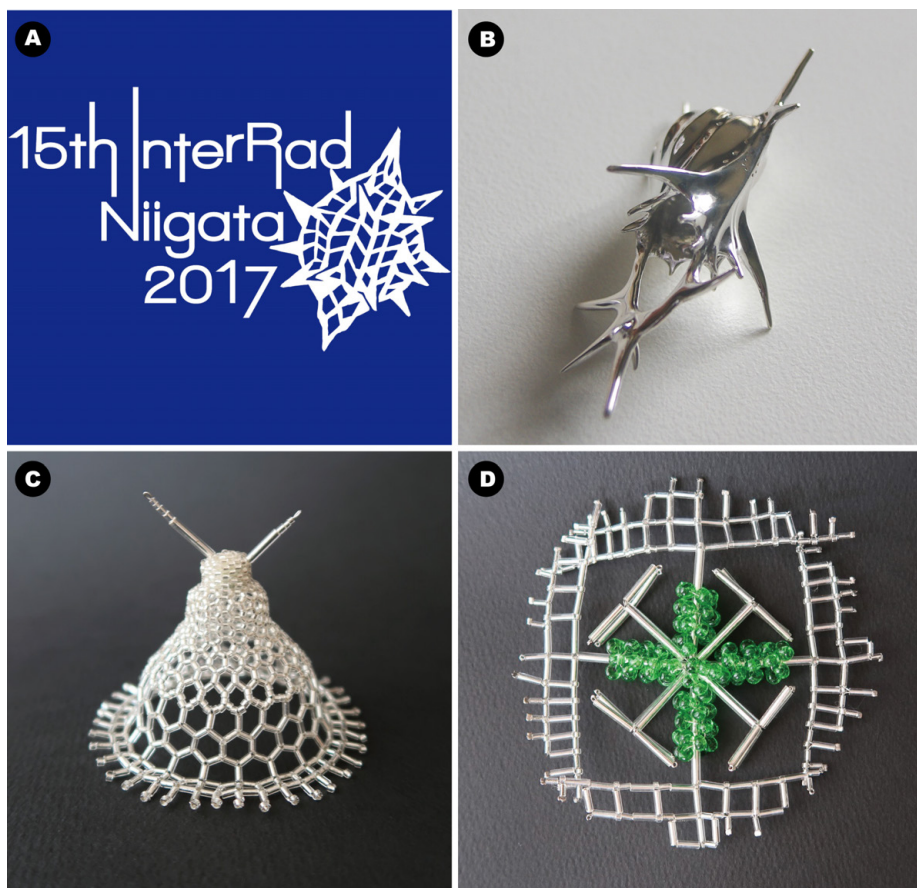


Fig. 1 Radiolarian-inspired artwork. A: Logo for InterRad XV in Niigata 2017 including simplified image of *Unuma echinatus* Ichikawa and Yao (created by Moria). B: Silver model of *Holoeciscus renzae* Schwartzapfel and Holdsworth, Devonian radiolaria (created by Yokoyama H.). C: Bead model of *Cycladophora pliocenica* (Hays), Neogene radiolaria (created by Ishiwata S.). D: Bead model of *Lithoptera muelleri* Haeckel, recent radiolaria (created by Ishiwata S.).

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have retained this scientific information, which is identifiable at the species level. In other words, Moria's simplified images were drawn through science-based selective simplification. Meanwhile, the radiolarian silver models created by Yokoyama, H. (Fig. 1B) were precisely-reproduced to retain as much of the original scientific information as possible. The bead models created by S. Ishiwata (Figs. 1C, D) were modified from original species under the limit of the materials, i.e. the models are composed of rod-shaped beads. Green beads of the model (Fig. 1D) expressed symbiotic algae of *Lithoptera muelleri* Haeckel.

Indeed, selective simplification is also important when conducting scientific activities (e.g. creating accurate sketches and schematic models). As such, this study discusses case examples of art designs that were inspired by both organisms and science-based simplification. Here, we introduce simplified images and precisely reproduced silver models involving two radiolarian species (i.e. *Unuma echinatus* and *Neoalbaillella pseudogrypus*).

Unuma echinatus Ichikawa and Yao

The Middle Jurassic *Unuma echinatus* Ichikawa and Yao was a symbolic radiolaria at InterRad XV in Niigata 2017 (the 15th Meeting of the International Association of Radiolarists). That is, the InterRad XV logo implemented this species into its design (Fig. 1A), which was used in the meeting's publication materials (e.g. Matsuoka and Ito, 2017; Ito *et al.*, 2017).

Ichikawa and Yao (1976) described this species as *Unuma (Spinunuma) echinatus* (Fig. 2A). Subsequent studies have generally not used the subgenus *Spinunuma*. Thus, the subgeneric diagnosis is currently an essential descriptor for this species. The diagnosis is as follows: "*Unuma* with well-developed apical horn, numerous stout

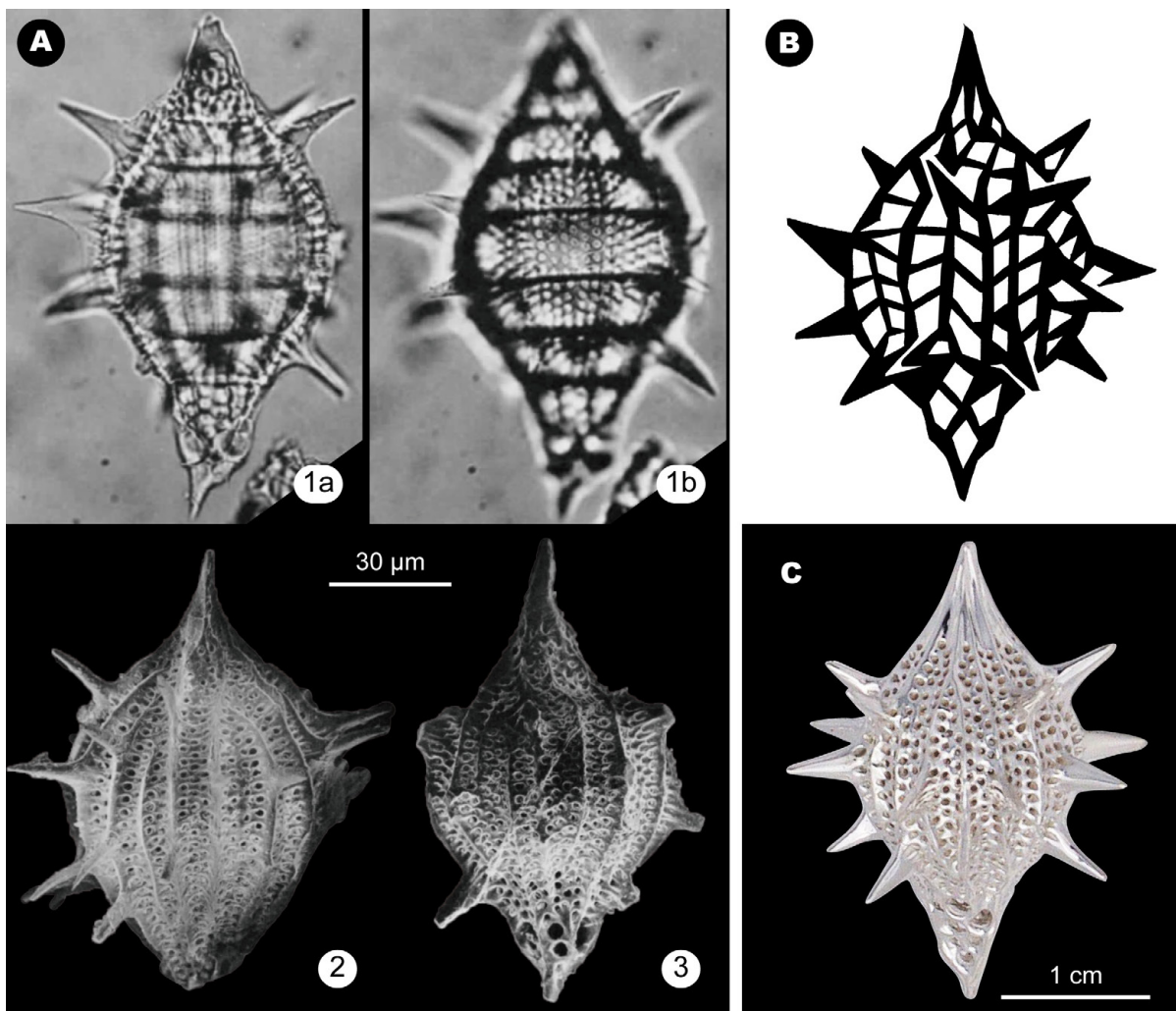


Fig. 2 *Unuma echinatus* Ichikawa and Yao. A: Reprinted type specimens from Ichikawa and Yao (1976). (1) Transmitted photomicrographs of holotype. (2) Scanning electron microscopy (SEM) image of paratype. (3) SEM image of paratype. B: Simplified image illustrated by Moria. C: Silver model created by H. Yokoyama.

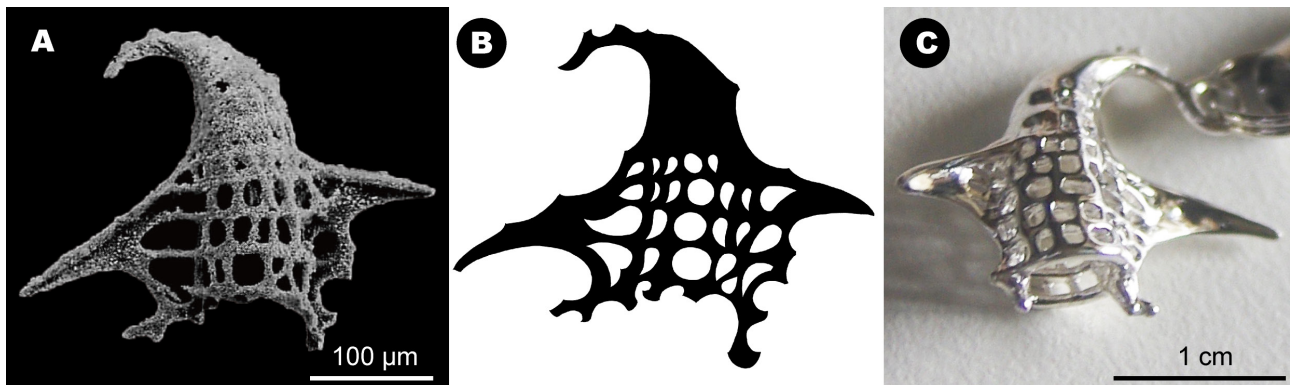


Fig. 3 *Neobaillella pseudogrypus* Sashida and Tonishi. A: Reprinted type specimen (SEM image of holotype) from Sashida and Tonishi (1988). B: Simplified image illustrated by Moria. C: Silver model created by H. Yokoyama.

radial spines, and distinct basal spine.” The silver model reproduced the diagnosis (Fig. 2C). The simplified image of *U. echinatus* (Fig. 2B) reflects this diagnosis as well, i.e. it contains a well-developed apical horn, numerous stout radial spines and a distinct basal spine. However, some points have been modified to differ from the original characteristics. For example, the surface pores of these specimens are small and circular (Figs. 2A, 2C), while those of the simplified image are large and polygonal (Fig. 2B).

***Neobaillella pseudogrypus* Sashida and Tonishi**

The diagnosis of the late Permian radiolaria *Neobaillella pseudogrypus* Sashida and Tonishi is “*Neobaillella* containing a bilaterally symmetrical shell with strongly curved apical cone and cylindrical pseudoabdomen having 3 to 4 horizontal rows of large square to rectangular windows” (Fig. 3A) (Sashida and Tonishi, 1988). The silver model reproduced the diagnosis as well (Fig. 3C). There are a few similar species, including *Neobaillella grypus* Ishiga, Kito and Imoto. However, *N. grypus* has a long pseudoabdomen (Ishiga *et al.*, 1982).

The simplified image of *N. pseudogrypus* (Fig. 3B) possesses the above characteristics (i.e. a strongly curved apical cone and cylindrical pseudoabdomen having 3 to 4 horizontal rows of large square to rectangular windows). However, the surface pores of these specimens and the silver model are grid-like (Figs. 3A, 3C), while those of the simplified images are circular and teardrop-shaped (Fig. 3B).

As is the case in *U. echinatus*, the shape of the surface pores of *N. pseudogrypus* differs between the original specimens and the simplified image. The simplified image of *U. echinatus* has polygonal pores (Fig. 2B) although the original specimens possess circular ones (Fig. 2A), i.e. the surface pores of the simplified image are more angular than those of the original specimens. Contrastively, the simplified image of *N. pseudogrypus* expressed more circular pores (Fig. 3B) compared to the original specimens (Fig. 3A).

Acknowledgments

This study’s radiolarian photomicrograph reprints were permitted for use by Prof. SASHIDA Katsuo (University of Tsukuba), Palaeontological Society of Japan and the Micropaleontology Press. The authors are most grateful to Dr. UCHINO Takayuki (Geological Survey of Japan, AIST) for the careful review and constructive comments on the manuscript.

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Early Oxfordian radiolarians from the ammonite-bearing Fludergraben section (Northern Calcareous Alps, Austria)

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SUZUKI Hisashi and GAWLICK Hans-Jürgen (2020) Early Oxfordian radiolarians from the ammonite-bearing Fludergraben section (Northern Calcareous Alps, Austria). *Bulletin of the Geological Survey of Japan*, vol. 71(4), p. 243–280, 8 figs, 3 plates, 2 appendices.

Abstract: A well-preserved and relatively rich radiolarian fauna is described from red to grey bedded radiolarites of the Fludergraben section in the Northern Calcareous Alps, Austria. These radiolarites were deposited just above the Klaus Formation, dated by ammonites as latest Callovian or the Callovian/Oxfordian boundary. The radiolarian fauna is therefore of an early Oxfordian age undoubtedly. Among long-lasting radiolarian species coming from the Callovian, we can distinguish some species that appeared in early Oxfordian time: *Kilinora spiralis*, *Fultacapsa sphaerica*, *Protunuma japonicus*, *Pseudoeuycyrtis reticularis*. We discuss these results in the light of existing radiolarian zonation for the middle Callovian to Oxfordian, and redefined the *Williriedellum dierschei* Zone (lower-middle Oxfordian), which was previously ranked as subzone in the *Zhamoidellum ovum* Zone, on the basis of the new index species. These new findings fill a gap in the definition of the Oxfordian by radiolarians and result in a better resolution of the radiolarian biostratigraphy.

In the chapter of systematic part, we describe 37 genera, 67 species and 2 subspecies including diagnosis emendations of 2 genera (*Loopus* and *Pseudodictyomitra*) and 1 species (*Protunuma japonicus*). The type species of the genus *Loopus* is examined and redesignated.

Keywords: Western Tethys, biostratigraphy, radiolarians, Oxfordian, Fludergraben section, Northern Calcareous Alps

1. Introduction

The existing Middle to Late Jurassic radiolarian zonation (e.g. Pessagno *et al.*, 1993 for western North America; Matsuoka, 1995 for Japan and western Circum-Pacific region; Baumgartner *et al.*, 1995b; Beccaro, 2004, 2006; Suzuki and Gawlick, 2003a for Tethyan and central Atlantic regions) have been controversially discussed and several attempts were made to refine the stratigraphic ranges of radiolarian taxa (O'Dogherty *et al.*, 2011, 2017). However, until today most radiolarian workers dealing with the Tethyan/Atlantic region have still used in general the Unitary Association Zonation of Baumgartner *et al.* (1995b) without or with only moderate modifications of the age ranges of several radiolarian species. The biostratigraphic resolution of Middle to Late Jurassic radiolarians is not high and the existing biostratigraphic radiolarian zones exhibit relatively long-time duration. A main problem for a stable and precise radiolarian

zonation with a much better biostratigraphic resolution is the worldwide scarcity of radiolaria-bearing sedimentary rocks in sections, where radiolarian associations can be correlated with other organisms, especially ammonoids.

In the Western Tethyan realm, and also in the Northern Calcareous Alps, radiolarian assemblages of the Callovian–Oxfordian contain species with relatively long biostratigraphic age ranges. Therefore, in most cases it cannot be decided, if a radiolarian assemblage is of Callovian or Oxfordian age, by use of the present radiolarian zonation.

Radiolarian species, which mark the beginning of the Oxfordian, are practically not known, because no successions, where radiolarian associations can be correlated with uppermost Callovian/lowermost Oxfordian ammonoids, have been worldwide known. In the radiolarian biozonation by Baumgartner *et al.* (1995b) the time span from middle Callovian to early Oxfordian is united in one radiolarian zone as the Unitary

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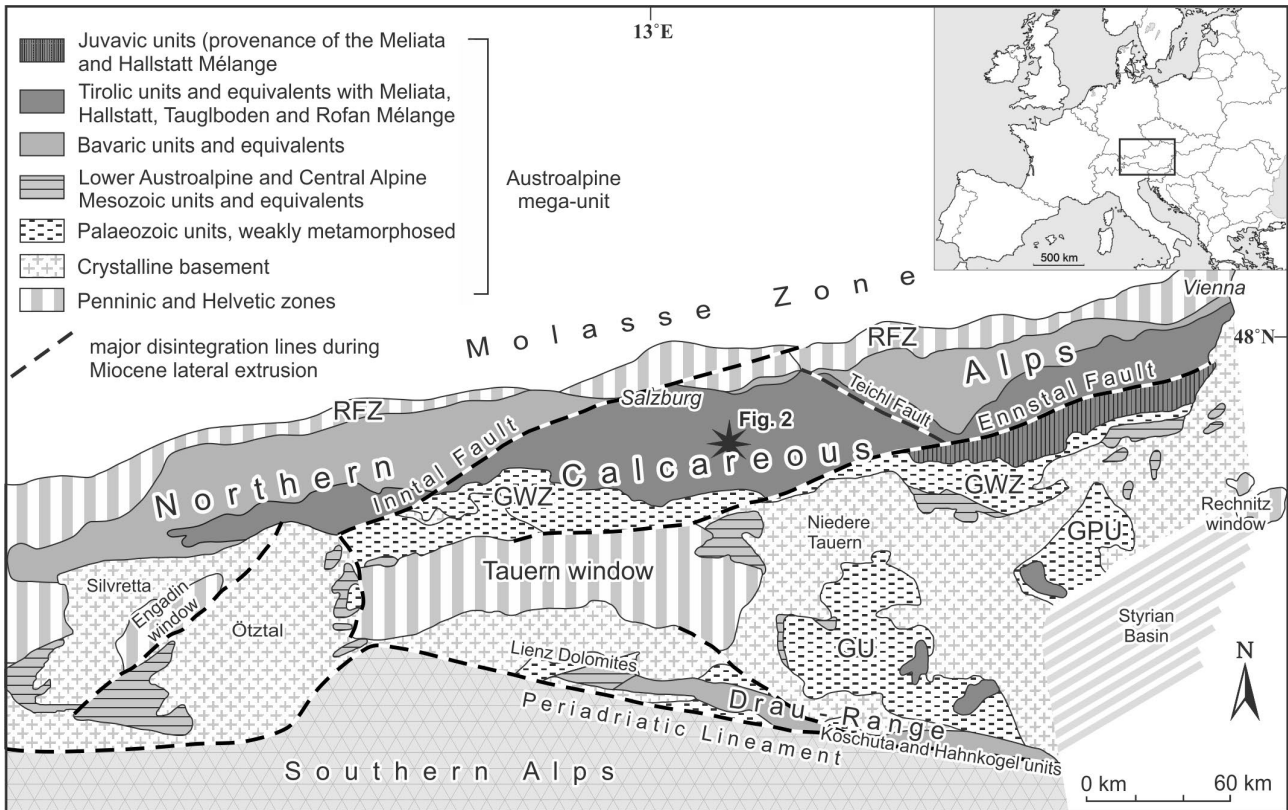


Fig. 1 Schematic tectonic map of the Eastern Alps (Tollmann, 1977; Frisch and Gawlick, 2003) and geographic position of the study area in the Northern Calcareous Alps. GPU: Graz Palaeozoic Unit, GU: Gurktal Unit, GWZ: Greywacke Zone, RFZ: Rhenodanubian Flysch Zone.

Association Zone 8. A more precise radiolarian zonation for the time around the Callovian/Oxfordian boundary is therefore highly needed. We analysed well-preserved Oxfordian radiolarian faunas from the base of a 900 m thick radiolarite succession (Gawlick *et al.*, 2007) in the Northern Calcareous Alps, i.e. the Fludergraben section near Altaussee, Austria (Figs. 1, 2). In the lowermost part of the section, red nodular limestones of the Klaus Formation were formed in the Middle Jurassic to the latest Callovian or to the Callovian/Oxfordian boundary, as proven by the following ammonites (Mandl, 1982): *Euaspidoceras* sp., *Holcophylloceras zignodianum* and fragments of *?Nebroditis* sp. Therefore, the radiolarite succession of the Fludergraben section provides the best opportunity to search for early Oxfordian marker of radiolarian species. Beside this, the age range of several radiolarian species occurring in these radiolarites must be prolonged, if they are so far known only from lower levels than the Oxfordian. In this paper we present the early Oxfordian radiolarian fauna, which helps to refine the radiolarian zonation for the Callovian and Oxfordian.

2. Geologic setting

The studied Fludergraben section is located in the Fludergraben valley in the central Northern Calcareous

Alps, southeast of Salzburg (Figs. 1, 2). The section belongs to the lowermost part of the Tauglboden Formation that overlies the Klaus Formation (Fig. 3). The Klaus Formation consists of red nodular limestone yielding ammonites of the latest Callovian to the Callovian/Oxfordian boundary. The Oxfordian to Tithonian Tauglboden Formation consists of up to 900 m thick grey to black siliceous to radiolaritic rocks (radiolarite) with intercalated simultaneous mass transport deposits (Gawlick and Frisch, 2003; Gawlick *et al.*, 2009). The base of the Tauglboden Formation starts with a red radiolarite followed by a grey to black radiolarite. The basal red radiolarite is up to 3 m thick and this part is distinguished from the main part of the Tauglboden Formation as the Fludergraben Member (Fig. 3a; Gawlick *et al.*, 2009). The sedimentary succession of the Tauglboden Formation was deposited in a trench-like foreland basin (Tauglboden Basin: Diersche, 1980) in front of a propagating nappe stack formed in Oxfordian time (Fig. 3b; Missoni and Gawlick, 2011; Gawlick and Missoni, 2019 and references therein). During the Middle to early Late Jurassic, the former passive continental margin of the Neo-Tethys attained a lower plate position due to ongoing ophiolite obduction. In the course of the ongoing ophiolite obduction, the former (Triassic–Middle Jurassic) outer passive margin became imbricated and a thin-skinned orogen was formed. In front of the

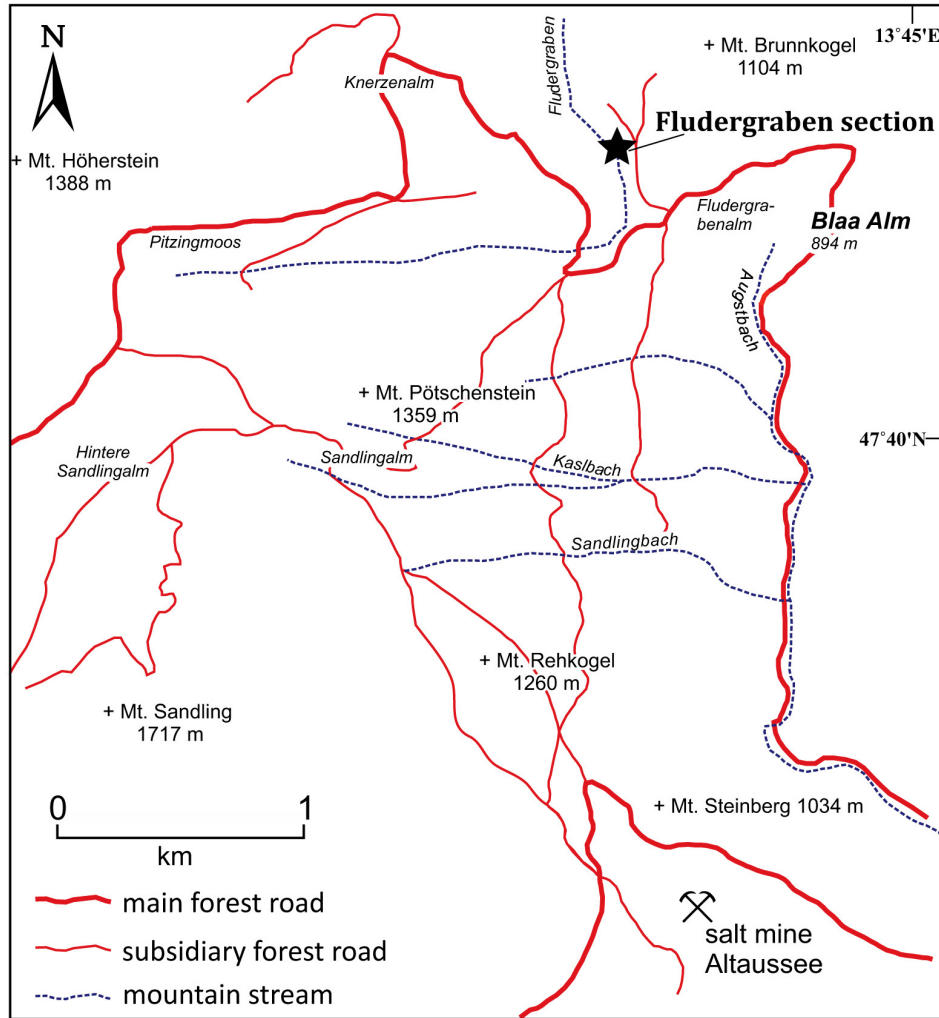


Fig. 2 Geographic position of the Fludergraben section (indicated by a star symbol) in the area of salt mine Altaussee–Mt. Sandling–Mt. Höherstein–Blaa Alm.

northwestward propagation thrust belt (nappe stack), the deep-water trench-like foreland basins were formed and incorporated into the thrust belt. In the early Oxfordian, the thrust belt reached the area of the Tauglboden Basin. Rapid deepening resulted first in the shift from carbonate to radiolarite deposition and later in deposition of mass transport deposits with its source in the adjacent nappe front (Trattberg Rise: Fig. 3b) (Gawlick and Missoni, 2019). A well-preserved section of the Tauglboden Formation is located in the Salzkammergut area, east of Salzburg.

3. Studied section and samples

The Fludergraben section in the Fludergraben valley (Fig. 2) consists of radiolarite, i.e. siliceous sedimentary rocks consisting of radiolarians. Radiolarite deposition of the Fludergraben section started almost instantaneously from the red nodular limestone containing ammonites (Fig. 4). The ammonite-bearing horizon of the uppermost Klaus

Formation is only 10 cm below occurrence of the first radiolarite bed. A short-lasting stratigraphic gap on top of the ammonite-bearing layer cannot be excluded because of the bad preservation of the ammonites without their original shells. This indicates that there was an enough time to solve ammonite shells. However, because a serious hardground is not detectable, long-lasting subsolution can be excluded.

The lowermost bed of the radiolarite sequence is originally a *Bositra*-radiolarian-bearing siliceous limestone (Fig. 5a), later completely silicified (sample D1051). The following red radiolarite is well-bedded. The thickness of each bed is 3–10 cm, in some cases intercalated by up to 5 mm-thick reddish siliceous claystones (Diersche, 1980). The radiolarite is completely silicified, but the preservation of the radiolarians is in cases rather good. The microfacies show bioturbated radiolarian wackestones to packstones (Fig. 5b, 5c). All radiolarite beds of up to 10 cm thickness are massive and without sedimentary lamination, as well visible in the higher part of the Tauglboden Formation

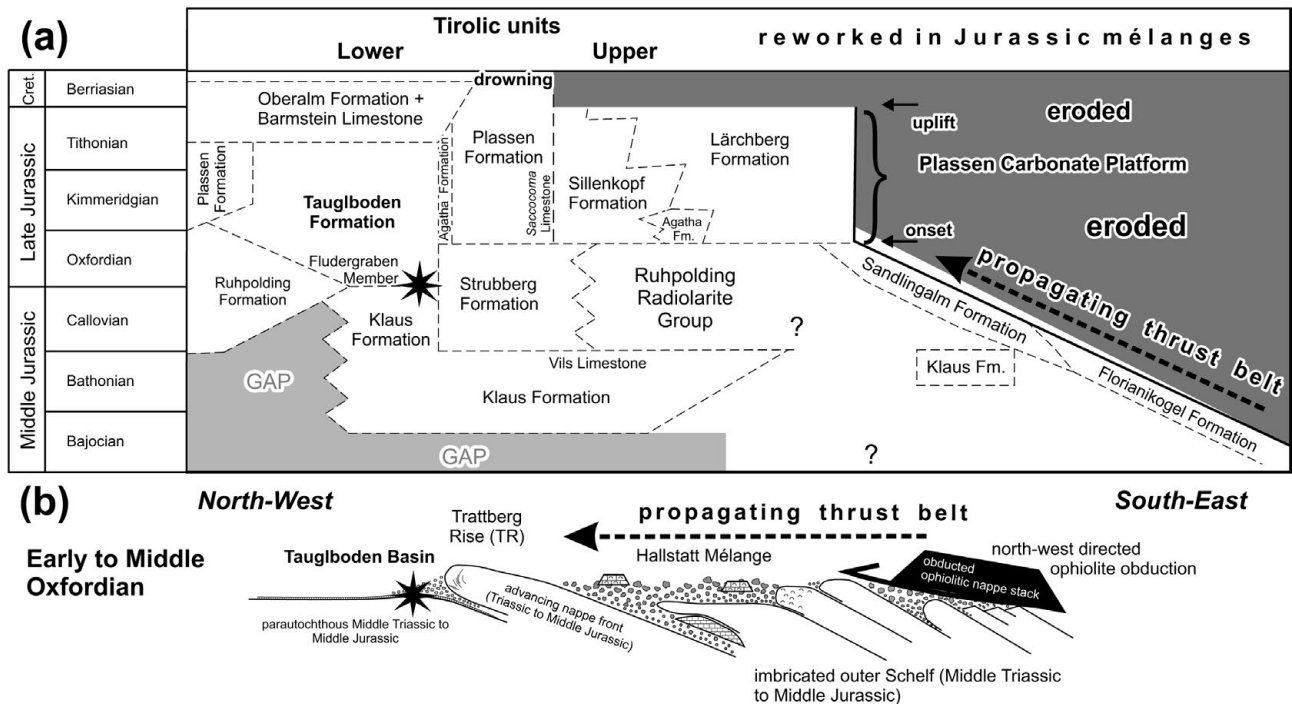


Fig. 3 (a) Simplified Middle to Late Jurassic stratigraphic table of the central Northern Calcareous Alps with an overview of common formation names after Gawlick *et al.* (2009) and stratigraphic and palaeotectonic position of the studied Fludergraben section (indicated by a star symbol). Cret.: Cretaceous, Fm.: Formation. (b) Early to Middle Oxfordian geodynamic reconstruction of the Northern Calcareous Alps according to Missoni and Gawlick (2011) and Gawlick and Missoni (2019). Due to ophiolite obduction since Middle Jurassic time the former northwestern passive continental margin attained a lower plate position and a thin-skinned orogen was formed. The Tauglboden Basin was generated in front of the propagating thrust belt (indicated by a star symbol).

(Gawlick *et al.*, 2012).

From the red radiolarites of the Fludergraben section, six radiolaria-bearing samples were collected in the first one metre just above the red condensed limestones with the ammonite horizon. The six samples are in ascending order as follows (Fig. 4): D1051, D1023, D1024, D1052, EW146, D1025.

4. Radiolarian fauna of the Fludergraben section

We have detected radiolarian species in all six samples with the methods of diluted hydrofluoric acid for decomposition and of hydrogen peroxide for residue cleaning. Their preservation is in some cases very poor, but also moderate to well-preserved radiolarians could be isolated. The radiolarian assemblages from all six samples are listed here, and are depicted in Plates 1–3.

D1051: *Archaeodictyomitra apiarium* (Rüst, 1885), *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Striatojaponocapsa* sp.

D1023: *Acanthocircus* cf. *suboblongus* (Yao, 1972), *Archaeospongoprimum* cf. *elegans* Wu, 1993, *Tritribs* cf. *exotica* (Pessagno, 1977a), *Archaeodictyomitra apiarium* (Rüst, 1885), *Archaeodictyomitra mirabilis* Aita, 1987, *Archaeodictyomitra rigida* Pessagno, 1977a, *Cinguloturris carpatica* Dumitrica, 1982, *Eucyrtidiellum*

circumperforatum Chiari, Marcucci and Prela, 2002, *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974), *Fultacapsa sphaerica* (Ozoldova, 1988), *Gongylothorax favosus favosus* Dumitrica, 1970, *Helvetocapsa matsukoi* (Sashida, 1999), *Hsuum brevicostatum* (Ozoldova, 1975), *Hsuum maxwelli* Pessagno, 1977a, *Loopus doliolum* Dumitrica, 1997, *Neorelumbra skenderbegi* Chiari, Marcucci and Prela, 2002, *Parahsuum* sp. S sensu Matsuoka, 1986, *Protunuma japonicus* Matsuoka and Yao, 1985, *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985, *Stichocapsa cicciana* Chiari, Marcucci and Prela, 2002, *Stichocapsa robusta* Matsuoka, 1984, *Stichomitra annibill* Kocher, 1981, *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006, *Kilinora* cf. *spiralis* (Matsuoka, 1982), *Tricolocapsa tetragona* Matsuoka, 1983, *Tricolocapsa undulata* (Heitzer, 1930), *Takemuraella hexagonata* (Heitzer, 1930), *Takemuraella hungarica* (Kozur, 1985), *Unuma gordus* Hull, 1997, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Zhamoidellum ovum* Dumitrica, 1970, *Zhamoidellum ventricosum* Dumitrica, 1970.

D1024: *Archaeospongoprimum* cf. *elegans* Wu, 1993, *Cinguloturris carpatica* Dumitrica, 1982, *Cyrtocapsa* sp. B, *Dictyomitrella kamoensis* Mizutani and Kido, 1983, *Eucyrtidiellum nodosum* Wakita, 1988, *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974), *Eucyrtidiellum*

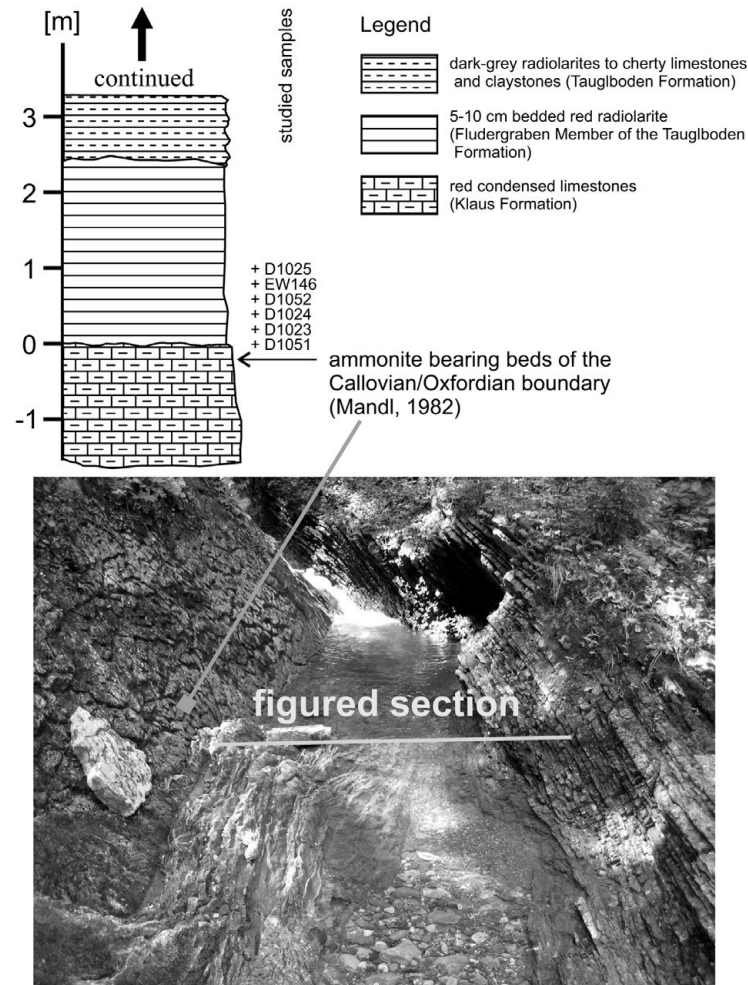


Fig. 4 Columnar section and photo show the lowermost part of the Fludergraben Member and position of studied samples.

unumaense (Yao, 1979), *Gongylothorax favosus favosus* Dumitrica, 1970, *Hsuum maxwelli* Pessagno, 1977a, *Loopus doliolum* Dumitrica, 1997, *Parahsuum* sp. S sensu Matsuoka, 1986, *Protunuma lanosus* Ozvoldova, 1996, *Striatojaponocapsa conexa* (Matsuoka, 1983), *Striatojaponocapsa riri* O'Dogherty *et al.*, 2006, *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006, *Tricolocapsa tetragona* Matsuoka, 1983, *Unuma typicus* Ichikawa and Yao, 1976, *Williriedellum crystallinum* Dumitrica, 1970, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Williriedellum marcucciae* Cortese, 1993, *Zhamoidellum ovum* Dumitrica, 1970.

D1052: *Trirabs exotica* (Pessagno, 1977a), *Cinguloturris carpatica* Dumitrica, 1982, *Dictyomitrella kamoensis* Mizutani and Kido, 1983, *Eucyrtidiellum circumperforatum* Chiari, Marcucci and Prela, 2002, *Eucyrtidiellum nodosum* Wakita, 1988, *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974), *Gongylothorax favosus oviformis* Suzuki and Gawlick, 2009, *Hsuum brevicostatum* (Ozvoldova, 1975), *Hsuum maxwelli* Pessagno, 1977a, *Loopus doliolum* Dumitrica, 1997, *Parahsuum* sp. S

sensu Matsuoka, 1986, *Podobursa nodosa* (Chiari, Marcucci and Prela, 2002), *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985, *Pseudoeucyrtis reticularis* Matsuoka and Yao, 1985, *Ristola altissima* (Rüst, 1885), *Stichocapsa robusta* Matsuoka, 1984, *Stichomitra annibill* Kocher, 1981, *Stichomitra* sp. A sensu Baumgartner *et al.*, 1995a, *Striatojaponocapsa conexa* (Matsuoka, 1983), *Striatojaponocapsa naradaniensis* (Matsuoka, 1984), *Striatojaponocapsa riri* O'Dogherty *et al.*, 2006, *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006, *Tetracapsa* sp. A sensu Suzuki and Gawlick, 2003b, *Tricolocapsa undulata* (Heitzer, 1930), *Unuma gordus* Hull, 1997, *Williriedellum carpathicum* Dumitrica, 1970, *Williriedellum crystallinum* Dumitrica, 1970, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Williriedellum marcucciae* Cortese, 1993, *Williriedellum sujkowski* Widz and De Wever, 1993, *Zhamoidellum ovum* Dumitrica, 1970.

EW146: *Archaeospongoprimum* cf. *imlayi* Pessagno, 1977a, *Archaeodictyomitra minoensis* (Mizutani, 1981), *Cinguloturris carpatica* Dumitrica, 1982, *Cinguloturris*

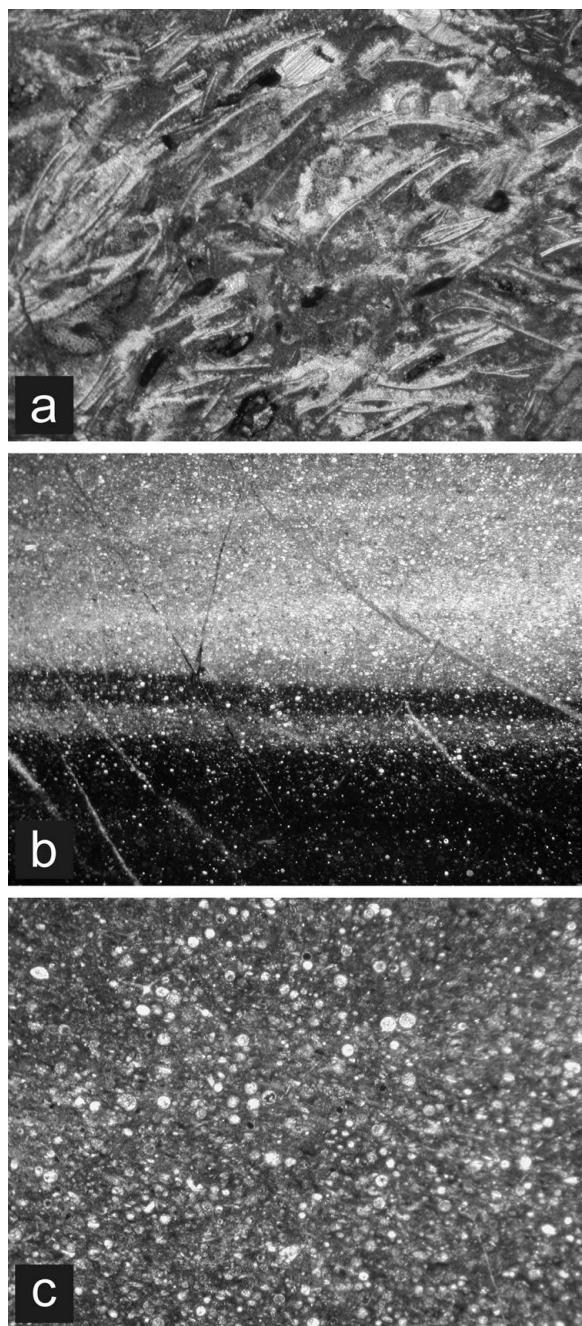


Fig. 5 Microfacies of the slightly siliceous red *Bositra*-bearing nodular limestone and the overlying red radiolarite of the Fludergraben section. (a) *Bositra* shells together with some crinoids and reworked hardground clasts. Width of the photo 0.5 cm. Sample D1051. (b) Red radiolarite above the red nodular limestone. Layered grey-red to red radiolarian wackestone to radiolarian packstone. In the basal radiolarian wackestone the radiolarians are well-preserved, in the upper radiolarian packstone the preservation of the radiolarians is moderate due to intense silification. Width of the photo 1.4 cm. Sample D1052. (c) Magnification of (b), upper part. The most radiolarians in this bioturbated red radiolarite are recrystallized and only some radiolarians are well-preserved. Width of the photo 0.5 cm.

primorika Kemkin and Taketani, 2004, *Dictyomitrella* cf. *kamoensis* Mizutani and Kido, 1983, *Eucyrtidiellum nodosum* Wakita, 1988, *Eucyrtidiellum* cf. *unumaense* (Yao, 1979), *Hsuum baloghi* Grill and Kozur, 1986, *Hsuum* cf. *brevicostatum* (Ozoldova, 1975), *Hsuum maxwelli* Pessagno, 1977a, *Parvicingula spinata* Vinassa, 1899, *Tricolocapsa undulata* (Heitzer, 1930), *Takemuraella hungarica* (Kozur, 1985), *Williriedellum carpathicum* Dumitrica, 1970.

D1025: *Archaeospongoprunum* cf. *elegans* Wu, 1993, *Archaeodictyomitra sixi* Yang, 1993, *Spongotropus* sp. D sensu Suzuki and Gawlick, 2003b, *Archaeodictyomitra mirabilis* Aita, 1987, *Archaeodictyomitra patricki*, Kocher, 1981, *Cinguloturris carpatica* Dumitrica, 1982, *Droltus galerus* Suzuki, 1995b, *Eucyrtidiellum nodosum* Wakita, 1988, *Eucyrtidiellum unumaense* (Yao, 1979), *Gongylothorax favosus favosus* Dumitrica, 1970, *Gongylothorax favosus oviformis* Suzuki and Gawlick, 2009, *Gongylothorax* sp. C sensu Suzuki and Gawlick, 2003b, *Helvetocapsa matsukoi* (Sashida, 1999), *Japonocapsa fusiformis* (Yao, 1979), *Praewilliriedellum* aff. *spinosum* Kozur, 1984, *Protunuma fusiformis* Ichikawa and Yao, 1976, *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985, *Saitoum pagei* Pessagno, 1977a, *Stichocapsa robusta* Matsuoka, 1984, *Japonocapsa tegiminis* (Yao, 1979), *Stichomitra annibill* Kocher, 1981, *Striatojaponocapsa naradaniensis* (Matsuoka, 1984), *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006, *Stylocapsa oblongula* Kocher, 1981, *Tetracapsa* sp. A sensu Suzuki and Gawlick, 2003b, *Tricolocapsa undulata* (Heitzer, 1930), *Takemuraella hexagonata* (Heitzer, 1930), *Unuma typicus* Ichikawa and Yao, 1976, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Williriedellum marcucciae* Cortese, 1993, *Williriedellum* sp. C sensu Gawlick *et al.*, 2018, *Zhamoidellum ovum* Dumitrica, 1970.

5. Systematic part

We describe radiolarian species from the Fludergraben section systematically. Radiolarian taxonomic classification shown here is in principle based on Takemura (1986), Suzuki *et al.* (2002), Suzuki and Gawlick (2003b) and Suzuki and Gawlick (2009). The familial classification of Nassellaria of these publications considers the cephalic skeletal elements which construct the fundamental structure of nassellarians (e.g. Takemura, 1986).

In the synonym lists, we use following mark and abbreviations. Astarisk: first description of taxon name, aff.: affinis, cf.: confer, non: not, pt.: partial.

Subclass **RADIOLARIA** Müller, 1858

Order **POLYCYSTIDA** Ehrenberg, 1839; emend. Riedel, 1967b

Suborder **ENTACTINARIA** Kozur and Mostler, 1982

Family SATURNALIDAE Deflandre, 1953

Genus *Acanthocircus* Squinabol, 1903; emend. Donofrio and Mostler, 1978

Type species: *Acanthocircus irregularis* Squinabol, 1903 (Campbell, 1954)

Acanthocircus cf. *suboblongus* (Yao, 1972) (Plate 1, fig. 14)

cf. *1972 *Spongosaturnalis? suboblongus* – Yao, p. 29, pl. 3, figs. 1–6, pl. 10, figs. 3a–3c.

Remarks: Only one part of the ring of this species preserved, so that we identify here with “cf.”

Suborder SPUMELLARIA Ehrenberg, 1876

Family SPONGULIDAE Haeckel, 1862

Genus *Archaeospongoprimum* Pessagno, 1973; emend. Kozur and Mostler, 1981

Type species: *Archaeospongoprimum venadoensis* Pessagno, 1973

Archaeospongoprimum cf. *elegans*, Wu, 1993 (Plate 1, fig. 3; Plate 2, fig. 1; Plate 3, fig. 21)

cf. 1930 *Ellipsoxiphus asper* Rüst – Heitzer, p. 389, pl. 27, fig. 17.

cf.*1993 *Archaeospongoprimum elegans* – Wu, p. 118, pl. 1, figs. 5, 7, 23.

Archaeospongoprimum cf. *imlayi* Pessagno, 1977a (Plate 3, fig. 8)

cf.*1977a *Archaeospongoprimum imlayi* – Pessagno, p. 73, pl. 3, figs. 2–4; ? pl. 3, fig. 1.

cf. 2003b *Archaeospongoprimum imlayi* Pessagno – Suzuki and Gawlick, p. 171, fig. 5.6; fig. 6.9. (detailed synonymy until 2003)

Genus *Spongotripus* Haeckel, 1881

Type species: *Spongotripus pauper* Rüst, 1888 (Kießling, 1999)

Spongotripus sp. D sensu Suzuki and Gawlick, 2003b (Plate 3, fig. 22)

*2003b *Spongotripus* sp. D – Suzuki and Gawlick, p. 172, fig. 5.7.

2018 *Spongotripus* sp. D sensu Suzuki and Gawlick – Gawlick *et al.*, fig. 18.29.

Family HAGIASTRIDAE Riedel, 1971; emend. Baumgartner, 1980

Genus *Tritrabs* Baumgartner, 1980

Type species: *Paronaella? casmaliaensis* Pessagno, 1977a

Tritrabs exotica Pessagno, 1977a (Plate 1, fig. 8; Plate 2, fig. 22)

*1977a *Paronaella? exotica* – Pessagno, p. 70, pl. 1, figs. 12, 13.

1980 *Tritrabs exotica* (Pessagno) – Baumgartner, p. 294, pl. 4, fig. 16.

1995a *Tritrabs exotica* (Pessagno) – Baumgartner *et al.*, p. 608, pl. 3119, figs. 1–3.

2006 *Tritrabs exotica* (Pessagno) – O’Dogherty *et al.*, p. 472, pl. 11, fig. 38.

2013 *Tritrabs exotica* (Pessagno) – Krische *et al.*, pl. 3, fig. 18.

Suborder NASSELLARIA Ehrenberg, 1876

Family POULPIDAE De Wever, 1981

Genus *Saitoum* Pessagno, 1977a

Type species: *Saitoum pagei* Pessagno, 1977a

Saitoum pagei Pessagno, 1977a (Plate 3, fig. 18)

*1977a *Saitoum pagei* – Pessagno, p. 98, pl. 12, figs. 11–14.

2003b *Saitoum pagei* Pessagno – Suzuki and Gawlick, p. 175, fig. 5.38.

2018 *Saitoum pagei* Pessagno – Gawlick *et al.*, fig. 12.18.

Family THEOPERIDAE Haeckel, 1881; emend. Takemura, 1986

Genus *Cinguloturris* Dumitrica, 1982

Type species: *Cinguloturris carpatica* Dumitrica, 1982

Cinguloturris carpatica Dumitrica, 1982 (Plate 1, fig. 15; Plate 2, fig. 3; Plate 3, figs. 6, 9)

*1982 *Cinguloturris carpatica* – Dumitrica in Dumitrica and Mello, p. 23, pl. 4, figs. 7–11.

1994 *Cinguloturris carpatica* Dumitrica – Ishida, fig. 3.2.

2003b *Cinguloturris carpatica* Dumitrica – Suzuki and Gawlick, p. 189, fig. 5.28; fig. 6.50. (detailed synonymy between 1994 and 2003)

2003 *Cinguloturris carpatica* Dumitrica – Wegerer *et al.*, fig. 7.13; fig. 11.5.

2006 *Cinguloturris carpatica* Dumitrica – Auer *et al.*, fig. 6.9.

2007 *Cinguloturris carpatica* Dumitrica – Auer *et al.*, fig. 6.14.

2009 *Cinguloturris carpatica* Dumitrica – Suzuki and Gawlick, p. 167, fig. 5.2; fig. 6.1A, 6.1B.

Remarks: *Cinguloturris carpatica* has tiny circular dents on the solid horizontal ridges of each post-thoracic segment.

- Cinguloturris primorika*** Kemkin and Taketani, 2004
(Plate 3, fig. 10)
2001 *Cinguloturris* cf. *cylindra* Kemkin and Rudenko – Missoni *et al.*, fig. 3.9.
*2004 *Cinguloturris primorika* – Kemkin and Taketani, p. 333, fig. 4.1–4.3.
2006 *Cinguloturris* cf. *cylindra* Kemkin and Rudenko – Gawlick *et al.*, fig. 8a.8.
2009 *Cinguloturris primorika* Kemkin and Taketani – Suzuki and Gawlick, p. 167, fig. 5.3A, 5.3B.
2011 *Cinguloturris primorika* Kemkin and Taketani – Gawlick *et al.*, fig. 3.11.
Remarks: *Cinguloturris primorika* has short costae- or node-like structures on the solid horizontal ridges of each post-thoracic segment, which are not arranged regularly.

Genus ***Parahsuum*** Yao, 1982

Type species: ***Parahsuum simplum*** Yao, 1982

- Parahsuum*** sp. S sensu Matsuoka, 1986
(Plate 1, fig. 7; Plate 2, figs. 2, 24)
*1986 *Parahsuum* sp. S – Matsuoka, pl. 2, fig. 13; pl. 3, fig. 14.
pt. 1995a *Parahsuum* sp. S – Baumgartner *et al.*, p. 384, pl. 3240, figs. 2, 4, 5; non pl. 3240, figs. 1, 3 [= *Parahsuum carpathicum* Widz and De Wever, 1993].
2003b *Parahsuum* sp. S sensu Matsuoka – Suzuki and Gawlick, p. 182, fig. 6.70. (detailed synonymy between 1994 and 2002)
2004 *Parahsuum?* sp. – Ishida, fig. 7.4.
2009 *Parahsuum* sp. S sensu Matsuoka – Suzuki and Gawlick, p. 167, fig. 5.5.
Remarks: *Parahsuum* sp. S has a short conical test and a slender, short apical horn.

Genus ***Hsuum*** Pessagno, 1977a

Type species: ***Hsuum cuestaensis*** Pessagno, 1977a

- Hsuum brevicostatum*** (Ozvodova, 1975)
(Plate 1, fig. 12; Plate 2, fig. 37)
*1975 *Lithostrobos brevicostatus* – Ozvodova, p. 84, pl. 102, fig. 1.
1994 *Transhsuum brevicostatum* (Ožvoldová) gr. – Goričan, p. 91, pl. 18, figs. 6–8. (detailed synonymy until 1993)
2003b *Hsuum brevicostatum* (Ozvodova) – Suzuki and Gawlick, p. 184; fig. 5.33; fig. 6.62. (detailed synonymy between 1994 and 2002)
2004 *Hsuum brevicostatum* (Ozvodova) – Gawlick *et al.*, fig. 3a.11.
2004 *Hsuum brevicostatum* (Ozvodova) – Ishida, fig. 7.2; fig. 8.8.
2005 *Hsuum brevicostatum* (Ozvodova) – Missoni *et al.*, fig. 10.16.
2006 *Hsuum brevicostatum* (Ozvodova) – Gawlick *et al.*,

fig. 8.18; fig. 9.15.

- 2009 *Hsuum brevicostatum* (Ozvodova) – Suzuki and Gawlick, p. 168, fig. 5.6.
2014 *Hsuum brevicostatum* (Ozvodova) – Suzuki *et al.*, p. 11, pl. 4, fig. 11.

Hsuum maxwelli Pessagno, 1977a

(Plate 1, fig. 4; Plate 2, fig. 15)

- *1977a *Hsuum maxwelli* – Pessagno, p. 81, pl. 7, figs. 14–16.
1994 *Transhsuum maxwelli* (Pessagno) gr. – Goričan, p. 92, pl. 18, figs. 1–4. (detailed synonymy until 1993)
2003b *Hsuum maxwelli* Pessagno – Suzuki and Gawlick, p. 183, fig. 5.32; fig. 6.64. (detailed synonymy between 1994 and 2002)
2004 *Hsuum maxwelli* Pessagno – Gawlick *et al.*, fig. 3b.26.
2004 *Hsuum maxwelli* Pessagno – Ishida, fig. 7.1; fig. 8.7.
2005 *Hsuum maxwelli* Pessagno – Missoni *et al.*, fig. 7.11; fig. 13.3.
2006 *Hsuum maxwelli* Pessagno – Gawlick *et al.*, fig. 8b.19; fig. 9a.16.
2009 *Hsuum maxwelli* Pessagno – Suzuki and Gawlick, p. 168, fig. 5.7.
2018 *Hsuum maxwelli* Pessagno – Gawlick *et al.*, fig. 12.11; fig. 18.11.

Hsuum baloghi Grill and Kozur, 1986

(Plate 3, fig. 7)

- *1986 *Hsuum baloghi* – Grill and Kozur, p. 254, pl. 3, figs. 3–6.
2003b *Hsuum baloghi* Grill and Kozur – Suzuki and Gawlick, p. 182, fig. 5.31.
Remarks: *Hsuum baloghi* has weakly developed longitudinal costae on the post-abdominal segments. In case of *Hsuum maxwelli*, longitudinal costae are strongly developed.

Genus ***Dictyomitrella*** Haeckel, 1887

Type species: ***Eucyrtidium articulatum*** Ehrenberg, 1876 (Campbell, 1954)

Dictyomitrella kamoensis Mizutani and Kido, 1983

(Plate 2, fig. 4; Plate 3, figs. 5, 15)

- *1983 *Dictyomitrella? kamoensis* – Mizutani and Kido, p. 258, pl. 53, figs. 2–4b.
1994 *Dictyomitrella? kamoensis* Mizutani and Kido – Goričan, p. 66, pl. 24, fig. 1. (detailed synonymy until 1993)
2003b *Dictyomitrella kamoensis* Mizutani and Kido – Suzuki and Gawlick, p. 188, fig. 6.49. (detailed synonymy between 1994 and 2002)
2015 *Dictyomitrella? kamoensis* Mizutani and Kido – Ishida, pl. 4, figs. 37–42; pl. 11, figs. 1–5.
2018 *Dictyomitrella kamoensis* Mizutani and Kido – Gawlick *et al.*, fig. 12.5.

Genus *Archaeodictyomitra* Pessagno, 1976

Type species: *Archaeodictyomitra squinaboli* Pessagno, 1976

Archaeodictyomitra apiarium (Rüst, 1885)

(Plate 1, figs. 2, 11)

*1885 *Litocampium apiarium* – Rüst, p. 314, pl. 39, fig. 8.

1977b *Archaeodictyomitra apiara* (Rüst) – Pessagno, p. 41, pl. 6, figs. 6, 14.

1981 *Archaeodictyomitra apiarium* (Rüst) – Kocher, p. 56, pl. 12, fig. 13.

1985 *Archaeodictyomitra apiara* (Rüst) – Matsuoka and Yao, pl. 2, fig. 4.

1999 *Archaeodictyomitra apiarium* (Rüst) – Gawlick and Suzuki, fig. 12.4.

2004 *Archaeodictyomitra apiarium* (Rüst) – Ishida, fig. 10.7.

2004 *Archaeodictyomitra apiarium* (Rüst) – Gawlick *et al.*, fig. 3a.10.

2014 *Archaeodictyomitra apiarium* (Rüst) – Suzuki *et al.*, p. 10, pl. 4, fig. 10; pl. 5, fig. 10.

2020 *Archaeodictyomitra apiarium* (Rüst) – Suzuki *et al.*, p. 107, fig. 3.5.

Archaeodictyomitra minoensis (Mizutani, 1981)

(Plate 3, fig. 11)

*1981 *Pseudodictyomitra minoensis* – Mizutani, p. 178, pl. 58, fig. 4; pl. 63, figs. 9, 10.

1985 *Archaeodictyomitra minoensis* (Mizutani) – Matsuoka and Yao, pl. 2, fig. 5.

1999 *Archaeodictyomitra minoensis* (Mizutani) – Gawlick and Suzuki, fig. 12.2.

1999 *Archaeodictyomitra minoensis* (Mizutani) – Gawlick *et al.*, fig. 8.5.

2006 *Archaeodictyomitra minoensis* (Mizutani) – Auer *et al.*, fig. 6.3.

2009 *Archaeodictyomitra minoensis* (Mizutani) – Auer *et al.*, fig. 9.4.

Archaeodictyomitra mirabilis Aita, 1987

(Plate 1, fig. 17; Plate 3, fig. 20)

*1987 *Archaeodictyomitra? mirabilis* – Aita, p. 71, pl. 1, figs. 14a, 14b; pl. 9, figs. 7, 8.

1995a *Archaeodictyomitra? mirabilis* Aita – Baumgartner *et al.*, p. 104, pl. 3236, figs. 1–4.

2001 *Archaeodictyomitra? mirabilis* Aita – Nishizonon, pl. 2, fig. 2.

2003b *Archaeodictyomitra mirabilis* Aita – Suzuki and Gawlick, p. 178, fig. 6.21.

2009 *Archaeodictyomitra mirabilis* Aita – Auer *et al.*, fig. 11.1.

Archaeodictyomitra patricki Kocher, 1981

(Plate 3, fig. 23)

* 1981 *Archaeodictyomitra patricki* – Kocher, p. 57, pl. 12, figs. 14–17.

1997 *Archaeodictyomitra* sp. – Suzuki and Nakae, pl. 1, fig. 7.

2003b *Archaeodictyomitra patricki* Kocher – Suzuki and Gawlick, p. 178, fig. 5.19. (detailed synonymy until 2002)

Archaeodictyomitra rigida Pessagno, 1977a

(Plate 1, fig. 10)

*1977a *Archaeodictyomitra rigida* – Pessagno, p. 81, pl. 7, figs. 10, 11.

2003b *Archaeodictyomitra rigida* Pessagno – Suzuki and Gawlick, p. 179, fig. 5.18; fig. 6.20. (detailed synonymy until 2002)

2004 *Archaeodictyomitra rigida* Pessagno – Gawlick *et al.*, fig. 3b.17.

2004 *Archaeodictyomitra* sp. – Ishida, fig. 7.8; fig. 10.10.

2005 *Archaeodictyomitra rigida* Pessagno – Missoni *et al.*, fig. 7.9; fig. 10.9

2006 *Archaeodictyomitra rigida* Pessagno – Gawlick *et al.*, fig. 8.3; fig. 9.4.

2006 *Archaeodictyomitra rigida* Pessagno – Auer *et al.*, fig. 6.4.

2007 *Archaeodictyomitra rigida* Pessagno – Auer *et al.*, fig. 6.10.

2009 *Archaeodictyomitra rigida* Pessagno – Suzuki and Gawlick, fig. 5.9.

Archaeodictyomitra sixi Yang, 1993

(Plate 3, fig. 50)

*1993 *Archaeodictyomitra sixi* – Yang, p. 122, pl. 19, figs. 3, 19; pl. 20, figs. 9, 10, 19.

2003b *Archaeodictyomitra sixi* Yang – Suzuki and Gawlick, p. 180, fig. 5.17; fig. 6.23. (detailed synonymy until 2003)

2007 *Archaeodictyomitra sixi* Yang – Auer *et al.*, fig. 6.11.

2007 *Archaeodictyomitra sixi* Yang – Gawlick *et al.*, fig. 17.5.

2010 *Archaeodictyomitra sixi* Yang – Gawlick *et al.*, fig. 22.2.

2011 *Archaeodictyomitra sixi* Yang – Gawlick *et al.*, fig. 1.6; fig. 2.5.

Genus *Neorelumbra* Kiessling, 1995

Type species: *Neorelumbra tippitae* Kiessling, 1995

Neorelumbra skenderbegi Chiari, Marcucci and Prela, 2002 (Plate 1, fig. 9)

*2002 *Neorelumbra skenderbegi*. – Chiari *et al.*, p. 68, pl. 1, figs. 14–21.

2003b *Neorelumbra skenderbegi* Chiari, Marcucci and Prela – Suzuki and Gawlick, p. 190, fig. 6.32. (detailed synonymy until 2002)

2007 *Neorelumbra skenderbegi* Chiari, Marcucci and Prela – Auer *et al.*, fig. 6.48.

2009 *Neorelumbra skenderbegi* Chiari, Marcucci and Prela – Suzuki and Gawlick, p. 169, fig. 5.11.

2011 *Neorelumbra skenderbegi* Chiari, Marcucci and Prela – Gawlick *et al.*, fig. 2.23.

Genus *Parvicingula* Pessagno, 1977a

Type species: *Parvicingula santabarbaraensis* Pessagno, 1977a

Parvicingula spinata Vinassa, 1899

(Plate 3, fig. 13)

*1899 *Lithocampe spinata* – Vinassa, p. 237, pl. 2, fig. 40.

1995a *Parvicingula? spinata* (Vinassa) – Baumgartner *et al.*, p. 412, pl. 3187, figs. 1–3.

2003b *Parvicingula spinata* (Vinassa) – Suzuki and Gawlick, p. 187, fig. 5.34. (detailed synonymy until 2002)

2007 *Parvicingula spinata* (Vinassa) – Auer *et al.*, fig. 6.56.

2014 *Parvicingula spinata* (Vinassa) – Suzuki *et al.*, p. 13, pl. 4, fig. 9.

Genus *Loopus* Yang, 1993; emend. herein

*1993 *Loopus* – Yang, p. 123.

1997 *Loopus* Yang – Dumitrica *et al.*, p. 30.

2003b *Loopus* Yang – Suzuki and Gawlick, p. 185.

2009 *Loopus* Yang – Suzuki and Gawlick, p. 170.

Type species: *Loopus doliolum* Dumitrica, 1997 (redesignation herein)

Emended diagnosis: Conical to subcylindrical multicyrtilid test, in case more or less constricted in distal portion. Cephalis with or without horn. Each segment of abdomen and postabdominal chambers is divided by single transverse row of pores. Boundary of each segment is constricted or not. Rims of pores extend on to the surface of each chamber to make short discontinuous costae. Each costa is usually not highly relieved and sometimes no costae are developed on the surface of chambers. In the latter case, test surface is smooth.

Remarks: *Pseudodictyomitra primitiva*, the type species of the genus *Loopus* Yang, 1993, should be attributed to the genus *Pseudodictyomitra*, to which Matsuoka and Yao (1985) assigned the species in their original description. Dumitrica *et al.* (1997) stated that fine bifurcating costae just above single row of pores on each segment is too detailed structure to be of a generic diagnosis. We agree with the opinion of Dumitrica *et al.* (1997), and the genus *Loopus* is used in the sense of Dumitrica *et al.* (1997), namely single row of pores on each segment with or without short costae that are not bifurcate above each pore. In these generic features, we redesignate the type species here, *Loopus doliolum* Dumitrica, 1997.

Loopus doliolum Dumitrica, 1997

(Plate 1, fig. 5; Plate 2, fig. 29)

1982 *Dictyomitra* sp. C – Yao *et al.*, pl. 4, fig. 28.

*1997 *Loopus doliolum* – Dumitrica in Dumitrica *et al.*,

p. 30, pl. 5, figs. 3, 5, 14.

2003b *Loopus doliolum* Dumitrica – Suzuki and Gawlick, p. 186, fig. 6.92, 6.93. (detailed synonymy until 2002)

2004 *Loopus nudus* (Schaaf) – Ishida, fig. 8.4; fig. 10.3.

2009 *Loopus doliolum* Dumitrica – Suzuki and Gawlick, p. 170, fig. 6.5.

2011 *Loopus doliolum* Dumitrica – Gawlick *et al.*, fig. 3.24.

2014 *Loopus doliolum* Dumitrica – Suzuki *et al.*, p. 12, pl. 5, fig. 11.

Remarks: *Loopus doliolum* differs from *Pseudodictyomitra primitiva* in having no distinct short costae or very weak short costae, which don't bifurcate just above pores on each segment.

Genus *Pseudodictyomitra* Pessagno, 1977b; emend. herein

Type species: *Pseudodictyomitra pentacolaensis* Pessagno, 1977b

Emended diagnosis: Multicyrtid test is conical or subcylindrical, in case more or less constricted in distal portion. Cephalis with or without horn. Thorax or abdomen and postabdominal chambers are divided each other by single or double transverse row of pores. In case of single pore rows, imperforate circular dents are arranged below perforate pore rows. Boundary of each postabdominal segment is constricted or not. On the surface of each chamber short discontinuous costae are developed. Each costa is bifurcating downwards to form a rim of pores. Such bifurcating structure is not conspicuous, when chamber surface has robust costae or no costae and smooth.

Remarks: After the original generic definition of Pessagno (1977b) *Pseudodictyomitra* has two transverse rows of primary pores. But many species which can be attributed to the genus *Pseudodictyomitra* has single row of pores with imperforate circular dents. Such character is visible in such species as *Pseudodictyomitra venusta* (Chiari *et al.*, 1997) [as *Cinguloturris? venusta*], *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985, *Pseudodictyomitra conicostriata* Dumitrica, 1997, *Pseudodictyomitra lilyae* (Tan, 1927) in sense of Dumitrica *et al.* (1997) etc. Therefore, we change the type species of the genus *Loopus* from *Pseudodictyomitra primitiva* to *Loopus doliolum* (see remarks of the genus *Loopus*).

Pseudodictyomitra primitiva Matsuoka and Yao, 1985

(Plate 1, fig. 6; Plate 2, fig. 38; Plate 3, fig. 19)

*1985 *Pseudodictyomitra primitiva* – Matsuoka and Yao, p. 131, pl. 1, figs. 1–6; pl. 3, figs. 1–4.

1996 *Pseudodictyomitra primitiva* Matsuoka and Yao – Nishizono, pl. 29, figs. 16–19.

2001 *Loopus primitivus* (Matsuoka and Yao) – Nishizono, pl. 2, fig. 10.

2002 *Loopus primitivus* (Matsuoka and Yao) – Hori *et al.*, pl. 11, fig. 25.

2004 *Loopus primitivus* (Matsuoka and Yao) – Ishida and

Kozai, fig. 6.5, 6.9, 6.10.

2004 *Loopus primitivus* (Matsuoka and Yao) – Kozai *et al.*, fig. 7.13, 7.14.

2007 *Pseudodictyomitra primitiva* Matsuoka and Yao – Auer *et al.*, fig. 6.65.

2011 *Pseudodictyomitra primitiva* Matsuoka and Yao – Gawlick *et al.*, fig. 3.29.

2014 *Pseudodictyomitra primitiva* Matsuoka and Yao – Suzuki *et al.*, p. 11, pl. 5, fig. 1.

Remarks: We place this species not in the genus *Loopus*, but in the genus *Pseudodictyomitra*, as mentioned above.

Genus *Pseudoecyrtis* Pessagno, 1977b

Type species: *Eucyrtis? zhamoidai* Foreman, 1973

Pseudoecyrtis reticularis Matsuoka and Yao, 1985 (Plate 2, fig. 18)

*1985 *Pseudoecyrtis reticularis* – Matsuoka and Yao, p. 132, pl. 1, figs. 16–21; pl. 3, figs. 14–17.

2001 *Pseudoecyrtis reticularis* Matsuoka and Yao – Missoni *et al.*, fig. 3.12.

2007 *Pseudoecyrtis reticularis* Matsuoka and Yao – Gawlick *et al.*, fig. 19.31.

Genus *Ristola* Pessagno and Whalen, 1982; emend. Baumgartner, 1984

Type species: *Parvicingula? procera* Pessagno, 1977a

Ristola altissima (Rüst, 1885) (Plate 2, fig. 17)

*1885 *Lithocampe altissima* – Rüst, p. 315, pl. 40, fig. 2.
1984 *Ristola altissima* (Rüst) – Baumgartner, p. 783, pl. 8, figs. 3, 4, 9.

2001 *Ristola altissima* (Rüst) – Missoni *et al.*, p. 783, fig. 3.1.

2001 *Ristola altissima* (Rüst) – Nishizono, pl. 3, fig. 9.

2015 *Ristola altissima* (Rüst) – Ishida, pl. 5, figs. 17, 18.

Family AMPHIPYNDACIDAE Riedel, 1967a

Genus *Takemuraella* O’Dogherty, Goričan and Gawlick, 2017

non 1974 *Triversus* – Sher, p. 323. (Nematoda)

1986 *Triversus* – Takemura, p. 62.

2003b *Triversus* Takemura – Suzuki and Gawlick, p. 194.

*2017 *Takemuraella* – O’Dogherty, Goričan and Gawlick, p. 57.

Type species: *Triversus japonicus* Takemura, 1986

Remarks: O’Dogherty *et al.* (2017) pointed out that the genus name “*Triversus*” is preoccupied by the nematoid genus *Triversus* Sher, and they renamed *Takemuraella*.

Takemuraella hungarica (Kozur, 1985) (Plate 1, fig. 16; Plate 3, fig. 12)

*1985 *Eoxitus hungaricus* – Kozur, p. 216, figs. 1a, 1b, 1d, 1e.

1986 *Triversus spinifer* – Takemura, p. 63, pl. 10, figs. 21–23; pl. 11, figs. 1, 2.

1995a *Parvicingula dhimenaensis* ssp. A – Baumgartner *et al.*, p. 406, pl. 4071, figs. 1–4.

2003b *Triversus hungaricus* (Kozur) – Suzuki and Gawlick, p. 195, fig. 60.58–60.60. (detailed synonymy until 2002)

pt. 2004 *Parvicingula dhimenaensis* Baumgartner – Ishida, fig. 7.9, 7.10; fig. 8.20; non fig. 10.13 [= *Parvicingula dhimenaensis* Baumgartner].

2007 *Triversus hungaricus* (Kozur) – Gawlick *et al.*, fig. 7.10; fig. 8.26; fig. 18.7.

2009 *Triversus hungaricus* (Kozur) – Suzuki and Gawlick, p. 170, fig. 5.14; fig. 6.6–6.8.

Takemuraella hexagonata (Heitzer, 1930)

(Plate 1, fig. 18; Plate 3, figs. 28, 29)

*1930 *Cyrtocalpis hexagonata* – Heitzer, p. 391, pl. 28, fig. 26.

1986 *Pseudodictyomitrella hexagonata* (Heitzer) – Grill and Kozur, pl. 4, figs. 2, 4.

2003b *Triversus hexagonatus* (Heitzer) – Suzuki and Gawlick, p. 194, fig. 5.48; fig. 6.61. (detailed synonymy until 2002)

2004 *Parvicingula* sp. – Ishida, fig. 7.13; non 12.20.

2005 *Triversus hexagonatus* (Heitzer) – Suzuki and Kuwahara, p. 50, pl. 1, fig. 8.

2006 *Triversus hexagonatus* (Heitzer) – Gawlick *et al.*, fig. 8c.40; fig. 9b.20.

2006 *Triversus hexagonatus* (Heitzer) – Auer *et al.*, fig. 6.48.

2009 *Triversus hexagonatus* (Heitzer) – Suzuki and Gawlick, p. 170, fig. 5.15; fig. 6.11A, 6.11B.

2009 *Stichomitra?* spp. – Ishida *et al.*, fig. 6.12, 6.13.

2011 *Triversus hexagonatus* (Heitzer) – Gawlick *et al.*, fig. 1.24; fig. 3.38.

Genus *Stichomitra* Cayeux, 1897

Type species: *Stichomitra bertrandi* Cayeux, 1897. The type species was subsequently designated by O’Dogherty (1994).

Stichomitra annibill Kocher, 1981; emend. Suzuki and Gawlick, 2003b

(Plate 1, fig. 13; Plate 2, figs. 19, 25; Plate 3, fig. 24)

*1981 *Stichomitra annibill* – Kocher, p. 96, pl. 16, figs. 24–26.

1987 *Stichomitra? tairai* – Aita, p. 72, pl. 3, figs. 7–9; pl. 10, figs. 3, 4.

1997 *Xitus singularis* – Hull, p. 138, pl. 47, figs. 1, 7, 20.

1999 *Xitus reticulatus* – Hori, p. 76, fig. 7.1–7.5.

1999 *Xitus singularis* Hull – Hori, p. 76, fig. 7.6.

2003a *Stichomitra annibill* Kocher – Suzuki and Gawlick, p. 119, pl. 1, fig. 14.

2003b *Stichomitra annibill* Kocher – Suzuki and Gawlick,

- p. 192, fig. 6.35, 6.36. (detailed synonymy until 2002)
2004 *Xitus spicularius* (Aliev) – Ishida, fig. 7.19; fig. 8.25.
2004 *Xitus* sp. – Ishida, fig. 8.26; ? fig. 7.18.
2005 *Stichomitra annibill* Kocher – Missoni *et al.*, fig. 13.4.
2006 *Stichomitra annibill* Kocher – Gawlick *et al.*, fig. 8b.30.
2006 *Stichomitra annibill* Kocher – Auer *et al.*, fig. 6.37.
2009 *Stichomitra annibill* Kocher – Suzuki and Gawlick, p. 176, fig. 5.16; fig. 6.16A, 6.16B.
2011 *Stichomitra annibill* Kocher – Gawlick *et al.*, fig. 3.32.
2014 *Stichomitra annibill* Kocher – Suzuki *et al.*, p. 15, pl. 5, figs. 5, 9.
2015 *Stichomitra annibill* Kocher – Ishida, pl. 10, figs. 30–36.

***Stichomitra* sp. A** sensu Baumgartner *et al.*, 1995a
(Plate 2, fig. 34)

- *1995a *Stichomitra* sp. A – Baumgartner *et al.*, p. 528, pl. 3192, figs. 1–3.

Genus ***Unuma*** Ichikawa and Yao, 1976

Type species: ***Unuma typicus*** Ichikawa and Yao, 1976

Unuma typicus Ichikawa and Yao, 1976
(Plate 2, fig. 6)

- *1976 *Unuma (Unuma) typicus* – Ichikawa and Yao, p. 112, pl. 1, figs. 1–3.
1994 *Unuma typicus* Ichikawa and Yao – Goričan, p. 96, pl. 10, fig. 13.
1995a *Unuma typicus* Ichikawa and Yao – Baumgartner *et al.*, p. 622, pl. 4059, figs. 1, 2. (detailed synonymy until 1991)
2009 *Unuma typicus* Ichikawa and Yao – Suzuki and Gawlick, p. 177, fig. 5.19.
cf. 2016 *Unuma cf. typicus* Ichikawa and Yao – Suzuki and Nakai, pl. 1, figs. 4a, 4b.

Unuma gordus Hull, 1997

(Plate 1, fig. 29; Plate 2, fig. 41)

- *1997 *Unuma gorda* – Hull, p. 172, pl. 43, figs. 9, 11, 12.
2003b *Unuma gorda* Hull – Suzuki and Gawlick, p. 198, fig. 5.36; fig. 6.68. (detailed synonymy until 2002)
2007 *Unuma gorda* Hull – Gawlick *et al.*, fig. 7.21; fig. 8.44; fig. 17.30; fig. 18.13.
2009 *Unuma gordus* Hull – Suzuki and Gawlick, p. 177, fig. 6.2A, 6.2B.

Genus ***Protunuma*** Ichikawa and Yao, 1976

Type species: ***Protunuma fusiformis*** Ichikawa and Yao, 1976

Protunuma fusiformis Ichikawa and Yao, 1976
(Plate 3, fig. 27)

- *1976 *Protunuma fusiformis* – Ichikawa and Yao, p. 116, pl. 2, figs. 1–4b.

Protunuma lanosus Ozvoldova, 1996

(Plate 2, fig. 8)

- *1996 ?*Protunuma lanosus* – Ožvoldová in Sykora and Ozvoldova, p. 23, pl. 2, fig. 13; pl. 3, figs. 1–6.
2003a *Protunuma lanosus* Ozvoldova – Suzuki and Gawlick, p. 119, pl. 1, fig. 12.
2007 *Protunuma lanosus* Ozvoldova – Gawlick *et al.*, fig. 7.12.

Protunuma japonicus Matsuoka and Yao, 1985; emend. herein

(Plate 1, fig. 30)

non 1930 *Cenellipsis multicostatus* – Heitzer, p. 388, pl. 17, fig. 13.

- *1985 *Protunuma japonicus* – Matsuoka and Yao, p. 130, pl. 1, figs. 11–15; pl. 3, figs. 6–9.

2001 *Protunuma japonicus* Matsuoka and Yao – Wegerer *et al.*, fig. 4b.16; fig. 5.11.

2007 *Protunuma multicostatus* (Heitzer) – Gawlick *et al.*, fig. 7.13; ? fig. 19.30.

2011 *Protunuma multicostatus* (Heitzer) – Gawlick *et al.*, fig. 3.28; ? fig. 2.28.

2013 *Protunuma multicostatus* (Heitzer) – Krische *et al.*, pl. 3, fig. 6.

non 2015 *Protunuma japonicus* Matsuoka and Yao – Ishida, pl. 3, fig. 16; pl. 8, fig. 15 [= *Protunuma multicostatus*].

Emended diagnosis: *Protunuma* species, which possesses not only two, but also three or four rows of pores between neighbouring two longitudinal plicae.

Remarks: Suzuki and Gawlick (2003b) regarded *Protunuma japonicus* as a younger synonym of *Protunuma multicostatus* (Heitzer, 1930) (= *Cenellipsis multicostatus*). If we follow the original description of Matsuoka and Yao (1985) “Two to four rows of pores present between neighbouring two longitudinal plicae”, namely including a specimen having “only two rows of pores between neighbouring two longitudinal plicae”, *Protunuma japonicus* should be a younger synonym of *Protunuma multicostatus* (Heitzer). Our careful observation of specimens of *Protunuma multicostatus* clarifies that it has only two rows of pores between neighbouring two longitudinal plicae (Fig. 6a). If a specimen having three rows of pores between two longitudinal plicae even in one portion, it should be *Protunuma japonicus* (Fig. 6b; Plate 1, fig. 30). Therefore, we separate *Protunuma japonicus* from the previously synonymized “*Protunuma multicostatus*”.

Genus ***Podobursa*** Wisniowski, 1889; emend. Foreman, 1973

Type species: ***Podobursa dunikowskii*** Wisniowski, 1889. Monotype.

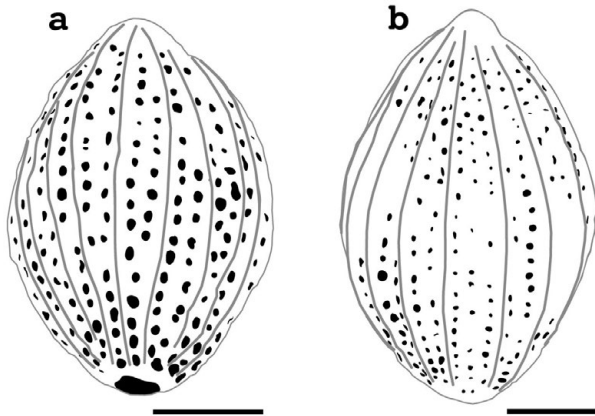


Fig. 6 Sketches of two *Protunuma* species. a: *Protunuma multicostatus* (Heitzer, 1930), from the Brielgraben section, b: *Protunuma japonicus* Matsuoka and Yao, 1985, from the Fludergraben section (Plate 1, fig. 30). *Protunuma japonicus* differs from *P. multicostatus* in having not only two, but also three longitudinal pore rows. Each scale bar is 30 μ m.

Podobursa nodosa (Chiari, Marcucci and Prela, 2002)
(Plate 2, fig. 31)

- 1997 *Podobursa?* sp. B. – Hull, p. 108, pl. 43, figs. 5, 18, 19.
*2002 *Williriedellum nodosum* – Chiari *et al.*, p. 84, pl. 5, figs. 15–19.
2009 *Podobursa nodosa* (Chiari, Marcucci and Prela) – Suzuki and Gawlick, p. 178, fig. 5.20, 5.21.

Genus ***Droltus*** Pessagno and Whalen, 1982

Type species: ***Droltus lyellensis*** Pessagno and Whalen, 1982.

Remarks: Suzuki *et al.* (2002) demonstrated a VB (branch of vertical spine) ring as the cephalic skeletal elements in their specimen of *Droltus hecatensis* Pessagno and Whalen. We, therefore, classify the genus *Droltus* into the family Amphipyndacidae.

Droltus galerus Suzuki, 1995b
(Plate 3, fig. 49)

- 1995a *Droltus* sp. – Suzuki, fig. 4.15.
*1995b *Droltus galerus* – Suzuki, p. 284, fig. 5.5–5.7; fig. 7.1a, 7.1b.
2006 *Droltus galerus* Suzuki – Auer *et al.*, fig. 6.11.
2007 *Droltus galerus* Suzuki – Auer *et al.*, fig. 6.17.
2009 *Droltus galerus* Suzuki – Suzuki and Gawlick, p. 177, fig. 6.3A–6.4B.

Remarks: Our specimen from the Fludergraben section exhibits sharp pointed cephalis rather than rounded one that seen in type specimens from the Lower Jurassic chert in the Umenoki Unit of Shikoku, Japan (Suzuki, 1995b).

Family **WILLIRIEDELLIDAE** Dumitrica, 1970

Genus ***Williriedellum*** Dumitrica, 1970

Type species: ***Williriedellum crystallinum*** Dumitrica, 1970

Williriedellum crystallinum Dumitrica, 1970

(Plate 2, figs. 16, 36)

- *1970 *Williriedellum crystallinum* – Dumitrica, p. 69, pl. 10, figs. 60a–60c, 62, 63.
1994 *Williriedellum crystallinum* Dumitrica – Goričan, p. 96, pl. 12, figs. 1, 2a–2c. (detailed synonymy until 1993)
2003b *Williriedellum crystallinum* Dumitrica – Suzuki and Gawlick, p. 199, fig. 6.76.
2005 *Williriedellum crystallinum* Dumitrica – Missoni *et al.*, fig. 7.23.
2006 *Williriedellum crystallinum* Dumitrica – Gawlick *et al.*, fig. 8c.41.
2006 *Williriedellum crystallinum* Dumitrica – Auer *et al.*, fig. 6.51.
2009 *Williriedellum crystallinum* Dumitrica – Suzuki and Gawlick, p. 178, fig. 5.24.
2011 *Williriedellum crystallinum* Dumitrica – Gawlick *et al.*, fig. 1.26; fig. 2.36; fig. 3.40.

Williriedellum sujkowski Widz and De Wever, 1993
(Plate 2, fig. 32)

- *1993 *Williriedellum sujkowski* – Widz and De Wever, p. 88, pl. 2, figs. 7–10.
2007 *Williriedellum sujkowski* Widz and De Wever – Auer *et al.*, fig. 6.123.
2010 *Williriedellum sujkowski* Widz and De Wever – Gawlick *et al.*, fig. 27.22.
2011 *Williriedellum sujkowski* Widz and De Wever – Gawlick *et al.*, fig. 1.27; fig. 3.42.

Williriedellum carpathicum Dumitrica, 1970
(Plate 2, fig. 20; Plate 3, fig. 16)

- *1970 *Williriedellum carpathicum* – Dumitrica, p. 70, pl. 9, figs. 56a, 56b, 57–59; pl. 10, fig. 61.
2003b *Williriedellum carpathicum* Dumitrica – Suzuki and Gawlick, p. 200, fig. 6.74. (detailed synonymy until 2003)
2004 *Tricolocapsa yaoi* Matsuoka – Ishida, fig. 8.33.
2007 *Williriedellum carpathicum* Dumitrica – Auer *et al.*, fig. 6.120.
2010 *Williriedellum carpathicum* Dumitrica – Gawlick *et al.*, fig. 16A.8; fig. 16B.13; fig. 19.43; fig. 22.6; fig. 50.3.
2011 *Williriedellum carpathicum* Dumitrica – Gawlick *et al.*, fig. 1.25; fig. 3.39.
2015 *Williriedellum* sp. 2 – Ishida, pl. 6, fig. 50.

Williriedellum marcucciae Cortese, 1993
(Plate 2, fig. 10; Plate 3, fig. 44)

- 1983 *Williriedellum* sp. A gr. – Matsuoka, p. 23, pl. 4, figs. 1–3; pl. 8, figs. 11–15.

- *1993 *Williriedellum marcuccii* – Cortese, p. 180, pl. 7, figs. 6, 7.
1994 *Williriedellum* sp. A sensu Matsuoka – Goričan, p. 96, pl. 12, figs. 9a–9c, 10a–10c, 11a, 11b. (detailed synonymy until 1993)
2003b *Williriedellum* sp. A sensu Matsuoka – Suzuki and Gawlick, p. 201, fig. 6.77. (detailed synonymy between 1994 and 2003)
2004 *Williriedellum* sp. A sensu Matsuoka – Gawlick *et al.*, fig. 3b.19.
2005 *Williriedellum* sp. A sensu Matsuoka – Missoni *et al.*, fig. 7.25.
2006 *Williriedellum* sp. A sensu Matsuoka – Auer *et al.*, fig. 6.53.
2006 *Williriedellum* sp. A sensu Matsuoka – Gawlick *et al.*, fig. 8c.43; fig. 9b.24.
2009 *Williriedellum marcucciae* Cortese – Suzuki and Gawlick, p. 179, fig. 5.25; fig. 6.49A, 6.49B.
2015 *Williriedellum marcucciae* Cortese – Ishida, pl. 1, figs. 51, 52; pl. 6, figs. 46–48.
2016 *Williriedellum marcucciae* Cortese – Suzuki and Nakai, pl. 1, figs. 1a, 1b
2018 *Williriedellum marcucciae* Cortese – Gawlick *et al.*, fig. 14.10; fig. 18.39.

- Williriedellum dierschei*** Suzuki and Gawlick, 2004 (Plate 1, figs. 1, 27; Plate 2, figs. 12, 30; Plate 3, fig. 45)
*2004 *Williriedellum dierschei* – Suzuki and Gawlick in Gawlick *et al.*, p. 311, fig. 4.1–4.6. (detailed synonymy until 2001)
2005 *Williriedellum dierschei* Suzuki and Gawlick – Missoni *et al.*, fig. 7.24; fig. 10.35.
2005 *Williriedellum dierschei* Suzuki and Gawlick – Suzuki and Kuwahara, p. 52, pl. 1, figs. 18, 19.
2006 *Williriedellum dierschei* Suzuki and Gawlick – Auer *et al.*, fig. 6.52.
2006 *Williriedellum dierschei* Suzuki and Gawlick – Gawlick *et al.*, fig. 9b.23.
2009 *Williriedellum dierschei* Suzuki and Gawlick – Suzuki and Gawlick, p. 179, fig. 5.27A, 5.27B, 5.28; fig. 6.48A, 6.48B.
2015 *Williriedellum dierschei* Suzuki and Gawlick – Ishida, pl. 1, figs. 47, 48; pl. 6, figs. 43–45.
2018 *Williriedellum dierschei* Suzuki and Gawlick – Gawlick *et al.*, fig. 14.9; fig. 18.37; cf. fig. 25.4.

- Williriedellum* sp. C** sensu Gawlick *et al.*, 2018 (Plate 3, fig. 43)
1992 *Tricolocapsa* sp. A – Ozvoldova, p. 115, pl. 2, figs. 6, 7.
2007 *Tricolocapsa* sp. A sensu Ozvoldova – Auer *et al.*, fig. 6.109.
*2018 *Williriedellum* sp. C – Gawlick *et al.*, fig. 18.40.
Remarks: Depicted specimen exhibits a three-chambered test with a large globose abdomen, which possesses a projected short tube-like aperture on its base. Somewhat large pores are scattered on a smooth surfaced abdomen.

Genus ***Praewilliriedellum*** Kozur, 1984

Type species: ***Praewilliriedellum cephalospinosum*** Kozur, 1984
Remarks: Kozur (1984) mentioned that the thorax of this genus is not or very slightly depressed into the abdomen, although the genus is classified into the family Williriedellidae by Kozur (1984). If the thorax is not depressed into the abdomen commonly, this genus should be classified into the family Arcanicapsidae.

Praewilliriedellum aff. spinosum Kozur, 1984 (Plate 3, fig. 46)
aff. *1984 *Praewilliriedellum spinosum* – Kozur, p. 52, pl. 1, figs. 1–3.
Remarks: Our specimens from the Fludergraben section have a slightly elongated test in comparison with the type specimens depicted by Kozur (1984). Thus, we describe here as *Praewilliriedellum aff. spinosum*.

Genus ***Zhamoidellum*** Dumitrica, 1970

Type species: ***Zhamoidellum ventricosum*** Dumitrica, 1970
Zhamoidellum ventricosum Dumitrica, 1970 (Plate 1, fig. 25)
*1970 *Zhamoidellum ventricosum* – Dumitrica, p. 79, pl. 9, figs. 55a, 55b.
2003b *Zhamoidellum ventricosum* Dumitrica – Suzuki and Gawlick, p. 202, fig. 6.57. (detailed synonymy until 2002)
2005 *Zhamoidellum ventricosum* Dumitrica – Missoni *et al.*, fig. 13.6.
2006 *Zhamoidellum ventricosum* Dumitrica – Auer *et al.*, fig. 6.57.
2009 *Zhamoidellum ventricosum* Dumitrica – Suzuki and Gawlick, p. 179, fig. 5.29.
2018 *Zhamoidellum ventricosum* Dumitrica – Gawlick *et al.*, fig. 18.41.
Remarks: A depicted specimen shows larger pores and pore frames on globous abdomen than those of other specimens showed previously.

Zhamoidellum ovum Dumitrica, 1970 (Plate 1, fig. 20; Plate 2, figs. 13, 35; Plate 3, fig. 26)
*1970 *Zhamoidellum ovum* – Dumitrica, p. 79, pl. 9, figs. 52a, 52b, 53, 54.
1994 *Zhamoidellum ovum* Dumitrica – Goričan, p. 97, pl. 13, figs. 3–7. (detailed synonymy until 1993)
2003b *Zhamoidellum ovum* Dumitrica – Suzuki and Gawlick, p. 203, fig. 6.56.
2004b *Zhamoidellum ovum* Dumitrica – Suzuki *et al.*, p. 385, fig. 5.3. (detailed synonymy between 1994 and 2003)
2004 *Zhamoidellum ovum* Dumitrica – Gawlick *et al.*, fig. 3b.27.
2004 *Zhamoidellum ovum* Dumitrica – Ishida, fig. 8.32;

fig. 10.22.

2005 *Zhamoidellum ovum* Dumitrica – Missoni *et al.*, fig. 7.28; fig. 13.7.

2006 *Zhamoidellum ovum* Dumitrica – Auer *et al.*, fig. 6.56.

2006 *Zhamoidellum ovum* Dumitrica – Gawlick *et al.*, fig. 8c.45.

2009 *Zhamoidellum ovum* Dumitrica – Suzuki and Gawlick, p. 179, fig. 5.30A, 5.30B; fig. 6.33A, 6.33B.

2009 *Williriedellum yaoui* (Kozur) – Ishida *et al.*, fig. 6.2.

2011 *Zhamoidellum ovum* Dumitrica – Gawlick *et al.*, fig. 1.28; fig. 2.39; fig. 3.45.

2014 *Zhamoidellum ovum* Dumitrica – Suzuki *et al.*, p. 16, pl. 4, fig. 2; pl. 5, fig. 16.

2015 *Zhamoidellum ovum* Dumitrica – Ishida, pl. 1, fig. 62; pl. 6, figs. 59, 60.

Family **ARCANICAPSIDAE** Takemura, 1986

Genus ***Stylocapsa*** Principi, 1909; emend. Tan, 1927

Type species: ***Stylocapsa exagonata*** Principi, 1909

Stylocapsa oblongula Kocher, 1980

(Plate 3, fig. 34)

* 1980 *Stylocapsa oblongula* – Kocher in Baumgartner *et al.*, p. 62, pl. 6, fig. 1.

2001 *Stylocapsa oblongula* Kocher – Suzuki *et al.*, fig. 5.10.

2001 *Stylocapsa oblongula* Kocher – Wegerer *et al.*, fig. 4a.18; fig. 6.3.

2007 *Stylocapsa oblongula* Kocher – Auer *et al.*, fig. 6.86.

2015 *Kilinora? oblongula* (Kocher) – Ishida, pl. 1, figs. 7, 8.

Genus ***Kilinora*** Hull, 1997

Type species: ***Stylocapsa? spiralis*** Matsuoka, 1982

Remarks: We agree with the establishment of the genus *Kilinora* by Hull (1997), to separate the species having a thorax with costae ornamentation from that with a latticed thorax.

Kilinora cf. spiralis (Matsuoka, 1982)

(Plate 1, fig. 31)

cf. *1982 *Stylocapsa? spiralis* – Matsuoka, p. 77, pl. 3, figs. 1–8.

Remarks: Our single specimen is poorly preserved and only a part of peculiar ornamentation, i.e. oblique plicae, can be observed.

Genus ***Gongylothorax*** Foreman, 1968; emend. Dumitrica, 1970

Type species: ***Dicolocapsa verbeeki*** Tan, 1927. Suzuki and Gawlick (2003b) discussed in detail.

Gongylothorax favosus Dumitrica, 1970

Remarks: *Gongylothorax favosus* is subdivided into two subspecies, namely the nominate subspecies *Gongylothorax favosus favosus* Dumitrica and the subspecies *Gongylothorax favosus oviformis* Suzuki and Gawlick.

Gongylothorax favosus favosus Dumitrica, 1970

(Plate 1, fig. 26; Plate 2, figs. 7, 28; Plate 3, fig. 35)

*1970 *Gongylothorax favosus* – Dumitrica, p. 56, pl. 1, figs. 1a–1c, 2.

1994 *Gongylothorax favosus* Dumitrica – Ishida, fig. 3.5.

2003a *Gongylothorax favosus* Dumitrica – Suzuki and Gawlick, p. 119, pl. 1, fig. 13.

2003b *Gongylothorax favosus* Dumitrica – Suzuki and Gawlick, p. 205, fig. 6.96. (detailed synonymy until 2002)

2005 *Gongylothorax favosus* Dumitrica – Missoni *et al.*, fig. 7.30; fig. 13.8.

2006 *Gongylothorax favosus* Dumitrica – Auer *et al.*, fig. 6.17.

2006 *Gongylothorax favosus* Dumitrica – Gawlick *et al.*, fig. 8a.16; fig. 9a.13.

2009 *Gongylothorax favosus favosus* Dumitrica – Suzuki and Gawlick, p. 180, fig. 5.31A–5.31C, 5.32A, 5.32B; fig. 6.21A, 6.21B.

2009 *Gongylothorax favosus* Dumitrica – Ishida *et al.*, fig. 6.9, 6.10.

2014 *Gongylothorax favosus favosus* Dumitrica – Suzuki *et al.*, p. 17, pl. 4, fig. 8; pl. 5, fig. 14.

Remarks: *Gongylothorax favosus favosus* differs from *Gongylothorax favosus oviformis* in having a spherical thorax with a depressed cephalis.

Gongylothorax favosus oviformis Suzuki and Gawlick, 2009 (Plate 2, fig. 23; Plate 3, fig. 36)

1994 *Gongylothorax* aff. *favosus* Dumitrica – Goričan, p. 70, pl. 13, figs. 9a–9c, 11a–11c. (detailed synonymy until 1993)

cf. 2005 *Gongylothorax* aff. *favosus* Dumitrica – Suzuki and Kuwahara, p. 55, pl. 2, figs. 9, 10. (detailed synonymy between 1994 and 2004)

2006 *Gongylothorax* aff. *favosus* Dumitrica – Gawlick *et al.*, fig. 8a.17; fig. 9a.12.

*2009 *Gongylothorax favosus oviformis* – Suzuki and Gawlick, p. 180, fig. 5.33A–5.34C; fig. 6.22A–6.26B.

Remarks: *Gongylothorax favosus oviformis* differs from *Gongylothorax favosus favosus* in having an elliptical test outline with a not so depressed cephalis. In case of *Gongylothorax favosus oviformis*, penta- or hexagonal pore frames become lager down to thoracic base.

***Gongylothorax* sp. C** sensu Suzuki and Gawlick, 2003b (Plate 3, fig. 42)

1997 *Gongylothorax siphonifer* Dumitrica – Yao, pl. 9, fig. 417.

*2003b *Gongylothorax* sp. C – Suzuki and Gawlick, p. 206, fig. 6.98.

2009 *Gongylothorax* sp. C sensu Suzuki and Gawlick–Suzuki and Gawlick, p. 181, fig. 5.35, 5.36.

2016 *Gongylothorax* sp. C sensu Suzuki and Gawlick – Gawlick *et al.*, fig. 11g.

Remarks: Our single specimen possesses a projected tube-like aperture on a base of bulbous thorax, on which somewhat larger pores are more sparsely distributed in comparison with the materials from north side of Mt. Loser (Suzuki and Gawlick, 2003b) and Hallstatt salt mine (Suzuki and Gawlick, 2009).

Genus *Tricolocapsa* Haeckel, 1881

Type species: *Tricolocapsa theophrasti* Haeckel, 1887

Tricolocapsa tetragona Matsuoka, 1983

(Plate 1, fig. 32; Plate 2, fig. 9)

*1983 *Tricolocapsa tetragona* – Matsuoka, p. 22, pl. 3, figs. 8–12; pl. 8, figs. 4–10.

cf. 1994 *Tricolocapsa cf. tetragona* Matsuoka – Ishida, fig. 3.13.

1994 *Tricolocapsa tetragona* Matsuoka – Goričan, p. 94, pl. 13, figs. 8, 10. (detailed synonymy until 1993)

1999 *Tricolocapsa tetragona* Matsuoka – Wegerer *et al.*, fig. 5.1.

2007 *Tricolocapsa tetragona* Matsuoka – Gawlick *et al.*, fig. 18.40.

2009 *Tricolocapsa tetragona* Matsuoka – Suzuki and Gawlick, p. 183, fig. 5.43.

2010 *Tricolocapsa tetragona* Matsuoka – Gawlick *et al.*, fig. 19.40; fig. 27.19.

2011 *Tricolocapsa tetragona* Matsuoka – Gawlick *et al.*, fig. 3.36.

Tricolocapsa undulata (Heitzer, 1930)

(Plate 1, fig. 22; Plate 2, fig. 27; Plate 3, figs. 17, 33)

*1930 *Lithobotrys undulata* – Heitzer, p. 390, pl. 28, fig. 22.

1987 *Sethocapsa funatoensis* – Aita, p. 73, pl. 2, figs. 6a–b, 7a–b; pl. 9, figs. 14, 15.

1987 *Sethocapsa yahazuensis* – Aita, p. 73, pl. 2, figs. 8a–b, 9a–b; pl. 9, figs. 16, 17.

1993 *Tricolocapsa undulata* (Heitzer) – Ozvoldova and Faupl, pl. 3, fig. 12.

2005 *Tricolocapsa undulata* (Heitzer) – Suzuki and Kuwahara, p. 59, pl. 2, fig. 3. (detailed synonymy until 2004)

2005 *Tricolocapsa undulata* (Heitzer) – Missoni *et al.*, fig. 7.37; fig. 10.45.

2006 *Tricolocapsa undulata* (Heitzer) – Auer *et al.*, fig. 6.44.

2006 *Tricolocapsa undulata* (Heitzer) – Gawlick *et al.*, fig. 8c.36; fig. 9b.21.

2009 *Tricolocapsa undulata* (Heitzer) – Suzuki and Gawlick, p. 183, fig. 5.44A, 5.44B, 5.45A, 5.45B; fig. 6.18A, 6.18B, 6.19A, 6.19B.

2011 *Tricolocapsa undulata* (Heitzer) – Gawlick *et al.*,

fig. 2.34; fig. 3.37.

2015 *Zhamoidellum undulata* (Heitzer) – Ishida, pl. 1, figs. 55–59; pl. 6, figs. 52–55.

Remarks: We integrate two species of Aita (1987), i.e. *Sethocapsa funatoensis* and *Sethocapsa yahazuensis*, into *Tricolocapsa undulata* (Heitzer, 1930) as younger synonyms (see Suzuki and Gawlick, 2003b; Suzuki and Kuwahara, 2005).

Genus *Striatojaponocapsa* Kozur, 1984

Type species: *Tricolocapsa plicarum* Yao, 1979

Striatojaponocapsa conexa (Matsuoka, 1983)

(Plate 2, fig. 39; Plate 3, fig. 31)

*1983 *Tricolocapsa conexa* – Matsuoka, p. 20, pl. 3, figs. 3–7; pl. 7, figs. 11–14.

1994 *Tricolocapsa conexa* Matsuoka – Goričan, p. 94, pl. 11, figs. 7a–b, 8, 9, 10a–b. (detailed synonymy until 1993)

1997 *Striatojaponocapsa conexa* (Matsuoka) – Hull, p. 166, pl. 37, fig. 20.

2003b *Tricolocapsa conexa* Matsuoka – Suzuki and Gawlick, p. 208, fig. 5.42; fig. 6.43–6.45.

2005 *Tricolocapsa conexa* Matsuoka – Missoni *et al.*, fig. 10.44.

2007 *Striatojaponocapsa conexa* (Matsuoka) – Hatakeda *et al.*, p. 54, pl. 2, figs. 1–10.

2009 *Striatojaponocapsa conexa* (Matsuoka) – Suzuki and Gawlick, p. 182, fig. 5.40; fig. 6.32A, 6.32B.

2015 *Striatojaponocapsa conexa* (Matsuoka) – Ishida, pl. 1, figs. 16–19; pl. 6, figs. 21–25.

Striatojaponocapsa riri O’Dogherty, Goričan and Dumitrica, 2006

(Plate 2, figs. 11, 40)

1994 *Tricolocapsa* sp. A – Goričan, p. 9, pl. 11, figs. 11–13.

*2006 *Striatojaponocapsa riri* – O’Dogherty, Goričan and Dumitrica, p. 447, pl. 8, figs. 14, 15.

2007 *Striatojaponocapsa riri* O’Dogherty, Goričan and Dumitrica – Hatakeda *et al.*, p. 55, pl. 2, figs. 11–20.

2007 *Tricolocapsa* sp. A sensu Goričan – Auer *et al.*, fig. 6.108.

2015 *Striatojaponocapsa riri* O’Dogherty, Goričan and Dumitrica – Ishida, pl. 1, figs. 20–24; pl. 6, figs. 26–32.

Striatojaponocapsa synconexa O’Dogherty, Goričan and Dumitrica, 2006

(Plate 1, fig. 24; Plate 2, fig. 33; Plate 3, fig. 30)

*2006 *Striatojaponocapsa synconexa* – O’Dogherty, Goričan and Dumitrica, p. 447, pl. 10, figs. 9–17. (Detailed synonymy)

2007 *Striatojaponocapsa synconexa* O’Dogherty, Goričan and Dumitrica – Hatakeda *et al.*, p. 54, pl. 1, figs. 11–20.

2015 *Striatojaponocapsa synconexa* O’Dogherty, Goričan and Dumitrica – Ishida, pl. 1, figs. 13–15; pl. 6, figs.

19, 20.

Striatojaponocapsa naradaniensis (Matsuoka, 1984)
(Plate 2, fig. 21; Plate 3, fig. 40)

*1984 *Stichocapsa naradaniensis* – Matsuoka, p. 145, pl. 1, figs. 1–5; pl. 2, figs. 1–6.

1994 *Stichocapsa naradaniensis* Matsuoka – Goričan, p. 88, pl. 11, fig. 6. (detailed synonymy until 1993)

2003b *Stichocapsa naradaniensis* Matsuoka – Suzuki and Gawlick, p. 213, fig. 6.53, 6.54a, 6.54b. (detailed synonymy between 1994 and 2002)

2005 *Stichocapsa naradaniensis* Matsuoka – Missoni *et al.*, fig. 7.43; fig. 10.55; fig. 13.12.

2009 *Stichocapsa naradaniensis* Matsuoka – Suzuki and Gawlick, p. 186, fig. 5.57A, 5.57B, 5.58; fig. 6.38A, 6.38B, 6.42A, 6.42B.

2009 *Stichocapsa naradaniensis* Matsuoka – Ishida *et al.*, fig. 6.3; fig. 7.9.

Genus ***Japonocapsa*** Kozur, 1984

Type species: ***Tricolocapsa fusiformis*** Yao, 1979

Japonocapsa fusiformis (Yao, 1979)

(Plate 3, figs. 47, 48)

*1979 *Tricolocapsa? fusiformis* – Yao, p. 33, pl. 4, figs. 12–18; pl. 5, figs. 1–4.

1994 *Tricolocapsa? fusiformis* Yao – Goričan, p. 94, pl. 9, fig. 14. (detailed synonymy until 1993)

2009 *Tricolocapsa fusiformis* Yao – Suzuki and Gawlick, p. 183, fig. 5.41, 5.42A, 5.42B, 5.57A, 5.57B; fig. 6.13A, 6.13B, 6.14, 6.17.

Remarks: In case of depicted specimens, a basal dish-like appendage is torn off.

Japonocapsa tegiminis (Yao, 1979)

(Plate 3, fig. 41)

*1979 *Stichocapsa tegiminis* – Yao, p. 34, pl. 5, figs. 5–13.

2002 *Stichocapsa tegiminis* Yao – Nakae, fig. 3m.

2009 *Stichocapsa tegiminis* Yao – Suzuki and Gawlick, p. 186, fig. 5.55A, 5.55B.

2018 *Stichocapsa tegiminis* Yao – Gawlick *et al.*, fig. 12.24.

Remarks: *Japonocapsa tegiminis* differs from *Japonocapsa fusiformis* in having four chambers (exclusive of an appendage). A depicted specimen has a wide basal dish-like appendage.

Genus ***Tetracapsa*** Haeckel, 1881

*1881 *Tetracapsa* – Haeckel, p. 438.

pt. 1887 *Stichocapsa* – Haeckel, p. 1515.

pt. 1981 *Tetracapsa* Haeckel – Petrushevskaya, p. 185.

1993 *Tetracapsa* Haeckel – Widz and De Wever, p. 86.

2003b *Tetracapsa* Haeckel – Suzuki and Gawlick, p. 211.

2004b *Tetracapsa* Haeckel – Suzuki *et al.*, p. 387.

2014 *Tetracapsa* Haeckel – Suzuki *et al.*, p. 18.

Type species: ***Tetracapsa pilula*** Rüst, 1885. This type

species was subsequently designated by Campbell (1954) (Petrushevskaya, 1981).

Remarks: Morphotypes having latticed four-chambered test with closed base appeared frequently in Middle and Late Jurassic time. These morphotypes have been described under the genus *Sethocapsa* or *Stichocapsa*. However, their four-chambered feature is conspicuous to separate from two-chambered *Sethocapsa* and five- or more chambered *Stichocapsa*.

***Tetracapsa* sp. A** sensu Suzuki and Gawlick, 2003b

(Plate 3, figs. 1, 32)

1997 *Stichocapsa* sp. A sensu Matsuoka and Yao – Suzuki and Nakae, pl. 2, fig. 11.

2001 *Stichocapsa* sp. A sensu Matsuoka and Yao – Miyamoto *et al.*, pl. 7, fig. 8.

2002 *Arcanicapsa* sp. 2 – Hori *et al.*, pl. 8, fig. 24.

*2003b *Tetracapsa* sp. A – Suzuki and Gawlick, p. 211, fig. 5.24.

2004b *Tetracapsa* sp. A – Suzuki *et al.*, p. 387, fig. 5.1a, 5.1b.

2007 *Tetracapsa* sp. A sensu Suzuki and Gawlick – Auer *et al.*, fig. 6.92.

2009 *Tetracapsa* sp. A sensu Suzuki and Gawlick – Suzuki and Gawlick, p. 185, fig. 6.37A, 6.37B.

Genus ***Stichocapsa*** Haeckel, 1881

Type species: ***Stichocapsa jaspidea*** Rüst, 1885 (Campbell, 1954)

Stichocapsa cicciona Chiari, Marcucci and Praela, 2002
(Plate 1, fig. 28)

*2002 *Stichocapsa cicciona* – Chiari *et al.*, p. 76, pl. 3, figs. 8–12.

2007 *Stichocapsa cicciona* Chiari, Marcucci and Praela – Auer *et al.*, fig. 6.78.

2011 *Stichocapsa cicciona* Chiari, Marcucci and Praela – Gawlick *et al.*, fig. 3.31.

Remarks: This species has a test with a wide basal aperture, so that its generic attribution to the genus *Stichocapsa*, which has a closed base, is questionable. Here we tentatively attribute the species to the genus *Stichocapsa*.

Stichocapsa robusta Matsuoka, 1984

(Plate 1, fig. 23; Plate 2, fig. 26; Plate 3, fig. 25)

*1984 *Stichocapsa robusta* – Matsuoka, p. 146, pl. 1, figs. 6–13; pl. 2, figs. 7–12.

2007 *Stichocapsa robusta* Matsuoka – Auer *et al.*, fig. 6.81.

Genus ***Cyrtocapsa*** Haeckel, 1881

Type species: ***Cyrtocapsa ovalis*** Rüst, 1885

***Cyrtocapsa* sp. B**

(Plate 2, fig. 14)

2003 *Cyrtocapsa* sp. – Wegerer *et al.*, fig. 9.18.

Remarks: Four or five chambered tests with a robust horn. Proximal three or four segments make a conical portion, and a final segment exhibits a globous ball-form with larger pores than those of conical portion.

Genus *Fultacapsa* Ozvoldova, 1997

Type species: *Acotripus sphericus* Ozvoldova, 1988

Fultacapsa sphaerica (Ozvoldova, 1988)

(Plate 1, fig. 21)

* 1988 *Acotripus sphericus* – Ozvoldova, p. 376, pl. 5, figs. 1–5, 7.

1997 *Fultacapsa sphaerica* (Ozvoldova) – Ozvoldova and Frantova, p. 59, pl. 5, figs. 1, 2.

cf. 2003b *Acotripus* cf. *sphaericus* Ozvoldova – Suzuki and Gawlick, p. 191, fig. 5.29.

2010 *Fultacapsa sphaerica* (Ozvoldova) – Gawlick *et al.*, fig. 37B.1.

Remarks: A specimen from the Fludergraben section differs from specimens of Ozvoldova (1988) and Ozvoldova and Frantova (1997) in having weak constriction between a proximal part and a last globous segment.

Genus *Helvetocapsa* O’Dogherty, Goričan and Dumitrica, 2006

Type species: *Tricolocapsa matsuokai* Sashida, 1999

Helvetocapsa matsuokai (Sashida, 1999); emend. Suzuki and Gawlick, 2009

(Plate 1, fig. 19; Plate 3, fig. 39)

1930 *Cenellipsis* aff. *perspicua* Rüst – Heitzer, p. 388, pl. 27, fig. 11.

*1999 *Tricolocapsa matsuokai* – Sashida in Sashida *et al.*, p. 566, pl. 1, figs. 4, 5.

2003b *Tricolocapsa matsuokai* Sashida – Suzuki and Gawlick, p. 209, fig. 6.38. (detailed synonymy until 2002)

2006 *Helvetocapsa matsuokai* (Sashida) – O’Dogherty *et al.*, p. 452, pl. 7, figs. 19–24.

2009 *Helvetocapsa matsuokai* (Sashida) – Suzuki and Gawlick, p. 187, fig. 5.61A, 5.61B; fig. 6.40, 6.46A, 6.46B.

2018 *Helvetocapsa matsuokai* (Sashida) – Gawlick *et al.*, fig. 14.3.

Remarks: Sashida *et al.* (1999) described this species for the first time under the genus *Tricolocapsa*, a three-chambered genus. O’Dogherty *et al.* (2006) erected a new genus *Helvetocapsa* and attributed this species to their new genus, although the number of the segments of this species were not observed. Suzuki and Gawlick (2009) observed the inner structure of it with a transmitted light microscope and clarified that *Helvetocapsa matsuokai* has five segments.

Family EUCYRTIDIELLIDAE Takemura, 1986

Genus *Eucyrtidiellum* Baumgartner, 1984

Type species: *Eucyrtidium? unumaensis* Yao, 1979

Eucyrtidiellum circumperforatum Chiari, Marcucci and Prela, 2002

(Plate 1, fig. 33; Plate 3, fig. 3)

*2002 *Eucyrtidiellum? circumperforatum* – Chiari *et al.*, p. 65, pl. 1, figs. 2–9.

2007 *Eucyrtidiellum circumperforatum* Chiari, Marcucci and Prela – Auer *et al.*, fig. 6.22.

2007 *Eucyrtidiellum circumperforatum* Chiari, Marcucci and Prela – Gawlick *et al.*, fig. 8.12.

2009 *Eucyrtidiellum circumperforatum* Chiari, Marcucci and Prela – Suzuki and Gawlick, p. 189, fig. 5.64.

Eucyrtidiellum unumaense (Yao, 1979)

(Plate 1, figs. 35, 36; Plate 3, figs. 14, 37)

*1979 *Eucyrtidium? unumaensis* – Yao, p. 39, pl. 9, figs. 1–11.

1994 *Eucyrtidiellum unumaense* (Yao) – Goričan, p. 69, pl. 9, figs. 5, 6. (detailed synonymy until 1993)

2003a *Eucyrtidiellum unumaense* (Yao) – Suzuki and Gawlick, p. 119, pl. 1, fig. 9.

2003b *Eucyrtidiellum unumaense* (Yao) – Suzuki and Gawlick, p. 215, fig. 5.21. (detailed synonymy between 1994 and 2002)

2005 *Eucyrtidiellum unumaense* ssp. (Yao) – Missoni *et al.*, fig. 10.62.

2006 *Eucyrtidiellum unumaense* ssp. (Yao) – Gawlick *et al.*, fig. 8a.14; fig. 9a.8.

2009 *Eucyrtidiellum unumaense* (Yao) – Suzuki and Gawlick, p. 188, fig. 5.62.

Remarks: *Eucyrtidiellum unumaense* is subdivided into the three subspecies, i.e. *E. unumaense unumaense* Yao, *E. unumaense dentatum* Baumgartner and *E. unumaense pustulatum* Baumgartner (Baumgartner *et al.*, 1995a; Suzuki and Gawlick, 2003b). Because our specimens possess not so conspicuous features of ornamentation on upper abdomen surface to identify subspecies, we describe them only as *Eucyrtidiellum unumaense*.

Eucyrtidiellum ptyctum (Riedel and Sanfilippo, 1974)

(Plate 1, fig. 34; Plate 3, fig. 4)

*1974 *Eucyrtidium ptyctum* – Riedel and Sanfilippo, p. 778, pl. 5, fig. 7; pl. 12, fig. 14; non pl. 12, fig. 15.

2003b *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Suzuki and Gawlick, p. 218, fig. 6.26, 6.27. (detailed synonymy between 1998 and 2002)

2005 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Suzuki and Kuwahara, p. 65, pl. 2, fig. 17.

2005 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Missoni *et al.*, fig. 7.48; fig. 10.61; fig. 13.5.

2006 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Gawlick *et al.*, fig. 8.10; fig. 9.7.

2006 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Auer *et al.*, fig. 6.14.

- 2009 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Suzuki and Gawlick, p. 188, fig. 5.63.
2014 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Suzuki *et al.*, p. 19, pl. 5. fig. 4.
2018 *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) – Gawlick *et al.*, fig. 14.12.

***Eucyrtidiellum nodosum* Wakita, 1988**

(Plate 2, fig. 5; Plate 3, figs. 2, 38)

- *1988 *Eucyrtidiellum nodosum* – Wakita, p. 408, pl. 4, fig. 29; pl. 5, fig. 16.
2001 *Eucyrtidiellum nodosum* Wakita – Nishizono, pl. 2, fig. 8.
2003b *Eucyrtidiellum nodosum* Wakita – Suzuki and Gawlick, p. 217, fig. 6.30. (detailed synonymy between 1994 and 2003)
2007 *Eucyrtidiellum nodosum* Wakita – Auer *et al.*, fig. 6.23.
2009 *Eucyrtidiellum nodosum* Wakita – Auer *et al.*, fig. 9.22; cf. fig. 13.3.

6. Discussion – Radiolarian zonation for the lower Oxfordian and correlation

Because radiolarian fauna from the lower Oxfordian that is calibrated by ammonite has hitherto not known all over the world, the Fludergraben fauna is a key for understanding Oxfordian marker species of radiolarians. Previously proposed radiolarian zonations have a relatively long-lasting period for the Callovian and Oxfordian. For example, the U. A. Zone 8 of Baumgartner *et al.* (1995b) ranges in age from middle Callovian to early Oxfordian. Thus, we can distinguish the Oxfordian radiolarian fauna from the Callovian one to make a comparison of faunal contents between Callovian and Oxfordian. In this chapter we discuss the first appearance horizons of possible marker species for the lower Oxfordian with descriptions of the middle and upper Callovian sections in the Northern Calcareous Alps.

6.1 Radiolarians from the middle Callovian Brielgraben section

In the Brielgraben section of the Northern Calcareous Alps, the Klaus Formation yields middle Callovian ammonites (Krystyn, 1971) from strata that underlie a radiolarite succession. We have detected radiolarians from the radiolarite of the Brielgraben section, which are partly listed in Suzuki and Gawlick (2006, 2009). We show the revised inventory of radiolarians from the sample BT1 in the appendix 1.

6.2 Radiolarians from the lower part of the Knallalm-Neualm section – upper Callovian

From the lower part of the Knallalm-Neualm section, Auer *et al.* (2007) reported radiolarian assemblages containing *Williriedellum carpathicum* from the samples MR149 and MR175. Gawlick *et al.* (2009) invented a

new subzone of the *Zhamoidellum ovum* Zone, i.e. the *Williriedellum carpathicum* Subzone, based on the lower part of the Knallalm-Neualm section that is situated below the *Kilinora spiralis*-bearing radiolarite. If the first appearance horizon of *Kilinora spiralis* can be placed in the lowermost Oxfordian, the *Williriedellum carpathicum* Subzone is correlated to the upper Callovian (see discussion in the section 6.4). We show the lists of radiolarian species from samples MR149 and MR175 in the appendix 2 (Auer *et al.*, 2007).

6.3 Marker species for the base of Oxfordian

To compare the above-mentioned radiolarian faunas from the middle and upper Callovian with the Fludergraben fauna, it should be made clear what are the marker species for the base of Oxfordian (Fig. 7). We choose four species, i.e. *Kilinora spiralis* (Matsuoka), *Fultacapsa sphaerica* (Ozoldova), *Protunuma japonicus* Matsuoka and Yao and *Pseudoeucyrtis reticularis* Matsuoka and Yao. *Kilinora spiralis* occurs, however, very rare in the Northern Calcareous Alps. From the Fludergraben section, we found a single specimen from the sample D1023, identified as *Kilinora cf. spiralis*. It is poorly preserved, and its surface ornamentation is ambiguous (Plate 1, fig. 31). Other three marker species, *Fultacapsa sphaerica* (Ozoldova), *Protunuma japonicus* Matsuoka and Yao and *Pseudoeucyrtis reticularis* Matsuoka and Yao, also occur as a single specimen, respectively. *Pseudodictyomitra primitiva* Matsuoka and Yao has also potential to be a marker, but a forerunner occurrence is known from the upper Callovian of the Knallalm-Neualm section (Auer *et al.*, 2007). In the following three sections, we discuss ranges of these species in detail.

6.4 Stratigraphic range of *Kilinora spiralis* — lower Oxfordian to lower Kimmeridgian

There is stratigraphical discrepancy of the first appearance horizon of *Kilinora spiralis* between Matsuoka (1995) and Baumgartner and Matsuoka (1995) (*Stylocapsa? spiralis* in their publications), although both used the same marker species of calcareous nannoplankton, *Stephanolithion hexum* Rood and Barnars, 1972, as discussed in Suzuki *et al.* (2004a). Matsuoka (1995) placed the first appearance horizon of *Kilinora spiralis* to the upper Callovian, based on the last occurrence of *Stephanolithion hexum* in the core 124 of the Site 534 in the Blake Bahama Basin (DSDP Leg 76). The last occurrence horizon of *Stephanolithion hexum*, which is correlated to the boundary between the middle and upper Callovian with the calibration of magnetostratigraphy (Roth, 1983), lies just above the first appearance horizon of *Kilinora spiralis* (Baumgartner and Matsuoka, 1995). On the other side, Baumgartner and Matsuoka (1995) reinterpreted the horizon of the last occurrence of *Stephanolithion hexum* in the core as a preservational bias, and its horizon was correlated to the upper Bathonian to lower Callovian (U. A. Zone 7) on the basis of a radiolarian age assignment. This is a circular

not frequently, but until now we have detected it only from the Oxfordian to Kimmeridgian.

6. 6 First appearance horizon of *Protunuma japonicus*, *Pseudoeucyrtis reticularis* and *Pseudodictyomitra primitiva*

Protunuma japonicus, *Pseudoeucyrtis reticularis* and *Pseudodictyomitra primitiva* were first described from the Torinosu Group of the Island Shikoku and Kii-Yura areas, Southwest Japan (Matsuoka and Yao, 1985). Matsuoka and Yao (1985) inferred the age of the *Pseudodictyomitra primitiva*–*Pseudodictyomitra* sp. A assemblage to the Tithonian, and this assemblage acts as the type of the *Pseudodictyomitra primitiva* Zone in Japan. According to Matsuoka (1995) the *Pseudodictyomitra primitiva* Zone is defined as the zone between the last occurrence horizon of *Hsuum maxwelli* and the first occurrence horizon of *Pseudodictyomitra carpatica*. Our early Oxfordian samples yield *Hsuum maxwelli* commonly, so that the correlation of our samples to the *Pseudodictyomitra primitiva* Zone of Japan cannot be made. However, some constituents of the *Pseudodictyomitra primitiva*–*Pseudodictyomitra* sp. A assemblage can be found in our samples, i.e. *Pseudodictyomitra primitiva*, *Pseudoeucyrtis reticularis*, *Protunuma japonicus*, *Archaeodictyomitra apiarium*, *Archaeodictyomitra minoensis*, *Cinguloturris carpatica*, *Eucyrtidiellum ptyctum* and *Zhamoidellum ovum* (= *Tricolocapsa* sp. A). Thus, the *Pseudodictyomitra primitiva*–*Pseudodictyomitra* sp. A assemblage contains many species determined in the Fludergraben fauna. It should pay attention that the first appearance horizon of *Pseudodictyomitra primitiva* is in the upper Callovian, as demonstrated in Fig. 7. Important is the absence of *Hsuum maxwelli* as the criterion, whether a radiolarian assemblage is attributed to the *Pseudodictyomitra primitiva* Zone or not. As *Protunuma japonicus* and *Pseudoeucyrtis reticularis* were found in our Fludergraben samples, these two species appeared already in early Oxfordian time.

6. 7 Shift of some radiolarian age ranges

Stratigraphic ranges of several species of the Fludergraben fauna, which are so far known in the Callovian or lower, have to be prolonged into the lower Oxfordian. These species are as follows (with previous age assignment).

Dictyomitrella kamoensis (U. A. Zone 3–7: Baumgartner *et al.*, 1995b)

Eucyrtidiellum circumperforatum (U. A. Zone 5–7: Chiari *et al.*, 2002)

Helvetocapsa matsukakai (*Striatojaponocapsa plicarum* Zone – upper Bajocian-lower Bathonian: Sashida *et al.*, 1999; U. A. Zone 6: O’Dogherty *et al.*, 2006)

Hsuum baloghi (lower *Unuma echinatus* Zone – Aalenian to lower Bajocian: Grill and Kozur, 1986)

Japonocapsa fusiformis (U. A. Zone 3–5: Baumgartner *et al.*, 1995b)

Neorelumbra skenderbegi (U. A. Zone 5–7: Chiari *et al.*

2002)

Protunuma fusiformis (Bajocian: Yao, 1997)

Protunuma lanosus (Callovian: Suzuki and Gawlick, 2003a)

Stichocapsa cicciona (U. A. Zone 5–7: Chiari *et al.* 2002)

Stichocapsa robusta (U. A. Zone 5–7: Baumgartner *et al.*, 1995b)

Japonocapsa tegiminis (Bajocian: Yao, 1979, 1997)

Tricolocapsa tetragona (upper *Striatojaponocapsa plicarum* Zone to lower *Striatojaponocapsa conexa* Zone – Bathonian: Matsuoka, 1995)

Unuma gordus (as *Unuma* sp. A, U. A. Zone 4–6: Baumgartner *et al.*, 1995b)

Unuma typicus (Bajocian: Yao, 1997; Callovian: Suzuki and Gawlick, 2009)

Among them we make comments on two important species, i.e. *Protunuma lanosus* and *Tricolocapsa tetragona*. *Protunuma lanosus*, which is the index species of the Callovian *Protunuma lanosus* Subzone of the *Zhamoidellum ovum* Zone of Suzuki and Gawlick (2003a), extends its range upwards into the Oxfordian. Consequently, the previous definition of the base of the *Williriedellum dierschei* Subzone, the last occurrence horizon of *Protunuma lanosus*, has to be changed. Another important species is *Tricolocapsa tetragona*, which was considered having a short stratigraphic range within the Bathonian (Matsuoka, 1983, 1995). As we demonstrate by the Fludergraben fauna, *Tricolocapsa tetragona* occurs in the lower Oxfordian strata. This stratigraphic range prolongation is supported by the occurrence of *Tricolocapsa tetragona* in the Torinosu-type limestone of east Shikoku, Japan (Ishida, 1994). This fauna yields also *Kilinora spiralis*, suggesting an Oxfordian age. Although Ishida (1994) mentioned that the stratigraphic range of *Tricolocapsa tetragona* was not consistent with those of other early Late Jurassic radiolarian species, its occurrence is now regarded not as an exception but as the reflection of its real stratigraphic range.

6. 8 Redefinition of the *Williriedellum dierschei* Zone

In the Jurassic radiolarian zonation of the Northern Calcareous Alps the *Williriedellum dierschei* Subzone of the *Zhamoidellum ovum* Zone was first established by Suzuki and Gawlick (2003a) as the partial-range zone of the species *Williriedellum dierschei* Suzuki and Gawlick, and it is defined by the last occurrence horizon of *Protunuma lanosus* for the base and the last occurrence horizon of *Eucyrtidiellum unumaense* for the top, indicating an early to middle Oxfordian age (Auer *et al.*, 2007). However, as we demonstrate here, *Protunuma lanosus* occurs also in the lower Oxfordian Fludergraben section, so that the base of the *Williriedellum dierschei* Subzone lies within the lower Oxfordian or higher, if we follow the above-mentioned definition. Our purpose of the radiolarian zonation is to distinguish the lower Oxfordian radiolarian zone from the Callovian one. And to make an age determination, it is better to take a positive criterion, i.e. the first appearance

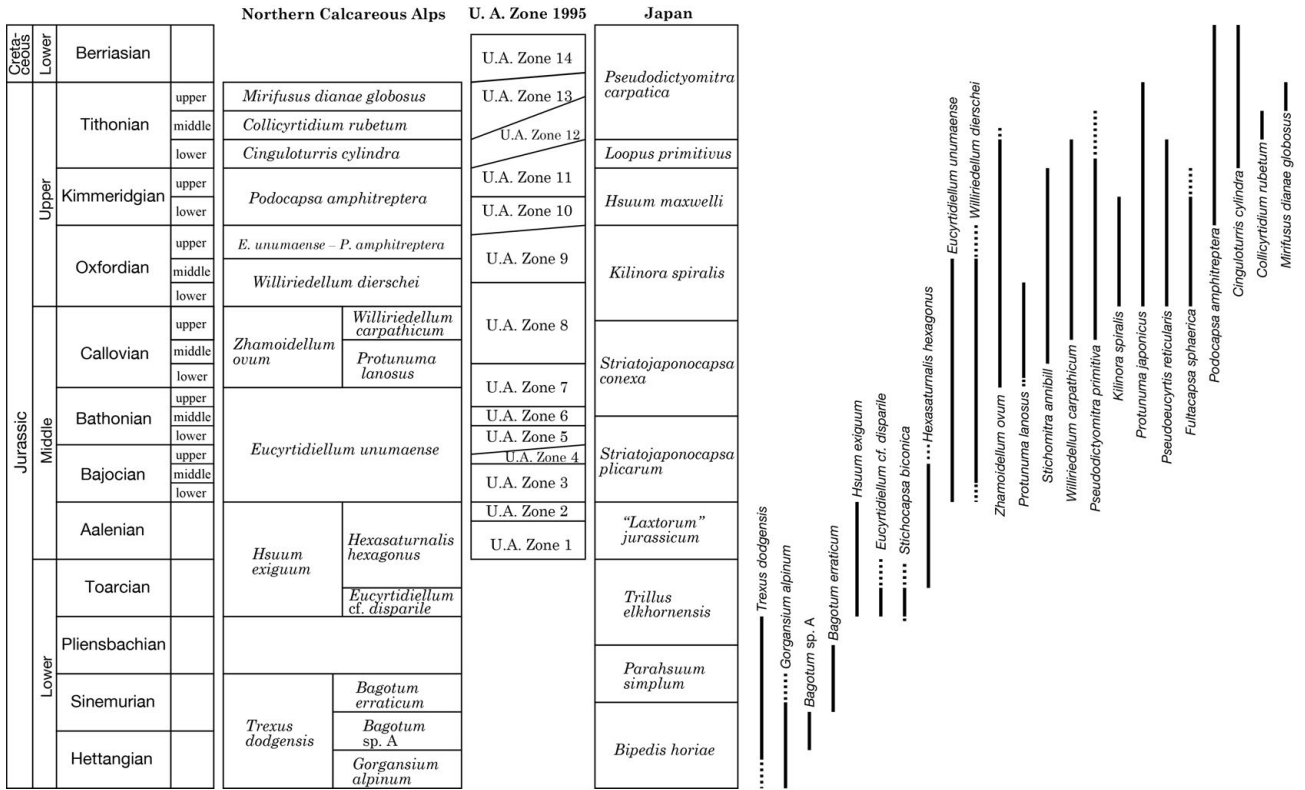


Fig. 8 Modified Jurassic radiolarian zonation for the Northern Calcareous Alps according to Suzuki and Gawlick (2003a), Steiger (1992), Gawlick *et al.* (2009) and this study. The U. A. Zone 1995 for the Western Tethyan realm of Baumgartner *et al.* (1995b) and the Japanese zonation of Matsuoka and Ito (2019) are shown on the side for comparison.

horizon, rather than a negative one, i.e. the last occurrence horizon. In this context, here we take the first appearance horizon of *Protunuma japonicus* as the definition of the base of the *Williriedellum dierschei* Subzone. *Fultacapsa sphaerica*, *Pseudoeucyrtis reticularis* and *Kilinora spiralis* are the subordinate marker species of this zone. Suzuki and Gawlick (2003a) and Gawlick *et al.* (2009) put it to the Subzone in the *Zhamoidellum ovum* Zone, because the faunal content of the Callovian-Oxfordian is very similar and no clear distinction was shown at that time. Because we can discriminate some early Oxfordian marker species among Callovian-Oxfordian-lasting species, we make this subzone ranked up as a zone apart from the *Zhamoidellum ovum* Zone of the Callovian, namely the *Williriedellum dierschei* Zone (Fig. 8). According as this, the overlying *Eucyrtidellum unumaense* – *Podocapsa amphitreptera* Interval Zone for the upper Oxfordian (Suzuki and Gawlick, 2003a) is also separated from the *Zhamoidellum ovum* Zone and it is here redefined as an independent zone (Fig. 8). And the upper limit of the *Williriedellum carpathicum* Subzone in the *Zhamoidellum ovum* Zone is also here emended as the first appearance horizon of *Protunuma japonicus*.

7. Conclusion

(1) 37 genera, 67 species and 2 subspecies of radiolarians

are systematically described from the lower Oxfordian Fludergraben section that is calibrated by ammonites.

(2) Four radiolarian species have a potential to be marker for the base of Oxfordian. These are *Kilinora spiralis* Matsuoka, *Fultacapsa sphaerica* (Ozoldova), *Protunuma japonicus* Matsuoka and Yao and *Pseudoeucyrtis reticularis* Matsuoka and Yao.

(3) The *Williriedellum dierschei* Zone is here redefined as the lower-middle Oxfordian radiolarian zone of the Northern Calcareous Alps.

(4) In the systematic part of radiolarians we have emended two genera and one species diagnoses, and redesignated of the type species of the genus *Loopus*.

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Plate 1 Scanning electron micrographs of radiolarians from the samples D1051 (1–2), D1023 (3–34) and D1024 (35–36), basal horizons of the Fludergraben section, Austria. A 50 µm scale bar applies to all photos.

1. *Williriedellum dierschei* Suzuki and Gawlick, 2004
2. *Archaeodictyomitra apiarium* (Rüst, 1885)
3. *Archaeospongoprunum* cf. *elegans* Wu, 1993
4. *Hsuum maxwelli* Pessagno, 1977a
5. *Loopus doliolum* Dumitrica, 1997
6. *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985
7. *Parahsuum* sp. S sensu Matsuoka, 1986
8. *Tritrabs* cf. *exotica* (Pessagno, 1977a)
9. *Neorelumbra skenderbegi* Chiari *et al.*, 2002
10. *Archaeodictyomitra rigida* Pessagno 1977a
11. *Archaeodictyomitra apiarium* (Rüst, 1885)
12. *Hsuum brevicostatum* (Ozoldova, 1975)
13. *Stichomitra annibill* Kocher, 1981
14. *Acanthocircus* cf. *suboblongus* (Yao, 1972)
15. *Cinguloturris carpatica* Dumitrica, 1982
16. *Takemuraella hungarica* (Kozur, 1985)
17. *Archaeodictyomitra mirabilis* Aita, 1987
18. *Takemuraella hexagonata* (Heitzer, 1930)
19. *Helvetocapsa matsukai* (Sashida, 1999)
20. *Zhamoidellum ovum* Dumitrica, 1970
21. *Fultacapsa sphaerica* (Ozoldova, 1988)
22. *Tricolocapsa undulata* (Heitzer, 1930)
23. *Stichocapsa robusta* Matsuoka, 1984
24. *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006
25. *Zhamoidellum ventricosum* Dumitrica, 1970
26. *Gongylothorax favosus favosus* Dumitrica, 1970
27. *Williriedellum dierschei* Suzuki and Gawlick, 2004
28. *Stichocapsa ciccionea* Chiari *et al.*, 2002
29. *Unuma gordus* Hull, 1997
30. *Protunuma japonicus* Matsuoka and Yao, 1985
31. *Kilinora* cf. *spiralis* (Matsuoka, 1982)
32. *Tricolocapsa tetragona* Matsuoka, 1983
33. *Eucyrtidiellum circumperforatum* Chiari *et al.*, 2002
34. *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974)
- 35–36. *Eucyrtidiellum unumaense* (Yao, 1979)

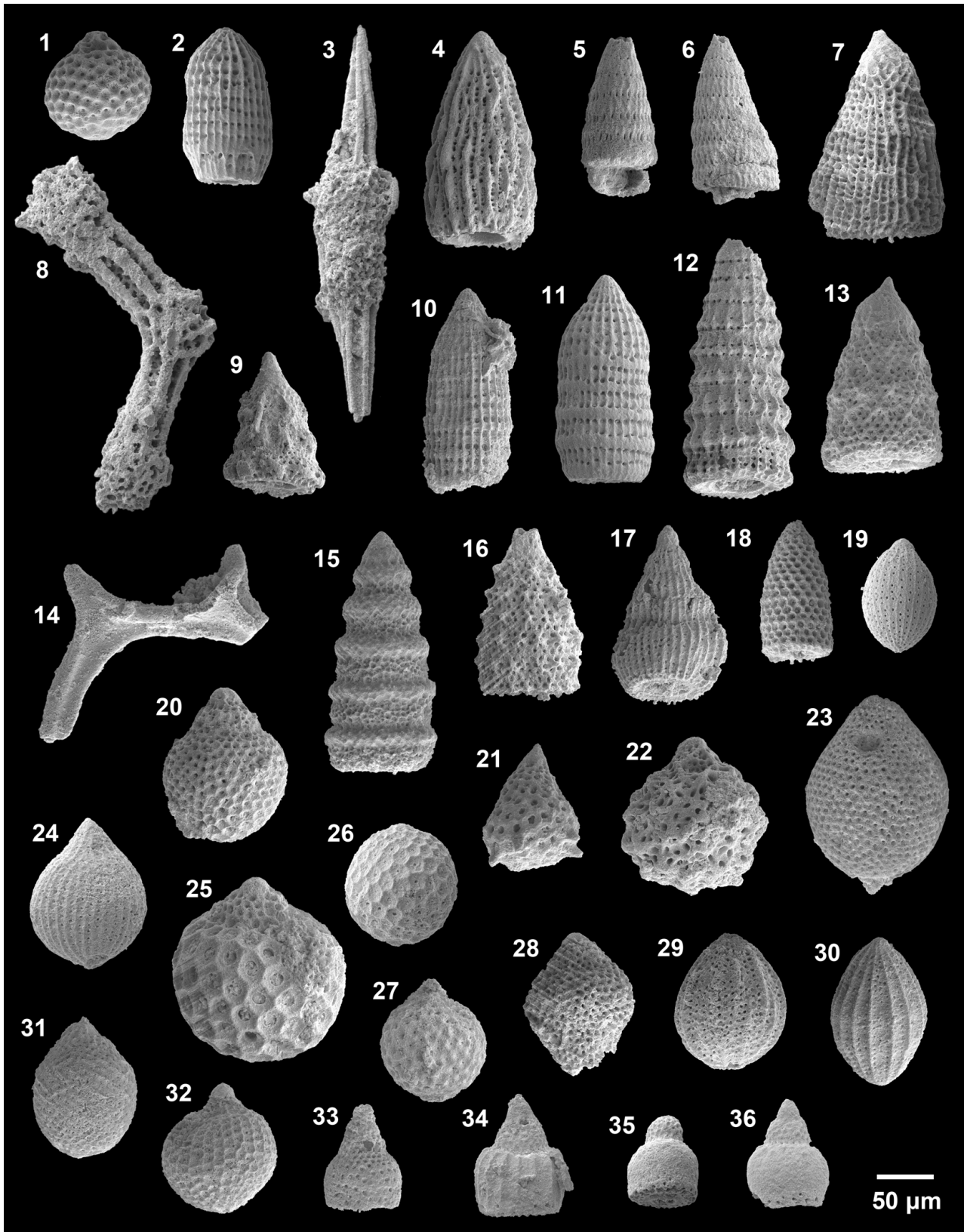


Plate 2 Scanning electron micrographs of radiolarians from the samples D1024 (1–16) and D1052 (17–41), basal horizons of the Fludergraben section, Austria. A 50 μm scale bar applies to all photos.

1. *Archaeospongoprimum* cf. *elegans* Wu, 1993
2. *Parahsuum* sp. S sensu Matsuoka, 1986
3. *Cinguloturris carpatica* Dumitrica, 1982
4. *Dictyomitrella kamoensis* Mizutani and Kido, 1983
5. *Eucyrtidiellum nodosum* Wakita, 1988
6. *Unuma typicus* Ichikawa and Yao, 1976
7. *Gongylothorax favosus favosus* Dumitrica, 1970
8. *Protunuma lanosus* Ozvoldova, 1996
9. *Tricolocapsa tetragona* Matsuoka, 1983
10. *Williriedellum marcucciae* Cortese, 1993
11. *Striatojaponocapsa riri* O'Dogherty *et al.*, 2006
12. *Williriedellum dierschei* Suzuki and Gawlick, 2004
13. *Zhamoidellum ovum* Dumitrica, 1970
14. *Cyrtocapsa* sp. B
15. *Hsuum maxwelli* Pessagno, 1977a
16. *Williriedellum crystallinum* Dumitrica, 1970
17. *Ristola altissima* (Rüst, 1885)
18. *Pseudoeucyrtis reticularis* Matsuoka and Yao, 1985
19. *Stichomitra annibill* Kocher, 1981
20. *Williriedellum carpathicum* Dumitrica, 1970
21. *Striatojaponocapsa naradaniensis* (Matsuoka, 1984)
22. *Tritrabs exotica* (Pessagno, 1977a)
23. *Gongylothorax favosus oviformis* Suzuki and Gawlick, 2009
24. *Parahsuum* sp. S sensu Matsuoka, 1986
25. *Stichomitra annibill* Kocher, 1981
26. *Stichocapsa robusta* Matsuoka, 1984
27. *Tricolocapsa undulata* (Heitzer, 1930)
28. *Gongylothorax favosus favosus* Dumitrica, 1970
29. *Loopus doliolum* Dumitrica, 1997
30. *Williriedellum dierschei* Suzuki and Gawlick, 2004
31. *Podobursa nodosa* (Chiari *et al.*, 2002)
32. *Williriedellum sujkowski* Widz and De Wever, 1993
33. *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006
34. *Stichomitra* sp. A sensu Baumgartner *et al.*, 1995a
35. *Zhamoidellum ovum* Dumitrica, 1970
36. *Williriedellum crystallinum* Dumitrica, 1970
37. *Hsuum brevicostatum* (Ozvoldova, 1975)
38. *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985
39. *Striatojaponocapsa conexa* (Matsuoka, 1983)
40. *Striatojaponocapsa riri* O'Dogherty *et al.*, 2006
41. *Unuma gordus* Hull, 1997

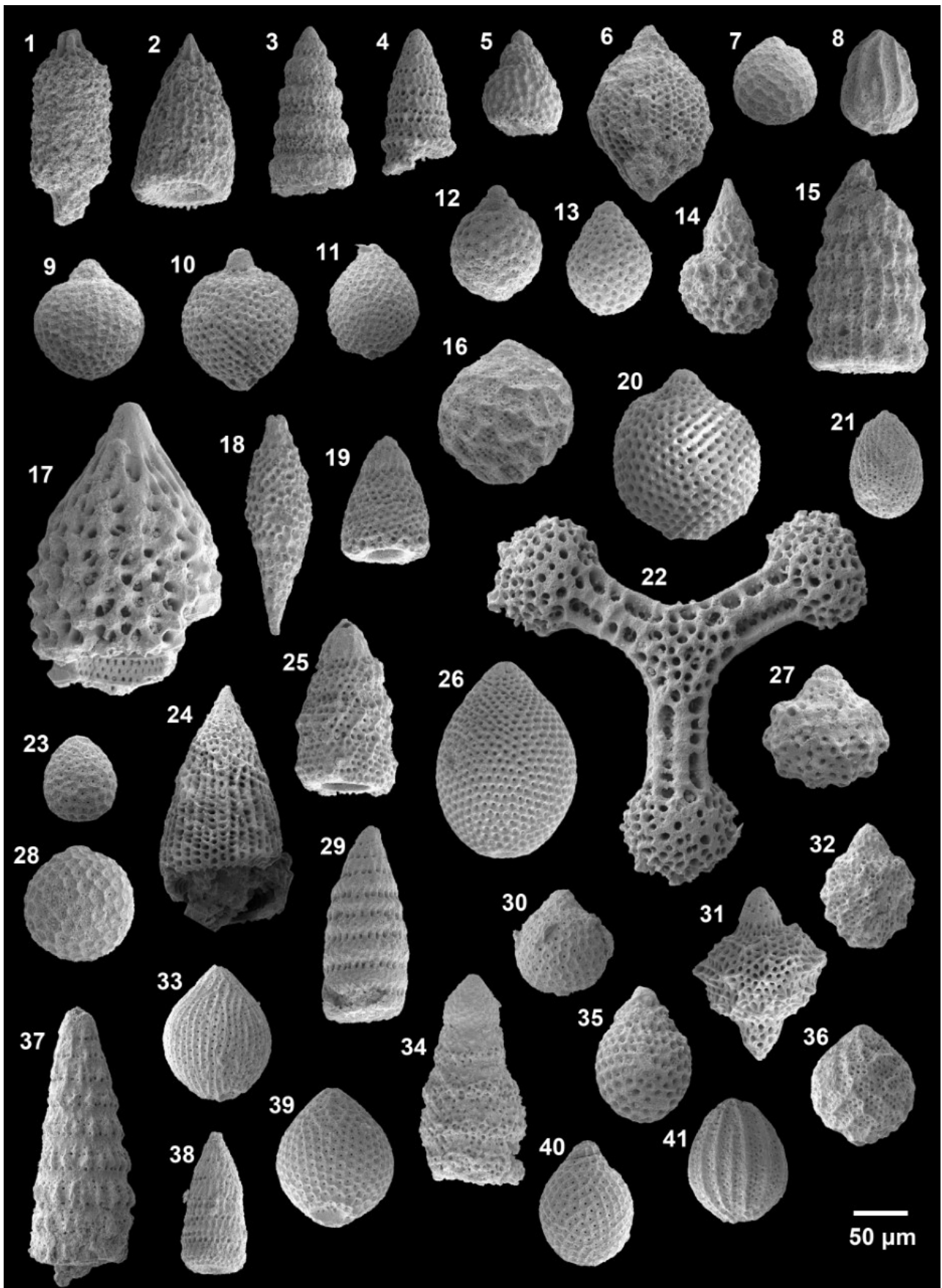
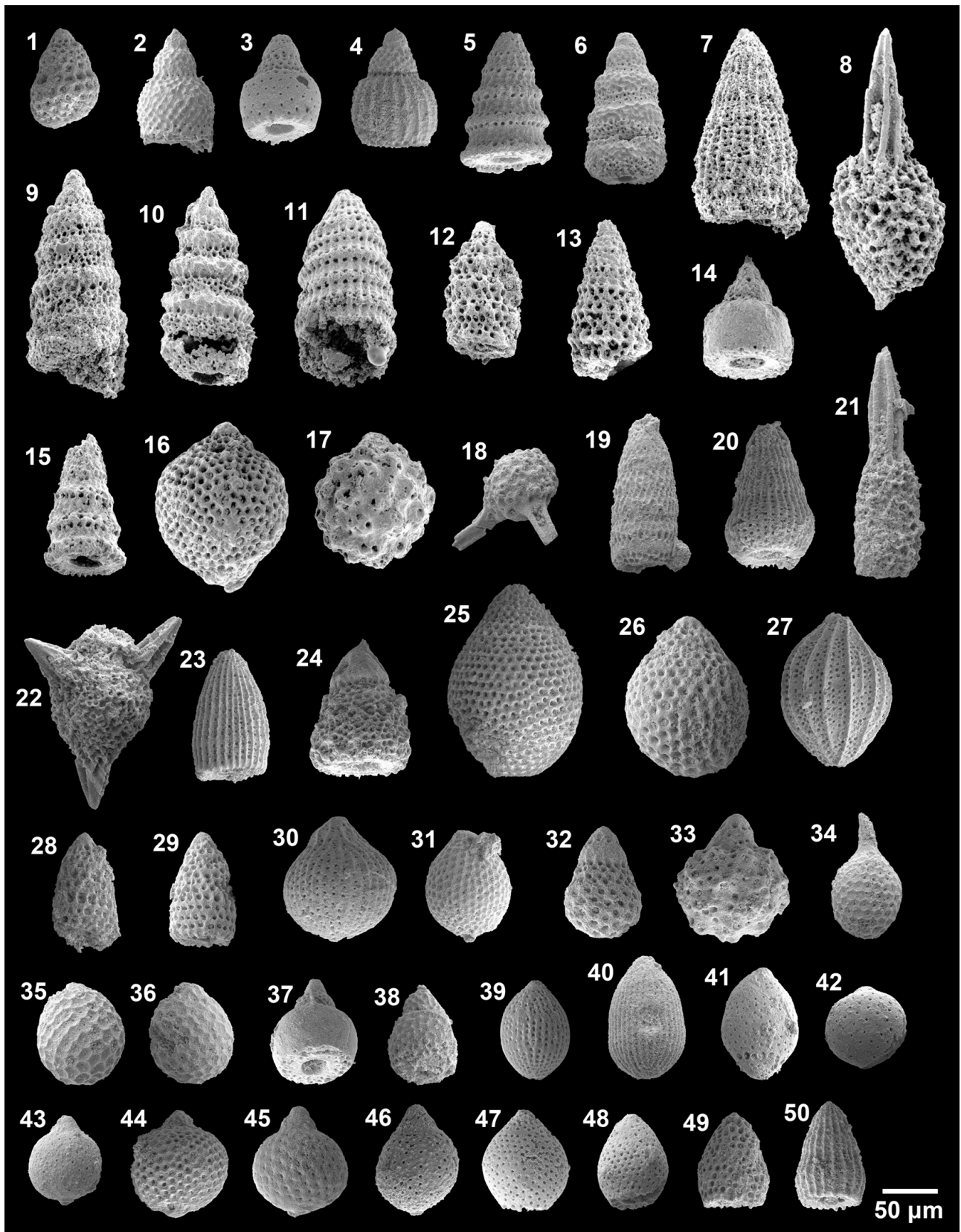


Plate 3 Scanning electron micrographs of radiolarians from the samples D1052 (1–6), EW146 (7–17) and D1025 (18–45), basal horizons of the Fludergraben section, Austria. A 50 µm scale bar applies to all photos.

1. *Tetracapsa* sp. A sensu Suzuki and Gawlick, 2003b
2. *Eucyrtidiellum nodosum* Wakita, 1988
3. *Eucyrtidiellum circumperforatum* Chiari *et al.*, 2002
4. *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974)
5. *Dictyomitrella kamoensis* Mizutani and Kido, 1983
6. *Cinguloturris carpatica* Dumitrica, 1982
7. *Hsuum baloghi* Grill and Kozur, 1986
8. *Archaeospongoprunum* cf. *imlayi* Pessagno, 1977a
9. *Cinguloturris carpatica* Dumitrica, 1982
10. *Cinguloturris primorika* Kemkin and Taketani, 2004
11. *Archaeodictyomitra minoensis* (Mizutani, 1981)
12. *Takemuraella hungarica* (Kozur, 1985)
13. *Parvicingula spinata* (Vinassa, 1899)
14. *Eucyrtidiellum unumaense* (Yao, 1979)
15. *Dictyomitrella* cf. *kamoensis* Mizutani and Kido, 1983
16. *Williriedellum carpathicum* Dumitrica, 1970
17. *Tricolocapsa undulata* (Heitzer, 1930)
18. *Saitoum pagei* Pessagno, 1977a
19. *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985
20. *Archaeodictyomitra mirabilis* Aita, 1987
21. *Archaeospongoprunum* cf. *elegans* Wu, 1993
22. *Spongotropus* sp. D sensu Suzuki and Gawlick, 2003b
23. *Archaeodictyomitra patricki* Kocher, 1981
24. *Stichomitra annibill* Kocher, 1981
25. *Stichocapsa robusta* Matsuoka, 1984
26. *Zhamoidellum ovum* Dumitrica, 1970
27. *Protunuma fusiformis* Ichikawa and Yao, 1976
- 28–29. *Takemuraella hexagonata* (Heitzer, 1930)
30. *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006
31. *Striatojaponocapsa conexa* (Matsuoka, 1983)
32. *Tetracapsa* sp. A sensu Suzuki and Gawlick, 2003b
33. *Tricolocapsa undulata* (Heitzer 1930)
34. *Stylocapsa oblongula* Kocher, 1981
35. *Gongylothorax favosus favosus* Dumitrica, 1970
36. *Gongylothorax favosus oviformis* Suzuki and Gawlick, 2009
37. *Eucyrtidiellum unumaense* (Yao, 1979)
38. *Eucyrtidiellum nodosum* Wakita, 1988
39. *Helvetocapsa matsuokai* (Sashida, 1999)
40. *Striatojaponocapsa naradaniensis* (Matsuoka, 1984)
41. *Japonocapsa tegiminis* (Yao, 1979)
42. *Gongylothorax* sp. C sensu Suzuki and Gawlick, 2003b
43. *Williriedellum* sp. C sensu Gawlick *et al.*, 2018
44. *Williriedellum marcucciae* Cortese, 1993
45. *Williriedellum dierschei* Suzuki and Gawlick, 2004
46. *Praewilliriedellum* aff. *spinosum* Kozur, 1984
- 47–48. *Japonocapsa fusiformis* (Yao, 1979)
49. *Droltus galerus* Suzuki, 1995b
50. *Archaeodictyomitra sixi* Yang, 1993



Appendix 1

Updated inventory of radiolarian species from the sample BT1 of the middle Callovian Brielgraben section.

BT1: *Gorgansium xigazeense* Wu, 1993, *Stylosphaera* cf. *lanceola* Parona, 1890, *Archaeodictyomitra amabilis* Aita, 1987, *Archaeodictyomitra* cf. *minoensis* (Mizutani, 1981), *Archaeodictyomitra mitra* Dumitrica, 1997, *Archaeodictyomitra rigida* Pessagno, 1977a, *Cinguloturris carpatica* Dumitrica, 1982, *Dictyomitrella kamoensis* Mizutani and Kido, 1983, *Droltus galerus* Suzuki, 1995b, *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974), *Eucyrtidiellum semifactum* Nagai and Mizutani, 1990, *Eucyrtidiellum takemurai* Hull, 1997, *Eucyrtidiellum unumaense dentatum* Baumgartner, 1995 in Baumgartner *et al.* (1995a), *Eucyrtidiellum unumaense unumaense* (Yao, 1979), *Gongylothorax favosus* Dumitrica, 1970, *Gongylothorax* sp. C sensu Suzuki and Gawlick (2003b), *Guexella nudata* (Kocher, 1980) in Baumgartner *et al.* (1980), *Helvetocapsa matsukoi* (Sashida, 1999), *Hiscocapsa magnipora* (Chiari *et al.*, 2002), *Hiscocapsa* cf. *acuta* Hull, 1997, *Hsuum brevicostatum* (Ozoldova, 1975), *Hsuum maxwelli* Pessagno, 1977a, *Japonocapsa* aff. *fusiformis* (Yao, 1979), *Loopus doliolum* Dumitrica, 1997, *Parvifavus* sp. A, *Praezhamoidellum buekkense* Kozur, 1984, *Praezhamoidellum* cf. *parvipora* (Tan, 1927), *Protunuma lanosus* Ozoldova, 1996, *Quarticella ovalis* Takemura, 1986, *Ristola procera* (Pessagno, 1977a), *Saitoum levium* De Wever, 1981, *Spongocapsula krahsteinensis* Suzuki and Gawlick, 2004, *Stichocapsa convexa* Yao, 1979, *Stichocapsa robusta* Matsuoka, 1984, *Striatojaponocapsa conexa* (Matsuoka, 1983), *Striatojaponocapsa naradaniensis* (Matsuoka, 1984), *Stylocapsa oblongula* Kocher, 1981, *Syringocapsa levis* (Hori, 1999), *Tetracapsa himedaruma* (Aita, 1987), *Tetracapsa* sp. A sensu Suzuki and Gawlick (2003b), *Theocapsomma* cf. *costata* Chiari *et al.*, 2002, *Theocapsomma cucurbitiformis* Baumgartner, 1995 in Baumgartner *et al.* (1995a), *Tricolocapsa tetragona* Matsuoka, 1983, *Tricolocapsa undulata* (Heitzer, 1930), *Tricolocapsa* sp. C sensu Auer *et al.* (2007), *Tricolocapsa* sp. M sensu Baumgartner *et al.* (1995a), *Takemuraella hexagonata* (Heitzer, 1930), *Takemuraella hungarica* (Kozur, 1985), *Unuma gordus* Hull, 1997, *Williriedellum crystallinum* Dumitrica, 1970, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Williriedellum marcucciae* Cortese, 1993, *Zhamoidellum ovum* Dumitrica, 1970.

Appendix 2

The inventory of radiolarian species from the samples MR149 and MR175 of the lower part of the Knallalm-Neualm section, described by Auer *et al.* (2007). The lower part of the Knallalm-Neualm section is the stratum typicum of the *Williriedellum carpaticum* Subzone in the *Zhamoidellum ovum* Zone.

MR149: *Acanthocircus* cf. *suboblongus* (Yao, 1972), *Alievium* sp., *Archaeodictyomitra amabilis* Aita, 1987, *Archaeodictyomitra apiarium* (Rüst, 1885), *Archaeodictyomitra* cf. *minoensis* (Mizutani, 1981), *Archaeodictyomitra mitra* Dumitrica, 1997, *Archaeodictyomitra rigida* Pessagno, 1977a, *Archaeospongoprunum* sp. (this specimen is reidentified here as *Archaeospongoprunum* cf. *elegans* Wu, 1993), *Cinguloturris carpatica* Dumitrica, 1982, *Dictyomitrella kamoensis* Mizutani and Kido, 1983, *Emiluvia* cf. *bisellea* Danelian, 1995, *Eucyrtidiellum* cf. *circumperforatum* Chiari *et al.*, 2002, *Eucyrtidiellum nodosum* Wakita, 1988, *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974), *Eucyrtidiellum unumaense pustulatum* Baumgartner, 1984, *Eucyrtidiellum unumaense* ssp. (Yao, 1979), *Gongylothorax favosus favosus* Dumitrica, 1970, *Gongylothorax favosus oviformis* Suzuki and Gawlick, 2009, *Gorgansium* sp., *Homoeoparonaella* sp., *Hsuum brevicostatum* (Ozoldova, 1975), *Hsuum hisuikyoense* Isozaki and Matsuda, 1985, *Hsuum maxwelli* Pessagno, 1977a, *Lithocampium* sp. C sensu Auer *et al.* (2007), *Loopus doliolum* Dumitrica, 1997, *Neorelumbra skenderbegi* Chiari *et al.*, 2002, *Napora* sp., *Paronaella* sp., *Parvicingula cappa* Cortese, 1993, *Parvifavus* sp., *Podobursa triacantha* (Fischli, 1916), *Praewilliriedellum spinosum*

Kozur, 1984, *Praezhamoidellum* cf. *parvipora* (Tan, 1927), *Protunuma lanosus* Ozvoldova, 1996, *Pseudodictyomitra primitiva* Matsuoka and Yao, 1985, *Stylosphaera lanceola* Parona, 1890, *Spongocapsula krahsteinensis* Suzuki and Gawlick, 2004, *Stichocapsa convexa* Yao, 1979, *Stichocapsa robusta* Matsuoka, 1984, *Stichomitra* sp., *Striatojaponocapsa conexa* (Matsuoka, 1983), *Striatojaponocapsa synconexa* O'Dogherty *et al.*, 2006, *Stylocapsa oblongula* Kocher, 1981, *Syringocapsa lata* Yang, 1993, *Syringocapsa suavis* Yang, 1993, *Tetracapsa* sp. A sensu Suzuki and Gawlick (2003b), *Tetraditryma* sp., *Theocapsomma bicornis* Baumgartner, 1995 in Baumgartner *et al.* (1995a) *Theocapsomma cordis* Kocher, 1981, *Theocapsomma costata* Chiari *et al.*, 2002, *Tricolocapsa leiostraca* (Foreman, 1973), *Tricolocapsa undulata* (Heitzer, 1930), *Tricolocapsium* sp. A sensu Auer *et al.* (2007), *Tricolocapsium* sp. B sensu Auer *et al.* (2007), *Tritrabs* cf. *casmaliaensis* (Pessagno, 1977a), *Tritrabs rhododactylus* Baumgartner, 1980, *Takemuraella hexagonata* (Heitzer, 1930), *Takemuraella hungarica* (Kozur, 1985), *Unuma gordus* Hull, 1997, *Williriedellum carpathicum* Dumitrica, 1970, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Williriedellum marcucciae* Cortese, 1993, *Xitus magnus* Baumgartner, 1995 in Baumgartner *et al.* (1995a), *Zhamoidellum ovum* Dumitrica, 1970.

MR175: *Amphipyndax* cf. *tsunoensis* Aita, 1987, *Archaeodictyomitra* cf. *apiarium* (Rüst, 1885), *Archaeodictyomitra minoensis* (Mizutani, 1981), *Archaeodictyomitra mitra* Dumitrica, 1997, *Archaeodictyomitra rigida* Pessagno, 1977a, *Archaeodictyomitra sixi* Yang, 1993, *Cinguloturris carpatica* Dumitrica, 1982, *Crucella* sp., *Droltus galerus* Suzuki, 1995b, *Eucyrtidiellum nodosum* Wakita, 1988, *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo, 1974), *Eucyrtidiellum semifactum* Nagai and Mizutani, 1990, *Eucyrtidiellum unumaense dentatum* Baumgartner, 1995 in Baumgartner *et al.* (1995a), *Eucyrtidiellum unumaense pustulatum* Baumgartner, 1984, *Eucyrtidiellum unumaense unumaense* (Yao, 1979), *Gongylothorax favosus oviformis* Suzuki and Gawlick, 2009, *Gongylothorax* aff. *siphonifer* Dumitrica, 1970, *Gorgansium* cf. *morganense* Pessagno and Blome, 1980, *Helvetocapsa matsukakai* (Sashida, 1999), *Hiscocapsa* cf. *hexagona* (Hori, 1999), *Homoeoparonaella* cf. *elegans* (Pessagno, 1977a), *Hsuum brevicostatum* (Ozvoldova, 1975), *Hsuum* cf. *exiguum* Yeh and Cheng, 1996, *Hsuum maxwelli* Pessagno, 1977a, *Lithocampium matsukakai* (Hull, 1997), *Loopus doliolum* Dumitrica, 1997, *Neorelumbra skenderbegi* Chiari *et al.*, 2002, *Parahsuum levicostatum* Takemura, 1986, *Parahsuum* aff. *simplum* Yao, 1982, *Parahsuum* sp. S sensu Matsuoka (1986), *Parvicingula cappa* Cortese, 1993, *Parvicingula spinata* (Vinassa, 1899), *Parvicingula dhimenaensis* Baumgartner, 1984, *Parvifavus wallacheri* (Grill and Kozur, 1986), *Parvifavus* sp. A sensu Auer *et al.* (2007), *Praewilliriedellum spinosum* Kozur, 1984, *Protunuma lanosus* Ozvoldova, 1996, *Protunuma ochiensis* Matsuoka, 1983, *Pseudodictyomitra venusta* (Chiari *et al.*, 1997) [= *Pseudodictyomitra* sp. D sensu Matsuoka and Yao (1985)], *Pseudoeucyrtis* sp. J sensu Baumgartner *et al.* (1995a), *Pseudodictyomitrella spinosa* Grill and Kozur, 1986, *Quarticella levis* Takemura, 1986, *Quarticella ovalis* Takemura, 1986, *Saitoum* cf. *pagei* Pessagno, 1977a, *Stylosphaera lanceola* Parona, 1890, *Spongotripus* sp. E, *Stichocapsa* aff. *biconica* Matsuoka, 1991, *Stichomitra* cf. *annibill* Kocher, 1981, *Stichomitra takanoensis* Aita, 1987, *Stylocapsa tecta* Matsuoka, 1983, *Striatojaponocapsa* cf. *conexa* (Matsuoka, 1983), *Striatojaponocapsa naradaniensis* (Matsuoka, 1984), *Striatojaponocapsa riri* O'Dogherty *et al.*, 2006 [= *Tricolocapsa* sp. A sensu Goričan (1994)], *Takemuraella hexagonata* (Heitzer, 1930), *Takemuraella hungarica* (Kozur, 1985), *Tetracapsa* sp. A sensu Suzuki and Gawlick (2003b), *Tetracapsa* sp. C sensu Auer *et al.* (2007), *Theocapsomma cordis* Kocher, 1981, *Theocapsomma* cf. *cucurbiformis* Baumgartner, 1995, *Tricolocapsa leiostraca* (Foreman, 1973), *Tricolocapsa undulata* (Heitzer, 1930), *Williriedellum* sp. C [= *Tricolocapsa* sp. A sensu Ozvoldova (1992)], *Tricolocapsa* sp. C sensu Auer *et al.* (2007), *Tritrabs* cf. *casmaliaensis* (Pessagno, 1977a), *Tritrabs simplex* Kito and De Wever, 1992, *Unuma gordus* Hull, 1997, *Williriedellum carpathicum* Dumitrica, 1970, *Williriedellum dierschei* Suzuki and Gawlick, 2004, *Williriedellum marcucciae* Cortese, 1993, *Xitus magnus* Baumgartner, 1995 in Baumgartner *et al.* (1995a), *Zhamoidellum kozuri* (Hull, 1997), *Zhamoidellum ovum* Dumitrica, 1970, *Zhamoidellum ventricosum* Dumitrica, 1970.

アンモナイト層準直上のジュラ系上部統基底フルダーグラール部層から産した
放散虫化石（北部石灰アルプス，オーストリア）

鈴木 寿志・ハンス・ユルゲン ガウリック

要 旨

北部石灰アルプスのフルダーグラール部（オーストリア）において、アンモナイトで年代決定されたクラウス層石灰岩（ジュラ系中部統最上部）の直上に累重する放散虫岩から放散虫群集を記載した。この放散虫群集はジュラ系上部統最下部（Oxfordian）からのものであり、放散虫生層序を考える上で重要である。ジュラ系中部統から得られる長期間生存種が多い中で、上部統最下部から初めて出現する指標種4種（*Kilinora spiralis*, *Fultacapsa sphaerica*, *Protunuma japonicus*, *Pseudoeucyrtis reticularis*）を識別した。得られた放散虫種の生存期間について再検討し、ジュラ系中部統から産する種が引き続き上部統からも産する例を明らかにした。その結果、北部石灰アルプスのジュラ紀放散虫化石帯において、これまで *Zhamoidellum ovum* 帯中に含まれていた *Williriedellum dierschei* 亜帯を、新たな指標種に基づき独立した帯として再定義した。古生物学的記載の章では37属67種2亜種を記載し、2属（*Loopus* 属, *Pseudodictyomitra* 属）1種（*Protunuma japonicus*）の標徴を改定するとともに、*Loopus* 属の模式種を再指定した。

Middle Jurassic radiolarians from the ammonite bearing Toyora Group, Yamaguchi Prefecture, Southwest Japan

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Abstract: The Toyora Group is one of the typical Lower to Middle Jurassic strata in Japan that is distributed throughout Yamaguchi Prefecture, Southwest Japan. It yields abundant ammonoids. Although, microfossils, such as radiolarians, have not been previously reported from the group, radiolarian fossils are first discovered at seven localities from the uppermost Toyora Group. Those correspond to the *Transsuum hisuikyoense* and *Striatojaponocapsa plicarum* zones and are determined to be from Aalenian to Bathonian in age. These radiolarian age determinations are a little older than those determined using ammonoids and inoceramids. According to previous studies, the assignment age of the first appearance of *Stj. plicarum* is only estimated to be near the Aalenian–Bajocian boundary. To discuss this issue, further study is required to correct the fossil data such as the Aalenian ammonoids that occur in the intervals of ammonoid and radiolarian localities.

Keywords: radiolaria, ammonoid, Jurassic, Toyora Group, Utano Formation, Yamaguchi Prefecture, Southwest Japan

1. Introduction

In Japan, most Jurassic strata are within accretionary complexes resulting from Mesozoic oceanic plate subduction. These strata are composed of various mixed rocks, such as oceanic plate-fragments, pelagic sediments and trench-filled clastic materials derived from the continent. The so-called shallow marine sediments are deposited on a continental shelf or in forearc basins composed of the accretionary strata. These shallow marine sediments overlie the accretionary strata on faulted or unconformable contacts. Before the rapid progress of research on the Jurassic accretionary complexes during the 1980s, these were a focus of stratigraphic research because of their abundant megafossils resulting from their comparatively limited mixing and weak deformation.

The geology of Southwest Japan is divided into the Inner (north) and Outer (south) zones by a major fault, the Median Tectonic Line, which formed in the Cretaceous. The Jurassic accretionary complexes are widely distributed in both zones.

The Lower to Middle Jurassic Toyora Group is one of the typical strata distributed in Yamaguchi Prefecture in the

Inner Zone of Southwest Japan (Fig. 1). The Toyora Group comprises stratified clastic rocks, namely, sandstones and mudstones, with a small amount of conglomerates. They are deposited under shallow marine to brackish water conditions based on their sedimentary facies features and fossil associations, including characteristic black shale that indicates anoxic sedimentary conditions.

Various fossils, mainly ammonoids, bivalves, gastropods and plants, have been reported from the Toyora Group. As this area is a type locality of the Early to Middle Jurassic ammonoid biostratigraphy in Japan, these ammonoids have been studied in detail since the early twentieth century. However, research on fossil radiolarians from the Paleozoic and Mesozoic in Japan started during the late 1970s when samples of radiolarians were collected and analyzed from the accretionary complexes. Sedimentary rocks in Mesozoic accretionary complexes of Japan that yield megafossils are very rare. This study aims to find radiolarians in the shallow marine sediments, such as the Toyora Group, for correlation with geologic ages assigned from ammonoids. Radiolarians from the Torinosu Group and its equivalent beds in the Outer Zone of Southwest Japan were the focus of previous studies on Mesozoic

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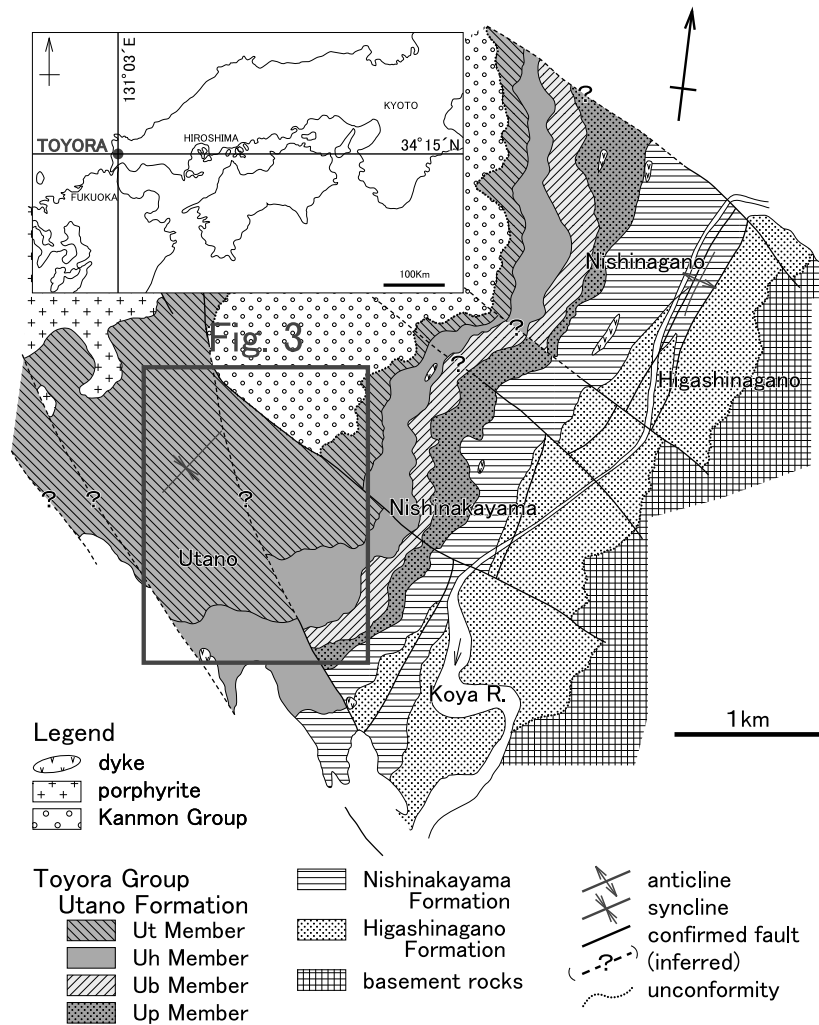


Fig. 1 Geological map of the Toyora Group around the Utano area, Yamaguchi Prefecture (modified from Hirano, 1971)

radiolarians (Matsumoto and Nishizono, 1985, Kozai *et al.*, 2006). Recently, Jurassic radiolarians were reported from shallow marine sediments of the Tetori Group on the Japan Sea side. On the basis of ammonoid biostratigraphy, radiolarians from this group are assigned to the Middle and Upper Jurassic (Callovian to Tithonian) (Hirasawa *et al.*, 2010). Sano and Kashiwagi (2015) explained that the high component ratio of *Spumellaria* (73–92 %) shows the Boreal element. This study, for the first time, describes radiolarians from the Toyora Group and the assignment of their geologic age and successfully correlates the geologic ages of the radiolarian biostratigraphy of the Middle Jurassic in Japan with the ammonoid one.

2. Geologic overview of the Toyora Group

Geological studies of the Toyora Group started with Yokoyama (1902) and Kobayashi (1926). Yokoyama (1902) described Jurassic ammonoids, while Kobayashi (1926) described inoceramids with lithostratigraphic notes. Toriyama (1938) showed the basic lithostratigraphic

framework of the group. In the current study, the geologic overview is described on the basis of the studies by Hirano (1971, 1973a, b) who established the lithostratigraphy and ammonoid biostratigraphy of the northern distribution area of the Toyora Group. Based on his research, distribution area of the Toyora Group is divided into the northern and southern areas by the Kikugawa Fault. The effect of a granitic rock intrusion is comparatively very weak in the northern area of this group. Nakada and Matsuoka (2009, 2011) established a detailed ammonoid biostratigraphy of the Nishinakayama Formation of the group and discussed the exact stratigraphic location of the Pliensbachian–Toarcian boundary in this formation.

2.1 Lithostratigraphy

The Toyora Group, comprising the Higashinagano, Nishinakayama, and Utano formations in ascending order, has a total thickness of 1800 m (Fig. 2). The Higashinagano Formation is 400 m in thickness and unconformably overlies the Sangun metamorphic rocks. It is composed of basal conglomerate, coarse sandstone, fine sandstone,

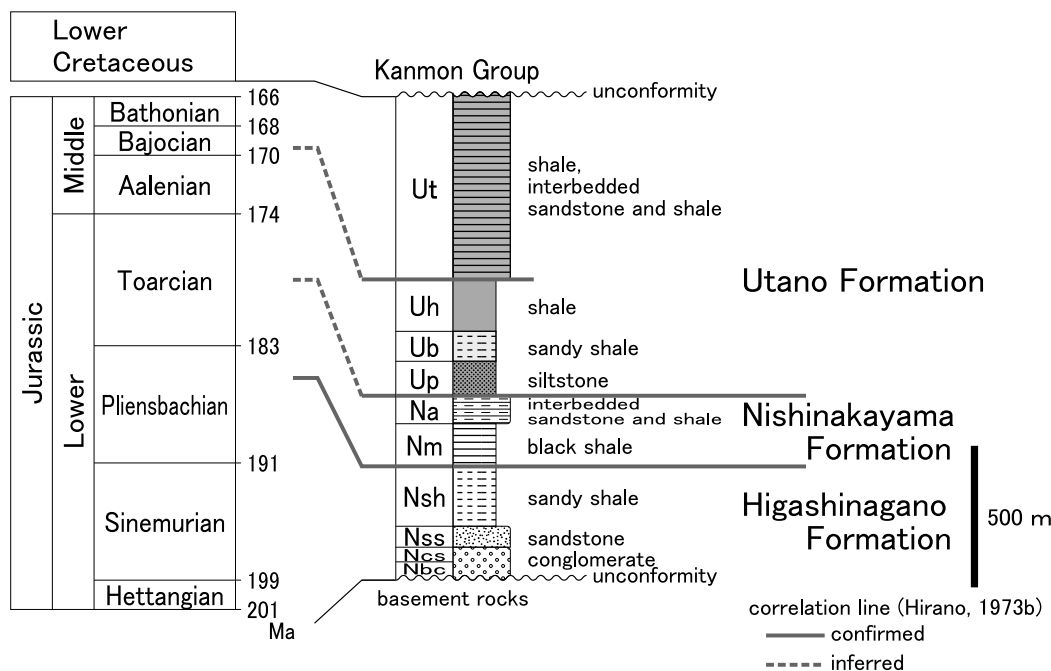


Fig. 2 Modified lithological succession and geologic age of the Toyora Group in the type locality after Hirano (1971, 1973a, b). Numerical ages are referred to the international chronostratigraphic chart ver. 2020/01 (Cohen *et al.*, 2013; updated in 2020). Abbreviations Nbc, Ncs, Nss and Nsh stand for the lower, middle, upper and uppermost members of the Higashinagano Formation; Nm and Na for the lower and upper members of the Nishinakayama Formation; Up, Ub, Uh and Ut for the lower, middle, upper and uppermost members of the Utano Formation.

and sandy shale in ascending order. This formation has been determined as a transgressive deposit because of its fining-upward sedimentary sequence. The Higashinagano Formation is subdivided into four members: the lower, middle, upper, and uppermost (abbreviated as the Nbc, Ncs, Nss and Nsh in Hirano, 1971, respectively) in ascending order. The Nishinakayama Formation is 250 m in thickness and is mainly composed of black shale, interbedding sandstone at its upper part. The black shale is recognized as the facies deposited under stagnant anoxic conditions because of the presence of sedimentary pyrite (Shikama and Hirano, 1970). This facies indicative of the shelf deposits records the early Toarcian oceanic anoxic event (Izumi *et al.*, 2012). The Nishinakayama Formation is divided into two members: the lower and upper (the Nm and Na, respectively) in ascending order. The Utano Formation is 1100 m in thickness and comprises silty shale, sandy shale, and interbedded sandstone and shale. In addition, the uppermost part of this formation has not been delimited because of a covering of the Lower Cretaceous Kanmon Group. The Utano Formation is recognized as a regressive sequence from its coarsening-upward sedimentation and is divided into four members: the lower, middle, upper, and uppermost (the Up, Ub, Uh, and Ut, respectively) in ascending order. At the Utano Dam locality, the maximum thickness of the uppermost member (Ut) is 650 m.

2.2 Biostratigraphy and geologic age based on megafossils

In studying the radiolarian assignment ages, it is effective to establish the biostratigraphy and geological age of each formation or member by the described megafossils, mainly ammonite (Fig. 2, Hirano, 1971, 1973a, b). Ammonoids are rarely found in the Higashinagano Formation but are abundant in the Nishinakayama Formation. The uppermost member of the Utano Formation (Ut) frequently yields ammonoids. Ammonoids have not been recognized in the lower member (Nbc) of the Higashinagano Formation. Abundant fossils, such as ammonoids, bivalves, gastropods, brachiopods, and corals, occur in the middle member (Ncs) of the Higashinagano Formation. The geologic age of Ncs is correlated with the early Sinemurian based on the occurrence of *Arietites* sp. That of the upper member (Nss), depending on its location, is correlated with the late Sinemurian to the Pliensbachian based on the age of Ncs in the northern area and the occurrences of the Pliensbachian ammonoid (*Amaltheus* cf. *stokes* (Sowerby) and *Arieticerat* aff. *apertum* Monestier) in the southern area. Based on the ages of Nss and the base of the Nishinakayama Formation above this member, the geological age of the uppermost member (Nsh) also depends on its location and is correlated with the latest Sinemurian to upper Pliensbachian.

Ammonoid fauna from the Nishinakayama Formation have been grouped into three zones: the *Fontanelliceras*

fontanellense, *Protogrammoceras nipponicum*, and *Dactyloceras helianthoides*, in ascending order, and the lower (Nm) and upper (Na) members of the Nishinakayama Formation are assigned to the upper Pliensbachian to lower Toarcian and lower Toarcian, respectively (Hirano, 1973a, b). Furthermore, Nakada and Matsuoka (2011) established four ammonoid zones in Nm and correlated these with European zonations: the *Canavaria japonica*, *Paltarpites paltus*, *Dactyloceras helianthoides*, and *Harpoceras inouyei* zones in ascending order.

In the Utano Formation, the lower (Up) and middle (Ub) members are correlated with the upper Toarcian given the occurrences of *Grammoceras* and *Phymatoceras* in abundance in the Up and *Phymatoceras* sp. in the Ub. In this formation, the upper member (Uh) yields ammonoids such as *Planammatoceras* cf. *kitakamiensis* Buckman, *Dumortieria?* sp. and *Calliphylloceras* sp. These are assigned to the uppermost Toarcian to partially lower Bajocian. The uppermost member (Ut) is correlated with the Bathonian because of the occurrence of *Harpophylloceras* sp. and *Inoceramus utanoensis* (Hirano, 1973b). Based on the assigned age of the megafossils, the border between the Uh and Ut is in the lower Bajocian to Bathonian with a concordant stratigraphic contact.

2.3 Pliensbachian–Toarcian boundary

Tanabe (1991) demonstrated that bituminous mudstones result from the deposition of marine sediments during anoxic events, which have occurred worldwide, based on the geochemical and sedimentological data and the extremely rare occurrence of benthic fossils. Nakada and Matsuoka (2011) determined that the Pliensbachian–Toarcian boundary is at the base of the *Paltarpites paltus* Zone based on the four established ammonoid zones and the lower member of the Nishinakayama Formation (Nm), which are correlated in detail with European zones.

3. Lithostratigraphy of the sampling section study route

Radiolarians were found from seven samples collected from the Utano B, C2, and D1 routes belonging to the Uh and Ut that crop out in the Utano Valley (along the current Utano Dam), where the thickest sequence of the Utano Formation is throughout distributed (Hirano, 1971). Hirano (1973a, b) has described ammonoids from the Utano A and D2 routes of this study (Fig. 3).

Utano A route

Along the Utano A route, the 400-m-thick Ut of the Utano Formation is exposed and mainly comprises interbedded sandstone and shale (Fig. 4). Sandy shale varying from 20 to 30 m in thickness occurs at three stratigraphic horizons. *Holcophylloceras* sp. described by Hirano (1973b; locality 59) occurs in the shale above the thick sandstone bed in the upper part of this route.

Utano B route

The lower part of the Utano B route (Fig. 4) is composed

of the 150-m-thick massive sandy shale of the Uh of the Utano Formation. The upper part of this route consists of the interbedded sandstone and shale of the Ut of the Utano Formation with sandy shale; these are repeated every tens of meters. Their total thickness is 400 m. Tuffaceous shales are infrequently intercalated within the upper part of the strata. Radiolarian samples UT-4, UT-5, and UT-6 were collected from three horizons in the upper half of the Ut.

Utano C1 and C2 routes

The thickness and sedimentary facies of the Uh and Ut of the Utano Formation along the Utano C1 and C2 (C2n and C2s) routes are similar to those of the Utano B route (Fig. 4). The Ut is overlain unconformably by the Cretaceous Kanmon Group at the stratigraphic top of the Utano C2s route. A radiolarian sample UT-7 was collected at the horizon 70 m beneath this unconformity.

Utano D1 and D2 routes

The lower part of the Utano D1 and D2 routes comprises the 250-m-thick sandy shale of the Uh of the Utano Formation. The upper 200 m part consists of the interbedded sandstone, shale, and sandy shale of the Ut (Fig. 4). These lithologies are repeated every tens of meters. In the Utano D1 route, a radiolarian sample UT-1 was collected from the Uh, and samples UT-2 and UT-3 were collected from the middle part of the Ut. Hirano (1973a) reported ammonoids (*Dumortieria?* sp.) from shale in the lower part of the Uh in the Utano D2 route.

4. Radiolarian assemblages and age assignment

Seven radiolarian samples were collected from the studied route. Radiolarians from four samples (UT-1, UT-2, UT-5 and UT-7) were identified, whereas those from the other three samples could not be identified because of poor preservation. Age assignments of these identified radiolarians are mainly discussed on the basis of the radiolarian zonations of Nishizono *et al.* (1997) and Matsuoka (1995).

4.1 Radiolarian assemblages

The locality and radiolarian assemblage of each sample are shown as follows.

Sample UT-1

Locality: Utano D1 route (sandy shale of the Uh).

Assemblage: Despite abundant Nassellaria and the poor preservation of the test surfaces in this sample, only *Praeparvicingula?* sp. A can be identified (Fig. 5m).

Sample UT-2

Locality: Utano D1 route (sandy shale of the Ut).

Assemblage: Abundant Spumellaria and Nassellaria are included in the sample; however, these are poorly preserved on the test surface. The following radiolarians were identified (Fig. 5b, e, j): *Canutus* sp., *Parahsuum* sp., and *Transhsuum* aff. *hisuikyoenense*.

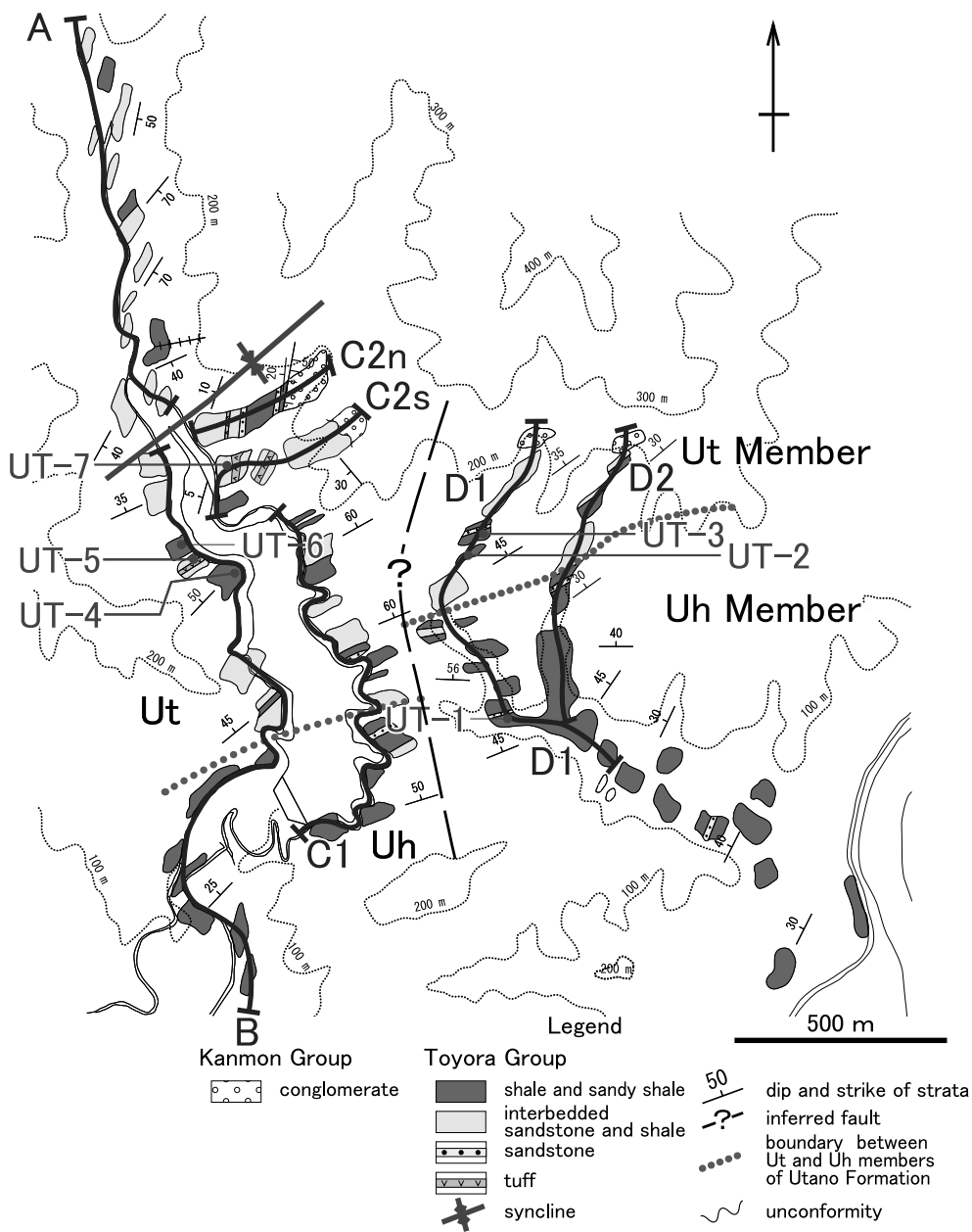


Fig. 3 Locality map for radiolarians and surveyed routes of Utano A, B, C1, C2, D1 and D2. UT-1 to Ut-7 are the locations of radiolarian occurrence.

Sample UT-3

Locality: Utano D1 route (sandy shale of the Ut located immediately above UT-2).

Assemblage: Despite the presence of abundant *Nassellaria*, no radiolarian species is identified because of their poor preservation on the test surfaces.

Sample UT-4

Locality: Utano B route (sandy shale of the Ut).

Assemblage: Despite abundant *Nassellaria* in the sample, no radiolarian species is identified due to their poor preservation on the test surfaces.

Sample UT-5

Locality: Utano B route (shale of the Ut).

Assemblage: Despite poor preservation, the following radiolarian species are identified (Fig. 5a, d, f, i, l, n–o, s, x–y): *Archicapsa pachyderma*, *Spongocapsula* sp. A, *Parahsuum*? *hiconocosta*, *Transhsuum* aff. *brevicostatum*, *Praeparvicingula aculeata*, *Wrangellium* aff. *burnsensis*, *Droltus hecatensis*, *Unuma typicus*, *Stichocapsa convexa* and *Stichocapsa magnipora*.

Sample UT-6

Locality: Utano B route (sandy shale of the Ut)

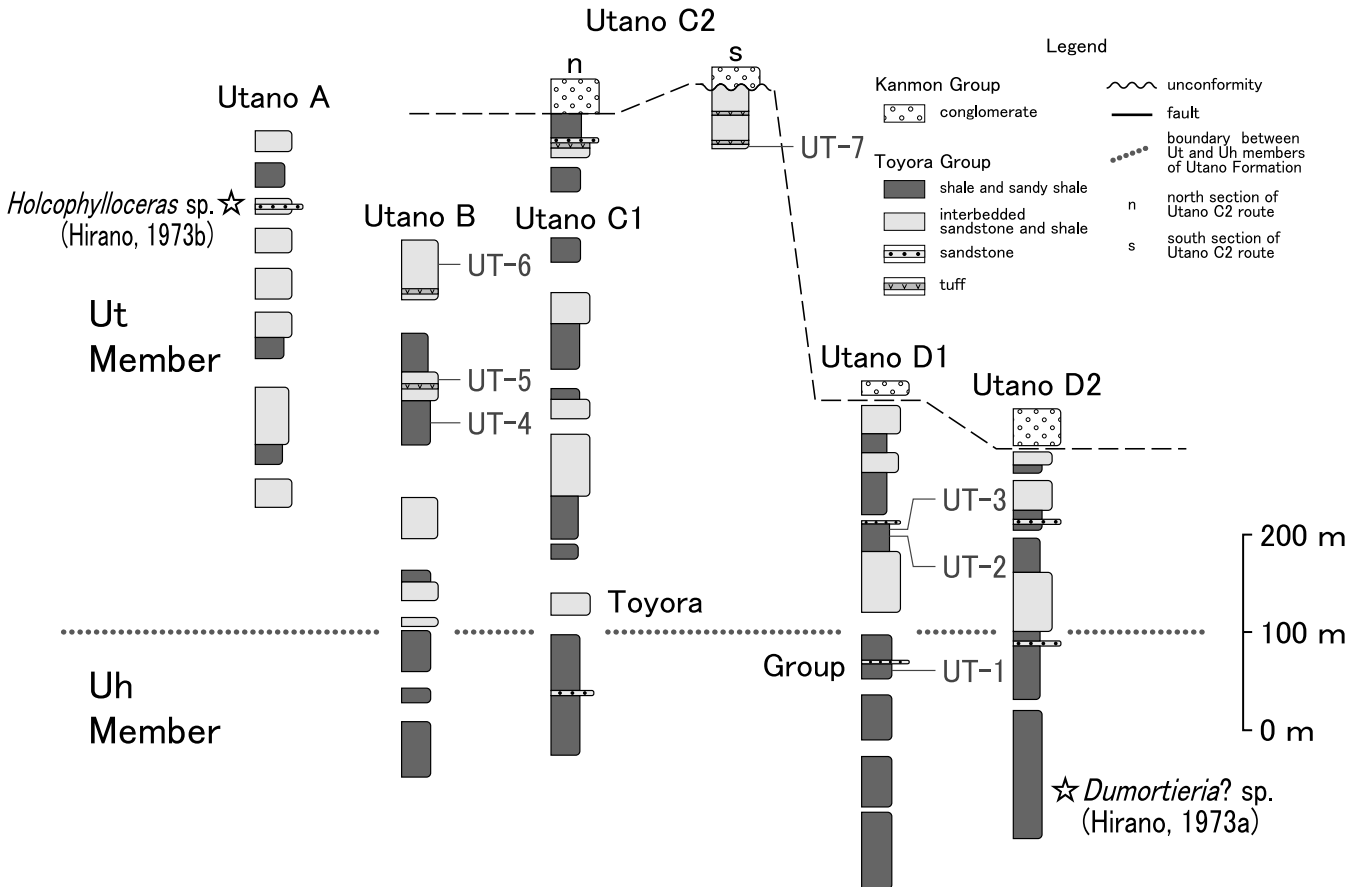


Fig. 4 Columnar sections of Utano A, B, C1, C2, D1 and D2 routes.

Assemblage: Despite the inclusion of radiolarians, no specimens is identified due to their poor preservation on the test surfaces.

Sample UT-7

Locality: Utano C2 route (nodule in the sandy shale of the Ut)

Assemblage: Well-preserved Nassellaria are included (Figs. 5c, g–h, k, p–r, t–w, z). The following radiolarians are identified: *Archicapsa pachyderma*, *Spongocapsula* aff. *krahsteinensis*, *Transhsuum maxwelli*, *Archaeodictyomitra* sp. H in Nishizono 1996, *Triversus hungaricus*, *Unuma laticostatus*, *Unuma typicus*, *Podobursa nodosa*, *Striatojaponocapsa plicarum*, *Tricolocapsa undulata*, and *Eucyrtidiellum unumaense*.

4.2 Correlation with radiolarian zonation and age assignment

According to Nishizono *et al.* (1997), the co-occurrence of *Archicapsa pachyderma*, *Stichocapsa convexa*, and *Striatojaponocapsa plicarum* in UT-7 collected from the Ut indicates the *Stj. plicarum* Zone (Nishizono *et al.*, 1997). UT-5 was collected from the stratigraphic level 240 m beneath sample UT-7 in the Ut. An assemblage with a co-occurrence of *A. pachyderma*, *Parahsuum?*

hiconocosta, *Unuma typicus*, and *Stichocapsa convexa* but no *Str. plicarum* occurred in UT-5, which was correlated with the *Transhsuum hisuikyoense* Zone (Nishizono *et al.*, 1997). This correlation does not contradict the assemblage of the *Th. hisuikyoense* Zone of sample UT-2, which is in the lower part of the Ut. The *Th. hisuikyoense* and *Stj. plicarum* zones are assigned to the Aalenian and Bajocian to Bathonian, respectively (Nishizono *et al.*, 1997). Therefore, it is estimated that the boundary between the two zones is in the middle part of the Ut, which is 360 m above the Uh–Ut boundary at the stratigraphic level between the samples UT-5 and UT-7 (Fig. 6).

According to Matsuoka (1995), the co-occurrence of *Stj. plicarum* and *Eucyrtidiellum unumaense*, as seen in UT-7, correlates with the *Stj. plicarum* and *Stj. conexa* zones (Matsuoka, 1995; Matsuoka and Ito, 2019). However, the co-occurrence of *Stj. plicarum*, *Unuma typicus*, *Unuma laticostatus*, *Stichocapsa convexa*, and *Eucyrtidiellum unumaense* but no *Stj. conexa* shows that the radiolarians from UT-7 could be correlated with the *Stj. plicarum* Zone. Furthermore, the radiolaria in UT-5 could be correlated with the *Laxtorum? jurassicum* Zone (Matsuoka, 1995) because the co-occurrence of *A. pachyderma*, *Unuma typicus*, and *Stichocapsa convexa* but no *Stj. plicarum* is found in UT-5. Matsuoka (1995) correlated the *L.?*

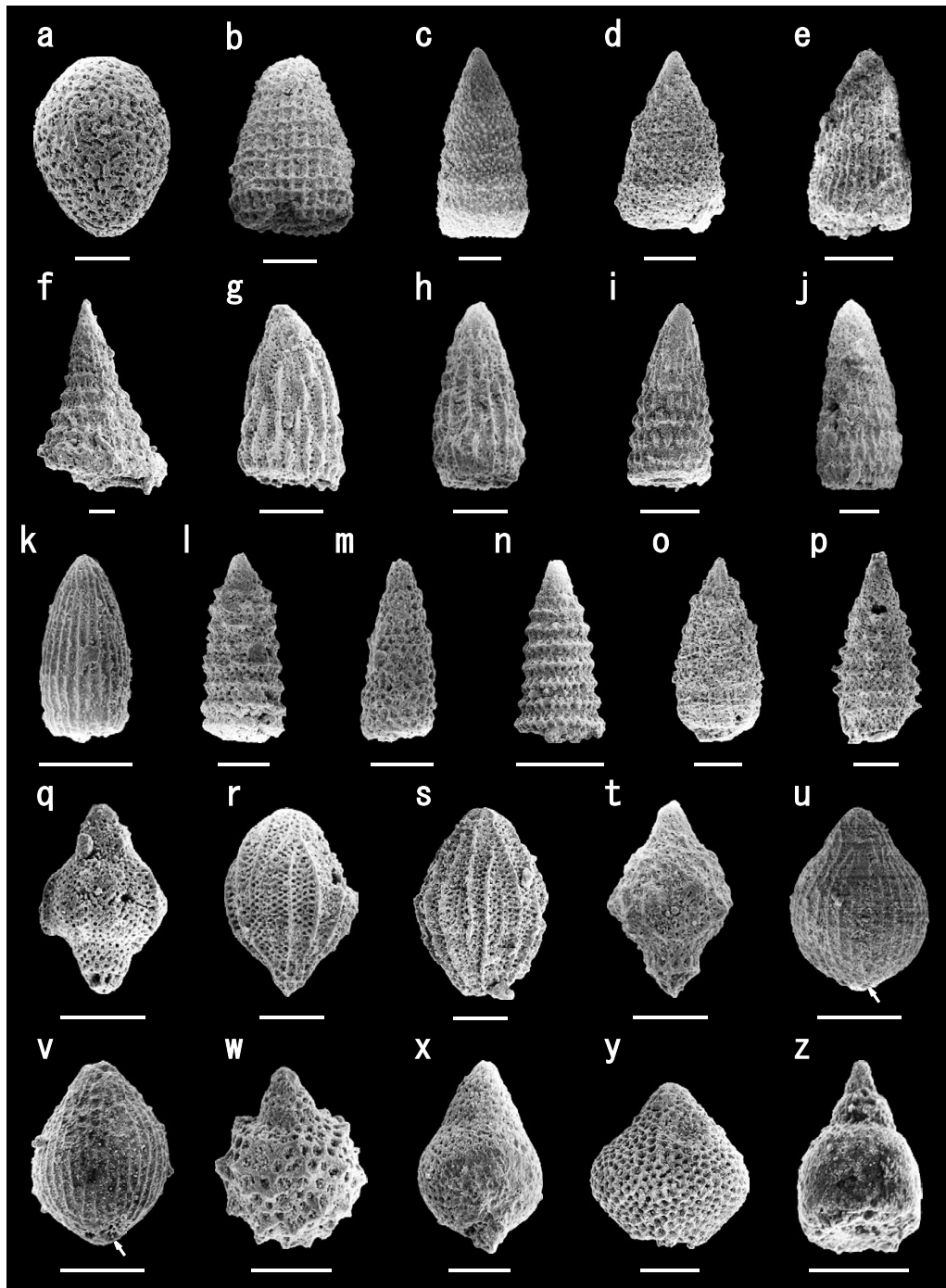


Fig. 5 SEM (Scanning Electron Microscope) photos of radiolarian.

a: *Archicapsa pachyderma* (Tan) (locality; UT-5). b: *Canutus* sp. (locality; UT-2). c: *Spongocapsula* aff. *S. krahsteinensis* Suzuki and Gawlick (UT-7). d: *Spongocapsula* sp. A (locality; UT-5). e: *Parahsuum* sp. (locality; UT-2). f: *Parahsuum?* *hiconocosta* Baumgartner and De Wever (locality; UT-5). g, h: *Transhsuum maxwelli* (Pessagno) (locality; UT-7). i: *Transhsuum* aff. *brevicostatum* (Ozoldova) (locality; UT-5). j: *Transhsuum* aff. *hisuikyoense* (Isozaki and Matsuda) (locality; UT-2). k: *Archaeodictyomitra* sp. H in Nishizono (1996) (locality; UT-7). l: *Praeparvicingula aculeata* (Carter) (locality; UT-5). m: *Praeparvicingula?* sp. A (locality; UT-1). n: *Wrangellium* aff. *burnsensis* (Pessagno and Whalen) (locality; UT-5). o: *Droltus hecatensis* Pessagno and Whalen (locality; UT-5). p: *Triversus hungaricus* (Kozur) (locality; UT-7). q: *Unuma laticostatus* (Aita) (locality; UT-7). r: *Unuma typicus* Ichikawa and Yao (locality; UT-7). s: *Unuma typicus* Ichikawa and Yao (locality; UT-5). t: *Podobursa nodosa* (Chiari, Maruccci and Prela) (locality; UT-7). u, v: *Striatojaponocapsa plicarum* (Yao) (locality; UT-7). Arrows of u and v show the circular area. w: *Tricolocapsa undulata* (Heitzer) (locality; UT-7). x: *Stichocapsa convexa* Yao (locality; UT-5). y: *Stichocapsa magnipora* Chiari, Maruccci and Prela (locality; UT-5). z: *Eucyrtidiellum unumaense* (Yao) (locality; UT-7). All scale bars indicate 50 μ m.

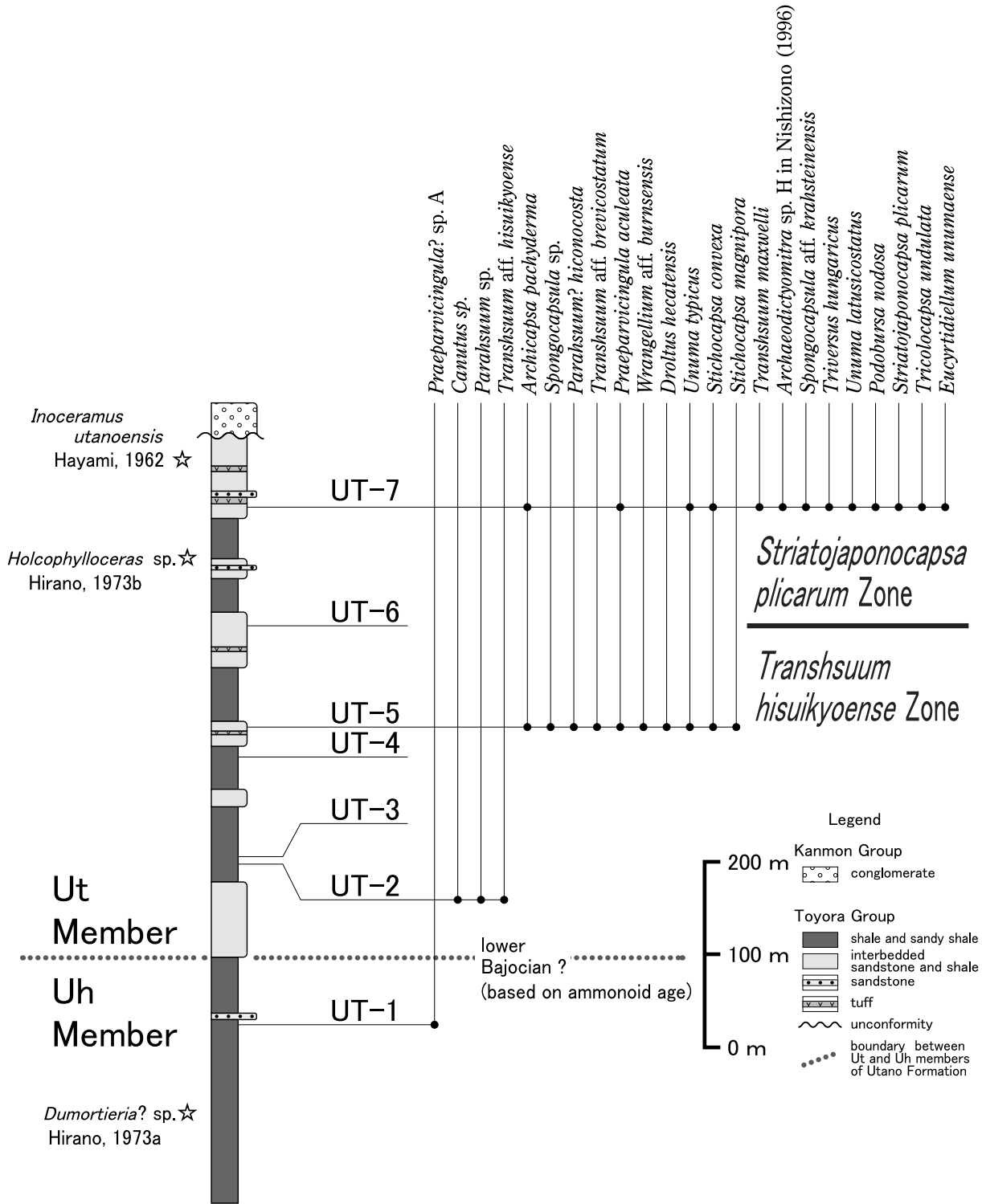


Fig. 6 Biostratigraphic distribution of radiolarian species in the Utano sections, the Toyora Group.

jurassicum and *Stj. plicarum* zones with the Aalenian and Bajocian to early Bathonian, respectively. Therefore, the geological ages assigned by radiolarians indicate that the boundary between the *Th. hisuikyoense* and *Stj. plicarum* zones (Nishizono *et al.*, 1997) is in the middle part of the Ut, which is the border of the Aalenian and Bajocian.

4. 3 Correlation between radiolarian and macrofossil age assignments

Find-spots of ammonoid are close to the two radiolarian sampling sites. *Dumortieria?* sp. occurs in the lower part of the upper member (Uh) in the Utano D2 route (Figs. 4 and 6), which is correlated with 280 m below UT-2; it

is assigned to the Toarcian Stage of the Lower Jurassic (Hirano, 1973a). Hirano (1973b) showed that the Uh is assigned to the horizon ranging from the uppermost Toarcian to the Bajocian (mainly Aalenian) on the basis of an ammonoid *Planammatoceras* cf. *kitakamiensis* in the Uh, occurred outside of the study area. The stratigraphic level of *Holcophylloceras* sp. (Hirano, 1973b: locality 59) is 180 m above UT-5 and 60 m below UT-7 (Figs. 4 and 6), which is the upper part of the Ut in the Utano A route. The genus *Holcophylloceras* ranges in age from the Bajocian to the Aptian in the Early Cretaceous (Sandoval *et al.*, 2001, Majidifard, 2003). Furthermore, Hayami (1962) showed that *Inoceramus utanoensis* (Kobayashi), which occurred in the Ut, is very similar to *I. kystatymensis* as reported from the Bathonian in the Lena River, Russia (Koschelkina, 1963; Hirano, 1973b). This indicates that the uppermost member (Ut) includes at least the Bajocian to Bathonian in the Middle Jurassic.

Previous studies have shown the Aalenian–Bajocian boundary is in the Ut based on radiolarians (e.g. Matsuoka, 1995; Nishizono *et al.*, 1997) and in the upper part of the Uh based on ammonoids (Hirano, 1973b). The difference in the stratigraphic interval of the two ammonoid localities is 580 m, while that of the radiolarian localities (UT-5: Aalenian and UT-7: Bajocian) is 240 m (Fig. 6).

The assigned ages of ammonoids from the Utano Formation, indicate that the first appearance age of *Striatojaponocapsa plicarum* could be redefined from the Aalenian to the early Bajocian. According to Nishizono *et al.* (1997), *Stj. plicarum* first appeared in the lower Bajocian based on calibration by ammonoids occurring in the Kitakami Mountains, Tohoku district, central Kyushu district and Canada. Matsuoka (1995) showed that *Stj. Plicarum* appeared in the Bajocian based on the correlation data by ammonoids occurring in Spain, Italy and Japan. Furthermore, Baumgartner *et al.* (1995) set the range of *Stj. plicarum plicarum* from the upper Bajocian to the lower Bathonian (UA Zones 4 and 5). However, in previous studies (Matsuoka, 1995; Nishizono *et al.*, 1997), the first appearance age of *Stj. plicarum* could not be established for the assigned ages of ammonoid and was only estimated to be around the Aalenian/Bajocian boundary.

Considering this difference, further study is required to collect fossil data such as the Aalenian ammonoids from these intervals.

5. Conclusions

Radiolarian fossils were first discovered from the Lower to Middle Jurassic Toyora Group containing abundant ammonoids. The radiolarian fossils, found in seven localities from the uppermost Toyora Group, are assigned to the Aalenian–Bathonian based on the presence of ammonoids and inoceramids. According to Nishizono *et al.* (1997), these radiolarian assemblages are correlated to the *Transsuum hisuikyoenense* and *Striatojaponocapsa plicarum* zones which have been assigned to the Aalenian

and Bajocian to lower Bathonian, respectively. Although the upper–uppermost member (Uh–Ut) boundary of the Utano Formation was correlated with the lower Bajocian based on megafossils (Hirano, 1973b), it should be assigned to the Aalenian based on the radiolarian zonation. The stratigraphic intervals of the two ammonoid localities and that of the radiolarian localities (UT-5: Aalenian and UT-7: Bajocian) are 580 m and 240 m, respectively. According to previous studies, the first appearance of *Stj. plicarum* could not be calibrated to the ammonoid assignment ages and is only estimated to be near the Aalenian/Bajocian boundary. To discuss this issue, further study is required to correct the fossil data such as the Aalenian ammonoid occurring in these intervals.

6. Systematic Paleontology

The familial classification system basically follows Takemura (1986), Suzuki *et al.* (2002), Suzuki and Gawlick (2003, 2009). The classification for genera *Archicapsa* and *Canutus* are based on Haeckel (1881), Pessagno and Whalen (1982), respectively.

Subclass **RADIOLARIA** Müller, 1858

Order **NASSELLARIA** Ehrenberg, 1875

Family **SETHOCAPSIDAE** Haeckel, 1881

Genus *Archicapsa* Rüst, 1885

Archicapsa pachyderma (Tan, 1927)

(UT-5, Fig. 5a)

1986 *Archicapsa pachyderma* (Tan) – Matsuoka and Yao, pl. 1, fig. 5.

1990 *Archicapsa pachyderma* (Tan) – Yao, pl. 2, fig. 15.

1990 *Archicapsa pachyderma* (Tan) – Hori, fig. 9.44.

1996 *Archicapsa pachyderma* (Tan) – Nishizono, pl. 12, fig. 4.

Remarks: Test is ellipsoidal with a spherical apical part and slightly pointed aperture side.

Range: This species occurs from the *Droltus?* sp. A–*Hsuum?* sp. G to *Striatojaponocapsa plicarum* zones in Outer zone of Southwest Japan (Nishizono, 1996).

Family **CANUTIDAE** Pessagno and Whalen, 1982

Genus *Canutus* Pessagno and Whalen, 1982

Canutus sp.

(UT-2, Fig. 5b)

Remarks: Test is short, inflated and conical. Test surface consists of square symmetrical pore frames with nodes. In present specimen, the two or three layered structure and apical part are unclear due to poor preservation.

Family **THEOPERIDAE** Haeckel 1881;

emend. Takemura, 1986

Genus *Spongocapsula* Pessagno, 1977

Spongocapsula sp. aff. *S. krahsteinensis* Suzuki and Gawlick in Gawlick *et al.* (2004)

(UT-7, Fig. 5c)

1996 *Spongocapsula?* sp. – Nishizono, pl. 27, fig. 11.

aff. 2004 *Spongocapsula krahsteinensis* n. sp. – Gawlick *et al.*, p. 313–315, abb. 4.7–4.10.

Remarks: Test is short, inflated and spindle-shaped. Gawlick *et al.* (2004) regarded that underneath the microgranular outer layer of this species is spongy test. This specimen has a resemble test of *S. krahsteinensis*.

Range: late Bajocian to Callovian (Gawlick *et al.*, 2004)

Spongocapsula sp. A

(UT-5, Fig. 5d)

Remarks: Test is conical shaped increasing slowly in height and moderately rapidly in width proximally, gradually decreasing in width distally. Cephalis and thorax are imperforate.

Genus *Parahsuum* Yao, 1982

Parahsuum sp.

(UT-2, Fig. 5e)

Remarks: Test is conical and lacking ornamentation at cephalis, thorax and abdomen due to poor preservation of this sample. In a side view, sixteen edged costae are visible on post-abdominal chambers. Single row of square pore frames is arranged with circular, primary pores between costae.

Parahsuum? *hiconocosta* Baumgartner and De Wever in Baumgartner *et al.* (1995)

(UT-5, Fig. 5f)

1985 *Andromeda?* sp. – De Wever *et al.*, pl. 1, figs. 12, 13, 16.

1995 *Parahsuum?* *hiconocosta* n. sp. – Baumgartner *et al.*, p. 378, pl. 3011, figs. 1 (H)–6.

1996 *Andromeda* sp. B – Nishizono, pl. 26, fig. 18.

Remarks: Test is elongated conical form with concave and wedge-shaped outline in lateral view. Segments are well marked by a nodose circumferential ridge. The present species has rectangular and elevated vertical pore-frames with a protruding nodose as characteristic structures of this species.

Range: This species occurs from the *Transhsuum hisuikyoense* Zone in Outer zone of Southwest Japan (Nishizono, 1996). UA Zones 2–4, late Aalenian–late Bajocian (Baumgartner *et al.*, 1995)

Genus *Transhsuum* Takemura, 1986

Transhsuum maxwelli (Pessagno, 1977)

(UT-7, Fig. 5g, h)

1977 *Hsuum maxwelli* n. sp. – Pessagno, p.81, pl. 7, figs. 14–16.

1995 *Transhsuum maxwelli* group (Pessagno) – Baumgartner *et al.*, p. 582, pl. 3180, figs. 4, 5.

1997 *Hsuum maxwelli* Pessagno – Nishizono *et al.*, pl. III, fig. 10.

2009 *Hsuum maxwelli* Pessagno – Suzuki and Gawlick, p. 168, fig. 5.7.

Remarks: Test is conical. Cephalis is perforate without long and massive apical horn. The specimen shown in Fig. 5h has the test of increasing rapidly in width. Short, massive and discontinuous costae are distributed on a test.

Range: UA Zones 3–10, early–middle Bajocian to late Oxfordian–early Kimmeridgian. (Baumgartner *et al.*, 1995). The first appearance of *Transhsuum maxwelli* group is located in the upper part of *Striatojaponocapsa plicarum* Zone in Outer zone of Southwest Japan (Matsuoka, 1995: Bajocian to middle Bathonian).

Transhsuum sp. aff. *T. brevicostatum* (Ozoldova, 1975) (UT-5, Fig. 5i)

aff. 1975 *Lithostrobos brevicostatus* n. sp. – Ozoldova, p. 84, pl. 102, fig. 1.

aff. 1979 *Lithostrobos brevicostatus* Ozoldova – Ozoldova, p. 259, pl. 5, fig. 2.

aff. 1995 *Transhsuum brevicostatum* group (Ozoldova) – Baumgartner *et al.*, p. 578, pl. 3181, figs. 2, 4.

aff. 1996 *Hsuum brevicostatum* Ozoldova – Nishizono, pl. 21, figs. 8, 9.

Remarks: Test is conical. Post-abdominal segments have short longitudinal ribs. The specimen shown in Fig. 5i could not be identified as *Transhsuum brevicostatum* because two longitudinal lines of pores are unclear due to poor preservation. *Th. brevicostatum* (Ozoldova) is rare in the *Striatojaponocapsa plicarum* Zone after Nishizono (1996).

Range: UA Zones 3–11, early–middle Bajocian to late Kimmeridgian–early Tithonian. (Baumgartner *et al.*, 1995).

Transhsuum sp. aff. *T. hisuikyoense* (Isozaki and Matsuda 1985)

(UT-2, Fig. 5j)

Measurements (in μm): height 200, maximum width 100.

Remarks: General form and surface ornamentations of post-abdominal segments are very similar to *Hsuum hisuikyoense* Isozaki and Matsuda, 1985. However, details of apical part are unclear due to poor preservation.

Range: UA Zones 2–4, late Aalenian–late Bajocian (Baumgartner *et al.*, 1995).

Genus *Archaeodictyomitra* Pessagno, 1976

Archaeodictyomitra sp. H in Nishizono, 1996

(UT-7, Fig. 5k)

1982 *Archaeodictyomitra* sp. A – Pessagno and Whalen,

p. 117, pl. 8, fig. 10.

1996 *Archaeodictyomitra* sp. H – Nishizono, pl. 24, fig. 4.

Remarks: This species has rounded test apically and constricted distally. Twelve costae are visible on post-abdominal chambers in a side view. This species occurs from the *Striatojaponocapsa plicarum* Zone in Outer zone of Southwest Japan (Nishizono, 1996).

Genus *Praeparvicingula* Pessagno, Blome and Hull in Pessagno *et al.*, 1993

Praeparvicingula aculeata (Carter in Carter *et al.*, 1988) (UT-5, Fig. 5l)

1988 *Parvicingula aculeata* n. sp. – Carter *et al.*, p. 54–55, pl. 18, figs. 1, 2, 7.

1997 *Parvicingula dhimenaensis dhimenaensis* Baumgartner – Yao, pl. 13, fig. 625.

Remarks: Test is subcylindrical and maintaining same width of post-abdominal chamber. Cephalis and thorax are sparsely perforate without horn. Post-abdominal chambers has three lateral rows of pore frames between ridges. Those are depressed in central row. Sharp pointed nodes are clear, that separate the abdomen and first few post-abdominal chambers. The narrow tube is lacking on the final post-abdominal chamber. Goričan *et al.* (2006) classified *Parvicingula aculeata* Carter to genus *Praeparvicingula* on the basis of the above test structures. **Range:** middle Toarcian–Early Bajocian (Carter *et al.*, 1988) (no-data earlier than middle Toarcian in Carter *et al.*, 1988)

Praeparvicingula? sp. A. (UT-1, Fig. 5m)

Remarks: Test is subcylindrical with dome-shaped cephalis. Horn and terminal tube are unknown due to poor preservation. The present specimen is different from *Praeparvicingula tlellensis* Carter in having the non-parallel pore alignment to circumferential ridges.

Family **Amphipyndacidae** Riedel, 1967

Genus *Wrangellium* Pessagno and Whalen, 1982

Wrangellium sp. aff. *W. burnsensis* (Pessagno and Whalen, 1982) (UT-5, Fig. 5n)

1988 *Parvicingula* sp. aff. *P. burnsensis* (Pessagno and Whalen) – Carter *et al.*, p. 55, pl. 18, figs. 10, 15.

Remarks: Test is characterized by having nodose circumferential ridges with H-linked structures. The specimen identified as *P.* aff. *burnsensis* Pessagno and Whalen (pl 18, figs 10 and 15 of Carter *et al.*, 1988) is considered to be classified as genus *Wrangellium* in having circumferential ridges with H-linked structures. **Range:** middle Toarcian to early Bajocian (Carter *et al.*, 1988)

Genus *Droltus* Pessagno and Whalen, 1982

Droltus hecatensis Pessagno and Whalen, 1982 (UT-5, Fig. 5o)

1982 *Droltus hecatensis* n. sp. – Pessagno and Whalen, p. 121, pl. 1, figs. 12, 13, pl. 4, figs. 1, 2, 6, 10.

1998 *Droltus hecatensis* Pessagno and Whalen – Carter *et al.*, p. 63, pl. 15, fig. 14.

2002 *Droltus hecatensis* Pessagno and Whalen – Suzuki *et al.*, 2002. p. 181–182, figs. 8G, 8H, 8L–8M.

2003 *Droltus hecatensis* Pessagno and Whalen – Suzuki and Gawlick, p. 191–192, fig. 6.72.

2009 *Droltus hecatensis* Pessagno and Whalen – Suzuki and Gawlick, p. 177, figs. 6.50A, 6.50B.

Remarks: Apical horn is ornamented with thick blades. Abdomen and several post-abdominal chambers have irregularly sized and shaped polygonal pore frames with solid small spines in somewhere. In lower one third, test consists of three longitudinal rows of pores between every adjacent pairs of costae.

Range: *Droltus hecatensis* occurs commonly in Lower Jurassic of west coast of Canada (Carter *et al.*, 1998) and Peru (Suzuki *et al.*, 2002). Suzuki and Gawlick (2003, 2009) described this species from the Callovian to Oxfordian in the Northern Calcareous Alps.

Genus *Triversus* Takemura, 1986

Triversus hungaricus (Kozur, 1985) (UT-7, Fig. 5p)

1984 *Parvicingula dhimenaensis* n. sp. – Baumgartner, p. 778, pl. 7, fig. 4.

1985 *Eoxitus hungaricus* n. sp. – Kozur, p. 216, figs. 1a, 1b, 1d, 1e.

1995 *Parvicingula dhimenaensis* Baumgartner ssp. A – Baumgartner *et al.*, p. 406, pl. 4071, figs. 1–4.

1996 *Parvicingula dhimenaensis* Baumgartner – Nishizono, pl. 25, figs. 11–13.

2003 *Triversus hungaricus* (Kozur) – Suzuki and Gawlick, p. 195–196, fig. 6.58–6.60.

2009 *Triversus hungaricus* (Kozur) – Suzuki and Gawlick, p. 170, fig. 5.14; figs. 6.6A, 6.6B, 6.7, 6.8.

Remarks: Test has an elongated cephalis and is spindle-shaped with pronounced spines on circumferential ridges. **Range:** UA Zones 3–8, early–middle Bajocian to middle Callovian (Baumgartner *et al.*, 1995).

Genus *Unuma* Ichikawa and Yao, 1976

Unuma latusicostatus (Aita, 1987) (UT-7, Fig. 5q)

1987 *Tricolocapsa latusicostata* n. sp. – Aita, p. 76, pl. 4, figs. 7a–8b; pl. 10, figs. 8, 9.

1995 *Unuma latusicostatus* (Aita) – Baumgartner *et al.*, p. 622, pl. 4058, figs. 1, 4.

1996 *Tricolocapsa latusicostata* (Aita) – Nishizono, pl. 13, fig. 11.

Remarks: The present specimen has seven longitudinal plicae in a half side view, and four longitudinal rows of pores between neighboring plicae. No nodes or spines on the plicae.

Range: UA Zones 2–5, late Aalenian to latest Bajocian–early Bathonian (Baumgartner *et al.*, 1995).

Unuma typicus Ichikawa and Yao, 1976

(UT-7, Fig. 5r; UT-5, Fig. 5s)

1976 *Unuma typicus* n. sp. – Ichikawa and Yao, p. 112, pl. 1, figs. 1–3.

1996 *Unuma typicus* Ichikawa and Yao – Nishizono, pl. 18, figs. 15, 16.

Remarks: The specimen shown in Fig. 5r (UT-7) has nine longitudinal plicae in a half side view, and four longitudinal rows of pores between neighboring plicae without spines. The specimen shown in Fig. 5s (UT-5) is very similar to *Unuma typicus* Ichikawa and Yao. Probably, this will be an immature (variation) specimen or basal appendage has been broken.

Range: This species occurs from the *Hsuum hisuikyoenense* to *Striatojaponocapsa plicarum* zones in Outer zone of Southwest Japan (Nishizono, 1996). UA Zones 3–4, early–middle Bajocian to late Bathonian (Baumgartner *et al.*, 1995).

Genus ***Podobursa*** Wisniewski 1889; emend. Foreman, 1973

Podobursa nodosa (Chiari, Marucucci and Prela, 2002)

(UT-7, Fig. 5t)

2002 *Williriedellum nodosum* n. sp. – Chiari *et al.*, p. 84, pl. 5, figs. 15–19.

2009 *Podobursa nodosa* (Chiari, Marucucci and Prela) – Suzuki and Gawlick, p. 178, figs. 5.20, 5.21.

Remarks: Cephalis and thorax are conical. Abdomen is large globose with nodes surrounded by irregular pores. Final segment terminates in a prolonged tube with elongate pores and solid pore flames.

Range: UA Zone 5, latest Bajocian to early Bathonian (Chiari *et al.*, 2002).

Family ***Arcanicsapsidae*** Takemura, 1986

Genus ***Striatojaponocapsa*** Kozur, 1984

Striatojaponocapsa plicarum (Yao, 1979)

(UT-7, Fig. 5u, v)

1979 *Tricolocapsa plicarum* n. sp. – Yao, p. 32, pl. 4, figs. 1–11.

1984 *Striatojaponocapsa plicarum* (Yao) – Kozur, p. 56, pl. 7, fig. 3

1996 *Tricolocapsa plicarum* Yao – Nishizono, pl. 13, figs. 14–16.

2007 *Striatojaponocapsa plicarum* (Yao) – Hatakeda *et al.*, p. 16, pl. 1, figs. 1–10

2009 *Striatojaponocapsa plicarum* (Yao) – Suzuki and Gawlick, p. 182, figs. 5.39A, 5.39B.

Measurements (in μm): Fig. 5u; height 92, width 80, width of basal appendage 30, Fig. 5v; height 95 (broken cephalis), width 75, width of basal appendage 28.

Remarks: Abdomen is spherical with eighteen longitudinal plicae along an equator in a side view. One row of small pores are arranged in neighboring two longitudinal plicae. The basal appendage of the specimen shown in Fig. 5u is small with unclear circular area. In the specimen of the shown in Fig. 5v, circular area without surrounding ridges (Hatakeda *et al.*, 2007) is wider than that of the specimen shown in Fig. 5u.

Range: *Stj. plicarum* with small appendage (30 to 35 μm) occurs in near the last horizon of this species (Hatakeda *et al.*, 2007). UA Zones 4–5, late Bajocian to latest Bajocian–early Bathonian (Baumgartner *et al.*, 1995). This species occurs from the *Striatojaponocapsa plicarum* to *Cinguloturris carpatica* zones in Outer zone of Southwest Japan (Nishizono, 1996).

Genus ***Tricolocapsa*** Haeckel, 1881

Tricolocapsa undulata (Heitzer, 1930)

(UT-7, Fig. 5w)

1930 *Lithobotrys undulata* n. sp. – Heitzer, p.390, pl. 28, fig. 22.

1987 *Sethocapsa funatoensis* n. sp. – Aita, p. 73, pl. 2, figs. 6a–11; pl. 7, figs. 14, 15.

1993 *Tricolocapsa undulata* (Heitzer) – Ozvoldova and Faupl, pl. 3, fig. 12.

1996 *Sethocapsa funatoensis* Aita – Nishizono, pl. 16, figs. 5, 6.

2003 *Tricolocapsa undulata* (Heitzer) – Suzuki and Gawlick, p. 210, fig. 5.41; fig. 6.39.

2009 *Tricolocapsa undulata* (Heitzer) – Suzuki and Gawlick, p. 183, figs. 5.44A, 5.44B, 5.45A, 5.45B; figs. 6.18A, 6.18B, 6.19A, 6.19B.

Remarks: This species was described by Aita (1987) as *Sethocapsa funatoensis*. This species differs from *Sethoc. yahazuensis* by having rather spinose or pointed nodes on the last segment (Aita, 1987). Suzuki and Gawlick (2003) regarded these two species as younger synonyms of *Lithobotrys undulata* Heitzer.

Range: This species occurs from the *Striatojaponocapsa plicarum* to *Stylocapsa? spiralis* zones in the Outer zone of Southwest Japan (Nishizono, 1996).

Genus ***Stichocapsa*** Haeckel, 1881

Stichocapsa convexa Yao, 1979

(UT-5, Fig. 5x)

1979 *Stichocapsa convexa* n. sp. – Yao, p. 35, pl. 5, figs. 14–16; pl. 6, figs. 1–7.

1986 *Stichocapsa convexa* Yao – Takemura, p. 55, pl. 7, figs. 9, 10.

1996 *Stichocapsa convexa* Yao – Nishizono, pl. 14, fig. 13.

Remarks: Test consists of four segments, conical at upper

half. Forth segment is a truncated sphere with small aperture.

Range: This species occurs mainly from middle of the *Transhsuum hisuikyense* to the uppermost of the *Striatojaponocapsa plicarum* zones in the Outer zone of Southwest Japan (Nishizono 1996).

Stichocapsa magnipora Chiari, Marucci and Prela, 2002 (UT-5, Fig. 5y)

2002 *Stichocapsa magnipora* – Chiari *et al.*, p. 76–77, pl. 3, figs. 13–17.

Remarks: Last segment is inflated with flattened base and aperture. Test has a large depression between the third chamber and the final one. Size of pores is smaller than holotype.

Range: UA Zones 4–7, late Bajocian to late Bathonian–early Callovian. (Chiari *et al.*, 2002).

Family **EUCYRTIDIELLIDAE** Takemura, 1986

Genus ***Eucyrtidiellum*** Baumgartner, 1984

Eucyrtidiellum unumaense (Yao, 1979)

(UT-7, Fig. 5z)

1979 *Eucyrtidium?* *unumaensis* n. sp. – Yao, p. 39, pl. 9, figs. 1–11.

1984 *Eucyrtidiellum putsulatum* Yao – Baumgartner, p. 765, pl. 4, figs. 4, 5.

1986 *Monosera unumaensis* (Yao) – Takemura and Nakaseko, p. 1022, figs. 4.1–4.9.

1990 *Eucyrtidiellum unumaense* (Yao) – Nagai and Mizutani, p. 597, figs. 4.6, 4.7.

1995 *Eucyrtidiellum unumaense putsulatum* Baumgartner – Baumgartner *et al.*, p. 220, pl. 3013, figs. 1, 2.

Remarks: Shell of four segments. Cephalis with a horn where the root remains. Thorax nodose with the sutural pores at distal part. Most of the abdomen is not preserved. However, the part of it with nodes in proximal portion. The fourth segment is lost due to poor preservation. These features show that this specimen similar to *Eucyrtidiellum unumaense putsulatum* Baumgartner *et al.*, 1995.

Range: UA Zones 3–8, early–middle Bajocian to middle Callovian – early Oxfordian (Baumgartner *et al.*, 1995).

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西南日本、山口県に分布するアンモナイトを含む豊浦層群から産出した放散虫

西園 幸久・米光 功雄

要 旨

豊浦層群は西南日本山口県に分布する日本における下部–中部ジュラ系模式地の一つであり、アンモナイト化石を多産する。しかし、放散虫のような微化石の報告は今まで知られていない。豊浦層群最上部7か所から *Transsuum hisuikyoense* 帯および *Striatojaponocapsa plicarum* 帯の放散虫を見出した。これらの放散虫は、中期ジュラ紀 Aalenian から Bathonian を指示すると考えられる。この放散虫指示年代は、アンモナイトやイノセラムスで決定された地質年代よりもやや古い。先行研究によれば *Stj. plicarum* の初出現年代は、Aalenian と Bajocian の境界付近と推定されているにすぎない。この課題を検討するためには、アンモナイトと放散虫の産出間隙から Aalenian を指示するアンモナイトのような化石資料のさらなる蓄積が必要である。

Radiolarian age of Triassic striped chert within the Jurassic accretionary complex of the Ashio terrane in the Ashikaga area, Tochigi Prefecture, central Japan

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ITO Tsuyoshi (2020) Radiolarian age of Triassic striped chert within the Jurassic accretionary complex of the Ashio terrane in the Ashikaga area, Tochigi Prefecture, central Japan. *Bulletin of the Geological Survey of Japan*, vol. 71(4), p. 297–312, 9 figs, 1 table.

Abstract: A striped structure within a single chert bed has been observed in the Tamba–Mino and Ashio terranes, Jurassic accretionary complexes of the Inner Zone of Southwest Japan. This study reports striped chert beds in four sections of the Ashio terrane in Ashikaga and Sano cities of Tochigi Prefecture, namely, Hikoma, Oiwa, Tsukiya and Orihime sections, and their radiolarian ages, except for the Hikoma section. The striped chert comprises streaks and spacing. The streak indicates a thin part of a pin-striped structure within a chert bed and consists mainly of clay minerals. The spacings indicate a thick part between the streaks and is composed mainly of cryptocrystalline quartz. The Oiwa section, which contains numerous striped chert beds, partially corresponds to the middle Carnian–middle Norian (Upper Triassic). The Tsukiya and Orihime sections, which include a few striped chert beds, partially correspond to the middle–upper Anisian (Middle Triassic) and upper Norian–lower Rhaetian (Upper Triassic), respectively.

Keywords: radiolaria, accretionary complex, Ashio terrane, striped chert, Triassic

1. Introduction

Chert is a hard and dense microcrystalline or cryptocrystalline sedimentary rock (Bates and Jackson, 1984) and is one of the major components of Jurassic accretionary complexes (ACs) of East Asia. The age of chert had not been determined because index fossils were unobtainable. Extraction methods for microfossils such as conodont and radiolaria had been proposed in the 1960s to 1970s (e.g. Hayashi, 1968, 1969; Pessagno and Newport, 1972). This development made microfossils valuable age indexes and allowed the age of the Jurassic ACs to be clarified in the strata which had been treated as the Paleozoic (e.g. Yao and Mizutani, 1993; Isozaki *et al.*, 2010; Agematsu-Watanabe and Kamata, 2018). The age of chert within the Jurassic ACs ranges from Pennsylvanian (Carboniferous) to Late Jurassic (e.g. Matsuoka *et al.*, 1998; Nakae, 2000).

As stated above, microfossils within chert such as radiolaria are valuable index fossils. However, they cannot be always obtained from chert because of several reasons such as the diagenetic effect. The determination of an alternative age index for chert would be valuable for the investigation of geologic units without fossils.

The author discovered a striped structure within a single chert bed in the Ashikaga area of Tochigi Prefecture,

Japan. The structure was reported from the Tamba–Mino and Ashio terranes in the previous studies (e.g. Iijima *et al.*, 1978; Kido, 1982; Yoshimura *et al.*, 1982; Kakuwa, 1991; Nikaido and Matsuoka, 2008, 2009, 2011). Based on these previous studies, the occurrences of the striped chert might be dominant in a specific age.

This study describes the striped chert beds in the Ashikaga area and determines their microfossil ages. Furthermore, the ages of striped chert reported in the previous studies are compiled for future references. This information would contribute to the discussion about the potential of the striped chert as an alternative age index for the Jurassic ACs in the Tamba–Mino and Ashio terranes.

2. Previous studies on the striped structure and terminology

Structures within chert beds were described in several international studies, particularly in the 1970s (e.g. Davis, 1918; Bastin, 1933; McBride and Thomson, 1970; Folk, 1973; Lowe, 1976; McBride and Folk, 1979). A research group from the University of Tokyo studied structures within the chert beds within the Jurassic ACs of Southwest Japan in the 1970s to early 1990s (e.g. Iijima *et al.*, 1978, 1979, 1985; Iijima and Utada, 1983; Kakuwa, 1991). Imoto (1983, 1984a, 1984b) focused on several characteristics

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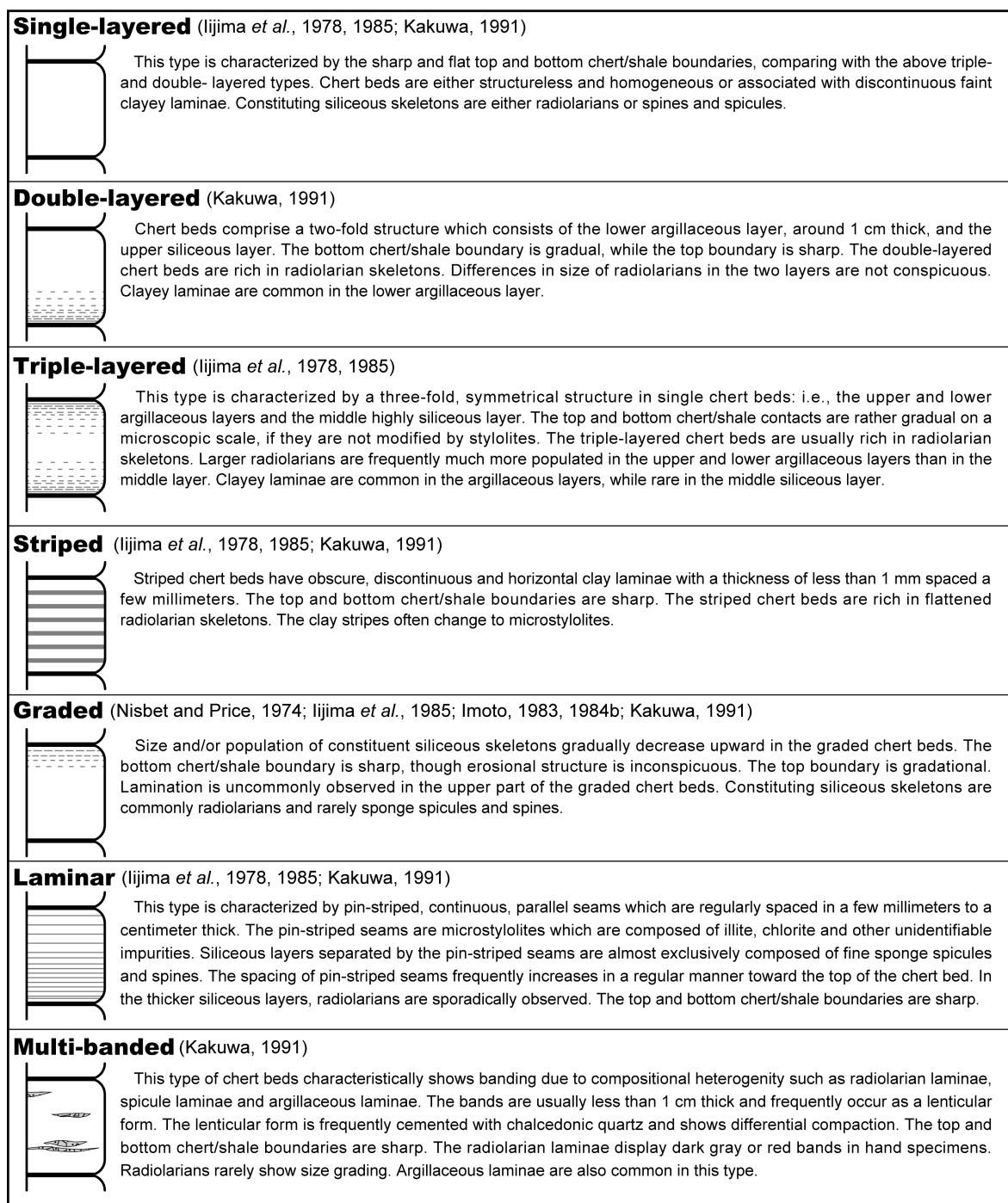


Fig. 1 Classification of the structure within a single chert bed based on previous studies with their characteristics and simplified images. The simplified images are based on Iijima and Utada (1983) and Kakuwa (1991). The characteristics of each structure are from Kakuwa (1991).

of chert, such as color and bed thickness, and classified their differences by age. Sugiyama (1997) classified chert into two types: B-type chert rather argillaceous and rich in siliceous organic remains; F-type chert highly silicified with very rare organic remains. The B- and F-types sensu Sugiyama (1997) can be corresponded to the B- and F-types sensu Imoto (1984b), respectively. Furthermore,

Sugiyama (1997) proposed A-type alternating occurrences of B- and F-type chert beds within short interval (ca. less than 1 m).

Figure 1 shows classification of structures within chert bed and their images in the previous studies. Iijima *et al.* (1978) recognized four types of single chert beds showing the Triassic age: single-layered, triple-layered, laminar

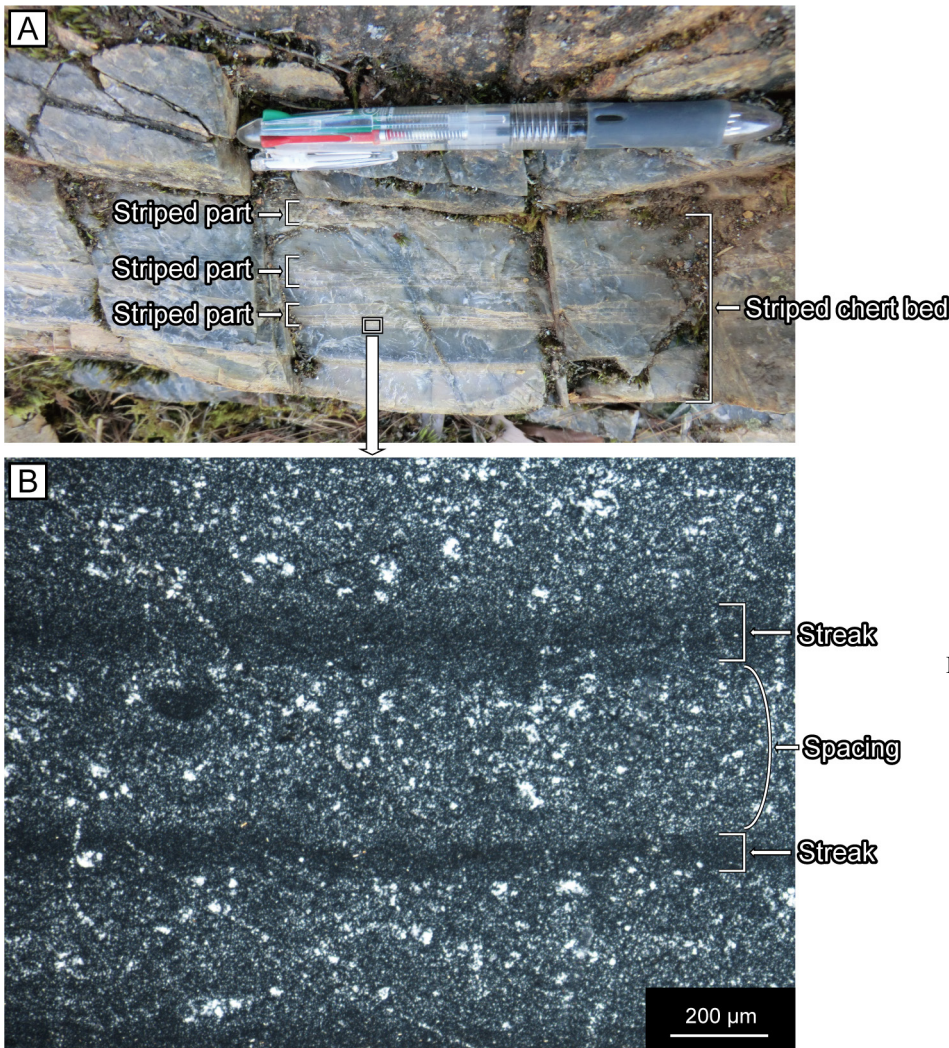


Fig. 2 Terminology of the description in this study. Field occurrence (A) and thin section (B) are from the lower part of the Tsukiya section. Striped chert bed: chert bed including striped part(s). Striped part: dominant part of the pin-striped structures. Streak: thin part of the pin-striped structure within the striped part. Spacing: a thick part between two streaks.

and striped. Iijima and Utada (1983) and Iijima *et al.* (1985) compiled siliceous rocks from Japan and reported a new graded type of single chert beds. Furthermore, Kakuwa (1991) recognized two additional types of chert beds: double-layered and multi-banded. Kakuwa (1991) synonymized the stylolite chert sensu Yoshimura *et al.* (1982) with the laminar chert. Sugiyama (1997) noted that the B-type is predominated by single-layered chert, whereas F-type by various degrees of triple-layered, striped and laminar types of Kakuwa (1991). Nikaido and Matsuoka (2009) described striped chert and classified them into three types according to the thickness changes of spacings between streaks in single chert beds: constant in thickness, thinning upward and thickening upward.

Both striped and laminar chert is characterized by having horizontal clay lamina-like structure (Kakuwa, 1991). Based on the description by Kakuwa (1991), the major difference between the striped and laminar chert is richness of the radiolarian skeletons. However, it can be greatly affected by secondary factors such as a surrounding igneous activity. The lamina-like structure of the striped chert often changes to microstylolite, and those

of the laminar chert do likewise. This article regards the laminar chert as a type of the striped chert, i.e. the striped chert that has stylolite and less radiolarian skeletons with thickening-upward spacing is the laminar chert. The previous studies also used the name of “varved chert” (Yamada *et al.*, 1985; Mizutani and Koido, 1992; Nakano *et al.*, 1995); however, the term of “varve” is an annual layer and has a petrogenetic meaning. For these reasons, the stylolite and varved chert are used synonymously with the striped chert in this study.

The terminology of the description in this article is shown in Fig. 2. Striped chert bed is a chert bed including the striped part(s) (Fig. 2A). The striped part is dominant part of the pin-striped structures composed of streaks and spacing (Fig. 2B). The streak indicates a thin part of the pin-striped structure. The spacing indicates a thick part between the streaks.

3. Geologic setting

The Tamba–Mino and Ashio terranes are distributed over the Inner Zone of Southwest Japan (Nakae, 2000;

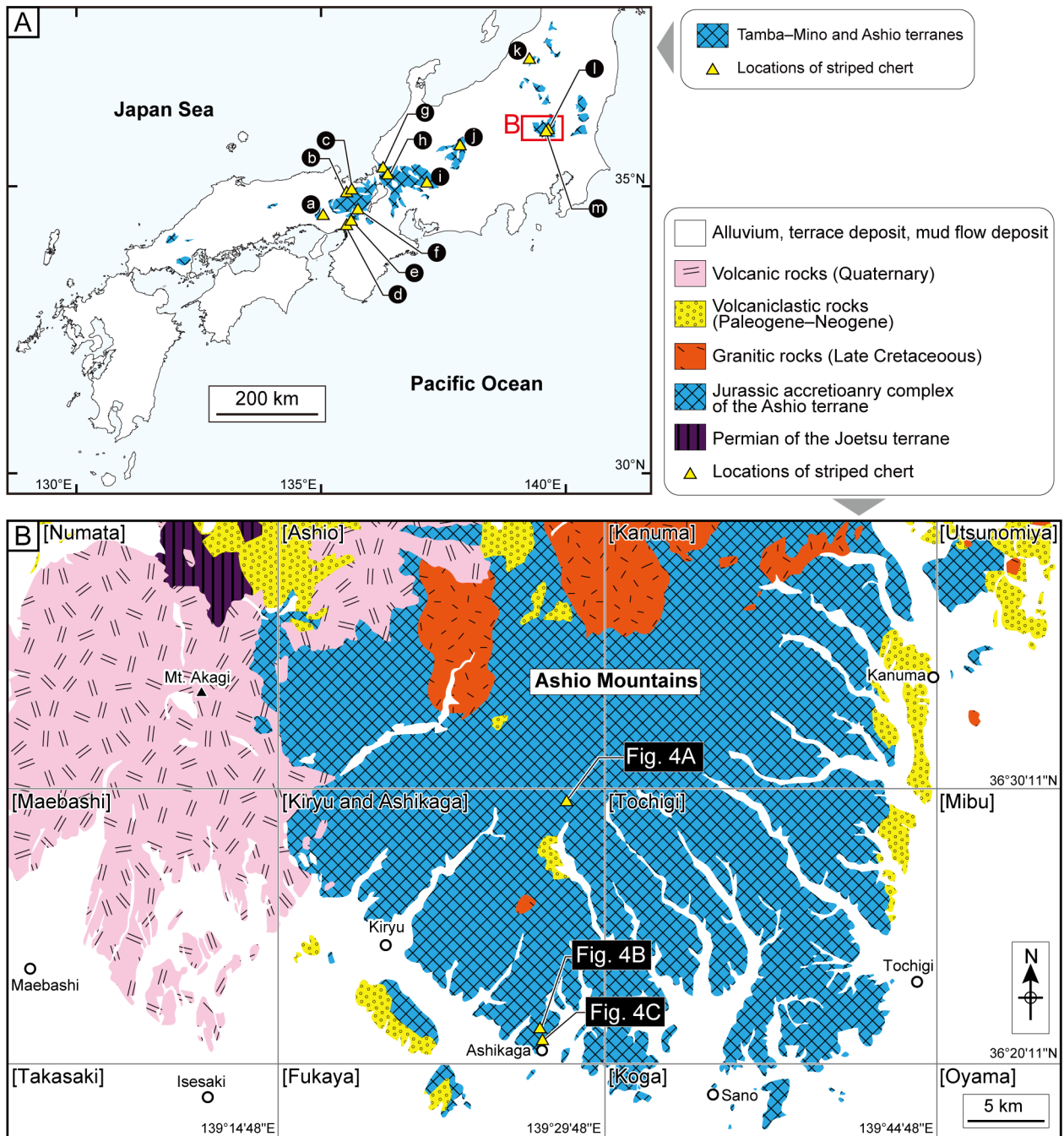


Fig. 3 Index maps of the study area. (A) Distribution of the Tamba-Mino and Ashio terranes (modified from Geological Survey of Japan, AIST, 2018) with locations of striped and related chert in the present and previous studies. The alphabets of lower-case within the black circle correspond to the locations (Loc.) shown in Table 1. (B) Simplified geologic map of the Ashio terrane in the Ashio Mountains (modified after Sudo *et al.*, 1991; Geological Survey of Japan, AIST, 2018). The geographical names in brackets indicate 1:50,000 topographic maps published by the Geospatial Information Authority of Japan.

Kojima *et al.*, 2016) (Fig. 3A). The Tamba-Mino terrane corresponds to the Ashio terrane. The terranes are mainly composed of late Carboniferous-Permian ocean ridge basalt and pelagic carbonate, late Carboniferous-Jurassic pelagic chert, Jurassic hemipelagic siliceous mudstone and trench-fill clastics, such as mudstone and sandstone (e.g. Nakae, 2000).

The components of the Ashio terrane are exposed in the Ashio Mountains in central Japan (Fig. 3B). The occurrences of microfossils such as radiolaria and conodont have been reported in this area since the 1960s (e.g. Hayashi, 1963, 1968; Koike *et al.*, 1971; Hayashi and Hasegawa, 1981; Aono, 1985; Kamata, 1996, 1997; Takayanagi *et al.*, 2001; Ito, 2019, 2020).

Kamata (1996) divided the Ashio terrane of the Ashio Mountains into three tectonostratigraphic units, namely, the Kuzu, Kurohone–Kiryu and Omama complexes. The Kuzu Complex is characterized by coherent facies and is mainly composed of repeated chert–clastic sequences with basalt–limestone blocks. The Kuzu Complex can be subdivided into three units: 1, 2 and 3 (Kamata, 1997). The Omama and Kurohone–Kiryu complexes are characterized by mixed facies that is represented by muddy mixed rocks including several types and sizes of blocks. Both complexes include chert, limestone and sandstone blocks. The Omama Complex contains large amounts of basalt and limestone, whereas the Kurohone–Kiryu Complex includes small amounts of these rocks.

4. Striped chert in the Ashikaga area

4.1 Study sections

This study investigated four sections in Sano and Ashikaga cities, Tochigi Prefecture (Fig. 4). The Hikoma, Tsukiya, Orihime and Oiwa sections include one to several striped chert beds (Fig. 5). According to the tectonostratigraphic division by Kamata (1996) and the geological maps of Sudo *et al.* (1991), the Hikoma section is located in the distributional area of the Kurohone–Kiryu Complex, whereas the Oiwa, Tsukiya and Orihime sections in unit 3 of the Kuzu Complex. The chert of all sections predominantly corresponds to the F-type chert sensu Imoto (1984b) and Sugiyama (1997).

The Hikoma section crops out along the prefectural road 208 in Hikoma, Sano City (Fig. 4A). The total stratigraphic thickness of this section is about 4 m. The chert in this section is generally dark-gray and weakly bedded. The single bed thickness of the individual chert is 5–15 cm and accompanies claystone with sizes less than 1 mm. One striped chert bed is interbedded.

The Oiwa section is exposed on a hiking trail in Oiwa-cho, Ashikaga City and is located above the Oiwa Tunnel (Fig. 4B). The total stratigraphic thickness of this section is about 15 m. The chert from this section is generally bright-gray and clearly bedded. The single bed thickness of the individual chert is 3–15 cm and accompanies claystone with sizes less than 1 mm. Numerous striped chert beds can be observed in this section. The lower part of this sequence is folded.

The Tsukiya section crops out along a road in Tsukiya-cho, Ashikaga City (Fig. 4B). The total stratigraphic thickness of this section is about 2 m. The chert of this section is generally gray and clearly bedded. The single bed thickness of the chert is 3–15 cm and accompanies claystone with sizes less than 1 mm. The gray chert includes a 1–2 cm wide black-colored part. Five striped chert beds are observed.

The Orihime section crops out along the road to the Orihime-jinja Shrine in Nishinomiya-cho, Ashikaga City (Fig. 4C). The total stratigraphic thickness of this section is about 3 m. The chert from this section is generally gray

to dark-gray and bedded. The single bed thickness of the gray chert is 5–10 cm and accompanies claystone with sizes less than 1 mm. Four striped chert beds are observed. Folded chert is recognized in the middle part.

4.2 Characteristics of striped chert

The striped chert is continuously and laterally distributed, at least at the outcrop (Fig. 6A, B). The maximum thickness of the streaks is about 1 mm (Fig. 6C). The streaks are almost parallel to the bed surface (Fig. 6B, C). A large amount of the striped chert is observed in the Oiwa section (Fig. 6D). The streaks are also folded by a small fault within the bed (Fig. 6E).

The color of the streaks are generally paler than the spacings at the outcrop. For example, the streaks in the Tsukiya and Oiwa sections are caramel color whereas the spacings are dark-gray (Fig. 6). However, marginal parts of the spacings along the streaks represent paler color like the streaks. Consequently, the color of the streaks may be due to fading during the diagenesis.

The thin section observations indicate that the striped chert is composed of cryptocrystalline quartz and a few clay minerals (Fig. 7). The streaks are almost parallel (Fig. 7A, B), and the thickness of the spacing between the streaks is several hundred micrometres. The streaks are composed mainly of clay minerals; the spacings consist mainly of cryptocrystalline quartz. The streaks are composed of finer material than that found in the spacing (Fig. 7B, C). The streaks are stylolitic in some samples (Fig. 7D, E). The stylolitic seam is about 5 μm in thickness and is slightly undulated.

5. Microfossil occurrence and age assignment

A total of 17 chert samples were collected from the study sections. The following methods were used to extract microfossils from the samples. The samples were crushed into about 1 cm fragments and then soaked in hydrofluoric acid (ca. 5 %) at room temperature (ca. 20 °C–25 °C) for 24 h. The residues were collected using a sieve with a mesh diameter of 0.054 mm and then enclosed within a slide prepared with a photocrosslinkable mounting medium (GJ-4006, Gluelabo Ltd.). The slides were analyzed using a transmitted light microscope and then photographed. Several specimens from the residues were mounted on stubs, analyzed and then photographed using scanning electron microscopy.

Six chert samples, including one striped chert sample, were collected from the Hikoma section. However, microfossils for age determination could not be obtained from these samples.

Four samples of striped chert were collected from the Oiwa section. Two samples (IT18101413 and IT18101415) yielded radiolarian remains. The twisted spine (Fig. 8A) from sample IT18101413 resembles a spine of *Capnuhosphaera deweveri* Kozur and Mostler. On the basis of the occurrence range of Sugiyama

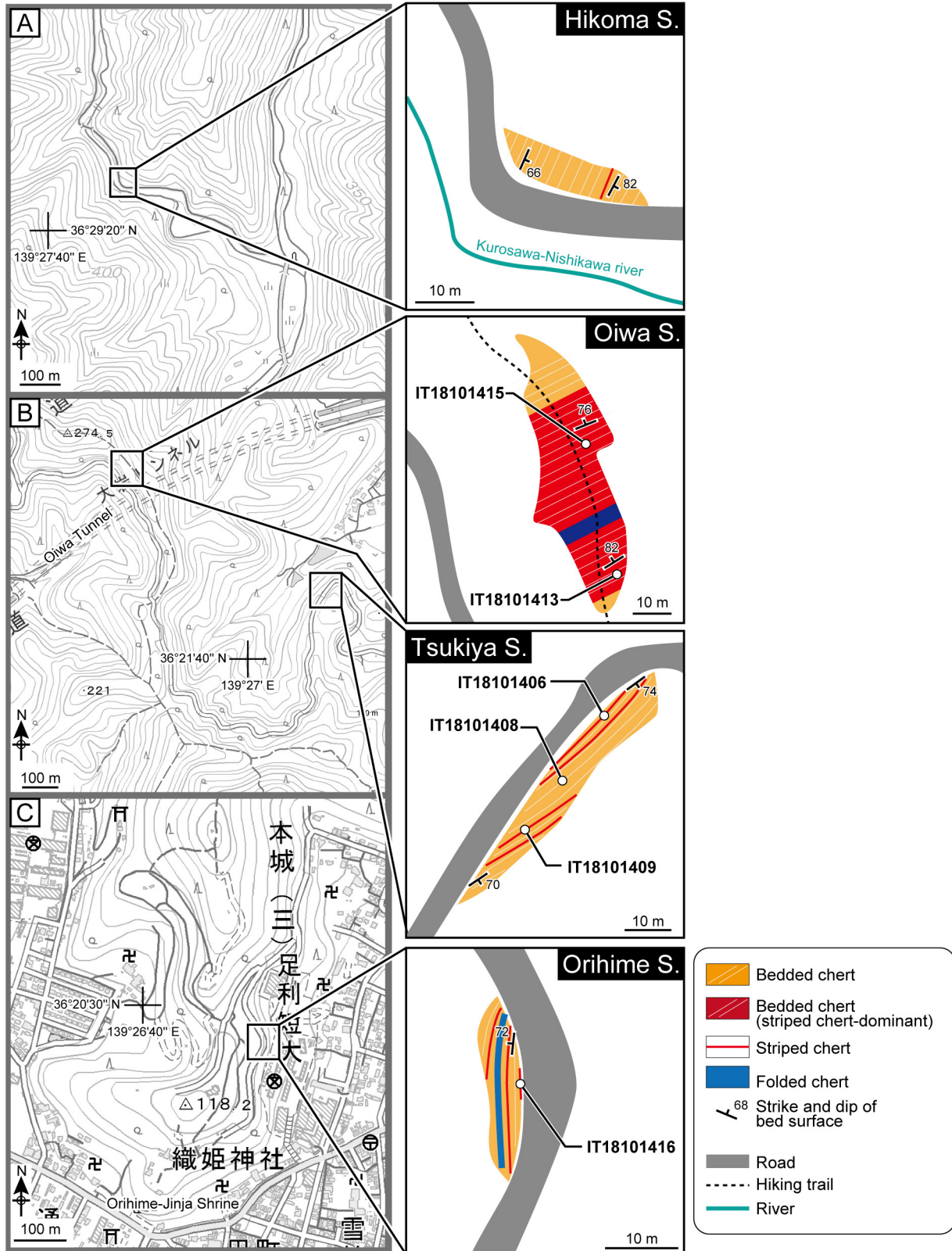


Fig. 4 Traverse maps of sections studied in this article. The maps are modified from the 1:25000 map of ‘Bamba’ and ‘Ashikaga-Hokubu’ by the Geospatial Authority of Japan. S.: section.

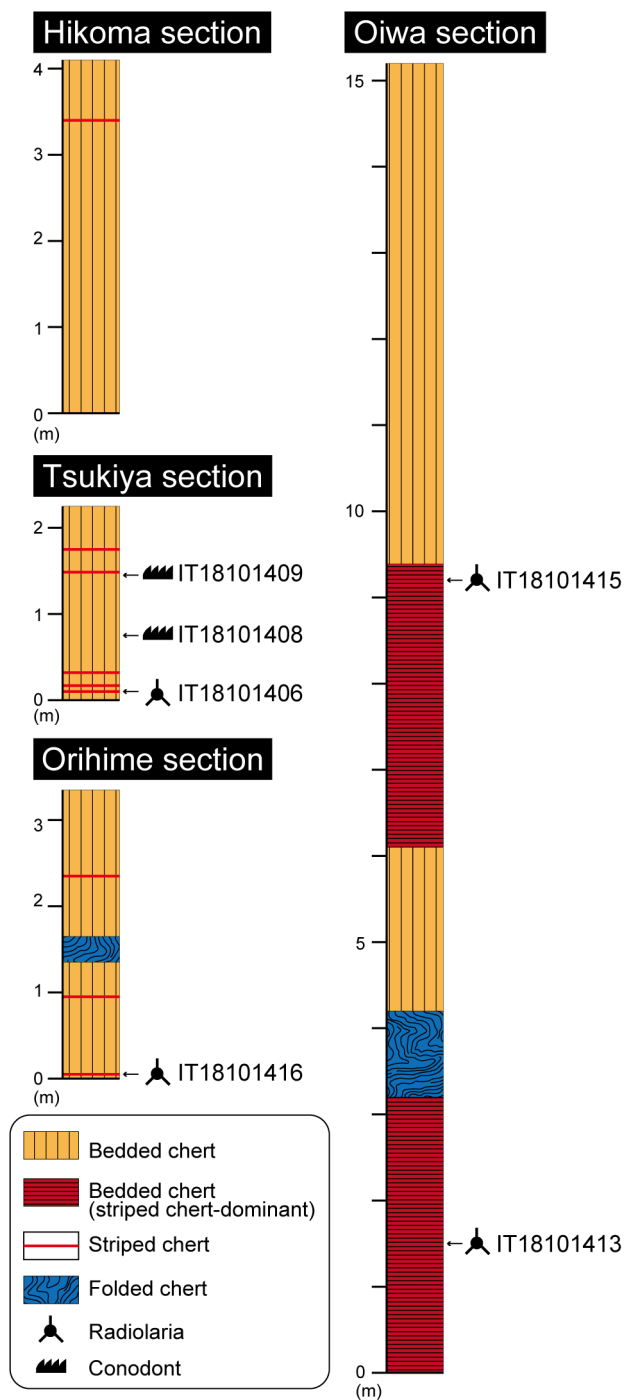


Fig. 5 Columnar illustration of the studied sections with microfossil horizons.

(1997), *C. deweveri* occurred in the *Capnuchosphaera* Lowest-Occurrence Zone (TR5A) to the *Trialatus robustus*–*Lysemelas olbia* Partial-Range Zone (TR6B). The results of this study indicate that sample IT 18101413 corresponds to the middle Carnian–middle Norian.

Six chert samples, including two striped chert samples, were collected from the Tsukiya section. Two chert samples (IT18101408 and IT18101409) yielded conodont fragments

(Fig. 8B, C). One striped chert sample (IT18101406) yielded radiolarian fossils: *Paroertlispongus*? sp. (Fig. 8D), *Pararchaeospongoprunum*? sp. (Fig. 8E), twisted spine (Fig. 8F) and *Spinotriassocampe*? sp. (Fig. 8G). According to O’Dogherty *et al.* (2009), *Paroertlispongus*, *Pararchaeospongoprunum* and *Spinotriassocampe* occurred in the middle Anisian–lower Carnian, upper Permian–upper Anisian and middle Anisian–lower Carnian, respectively. Consequently, sample IT 18101406 might correspond to the middle–upper Anisian.

Two chert samples, including one striped chert sample, were collected from the Orihime section. Only one sample from the striped chert bed (IT18101416) yielded *Lysemelas* sp. cf. *L. olbia* Sugiyama (Fig. 8H). On the basis of the occurrence range reported by Sugiyama (1997), *L. olbia* occurred in the *L. olbia* Lowest-Occurrence Zone (TR7) to Skirt F Lowest-Occurrence Zone (TR8C). Sample IT 18101416 might correspond to the upper Norian–lower Rhaetian.

6. Ages of striped chert in other areas

6.1 Tamba–Mino and Ashio terranes

Several researchers have noted the presence of the striped chert in the Tamba–Mino and Ashio terranes (Fig. 3A). In this chapter, the previous studies on the striped chert and their ages are reviewed (Table 1).

Kido (1982) described four sections including striped chert beds, namely, the Kashibara, Kamiaso Bridge, Hisuikyo and Hosobi-dani sections, and determined the ages of the former three sections. The striped chert bed in the Kashibara section is about 8 m below the sampling point of sample KC1 yielding *Praemesosaturnalis gracilis* Kozur and Mostler, which occurred in TR8A to TR8C (Sugiyama, 1997). One striped chert bed in the Kamiaso Bridge section is located between samples BC2 and BC3, which include several species of the genus *Pseudostylosphaera* Kozur and Mostler. *Pseudostylosphaera compacta* (Nakaseko and Nishimura) is observed in both samples, and *Pseudostylosphaera goestlingensis* (Kozur and Mostler) is observed in BC3. The occurrence range of the latter species is TR4A to TR5A (Sugiyama, 1997). Another chert bed in the Kamiaso Bridge section is located between samples BC4 and BC5. Sample BC5 yielded *Capnodoce sarisa* De Wever, which occurred in TR6A to TR7 (Sugiyama, 1997). One striped chert bed in the Hisuikyo section is about 5 m below sample HC1, which yielded *Yeharaia elegans* Nakaseko and Nishimura. The occurrence range of *Y. elegans* is TR3B to TR4A, Ladinian (Sugiyama, 1997). Another chert bed in the Hisuikyo section is located between samples HC2 and HC3. Sample HC2 yielded *Pentactinocarpus* sp. cf. *P. fusiformis* Dumitrica and sample HC3 yielded *Hexasaturnalis hexagonus* (Yao). *Pentactinocarpus fusiformis* occurred in TR3B (Sugiyama, 1997). Meanwhile, an occurrence range of *H. hexagonus* is from the late *Trillus elkhornensis* zone (JR2)

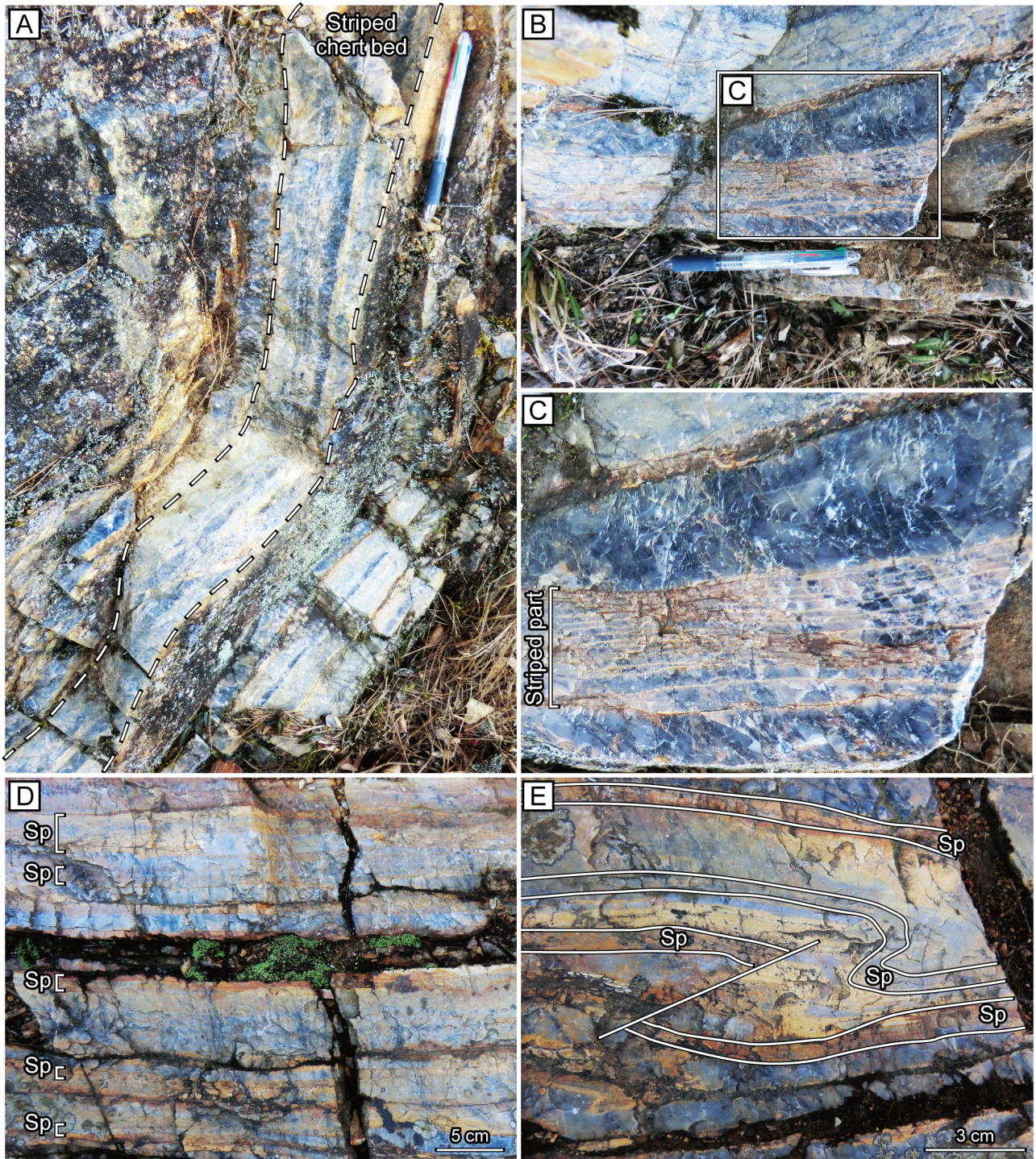


Fig. 6 Field occurrences of striped chert beds. (A) Laterally continuous striped chert beds from the lower part of the Tsukiya section. (B, C) Striped chert bed from the lower part of the Tsukiya section. The striped part is observed in the middle part within the striped chert bed. The bed surface and the streaks are almost parallel. (D) Bedded chert containing several striped parts (Sp) from the lower part of the Oiwa section. (E) Fold and fault within the chert bed and deformed striped parts (Sp) from the lower part of the Oiwa section. The streaks were deformed along the deformation of the chert beds by the fold and fault.

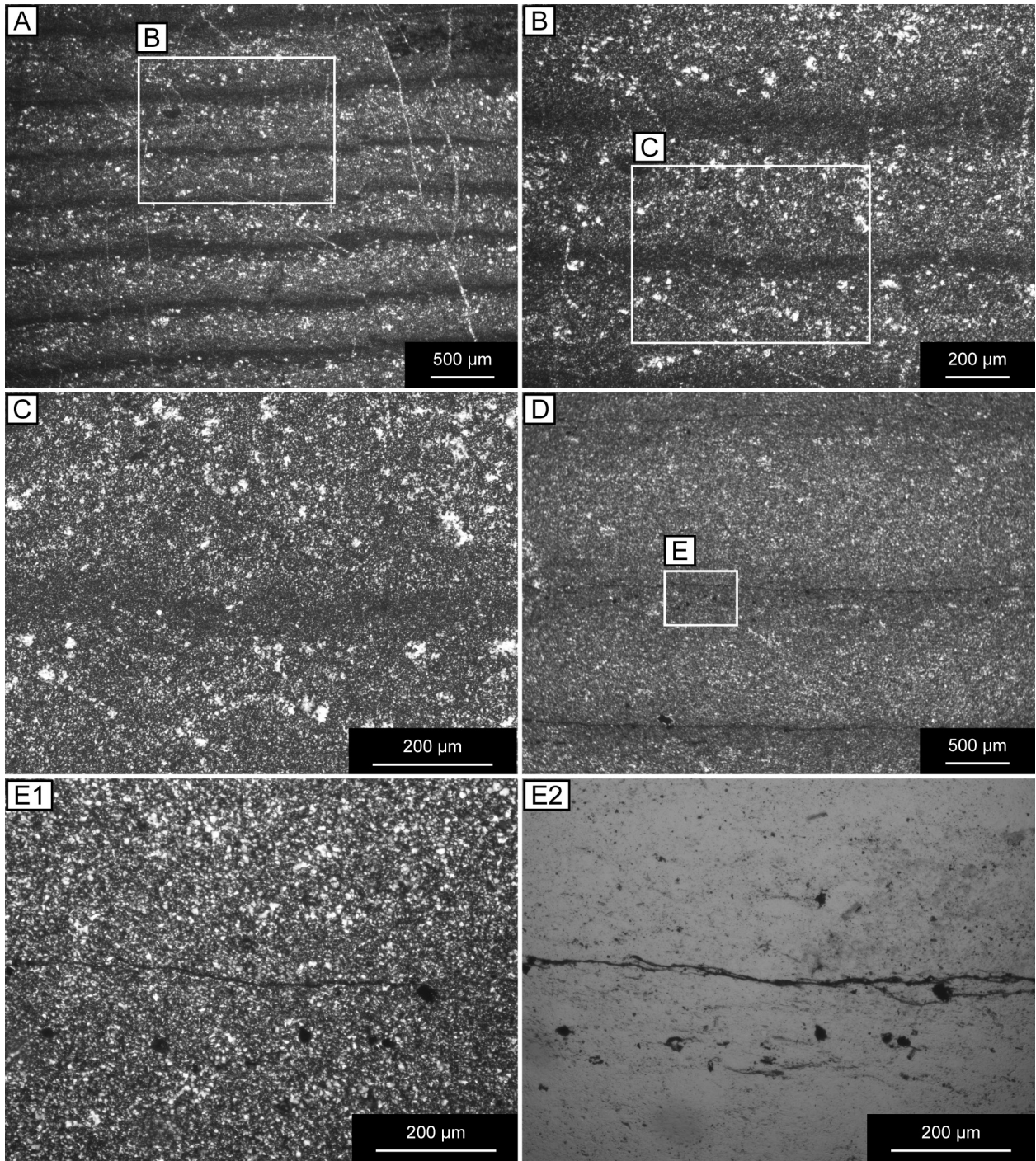


Fig. 7 Thin sections of the striped chert consisting of cryptocrystalline quartz and a few clay minerals. (A–C) Striped chert from the lower part of the Tsukiya section (sample IT18101406). The streaks are almost parallel. The thickness of the spacing is several hundred micrometres. The streaks are composed of finer material (mainly of clay minerals) than that found in the spacings. (D, E) Striped chert from the upper part of the Hikoma section (sample IT18101404). The streaks are stylolitic and slightly undulated in some samples. (A–D, E1) Crossed nicols. (E2) Open nicols.

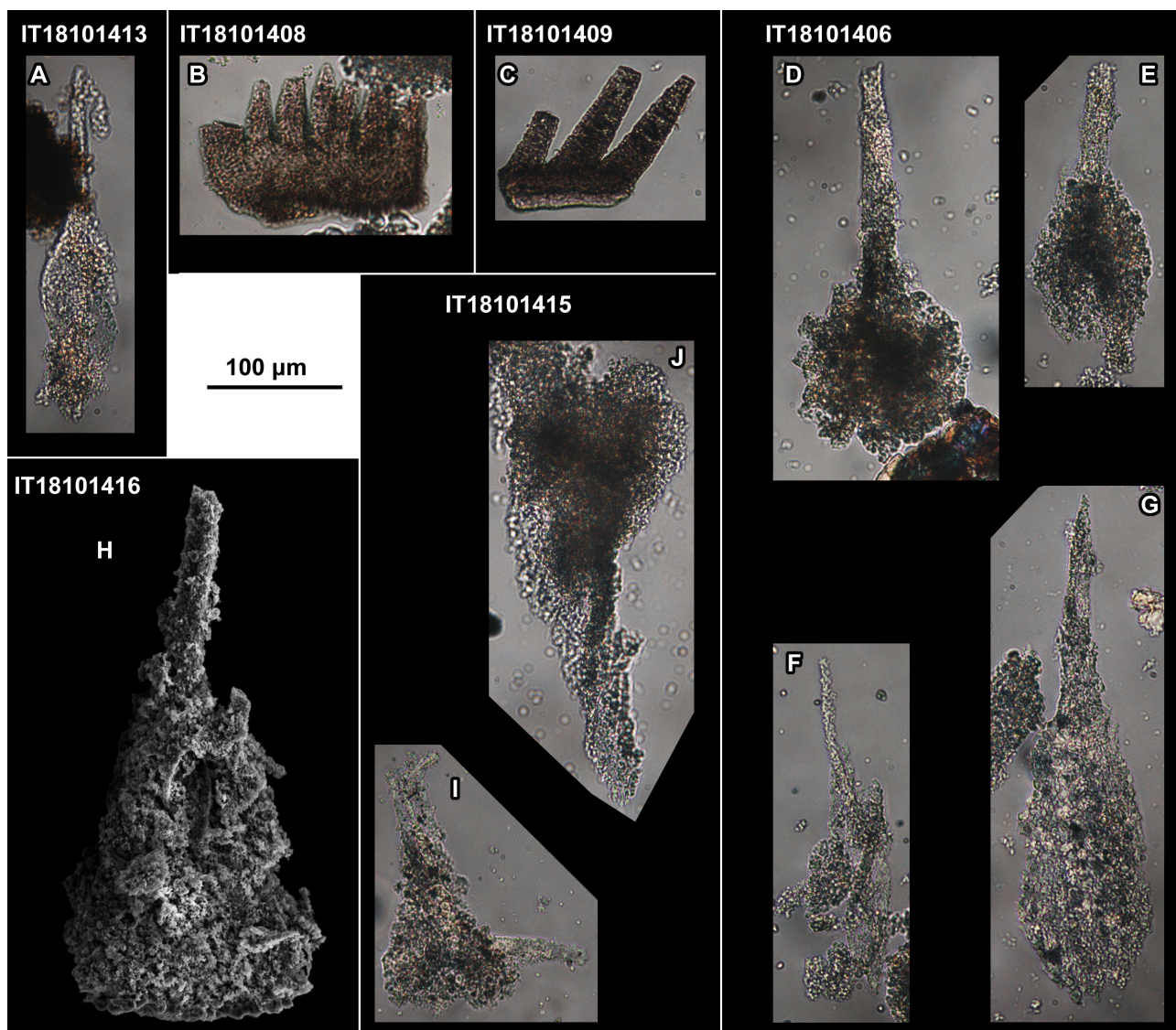


Fig. 8 Radiolarian and conodont fossils obtained from the studied sections. (A, F) Twisted spine. (B, C) Conodont fragment. (D) *Paroertlispongus?* sp. (E): *Pararchaeospongoprunum?* sp. (G) *Spinotriassocampe?* sp. (H) *Lysemelas* sp. cf. *L. olbia* Sugiyama. (I, J) *Spumellaria* gen. et sp. indet. with twisted spine.

to *Striatojaponocapsa plicarum* zone (JR4) (Matsuoka, 1995) corresponding to the Toarcian–middle Bathonian, late Early–Middle Jurassic (Matsuoka and Ito, 2019). Consequently, the age of the striped chert bed between samples HC2 and HC3 is some age between the Late Triassic and Early Jurassic.

Yoshimura *et al.* (1982) described “stylolitic chert” in the Imajo area of the Nanjo Massif, Fukui Prefecture. The age of the “stylolitic chert” was undetermined.

Imoto (1984a, b) studied Paleozoic and Mesozoic chert in the Tamba area. Imoto (1984b) noted that both Permian and Triassic–Jurassic chert contains sedimentary structures such as parallel lamination, and showed their photographs.

Yamada *et al.* (1985), Mizutani and Koido (1992) and Nakano *et al.* (1995) noted the presence of “varved chert”.

Yamada *et al.* (1985) stated that the age of the “varved chert” is the Triassic.

Kakuwa (1991) focused on structures within single chert beds and described them in several areas. Most of the striped chert beds of these sections correspond to the Triassic (Table 1). Among the sections including the striped chert beds studied by Kakuwa (1991), better ages were obtained from four sections by using conodont and radiolaria. The Koze section corresponds to the Spathian, Anisian and Norian; the Unuma section corresponds to the Norian; the Hisuikyo section corresponds to the Carnin–Rhaetian; and the Karasawa section corresponds to the Carnian.

The Sakahogi section, which is exposed along the Kiso River, mainly consists of successive Triassic bedded chert. Therefore, most researchers have studied this section and

Table 1 The description of the striped chert and related chert based on previous studies. The alphabets of lower-case in the locations (Loc.) correspond to Fig. 3A.

Reference	Loc.	Area or section	Chert description	Stage, Series and System
Kido (1982)	i	Kashibara section	Striped chert	Norian–Rhatian (Upper Triassic)
	i	Kamiaso Bridge section	Striped chert	Ladinian–Carnian (Middle–Upper Triassic)
	i	Hisuikyo section	Striped chert	Triassic–Jurassic?
	i	Hosobi-dani section	Striped chert	-
Yoshimura <i>et al.</i> (1982)	g	Imajo	Stylolitic chert	-
Imoto(1984b)	c	Yagi	Parallel lamination	middle Permian and Triassic
Yamada <i>et al.</i> (1985)	j	Takayama	Varved chert	Upper Triassic
Kakuwa (1991)	d	Kinzoji section	Laminar chert	Triassic–Jurassic
	e	Kurio section	Laminar chert	Upper Triassic
	a	Kuwahara section	Striped chert	Triassic–Jurassic
	f	Okuhatcho section	Laminar chert; Striped chert	Triassic–Jurassic
	b	Ohtaki-kita section	Laminar chert	Triassic–Jurassic
	d	Koze section	Striped chert	Spathian, Anisian, Norian (Lower, Middle and Upper Triassic)
	i	Unuma section	Laminar chert	Norian (Upper Triassic)
	i	Hisuikyo section	Laminar chert	Carnian–Rhaetian (Upper Triassic)
	h	Kammuriyama section	Laminar chert	Triassic–Jurassic
	m	Yamamae section	Laminar chert	Triassic–Jurassic
Mizutani and Koido (1992)	l	Karasawa section	Laminar chert; Striped chert	Carnian (Upper Triassic) and lower Jurassic
	k	Kuromatagawa section	Laminar chert	Triassic–Jurassic
	i	Kanayama	Varved chert	-
Nakano <i>et al.</i> (1995)	j	Norikuradake	Varved chert	-
Nikaido and Matsuoka (2009)	i	Sakahogi section	Striped chert	upper Ladinian–upper Norian (Middle–Upper Triassic)
This study	m	Hikoma section	Striped chert	-
	m	Oiwa section	Striped chert	middle Carnian–middle Norian (Upper Triassic)
	m	Tsukiya section	Striped chert	upper Anisian (Middle Triassic)
	m	Orihime section	Striped chert	upper Norian–lower Rhaetian (Upper Triassic)

the surrounding area in detail (e.g. Sugiyama, 1997; Onoue *et al.*, 2012, 2016, 2017; Nozaki *et al.*, 2019). Nikaido and Matsuoka (2009) studied the stratigraphic distribution of the striped chert beds of the Sakahogi section and correlated it with the radiolarian zonation proposed by Sugiyama (1997). On the basis of the correlation, the first occurrence of the striped chert is TR4A, their dominant

intervals are in TR5A to TR6B, and their last occurrence is observed in TR8A (Fig. 9).

6. 2 Permian and Jurassic ACs in other areas

Permian and Jurassic ACs other than the Tamba–Mino and Ashio terranes exist in the Japanese Islands such as the Akiyoshi terrane (Permian AC) of the Inner Zone

of Southwest Japan, Chichibu composite terrane of the Outer Zone of Southwest Japan and the North Kitakami terrane of Northeast Japan (e.g. Kojima *et al.*, 2016). Furthermore, the Jurassic ACs extend to adjacent regions of Japan, such as Northeast China, Far Eastern Russia and the Philippines (e.g. Kojima and Kametaka, 2000). The occurrence of the striped chert in geologic units other than the Tamba–Mino and Ashio terranes is noted in this section for future reference.

According to the results of fieldworks (e.g. Ito and Matsuoka, 2018 and reference therein), the presence of striped chert beds has rarely been noted in the Chichibu composite terrane. The author discovered striped chert beds near Mt. Gusuku on Ie Island, Okinawa Prefecture. Although the age of the striped chert beds has not been determined, the chert in Ie Island ranges from upper Permian to Lower Jurassic (Shen *et al.*, 1996; Ito and Matsuoka, 2017). The author could not find the description of stripe chert beds from the North Kitakami terrane and corresponding geologic units in Northeast China and Far East Russia. The author studied the Akiyoshi terrane (Ito and Matsuoka, 2015, 2016); however, according to field works and the literature, striped chert beds have never been found.

Matsuoka (2002) and Matsuoka *et al.* (2009) noted a large number of striped chert beds in the North Palawan Block of the Philippines which is the southwestern extension of the Jurassic ACs on the Japanese Islands. Matsuoka *et al.* (2009) found Early Jurassic radiolarians from intervals of dominant striped chert beds in the North Palawan Block.

Previous studies proposed that global cyclic phenomena, such as the 20 kyr- to Myr-scale Milankovitch cycle, led to the rhythmical bedding of chert (e.g. Hori *et al.*, 1993; Ikeda *et al.*, 2010, 2017). Matsuoka *et al.* (2009), however, noted that the mm-scale striped structure within a single chert bed is probably related to shorter periodic events on ~kyr timescale. Therefore, other factors might have affected the formation of striped chert beds.

7. Concluding remarks

The ages of the sections including the striped chert beds in the Ashikaga area were determined in this study. The Tsukiya and Orihime sections, including a few striped chert beds, correspond to the middle–upper Anisian and upper Norian, respectively. The Oiwa section, which contains numerous striped chert beds, corresponds to the middle Carnian–lower Norian. The results of the present study are consistent with the description of Nikaïdo and Matsuoka (2009) (Fig. 9). Likewise, the age determinations for other areas (e.g. Kido, 1982; Kakuwa, 1991) indicated that the striped chert beds are generally the Carnian–lower Norian.

Consequently, the striped chert might be used as an alternative age index, at least for studies in the Tamba–Mino and Ashio terranes. That is, several striped chert beds likely indicate the Triassic age, and the dominant interval

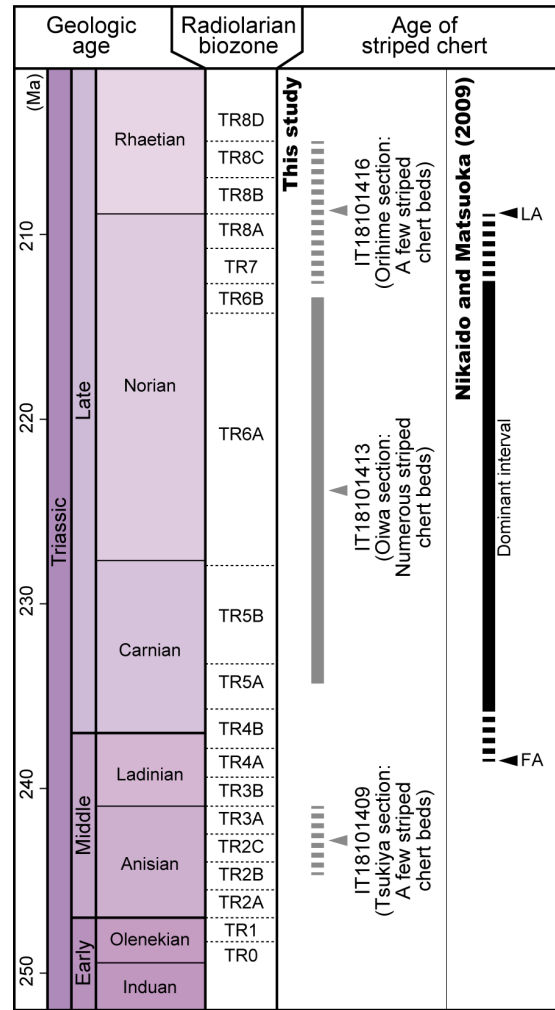


Fig. 9 Geologic time scale of the Triassic including the distribution of striped chert beds. Geologic time scale is after Ogg *et al.* (2016). Radiolarian biozones are after Sugiyama, 1997 and are partially modified based on the calibration by Yamashita *et al.* (2018).

of the striped chert beds is probably the Carnian–lower Norian. Meanwhile, a few striped chert beds correspond to the Permian (Imoto, 1984a) and Lower Jurassic (Kakuwa, 1991). Further descriptions and age assignments of striped chert beds, including petrogenetic work, will provide additional information on past ocean conditions.

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栃木県足利地域の足尾テレーンジュラ紀付加体に含まれる三畳系ストライプチャートの放散虫年代

伊藤 剛

要 旨

チャートの単層中に発達するストライプ構造は、西南日本内帯のジュラ紀付加体丹波 - 美濃テレーン及び足尾テレーンで見られる。本研究では、栃木県足利市及び佐野市の4セクション（飛駒・大岩・月谷・織姫セクション）中にみられるストライプチャートについて記載するとともに、飛駒セクションを除く3セクションの放散虫化石年代について検討する。ストライプチャートは、ストリークとスペーシングからなる。ストリークは、ピンストライプ状構造における薄い部分を指し、主に粘土鉱物からなる。スペーシングは、ストリークの間厚い部分であり、隠微晶質石英を主体とする。ストライプチャートに富む大岩セクションは、上部三畳系カーニアン階中部～ノーリアン階中部に部分的に対比される。ストライプチャートを部分的に含む月谷セクションと織姫セクションは、中部三畳系アニシアン階中部～上部と上部三畳系ノーリアン階上部～レーティアン階下部にそれぞれ対比される。

Late Jurassic radiolarians from mudstone near the U–Pb-dated sandstone of the North Kitakami Belt in the northeastern Shimokita Peninsula, Tohoku, Japan

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UCHINO Takayuki and SUZUKI Noritoshi (2020) Late Jurassic radiolarians from mudstone near the U–Pb-dated sandstone of the North Kitakami Belt in the northeastern Shimokita Peninsula, Tohoku, Japan. *Bulletin of the Geological Survey of Japan*, vol. 71(4), p. 313–330, 7 figs, 1 table, 2 plates.

Abstract: In the northeastern Shimokita Peninsula, Aomori Prefecture, a Jurassic accretionary complex (AC) belonging to the North Kitakami Belt is distributed in three hilly areas of the Kuwabatayama, Katasakiyama and Omori. Although the AC in the Kuwabatayama area has been extensively studied and subdivided into the Late Jurassic Iwaya Unit and the Early Cretaceous Shitsukari Unit, those in the other areas are not fully understood except for a recent report of the detrital zircon U–Pb age from sandstone in the Omori area.

We extracted radiolarian fossils such as *Eucyrtidiellum cf. pyramis* indicating the Late Jurassic (probably Kimmeridgian) from mudstone near the U–Pb-dated sandstone in the Omori area. Because this radiolarian age is close to the zircon U–Pb age, and the horizons of the mudstone and sandstone are close without any recognizable distinctive tectonic discontinuity between them, the clastic rocks in the Omori area may be stratigraphically continuous or contemporaneous sequences deposited around the Kimmeridgian.

The ACs in the Katasakiyama and Omori areas are correlative to the Iwaya Unit in the Kuwabatayama area based on the lithology, geologic structure and clastic rock age. Therefore, the ACs in the northeastern Shimokita Peninsula can be divided into the Late Jurassic and the Early Cretaceous units; the former is the Iwaya Unit in the Kuwabatayama area and the unnamed ACs in the Omori and Katasakiyama areas, while the latter is the Shitsukari Unit in the Kuwabatayama area.

Keywords: radiolarian fossil, Late Jurassic, Kimmeridgian, accretionary complex, Shimokita Peninsula, North Kitakami Belt, Northeast Japan

1. Introduction

The North Kitakami Belt, located in the northern region of the Kitakami Massif of Tohoku (Northeast Japan), is mainly occupied by a Jurassic accretionary complex (AC). This AC is also sparsely exposed in the northward and westward areas of the massif because of the broad-scale coverage of Cenozoic volcanic and sedimentary rocks; only small amounts of the AC crop out in the Shimokita Peninsula, the south of Hirosaki City, the west of Lake Towada, the north of Mt. Hachimantai and the north of Mt. Moriyoshi (Fig. 1). A northeastern part in the Shimokita Peninsula, Aomori Prefecture, is marked by the Shimokita Hill, which is bounded by an escarpment along its east coast on the Pacific Ocean. The ACs in the northeastern Shimokita Peninsula are distributed in the Kuwabatayama, Katasakiyama and Omori areas from the

north to south of Higashidori Village (Fig. 2a). The ACs in these areas are exposed as a coastal terrace with a 200- to 300-m elevated steep escarpment although they are mainly overlaid by the Neogene sediments of the Sunakomata or Tomari formations (Imai, 1961) at the foot of the terrace. This paper mainly focuses on the ACs in the Katasakiyama and Omori areas.

The AC in the Kuwabatayama area crops out well along the seashore of Cape Shiriya and contains huge limestone blocks with fossils. Although it has been the subject of study by many researchers, the Katasakiyama and Omori areas have been little studied because of poor and fragmental exposures inland. Because the ACs in these two areas have not been mapped in detail despite the existence of 1:50,000 quadrangle series geologic maps, named “Chikagawa” (Imai, 1961) and “Shiriyazaki” (Tsushima and Takizawa, 1977), their geologic ages and

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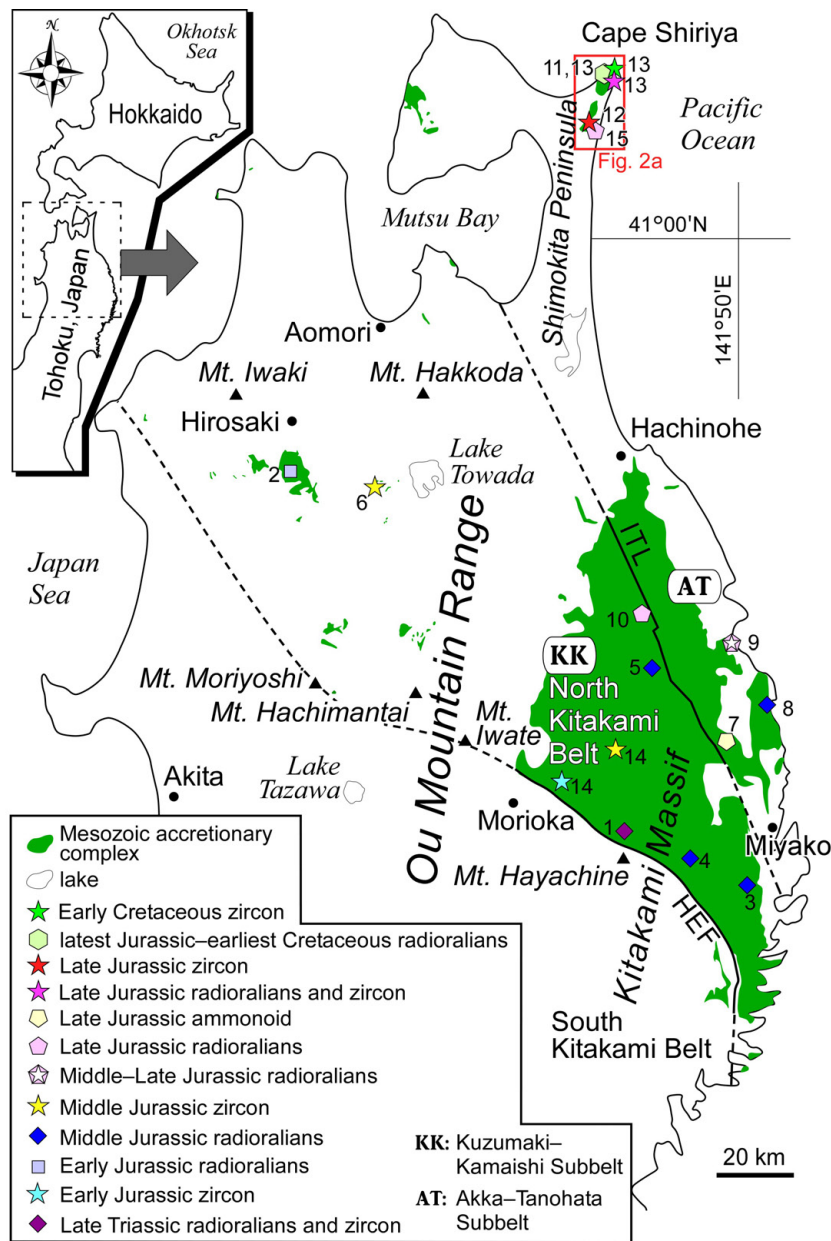


Fig. 1 Distribution map of Mesozoic accretionary complexes (green regions) in the North Kitakami Belt derived from the Seamless Digital Geological Map of Japan (1:200,000) V2 of the Geological Survey of Japan, AIST (2019). Symbols with numbers indicate the locations for which fossil and zircon ages from terrigenous clastic rocks were reported. 1: Kawamura *et al.* (2013); Uchino (2017), 2: Ueda *et al.* (2009), 3: Yoshihara *et al.* (2002); Suzuki and Ogane (2004), 4: Suzuki *et al.* (2007a), 5: Suzuki *et al.* (2007b); Ehiro *et al.* (2008), 6: Uchino (2018a), 7: Suzuki *et al.* (2007a), 8: Matsuoka and Oji (1990), 9: Minoura and Tsushima (1984), 10: Nakae and Kamada (2003), 11: Matsuoka (1987), 12: Uchino (2018b), 13: Ueda *et al.* (2018), 14: Uchino (2019), 15: this study.

HEF: Hayachine Eastern Marginal Fault, ITL: Iwaizumi Tectonic Line. The broken line shows an inferred fault.

tectonostratigraphy are not well understood. Recently, Uchino (2018b) obtained a zircon U-Pb age of ca. 155 Ma for the sandstone of the undated AC in the Omori area, indicating that it is slightly older than the latest Jurassic–earliest Cretaceous age for the clastic rocks (Matsuoka, 1987) in Cape Shiriya, which are regarded as the youngest in the North Kitakami Belt. To understand

the tectonostratigraphy of the Shimokita Peninsula, more data, such as radiolarian fossil ages, are required for this region. This paper reports radiolarians from mudstone near the location of sandstone dated using detrital zircon U-Pb geochronology, and discusses the local geologic correlation based on lithology and ages.

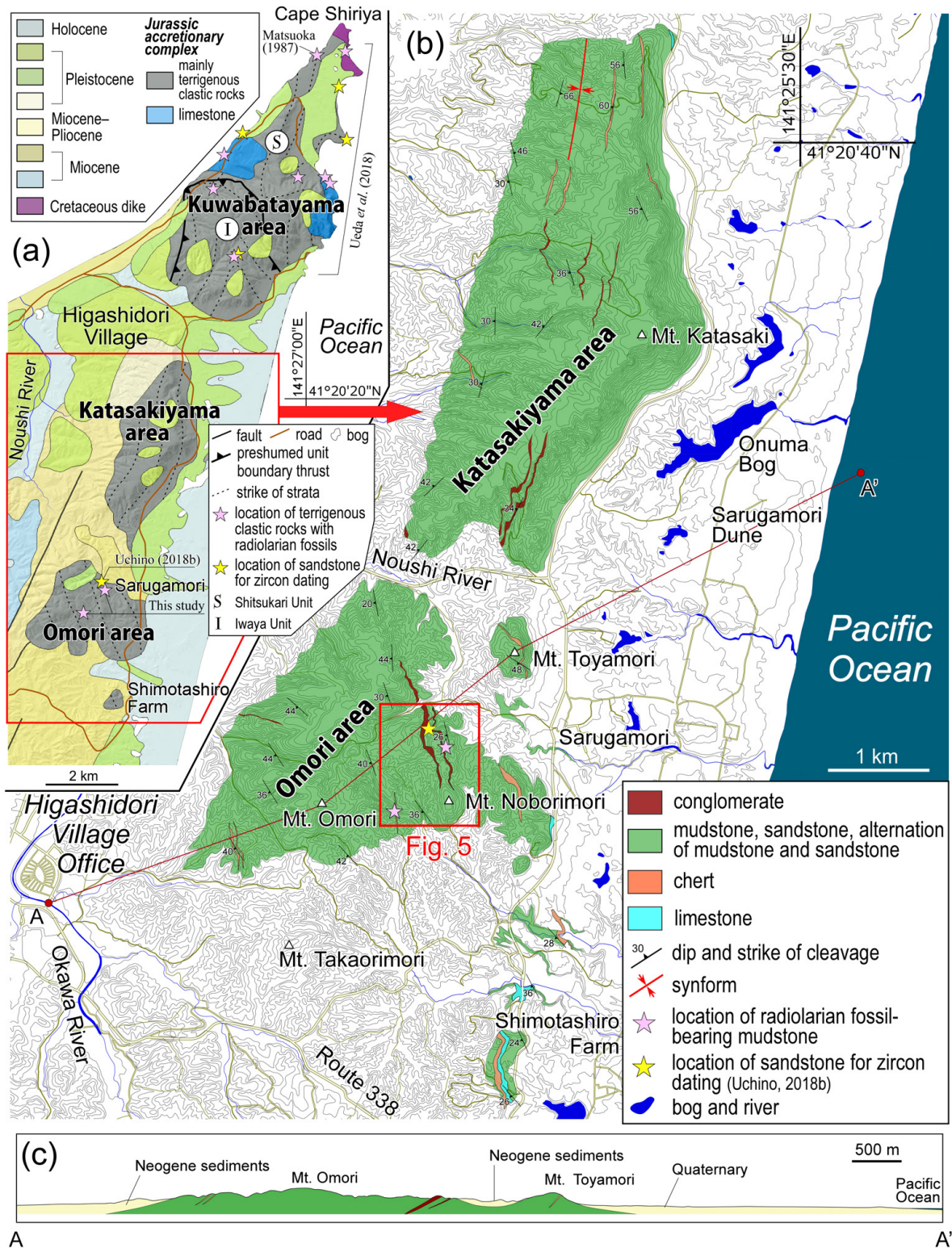


Fig. 2 (a) Index map of the northeastern Shimokita Peninsula, Aomori Prefecture. The geologic map was derived from the Seamless Digital Geological Map of Japan (1:200,000) V2 of the Geological Survey of Japan, AIST (2019). (b) Detailed geologic map of the accretionary complex in the Katsakiyama and Omori areas. The Quaternary deposits on the accretionary complexes in the two areas are not shown. (c) Geologic profile from the southwest (A) through Mt. Omori and Mt. Toyamori to the northeast (A').

Reference Area	Imai (1961)	Murata (1962)		Tsushima and Takizawa (1977)	Kawamura <i>et al.</i> (1994)	Kamada (2000)	Ueda <i>et al.</i> (2018)		
Kuwabata- yama	Shiriyu Formation	Shimokita Group	Tatemachijima Formation	Shiriyu Group	A formation	Shiriyu Complex	Shiriyu Complex	Shiriyu Complex	Shitsukari Unit
			Kuwabatakeyama Formation		B formation				Iwaya Unit
			Iwaya Formation		C formation				
			Katasakiyama Formation		C formation?				
Katasaki- yama									
Omori									

Fig. 3 Nomenclatural history of the stratigraphic units of the accretionary complexes in the northeastern Shimokita Peninsula, Aomori Prefecture. Tsushima and Takizawa (1977) indicates that the C formation in the Kuwabatayama area is probably correlative to the northern unit in the Katasakiyama area.

2. Geological outline

The North Kitakami Belt is divided into two sub-belts, namely the southwestward Kuzumaki–Kamaishi Subbelt and the northeastward Akka–Tanohata Subbelt, based on sandstone composition (dominated by plagioclase or K-feldspars), presence or absence of Paleozoic marine fossils and presence or absence of Late Jurassic coral-bearing limestone (Ehiro *et al.*, 2005; Kojima *et al.*, 2016) (Fig. 1). The AC in the Shimokita Peninsula belongs to the Akka–Tanohata Subbelt (e.g. Kamada, 2000; Ehiro *et al.*, 2008). The accretion age generally becomes younger from the southwest to the northeast (e.g. Suzuki *et al.*, 2007a; Ehiro *et al.*, 2008) although this trend does not always hold in the entire area (Nakae and Kamada, 2003). Of note, the youngest marine sediment in the North Kitakami Belt, from terrigenous clastic rocks in Cape Shiriyu, was dated to the Early Cretaceous (Matsuoka, 1987; Ueda *et al.*, 2018) (Fig. 1).

2.1 Nomenclature of stratigraphic units

The stratigraphic divisional scheme in the northeastern Shimokita Peninsula has changed over time (Imai, 1961; Murata, 1962; Tsushima and Takizawa, 1977; Kawamura *et al.*, 1994; Kamada, 2000; Kojima *et al.*, 2016; Ueda *et al.*, 2018) (Fig. 3). The pre-Neogene sedimentary rocks in this peninsula were previously thought to be “Paleozoic strata” and were named the Shiriyu Formation in the 1:50,000 quadrangle series geologic map “Chikagawa” (Imai, 1961). In consideration of Mesozoic hexacorals reported from the “Paleozoic strata” in the Kuwabatayama area (Onuki, 1959), the stratigraphic divisional scheme in the northeastern Shimokita Peninsula was revised to comprise the Katasakiyama Formation in the Katasakiyama area, and the Tatemachijima, Kuwabatakeyama and Iwaya formations in the Kuwabatayama area. These four formations were grouped as the Shimokita Group (Murata, 1962). However, the name Shimokita Group

was already used as a formal Miocene stratigraphic unit. Therefore, Tsushima and Takizawa (1977) renamed the four formations to the Shiriyu Group. Because the strata of the North Kitakami Belt were thought to be an AC (e.g. Minoura, 1985), the Shiriyu Group was renamed to the Shiriyu Complex. A tectonostratigraphic continuity of this complex to the ACs of the Oshima Belt in southern Hokkaido has been found (Kawamura *et al.*, 1994). Allocation of the Shiriyu Complex was expanded by Kawamura *et al.* (1994) to any ACs in the Shimokita Peninsula. Kamada (2000) confined this allocation to only the AC in the Kuwabatayama area. Recent geologic mapping, petrology, radiolarian ages and zircon U–Pb dating of the Shiriyu Complex by Ueda *et al.* (2018) have led to the proposal of using the Shitsukai and Iwaya units as sub-stratigraphic units of the Shiriyu Complex in the Kuwabatayama area.

We withhold specific geologic division names for the Katasakiyama and Omori areas because the first author is still engaged in a study of these areas, although the ACs in the Katasakiyama and Omori areas are likely to be correlative to the Iwaya Unit in the Kuwabatayama area, as discussed later.

2.2 Brief overview of geology and paleontology in study areas

2.2.1 AC in Kuwabatayama area

The Kuwabatayama area in the present paper covers nearly the entire study areas of Murata (1962), Oho and Iwamatsu (1986), Kamada (2000), Sano *et al.* (2009) and Ueda *et al.* (2018). The Shiriyu Complex (Kawamura *et al.*, 1994; Kamada, 2000) is characterized by imbricated stacks of coherent chert and clastic rock layers, and characteristic kilometer-long limestone blocks (Ueda *et al.*, 2018). The AC in Cape Shiriyu is intruded by diorite 122.1±1.4 (2σ) Ma in age (Ueda *et al.*, 2018). Oho and Iwamatsu (1986) considered this AC as an olistostrome of submarine landslide deposits based on a

slump with huge limestone olistoliths and the presence of a microbrecciated matrix. Recently, Ueda *et al.* (2018) divided the AC in this area into two tectonostratigraphic units (Iwaya and Shitsukari units) based on lithology and radiolarian and zircon U-Pb ages. They found that the Iwaya Unit is composed of mudstone, sandstone, chert, siliceous mudstone, and minor amounts of limestone and conglomerate. The Shitsukari Unit is composed of conglomerate, sandstone, mudstone, siliceous mudstone, chert, limestone and a minor amount of metabasalt. The Iwaya Unit, which has an approximately N-S to NW-SE trending synform, is supposed to lie structurally above the Shitsukari Unit, which has a NE-SW trending synform and antiform pair (Fig. 2a). Ueda *et al.* (2018) also found that the Shitsukari Unit formed as an imbricated accretionary wedge which was composed of debris from an inner trench slope.

The limestone was dated to the Late Jurassic using Hexacorallia and Stromatoporoidea fossils (Onuki, 1959; Murata, 1962) and another limestone was dated to the Late Triassic using megalodontoid bivalves (Sano *et al.*, 2009). The microfossil ages of chert range from the Middle Triassic (Anisian) to the Late Jurassic (Kimmeridgian) (Toyohara *et al.*, 1980; Oho and Iwamatsu, 1986; Matsuoka, 1987; Ueda *et al.*, 2018). Siliceous mudstone yields Middle Jurassic to latest Jurassic radiolarians (Ueda *et al.*, 2018), and tuffaceous mudstone yields latest Jurassic to earliest Cretaceous radiolarians (Matsuoka, 1987). The zircon U-Pb ages of clastic rocks, including intercalated tuff layers, are correlated to the Early Cretaceous (Berriasian to Barremian) (Ueda *et al.*, 2018).

2. 2. 2 AC in Katasakiyama area

The AC in the Katasakiyama area shows a 7 km × 2.5 km lenticular distribution pattern around Mt. Katasaki. The northern part of the Katasakiyama area was partly covered by Murata (1962), who mapped a synform in the southern margin of his map. A detailed geologic map of the AC is shown in Fig. 2b, along with that in the Omori area (see next section). The AC exceeds over 1,200 m in thickness. It consists mainly of mudstone with minor amounts of chert, alternating beds of sandstone and mudstone, sandstone, conglomerate and limestone. The sheets or lens of oceanic rocks such as limestone and chert are less than several meters in thickness. The AC trends in the NNW-SSE to NNE-SSW directions and dips at a low-middle angle to the west (Fig. 2c). A thin limestone sheet several meters in thickness occurs at the northeastern edge of this area.

2. 2. 3 AC in Omori area

The AC in the Omori area shows a 2.7 km × 2.9 km elliptical distribution pattern around Mt. Omori. The thickness of the AC exceeds over 2,000 m. It consists mainly of mudstone with minor amounts of limestone, chert, alternating beds of sandstone and mudstone, sandstone and conglomerate. Similar to the limestone

sheet at the northeastern edge of the Katasakiyama area, a limestone sheet <25 m in thickness is aligned, from place to place, along the eastern margin of the Omori area with a NNW-SSE trend, suggesting a possible southern extension of the limestone sheet from the Katasakiyama area. The AC trends in the NW-SE to NNW-SSE directions and dips at a low-middle angle to the west (Fig. 2b, c).

3. Lithology

The lithologies of the ACs in the Katasakiyama and Omori areas are very similar and are thus described collectively in this section. That in the Kuwabatayama area is not explained here because it has been described in previous reports.

Limestone, which accounts for ca. 0.4 % of the mapped distribution area of the AC, is gray to pale gray and mostly recrystallized. Macroscopic and microscopic calcite veins less than 1 cm in width frequently crosscut the limestone, and some of them are dark pink. Matrix-supported and poorly sorted calcirudite infrequently occurs with preserving as an original sedimentary structure. The limestone clasts in the calcirudite do not exceed 4 cm in length. Chert, which accounts for ca. 1 % of the mapped distribution area, is gray. It has a bedded structure, where several-centimeter-thick chert layers alternate with ca. 1-mm-thick claystone layers. The chert is mostly recrystallized. Dark gray mudstone is the most dominant rock type in the AC. The mudstone is more or less slaty, and frequently contains radiolarian fossil pseudomorphs. Quartz veins less than 1 cm in width occasionally intrude into the mudstone.

Sandstone, which accounts for a ca. 1 % of the mapped distribution area, occurs as feldspathic arenite (Fig. 4a) and lithic arenite to wacke. Quartz grains are more common than feldspar ones, which have nearly equal amounts of plagioclase and K-feldspar. This sandstone varies from fine to very coarse grains and is in general poorly sorted. It also contains many lithic fragments such as chert and siliceous mudstone, and frequently contains contemporaneous mud chips. It is occasionally intruded by quartz veins less than 5 cm in width. The sandstone sometimes alternates with mudstone in the form of layers with a thickness on the order of several millimeters. Rarely, pale purple tuffaceous sandstone appears.

Poorly sorted, granule to pebbly conglomerates (Fig. 4b, c) account for 1 % of the mapped distribution area of the AC in these two areas. They are mostly characterized by a clast-supported fabric with coarse quartz grains, angular to sub-rounded variable clasts of chert (gray, dark gray, white and rarely red), siliceous mudstone, mudstone and sandstone.

These sedimentary rocks have undergone extensive post-depositional alteration and deformation, with layer-parallel slaty cleavage in the fissile mudstone and considerably flattened clasts in the conglomerate (Fig. 4c). The mudstone sometimes exhibits whitish parts caused by

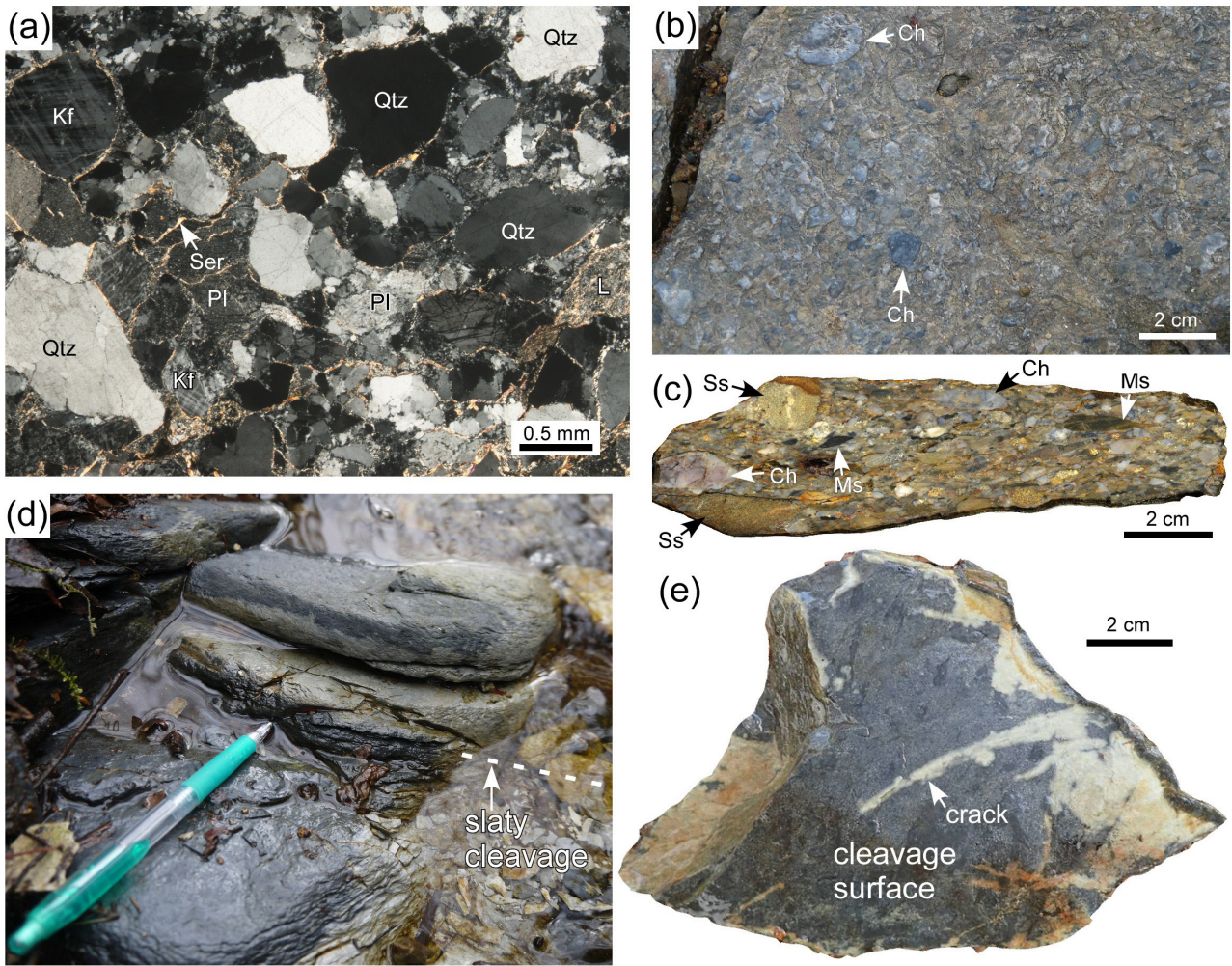


Fig. 4 (a) Photomicrograph of feldspathic arenite. Cross-polarized light. Mizunashi Stream, east ramp of Mt. Katasaki. Kf: K-feldspar, Pl: plagioclase, Qtz: quartz, Ser: sericite, L: lithic fragment (mudstone). (b) Conglomerate outcrop. Stream to the west of Sarugamori hamlet. Ch: chert. (c) Conglomerate specimen. Ch: chert, Ms: mudstone, Ss: sandstone. (d) Altered mudstone outcrop. The pen is 15 cm long. (e) Altered mudstone specimen. Stream to the west of Mt. Toyamori.

fluid flow along the slaty cleavages and cracks, and looks as if it was tuffaceous or calcareous mudstone (Fig. 4d, e). Fine clay minerals such as sericite, which are mostly in a unidirectional polarization extinction position, are found between the clastic grains (Fig. 4a). Very fine ferric oxyhydroxide minerals frequently develop along the cleavages and cracks or between the clastic grains. Calcite spots (<0.2 mm in diameter) that are not aligned with the cleavages are occasionally present in the clastic rocks, and quartz spots (<0.2 mm) containing very fine dusty inclusions appear in the matrix of limestone and calcite veins.

4. Radiolarian fossil

4.1 Sample location and extraction method

Samples containing radiolarian fossils were recovered from mudstone at two outcrops in the stream: Loc. 1

(41°17'05"N, 141°22'35"E) ca. 550 m west and Loc. 2 (41°17'25"N, 141°22'56"E) ca. 550 m north of Mt. Noborimori (Fig. 2b). The mudstone is dark gray and does not exhibit the apparent deformation or alteration. A route map and a columnar section around these outcrops are shown in Figs. 5 and 6, respectively. Of note, Loc. 2 is close to an outcrop of the Late Jurassic sandstone dated using detrital zircon U-Pb geochronology (Uchino, 2018b).

Rock chips were soaked in 5% HF for ca. 18 hours and then residual fractions were collected with sieves of #65 and #250 meshes. After these steps were repeated three times, radiolarian samples were picked with a brush under a stereomicroscope and subsequently examined with a scanning electron microscope.

Regarding taxonomy, the species was determined by strictly referring to the holotype or other type series images in original publications, and the genus was determined by

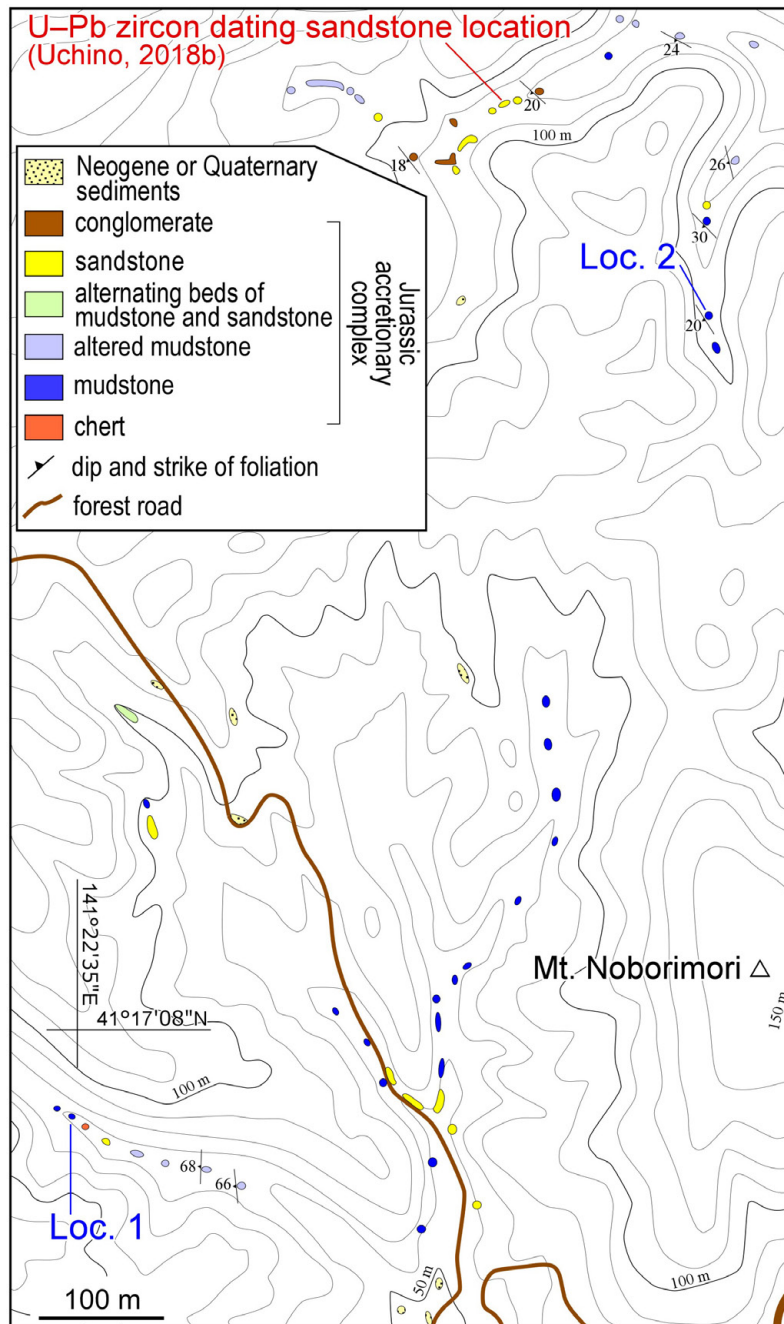


Fig. 5 Route map around the location of radiolarian fossil-bearing mudstone. Contour lines were derived from digital elevation model data (10-m mesh) from the Geospatial Information Authority of Japan.

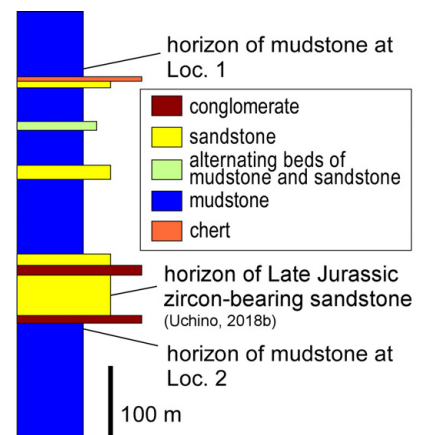


Fig. 6 Columnar section around the location of radiolarian fossil-bearing mudstone.

referring to updated concepts if possible (e.g. O’Dogherty *et al.*, 2017). The biostratigraphic correlation to the geologic time scale at the stage level was largely based on the Jurassic standard radiolarian zones proposed by Matsuoka (1995) and some revisions in subsequent papers (Hatakeda *et al.*, 2007; Ishida *et al.*, 2009). Because of a contradiction in the denoted geologic stages of the same fauna among Japan, Europe and North America, the assigned geologic time scale was converted from

biozones in other regions to those of Japan by referencing Baumgartner *et al.* (1995), Yang and Matsuoka (1997) and Goričan *et al.* (2018).

4. 2 Identification and age

4. 2. 1 Radiolarians at Loc. 1

Many of the radiolarian fossils were deformed and recrystallized. Although their surface structure is hard to be recognized, 33 species or species levels are identified.

Table 1 List of radiolarian fossils. Red and blue numbers indicate individuals from the mudstone at Loc. 1 and Loc. 2, respectively.

Fossil	Loc. 1	Loc. 2	Individual number in plate
<i>Archaeodictyomitra</i> cf. <i>inornata</i> Hull	*	*	21–23, 4
<i>Archaeodictyomitra</i> cf. <i>unica</i> Wu	*		15
<i>Archaeodictyomitra</i> aff. <i>prisca</i> Kozur and Mostler in Grill and Kozur (1986)	*		7
<i>Archaeodictyomitra</i> aff. <i>rigida</i> Pessagno	*	*	16, 17, 1, 2
<i>Archaeodictyomitra</i> <i>spelae</i> Chiari, Cortese and Marcucci	*		1–3
<i>Archaeodictyomitra</i> cf. <i>spelae</i> Chiari, Cortese and Marcucci	*		4–6
<i>Archaeodictyomitra</i> aff. <i>suzukii</i> Aita	*		18–20
<i>Archaeodictyomitra</i> cf. <i>tyaughtonensis</i> Cordey		*	3
<i>Archaeodictyomitra</i> aff. <i>vulgaris</i> Pessagno	*		11–14
<i>Archaeodictyomitra</i> sp. A	*		8
<i>Archaeodictyomitra</i> sp. B	*		9, 10
<i>Archaeodictyomitra</i> sp. C	*		24
<i>Archaeodictyomitra</i> sp. D	*		25
<i>Archaeospongoprimum</i> cf. <i>mizutanii</i> Ožvoldová in Ožvoldová <i>et al.</i> (2000)	*		109
<i>Archaeospongoprimum</i> sp.	*		108
<i>Bistarkum</i> aff. <i>mangartense</i> Goričan, Šmuc and Baumgartner	*		112, 113
<i>Cinguloturris</i> cf. <i>carpatica</i> Dumitrică in Dumitrică and Mello (1982)	*	*	38–43, 17
<i>Cinguloturris</i> cf. <i>floridicingula</i> (Li)		*	15
<i>Cinguloturris</i> cf. <i>getsensis</i> O'Dogherty, Goričan and Dumitrică in O'Dogherty <i>et al.</i> (2006)		*	16
<i>Cinguloturris</i> cf. <i>latiannulatum</i> (Grill and Kozur)	*		44
<i>Cinguloturris</i> sp.	*		45
<i>Crococapsa</i> aff. <i>truncata</i> (Wu)	*		63
<i>Eucyrtidellum</i> cf. <i>nodosum</i> Wakita	*		62
<i>Eucyrtidellum</i> cf. <i>pyramis</i> (Aita) in Aita and Okada (1986)	*		59–61
<i>Eucyrtidellum</i> sp.		*	29
<i>Favosyringium</i> cf. <i>affine</i> (Rüst) sensu Steiger (1992)	*		111
<i>Kilinora</i> sp.	*		69
<i>Loopus</i> cf. <i>venusta</i> (Chiari, Cortese and Marucci)	*		37
<i>Minutosolla</i> sp.	*		88, 89
<i>Parahsuum mudongensis</i> (Li) sensu lato	*		26–28
<i>Podobursa</i> sp.	*		110
<i>Praewilliriedellum</i> sp.	*		78
<i>Hiscocapsa robusta</i> (Matsuoka)	*		91, 92
<i>Hemicryptocapsa</i> cf. <i>yaoi</i> (Kozur)	*		93
<i>Hemicryptocapsa carpathica</i> (Dumitrică)	*		79
<i>Praezhamoidellum</i> sp.	*		94–98
<i>Praezhamoidellum</i> ? sp.	*		90
<i>Quarkus</i> sp.		*	30
<i>Spongocapsula palmerae</i> Pessagno	*		29–31
<i>Spongocapsula</i> sp.	*		32, 33
<i>Spongocapsula</i> ? sp.	*		34–36
<i>Striatojaponocapsa</i> cf. <i>conexa</i> (Matsuoka)	*		70
<i>Striatojaponocapsa synconexa</i> O'Dogherty, Goričan and Dumitrică in O'Dogherty <i>et al.</i> (2006)	*		71–76
<i>Striatojaponocapsa</i> sp.	*		77
<i>Svinitzium</i> sp.	*		46, 47
<i>Tetracapsa</i> sp.		*	31
<i>Transhsuum</i> sp.	*		49–55
<i>Wrangellium</i> sp.	*		48
<i>Williriedellum</i> sp.	*	*	80–87, 20–23
<i>Complexapora</i> aff. <i>kiesslingi</i> Hull	*		105, 106
<i>Zhamoidellum</i> cf. <i>mikamense</i> Aita		*	19
<i>Zhamoidellum ovum</i> Dumitrică	*		101–104
<i>Zhamoidellum</i> cf. <i>ventricosum</i> Dumitrică	*		100
<i>Zhamoidellum</i> sp.	*		99
Minocapsidae gen. et sp. indet.	*		107
Pyloniodea gen. et sp. indet.	*		58
Multisegmented nassellarians		*	5–14
Spherical radiolarians		*	18
Williriedelloidea gen. et sp. indet.		*	24–26
Syringocapsidae? gen. et sp. indet.		*	27, 28
Few segmented nassellarians		*	32–35
Four-armed flat Pyloniodea		*	36
Nassellaria gen. et sp. indet.	*		56, 57, 64–68, 114

A total of approximately 15 species coexisted in the early to middle Late Jurassic (roughly Oxfordian or Kimmeridgian): *Archaeodictyomitra* cf. *inornata* Hull, *Archaeodictyomitra* cf. *unica* Wu, *Archaeodictyomitra spelae* Chiari *et al.* (a single report from the Oxfordian), *Archaeospongoprimum* cf. *mizutani* Ožvoldová [known from Oxfordian and Tithonian], *Cinguloturris* cf. *carpatica* Dumitrică, *Cinguloturris* cf. *latiannulatum* (Grill and Kozur), *Eucyrtidiellum* cf. *nodosum* Wakita, *Eucyrtidiellum* cf. *pyramis* (Aita) (Kimmeridgian and later), *Loopus* cf. *venusta* (Chiari *et al.*), *Parahsuum mudongensis* (Li) sensu lato, *Hiscocapsa robusta* (Matsuoka), *Hemicryptocapsa* cf. *yaoi* (Kozur), *Spongocapsula palmerae* Pessagno, *Striatojaponocapsa* cf. *conexa* (Matsuoka), *Hemicryptocapsa carpathica* (Dumitrică), *Zhamoidellum ovum* Dumitrică, and *Zhamoidellum* cf. *ventricosum* Dumitrică. By contrast, 12 species, such as *Striatojaponocapsa synconexa* O'Dogherty *et al.*, are referable to the late Middle Jurassic (roughly Callovian). Eight species are common in the late Late Jurassic (Tithonian). These eight species, however, range down to the middle Late Jurassic (Kimmeridgian), and are thus not necessarily from the Tithonian. Therefore, the fauna from Loc. 1 is roughly dated to the early to middle Late Jurassic. Based on species whose ranges are well documented (Aita, 1987; Matsuoka, 1995; Nishizono, 1996; Hori, 1999; Hatakeda *et al.*, 2007), the co-occurrence of *Striatojaponocapsa* cf. *conexa* (Matsuoka) and *Striatojaponocapsa synconexa* O'Dogherty *et al.* is correlative to JR5 (*S. conexa* Zone) to JR6 (*Kilinora spiralis* Zone) of Matsuoka (1995). These two zones are generally correlated to the Bathonian and Callovian. However, Ishida *et al.* (2009) reported the earliest Kimmeridgian ammonoid, *Ataxioceras* (*Ataxioceras*) *kurisakense*, from bioturbated sandy mudstone 1.5 m above the horizon of the JR6 fauna with *S. conexa* and *S. synconexa* (= *Stichocapsa naradaniensis* Matsuoka shown in Fig. 7.9 of Ishida *et al.*, 2009) from the Kurisaka Formation of the Kurosegawa Belt in Shikoku, Southwest Japan. This suggests that an extension of geologic time ranges up to the earliest Kimmeridgian from the Callovian, or an age revision from the Callovian to a significantly younger age is needed. The contradiction of ranges of the radiolarian zone shown by Matsuoka (1995) was pointed out for stratigraphically important *Kilinora spiralis*, *Loopus primitivus* (Matsuoka and Yao), *Striatojaponocapsa plicarum* (Yao) and *S. conexa* by Hori *et al.* (2002). Because the occurrence of *S. conexa* and *S. synconexa* can also be explained by reworking to the lowest Kimmeridgian strata, we focus on *Eucyrtidiellum* cf. *pyramis*, whose ranges are the youngest among the fauna from Loc. 1. The genus *Eucyrtidiellum* is composed of 19 species. The form of *E. pyramis* is quite distinctive, with widely spaced longitudinal plicae on both the thorax and abdomen and a smooth cone shape, from any other *Eucyrtidiellum* species. The first occurrence of *E. pyramis* was correlative to the radiolarian *Ditrabs sansalvadorensis*

Zone of Aita (1987) and UAZ 12 of Baumgartner *et al.* (1995), indicating the Tithonian according to Aita (1987) and Goričan *et al.* (2018). The oldest occurrence of *E. pyramis* is from the Nusplingen Lithographic Limestone in Germany, which yielded late Kimmeridgian ammonoids indicative of the *Lithacoceras ulmense* Subzone of Zügel *et al.* (1998). Although it is unclear why *E. cf. pyramis*, *S. cf. conexa* and *S. synconexa* co-exist in the same sample, the sample can be roughly dated to the Kimmeridgian (157.3±1.0 Ma to 152.1±0.9 Ma).

4. 2. 2 Radiolarians at Loc. 2

Radiolarian fossils at Loc. 2 were more deformed and recrystallized than those at Loc. 1. Their surface structure was difficult to identify. A total of 11 taxa at the genus or species level were identified in this sample. Identified taxa at the species level are *Archaeodictyomitra* cf. *inornata*, *Archaeodictyomitra* aff. *rigida*, *Archaeodictyomitra* cf. *tyaughtonensis* Cordey, *Cinguloturris* cf. *carpatica*, *Cinguloturris* cf. *floridicingula* (Li), *Cinguloturris* cf. *getsensis* O'Dogherty *et al.* and *Zhamoidellum* cf. *mikamense* Aita. Although three of these seven species (*A. cf. inornata*, *A. aff. rigida* and *C. cf. carpatica*) were also found at Loc. 1, they are generally dated to the late Middle Jurassic (probably Callovian) to middle Late Jurassic (probably Kimmeridgian). There were no taxa that could be certainly assigned to a geologic age later than this suggested age. It was impossible to determine ages at higher resolution. Most of these species (e.g. *A. cf. tyaughtonensis*) are rarely reported after their first description or they cannot be used for age determination. For example, *Cinguloturris carpatica* belongs to the latter case. It is subdivided into several *Cinguloturris* species, such as *C. floridicingula* and *C. getsensis*. *Zhamoidellum mikamense* was combined with *Zhamoidellum ovum* Dumitrică as a junior synonym (Baumgartner *et al.*, 1995). Thus, taxonomic confirmation of these two species is needed to determine the precise age. In consideration of the insufficient number of reports for the rare taxa and incomplete updating of the taxonomic concept of age-index species, we retain the radiolarian age for the Loc. 2 sample as the late Middle Jurassic (probably Callovian) to middle Late Jurassic (probably Kimmeridgian).

5. Discussion

5. 1 Depositional age of clastic rocks

The youngest cluster U-Pb ages of detrital zircon from sandstone in the Omori area are 154.7±1.5 Ma (1σ) and 155.6±3.2 Ma (2σ), obtained using different calculation methods (Uchino, 2018b). These ages correspond to the latest Oxfordian–Kimmeridgian under the geologic time scale of Cohen *et al.* (2013). A detrital zircon age suggests only a possible lower limit of a depositional age of sandstone, although the zircon age can approximate the depositional age if volcanism in its hinterland was very active. However, volcanism in eastern Asia, including

the paleo-Japanese islands, was not very active during the Late Jurassic–earliest Cretaceous (e.g. Sagong *et al.*, 2005; Kiminami and Imaoka, 2013; Lee *et al.*, 2018). The low age spectrum peak of the youngest cluster of the detrital zircon from the sandstone in the Omori area shown by Uchino (2018b) probably reflects this weak volcanism in the hinterland. Therefore, the sandstone was deposited in the latest Oxfordian–Kimmeridgian or later.

The radiolarian fossil age from mudstone can directly indicate the depositional age of the mudstone if the fossils were not reworked. In consideration of the reliability of the assigned radiolarian ages for the mudstone at Locs. 1 and 2, the mudstone ages of the Kimmeridgian at Loc. 1 and the late Middle Jurassic (probably Callovian) to middle Late Jurassic (probably Kimmeridgian) at Loc. 2 are close to the U–Pb age of the sandstone in the Omori area. In addition, the horizons of the mudstone and sandstone are close to each other without any recognizable distinctive tectonic discontinuity between them (Figs. 5, 6), indicating that these clastic rocks may be stratigraphically continuous or contemporaneous sequences. Therefore, the clastic rocks in the Omori area were deposited in a trench around the Kimmeridgian (middle Late Jurassic).

5.2 Tectonostratigraphic correlation

The ACs identified in the Shimokita Peninsula are, from north to south, the Shiriya Complex (the Iwaya and Shitsukari units) in the Kuwabatayama area, an unnamed AC in the Katasakiyama area and an unnamed AC in the Omori area. Tsushima and Takizawa (1977) pointed out that the Iwaya Unit (“C formation” in the original paper) of the Shiriya Complex in the Kuwabatayama area is similar to the northernmost part of the AC in the Katasakiyama area (Fig. 3). In consideration of the trends in geologic structures, such as the NE–SW strikes of the Shitsukai Unit in the northern Kuwabatayama area, the N–S to NW–SE strikes of the Iwaya Unit in the southern Kuwabatayama area, the NNE–SSW to NNW–SSE strikes of the AC in the Katasakiyama area and the NW–SE to NNW–SSE strikes in the Omori area, the Iwaya Unit in the Kuwabatayama area and the ACs in the Katasakiyama and Omori areas are presumably tectonically continuous (Fig. 2a).

According to Ueda *et al.* (2018) and other previous papers, the Iwaya Unit differs from the Shitsukari Unit in that the latter exclusively contains basalt, huge limestone blocks and lithic sandstone, and is characterized by slump facies indicating debrite (Figs. 2a, 3). In addition, the age of the clastic rocks in the Iwaya Unit is the Late Jurassic, younger than the Early Cretaceous age of the clastic rocks in the Shitsukai Unit.

The lithologies of the ACs in the Katasakiyama and Omori areas are more similar to that in the Iwaya Unit than that in the Shitsukai Unit, in particular in terms of the presence of characteristic quartzo-feldspathic sandstone (Fig. 4a) and small amounts of limestone and conglomerate. The assigned age of the Kimmeridgian (the Late Jurassic)

to the clastic rocks of the AC in the Omori area overlaps the age of the Iwaya Unit in the southern Kuwabatayama area, which supports the probable tectonostratigraphic continuity of the ACs in the Katasakiyama and Omori areas to the Iwaya Unit in the Kuwabatayama area.

Although more detailed studies are needed to confirm this assumption, the ACs in the northeastern Shimokita Peninsula are likely to comprise a Late Jurassic tectonostratigraphic unit and an overlain Early Cretaceous unit. The former is the Iwaya Unit in the southern Kuwabatayama area and the unnamed ACs in the Katasakiyama and Omori areas, and the latter is the Shitsukari Unit, which has only been found in the northern Kuwabatayama area so far.

5.3 Younging polarity

It has been reported that a younging polarity of accretion ages for trench-fill terrigenous clastic rocks in the North Kitakami Belt is detectable from the southwest to the northeast, in a direction perpendicular to the general NW–SE to NNW–SSE distribution trends of the ACs in the belt (e.g. Suzuki *et al.*, 2007a; Ehiro *et al.*, 2008; Kojima *et al.*, 2016) (Figs. 1, 7). The geologic columns of the ocean plate stratigraphy reconstructed in each area are compiled in Fig. 7, although the terms “complex” and “unit” are mixed in the figure. The depositional ages of terrigenous rocks within the ACs in the “B” zone of Otoh and Sasaki (2003) range from the Late Triassic in the Kadoma Complex (Uchino, 2017), through the Early Jurassic in the Nishimatayama Unit and the Nakatsugawa Complex (Ueda *et al.*, 2009; Uchino, 2019), to the Middle Jurassic in the Shibamori Complex, possibly the Nakatsugawa Complex, and the Tsugaruishi Unit (Yoshihara *et al.*, 2002; Suzuki and Ogane, 2004; Suzuki *et al.*, 2007a; Uchino, 2018a; Uchino, 2019). No age data for the trench-fill terrigenous rocks has been reported in the “C” zone excluding the Late Jurassic coral fossils from possible shallow marine deposits. The “D” and “E” zones range from the Middle Jurassic to the Late Jurassic in the Kado–Akka areas (Nakae and Kamada, 2003; Suzuki *et al.*, 2007b; Ehiro *et al.*, 2008).

The “B”–“E” zones belong to the Kuzumaki–Kamaishi Subbelt, and the “A” zone belongs to the Nedamo Belt. The Akka–Tanohata Subbelt fully corresponds to the “F” and “G” zones, whose depositional ages range from the Middle–Late Jurassic in the Takayashiki and Magisawa units (Minoura and Tsushima, 1984; Matsuoka and Oji, 1990; Suzuki *et al.*, 2007a), through the Late Jurassic in the Iwaya Unit and its equivalent (Matsuoka, 1987; Ueda *et al.*, 2018; this study), to the late Late Jurassic to the early Early Cretaceous in the Shitsukari Unit (Matsuoka, 1987; Ueda *et al.*, 2018). The present study confirmed that the Late Jurassic ACs in the northeastern Shimokita Peninsula are distributed in the Kuwabatayama area as the Iwaya Unit and in the Omori area as an unnamed AC, and probably in the Katasakiyama area as an unnamed AC.

A well-ordered younging polarity was found in the

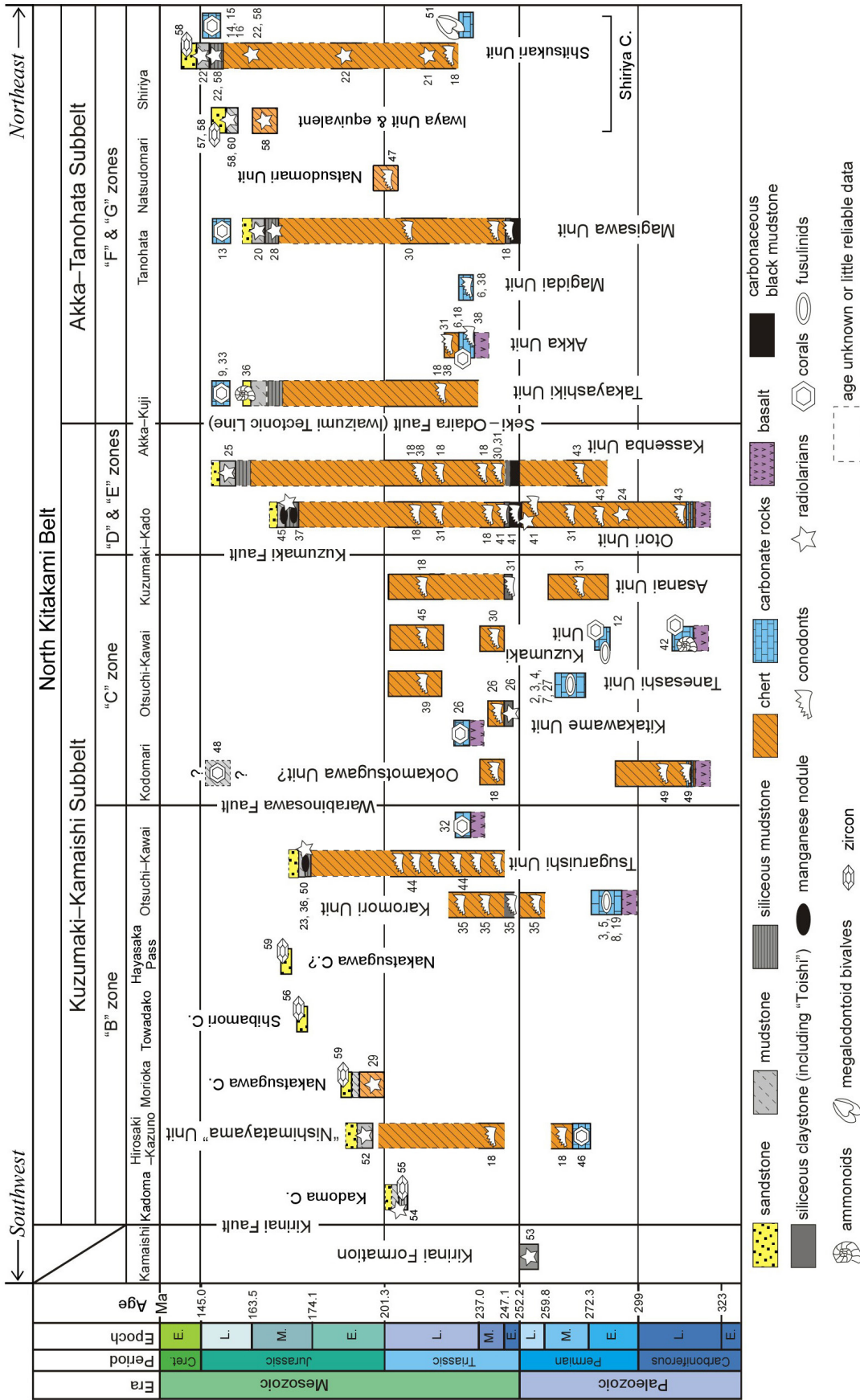


Fig. 7 Schematic compilation diagram of age and lithostratigraphic column of the North Kitakami Belt. This diagram was revised based on Fig. 42 of Suzuki *et al.* (2007a), Fig. 3 of Ehro *et al.* (2008) and Fig. 2b. 41 of Kojima *et al.* (2016). Numbers beside fossil and zircon symbols correspond to literature references. See Suzuki *et al.* (2007a) for nos. 1–35 and Ehro *et al.* (2008) for nos. 36–45. 46: Fujimoto and Kobayashi (1961), 47: Murata and Nagai (1972), 48: Kato (1972), 49: Murata *et al.* (1974), 50: Yoshihara *et al.* (2002); 51: Sano *et al.* (2009), 52: Ueda *et al.* (2009), 53: Nakae and Kurihara (2011), 54: Kawamura *et al.* (2013), 55: Uchino (2017), 56: Uchino (2018a), 57: Uchino (2018b), 58: Ueda *et al.* (2018), 59: Uchino (2019), 60: this study. See Otoh and Sasaki (2003) for "A"–"G" zones. L.: Late, M.: Middle, E.: Early, C.: Complex. Tectonostratigraphic names are taken from the original papers or Fig. 2b. 41 of Kojima *et al.* (2016), and the appropriateness of the terms "complex" and "unit" are not discussed in the present paper.

Kuzumaki–Kamaishi Subbelt. By contrast, the age polarity in the Akka–Tanohata Subbelt appears somewhat disordered, as shown in Fig. 7. Possible reasons for this disorder include wrong reconstruction of complexes/units, no age data for the terrigenous clastic rocks of the Magidai Unit in the Tanohata area, and repeated distribution of the same tectonostratigraphic units by folding or out-of-sequence thrust. Despite this, when viewing the younging polarity from a broad perspective, the trend from the Late Triassic of the southwesternmost AC (Kadoma Complex) of the Kuzumaki–Kamaishi Subbelt to the early Early Cretaceous of the northeasternmost AC (Shitsukari Unit of the Shiriya Complex) of the Akka–Tanohata Subbelt is well constrained in the North Kitakami Belt based on the data in this study.

6. Conclusion

Radiolarian fossils from around the Kimmeridgian (middle Late Jurassic) were extracted from mudstone in the AC in the Omori area, northeastern Shimokita Peninsula. The ACs in the Kasakiyama and Omori areas are correlative to the Iwaya Unit, the southern unit in the Kuwabatayama area, based on the lithology, geologic structure and clastic rock ages.

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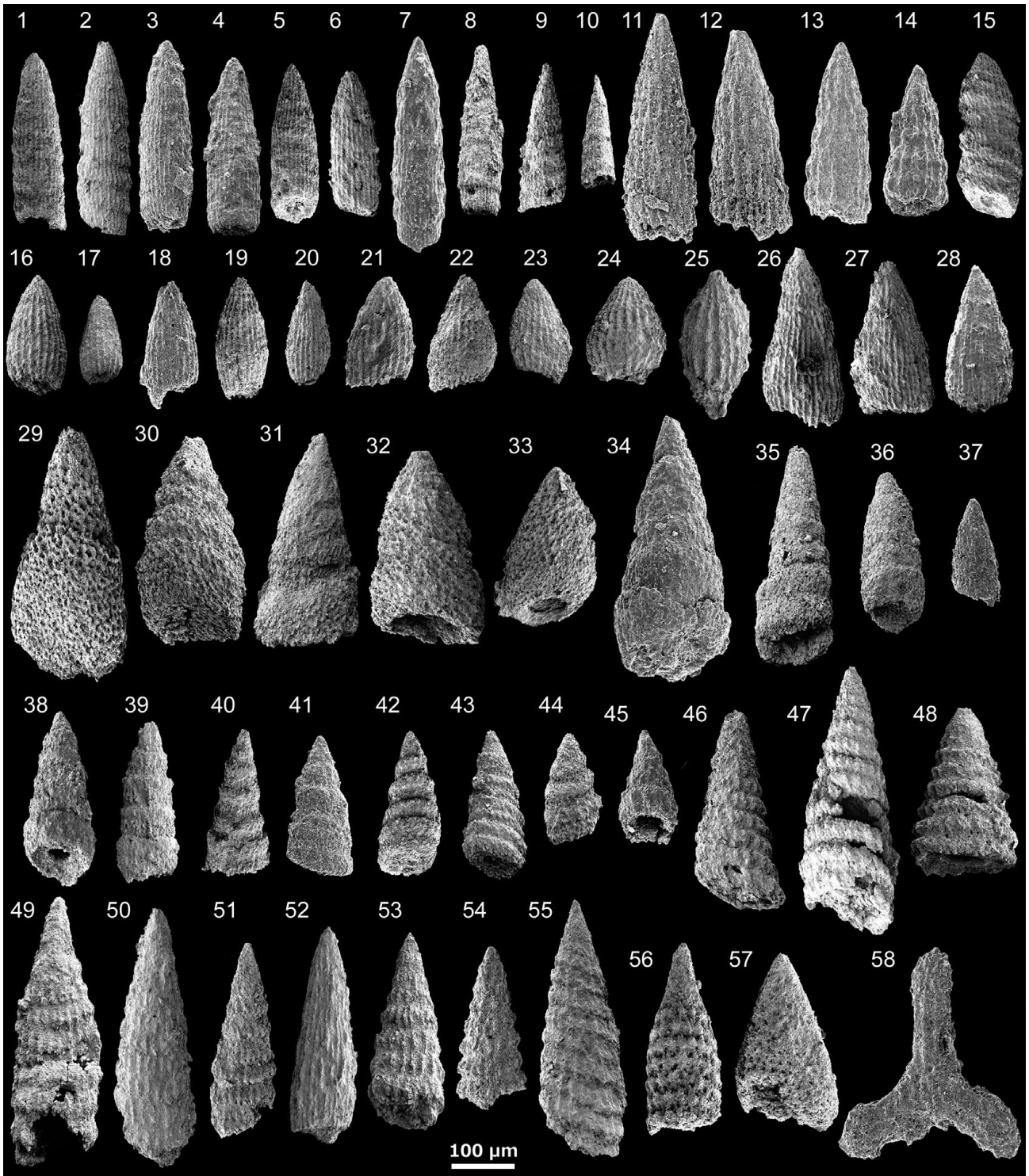


Plate 1 Scanning electron microscopy images of the Kimmeridgian (middle Late Jurassic) radiolarians extracted from mudstone at Loc. 1. 1–3: *Archaeodictyomitra spelae* Chiari, Cortese and Marcucci, 4–6: *Archaeodictyomitra* cf. *spelae* Chiari, Cortese and Marcucci, 7: *Archaeodictyomitra* aff. *prisca* Kozur and Mostler, 8: *Archaeodictyomitra* sp. A, 9–10: *Archaeodictyomitra* sp. B, 11–14: *Archaeodictyomitra* aff. *vulgaris* Pessagno, 15: *Archaeodictyomitra* cf. *unica* Wu, 16–17: *Archaeodictyomitra* aff. *rigida* Pessagno, 18–20: *Archaeodictyomitra* aff. *suzukii* Aita, 21–23: *Archaeodictyomitra* cf. *inornata* Hull, 24: *Archaeodictyomitra* sp. C, 25: *Archaeodictyomitra* sp. D, 26–28: *Parahsuum mudongensis* (Li) sensu lato, 29–31: *Spongocapsula palmerae* Pessagno, 32–33: *Spongocapsula* sp., 34–36: *Spongocapsula*? sp., 37: *Loopus* cf. *venusta* (Chiari, Cortese and Marucchi), 38–43: *Cinguloturris* cf. *carpatica* Dumitrică, 44: *Cinguloturris* cf. *latiannulatum* (Grill and Kozur), 45: *Cinguloturris* sp., 46–47: *Svinitzium* sp., 48: *Wrangellium* sp., 49–55: *Transhsuum* sp., 56–57: Nassellaria gen. et sp. indet., 58: Pyloniodea et sp. indet.

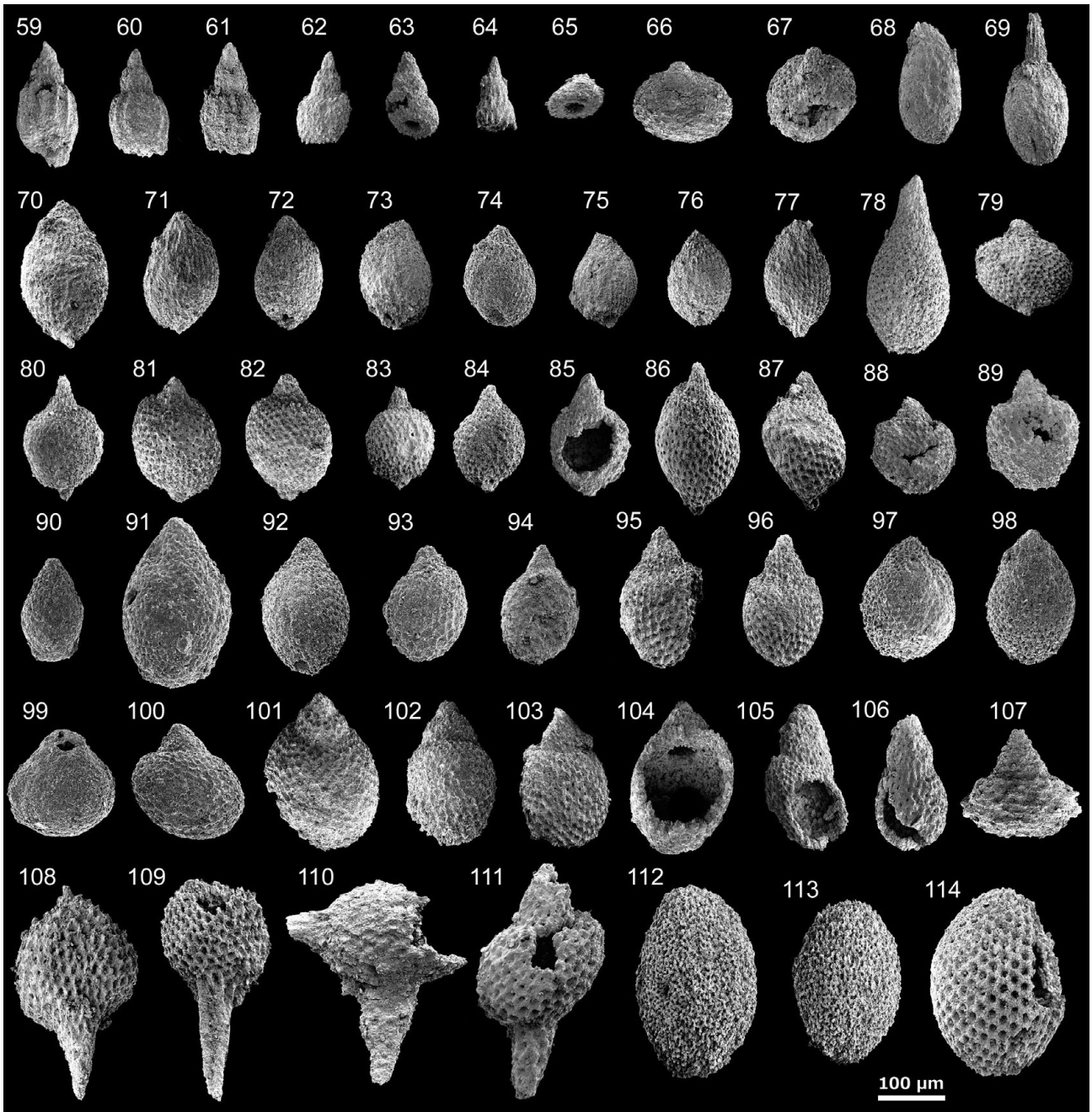


Plate 1 (Continued)

59–61: *Eucyrtidiellum* cf. *pyramis* (Aita), **62:** *Eucyrtidiellum* cf. *nodosum* Wakita, **63:** *Crococapsa* aff. *truncata* (Wu), **64–68:** *Nassellaria* gen. et sp. indet., **69:** *Kilinora* sp., **70:** *Striatojaponocapsa* cf. *conexa* (Matsuoka), **71–76:** *Striatojaponocapsa* *synconexa* O'Dogherty, Goričan and Dumitrică, **77:** *Striatojaponocapsa* sp., **78:** *Praewilliriedellum* sp., **79:** *Hemicryptocapsa* *carpathica* (Dumitrică), **80–87:** *Williriedellum* sp., **88–89:** *Minutosolla* sp., **90:** *Praezhamoidellum*? sp., **91–92:** *Hiscocapsa* *robusta* (Matsuoka), **93:** *Hemicryptocapsa* cf. *yaoi* (Kozur), **94–98:** *Praezhamoidellum* sp., **99:** *Zhamoidellum* sp., **100:** *Zhamoidellum* cf. *ventricosum* Dumitrică, **101–104:** *Zhamoidellum* *ovum* Dumitrică, **105–106:** *Complexapora* aff. *kiesslingi* Hull, **107:** *Minocapsidae* gen. et sp. indet., **108:** *Archaeospongoprimum* sp., **109:** *Archaeospongoprimum* cf. *mizutanii* Ožvoldová, **110:** *Podobursa* sp., **111:** *Favosyringium* cf. *affine* (Rüst) sensu Steiger (1992), **112–113:** *Bistarkum* aff. *mangartense* Goričan, Šmuc and Baumgartner, **114:** *Nassellaria* gen. et sp. indet.

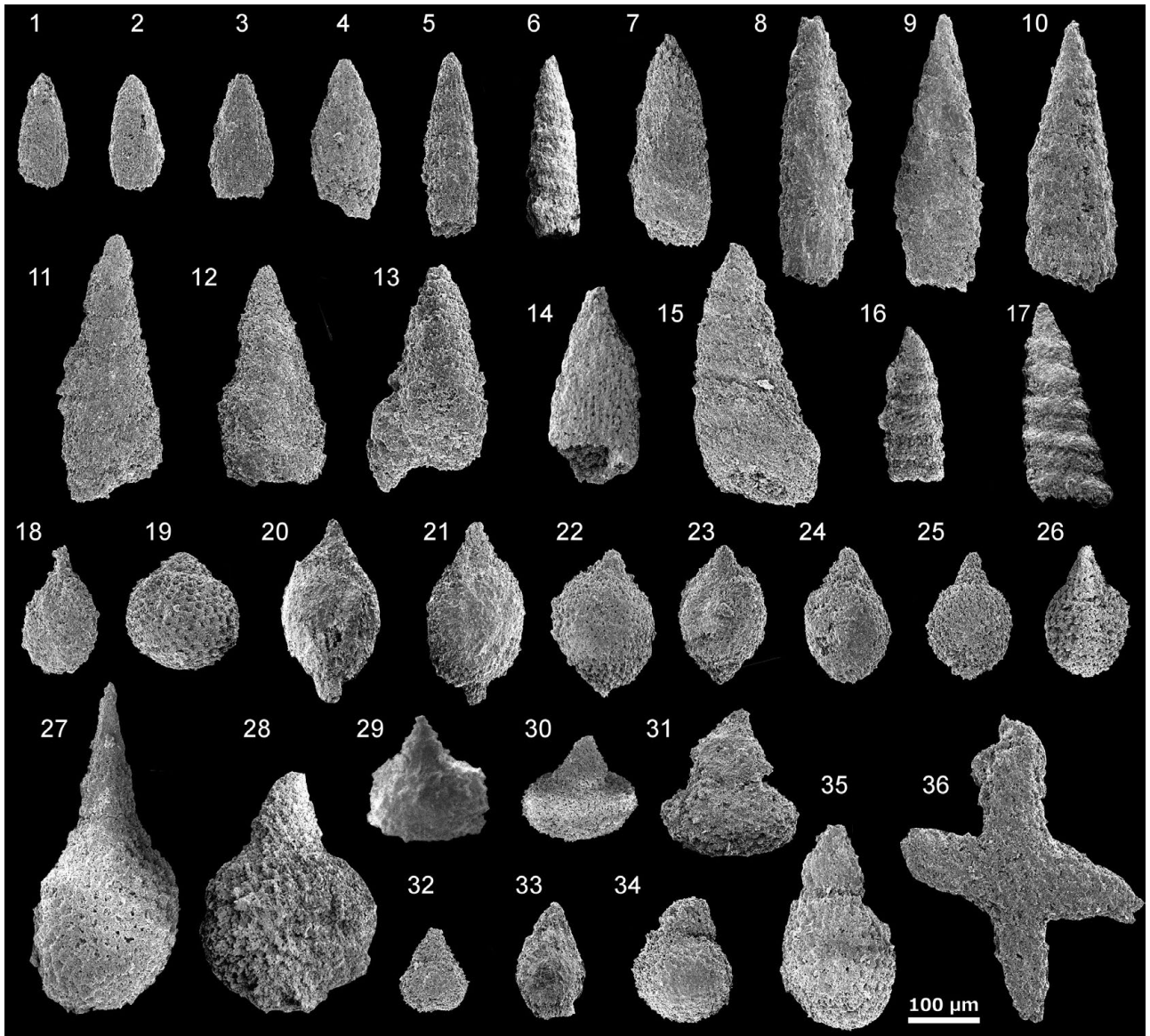


Plate 2 Scanning electron microscopy images of probably Callovian–Kimmeridgian radiolarians extracted from mudstone at Loc. 2. 1–2: *Archaeodictyomitra* aff. *rigida* Pessagno, 3: *Archaeodictyomitra* cf. *tyaughtonensis* Cordey, 4: *Archaeodictyomitra* cf. *inornata* Hull, 5–14: Multisegmented nassellarians, 15: *Cinguloturris* cf. *floridicingula* (Li), 16: *Cinguloturris* cf. *getsensis* O'Dogherty, Goričan and Dumitrică, 17: *Cinguloturris* cf. *carpatica* Dumitrică, 18: Spherical radiolarians, 19: *Zhamoidellum* cf. *mikamense* Aita, 20–23: *Williriedellum* sp., 24–26: *Williriedelloidea* gen. et sp. indet., 27–28: *Syringocapsidae*? gen. et sp. indet., 29: *Eucyrtidiellum* sp., 30: *Quarkus* sp., 31: *Tetracapsa* sp., 32–35: Few segmented nassellarians, 36: Four-armed flat *Pylonioidea*.

下北半島北東部，北部北上帯の U-Pb 年代測定砂岩近傍の泥岩から得られた 後期ジュラ紀放散虫化石

内野 隆之・鈴木 紀毅

要 旨

青森県下北半島の北東部では，北部北上帯に属する付加体が，^{くわばたやま}桑畑山地域，片崎山地域，大森地域に分布している。桑畑山地域の付加体については，後期ジュラ紀の岩屋ユニットと前期白亜紀前半の^{しつかり}尻労ユニットに区分されるなど，これまで多くの研究がなされているものの，片崎山・大森地域の付加体については，大森地域の砂岩から碎屑性ジルコン U-Pb 年代が得られているほかは，詳しいデータはほとんどない。

本研究ではジルコン年代が測定された砂岩近傍の泥岩から *Eucyrtidiellum* cf. *pyramis* をはじめとする後期ジュラ紀（おそらくキンメリッジアン期）の放散虫化石が見出された。この泥岩の化石年代と砂岩のジルコン年代とは大差なく，また泥岩と砂岩との層準の間に不連続構造面も確認されないことから，両者の堆積年代に大きな乖離はないと考えられる。

岩相・地質構造・放散虫化石年代から，片崎山・大森地域の付加体と，桑畑山地域の岩屋ユニットは対比可能である。つまり，下北半島北東部の付加体は，後期ジュラ紀に形成された桑畑山地域の岩屋ユニット及び片崎山・大森地域の未命名ユニットと，前期白亜紀に形成された桑畑山地域の尻労ユニットとに区分される。

難読・重要地名

Chikagawa：近川，Higashidori：東通，Iwaya：岩屋，Kadoma：門馬，Katasakiyama；Mt. Katasaki：片崎山，Kuwabatayama；Kuwabatakeyama；Mt. Kuwabata：桑畑山，Mt. Noborimori：登森，Mt. Omori；Omori：大森，Mt. Toyamori：トヤ森，Magisawa：榎木沢，Magidai：間木平，Sarugamori：猿ヶ森，Shimokita：下北，Shiriya：尻屋，Shiriyazaki：尻屋崎，Shitsukari：尻労，Takayashiki：高屋敷，Tatemachijima：立待島

SIMS analysis of Si isotope for radiolarian test in Mesozoic bedded chert, Inuyama, central Japan

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Maximilien BÔLE, IKEDA Masayuki, Peter O. BAUMGARTNER, HORI S. Rie and Anne-Sophie BOUVIER (2020) SIMS analysis of Si isotope for radiolarian test in Mesozoic bedded chert, Inuyama, central Japan. *Bulletin of the Geological Survey of Japan*, vol. 71(4), p. 331–353, 5 figs, 2 tables, 2 appendices.

Abstract: The global silica cycle is an important component of the long-term climate system, yet its controlling factors are largely uncertain due to poorly constrained proxy records. Because radiolarians and other organisms preferentially extract lighter ²⁸Si from the ocean, the $\delta^{30}\text{Si}$ of biosiliceous tests can thus be used for a potential proxy of productivity. Additionally, $\delta^{30}\text{Si}$ of oceanic silica could have reflected changes in the isotopic ratio of sources and sinks.

Here we show $\delta^{30}\text{Si}$ records measured by secondary ion mass spectrometer (SIMS) in radiolarian silica, precipitated inside radiolarian molds in early Mesozoic bedded chert of the Inuyama section, central Japan. Range of measured $\delta^{30}\text{Si}$ between -0.3 and 2 ‰ is consistent with that of modern and Cenozoic radiolarian tests. Relatively large intra-chert bed variability up to ~ 0.8 ‰ (1SD) support that $\delta^{30}\text{Si}$ of the Mesozoic radiolarian molds are not perfectly homogenized in a chert bed during diagenesis. We found an overall inverse correlation between 10-Myr scale $\delta^{30}\text{Si}$ and biogenic silica (BSi) burial flux, which contradicts with a conventional interpretation of $\delta^{30}\text{Si}$ as paleoproductivity proxy, despite the low-resolution and scattered our $\delta^{30}\text{Si}$ records. Although most of the factors controlling oceanic $\delta^{30}\text{Si}$ are difficult to be constrained, this inverse relation might be explained by changes in $\delta^{30}\text{Si}$ of mafic/felsic rock weathering ratio, which inferred from paleogeography. Further high-resolution $\delta^{30}\text{Si}$ records will allow a better understanding of the past silica cycle.

Keywords: Silicon isotopes, $\delta^{30}\text{Si}$, Radiolarites, Mesozoic oceanic silica cycle, SIMS

1. Introduction

The global silica cycle is linked to long-term changes in Earth's climate through feedback mechanisms between atmospheric CO₂, climate and the rate of silicate weathering, followed by carbonate and biogenic silica (BSi) deposition. Changes in Si and C cycle dynamics are linked to global climate changes throughout Earth's history, a relationship, which in turn, allows numerical models to reconstruct past atmospheric pCO₂ (Berner, 1991). Understanding the global silica cycle is therefore crucial to elucidate the response of Earth's surface system to changes in external (astronomical) and internal (tectonic and volcanic) forcings.

Silicate weathering and BSi burial are important to constrain the silica cycle as major source and sink,

respectively, but are difficult to quantify, and poorly understood their dynamic relation due to large uncertainties in the proxy records. Radiolarians dominated as producers of BSi during much of the Phanerozoic (Hein *et al.*, 1987), whereas siliceous sponges are largely restricted to marginal settings, and diatoms became quantitatively important only in the Cenozoic (Racki and Cordey, 2000; Kidder and Erwin, 2001). Radiolarites were deposited in a broad low-latitude belt, while radiolarian-bearing siliceous mudstones dominated in mid-latitudes (Baumgartner, 2013).

The volume of Paleozoic and Mesozoic Radiolarian-rich deposits is largely underestimated, because much of the ocean floor has been subducted. Plate tectonic reconstructions of Panthalassa and Tethys, based on accreted remnants preserved in Circum-Caribbean,

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Circum-Pacific and Himalayan terranes, suggest that radiolarian-rich sediments covered more than 80 % of the area of the Mesozoic ocean (Baumgartner *et al.*, 2018).

The modern oceanic silica cycle is relatively well-known, and is considered to be close to steady state (Tréguer and De La Rocha, 2013). Rivers are the main suppliers of silicic acid to oceans followed by seafloor weathering, groundwater, hydrothermal and aeolian inputs. The surface BSi production of diatoms overpasses by two orders of magnitude of the total silicon input to the ocean resulting in high dissolved silica (DSi) undersaturation and strong recycling. With depth, the undersaturation becomes weaker due to silicon recycling, but never reaches the saturation of any silica phases. Only about 3 % of BSi produced in the photic zone is trapped in sediments (Tréguer and De La Rocha, 2013). If steady state is assumed, the total BSi burial has to be proportional to the total input from all sources on timescales longer than the residence time of oceanic DSi (Tréguer and De La Rocha, 2013).

BSi of radiolarian silica in bedded cherts is potentially a unique proxy for past Si cycle, because the estimated radiolarian BSi burial flux in the low-latitude pelagic Panthalassa was comparable with the modern global BSi flux, and was the possible major sink of DSi (Ikeda *et al.*, 2017). This hypothesis is consistent with overall in-phase relation between radiolarian BSi flux and global silicate weathering flux calculated by GEOCARBSULFvolc model (Ikeda *et al.*, 2017), implying that the estimated BSi burial flux can be proportional to silicate weathering flux over timescales longer than residence time of oceanic Si (<100 kyr; Ritterbush *et al.*, 2015).

However, controlling factors for the BSi burial are still controversial. Although changes in oceanic upwelling intensity and consequent BSi productivity are proposed as potential controlling factors (Hori *et al.*, 1993; De Wever *et al.*, 2014), their temporal variations are also difficult to be understood due to large uncertainty in their proxy records, such as Al/Ti ratio (Murray *et al.*, 1993; Murray and Leinen, 1996; Dymond *et al.*, 1997). On the other hand, controlling factors for siliceous weathering are also still debated. Today, more than 70 % of silicate weathering occurs only in <10 % land area with highly-weatherable volcanic rock region under humid monsoonal climate (Hartmann *et al.*, 2014). Considering the Mesozoic paleogeography, wide distribution of the volcanic islands and large igneous provinces under intensified mega-monsoonal climate could have further modulate the global silicate weathering (Ikeda *et al.*, 2017), despite of lack of quantitative constraints.

Si isotope of BSi is a potential proxy to understand past Si cycle. Glacial-Interglacial scale $\delta^{30}\text{Si}$ variations have been documented (e.g. Brzezinski *et al.*, 2002), potentially due to an increase of the diatom productivity and extraction of light silicon by diatoms during interglacial periods (De La Rocha *et al.*, 1998).

Only few scattered data of $\delta^{30}\text{Si}$ from radiolaria are

published. (Wu *et al.*, 1997; Egan *et al.* 2012; Ding *et al.*, 1996; Hendry *et al.* 2014; Abelmann *et al.*, 2015; Fontorbe *et al.*, 2016). Silicon fractionation by modern radiolarians varies between -0.8 ‰ and -2.1 ‰ (Egan *et al.*, 2012; Abelmann *et al.*, 2015), which is similar to that by diatom (Frings *et al.*, 2016). Although factors controlling of $\delta^{30}\text{Si}$ records of radiolarian test are still debated, even for Cenozoic (e.g. Fontorbe *et al.*, 2016), early Mesozoic Si cycle seems to be a simpler system due to lack of diatom in continent and ocean. In this paper, we investigated the past oceanic silica cycle through in situ $\delta^{30}\text{Si}$ in radiolarian molds of Mesozoic bedded cherts. Then we compared our $\delta^{30}\text{Si}$ records with BSi burial flux (Ikeda *et al.*, 2017), to constrain the early Mesozoic Si cycle.

2. Material

We sampled material from bedded cherts from the Inuyama area, central Japan (Fig. 1). These cherts are part of an accretionary prism and are incorporated into several tectonic imbricates (Matsuda and Isozaki, 1991; Kimura and Hori, 1993). High-resolution radiolarian and conodont biostratigraphy, chemo-cyclostratigraphy in this succession have allowed to reconstruct the best studied Early Triassic to Early Jurassic bedded chert sequence (Yao *et al.*, 1980, Hori, 1990; Sugiyama, 1997; Ikeda *et al.*, 2010; Ikeda and Tada, 2013, 2014). Based on biostratigraphic age constraints, average duration of a chert-shale couplet are ~20 kyr throughout the early Mesozoic (Ikeda *et al.*, 2010; Ikeda and Tada, 2014), which is consistent with the precession-scale changes in the accumulation rate of BSi under the extremely slow accumulation of shale mostly composed of aeolian dust (e.g. Hori *et al.*, 1993). Estimated BSi fluctuations should be proportional to DSi input from chemical weathering paced with the monsoon dynamics, over timescales longer than the residence time of oceanic DSi (20 kyr; Tréguer and De La Rocha, 2013; <~100 kyr; Ritterbush *et al.*, 2015), because low-mid-latitude BSi burial flux (Ikeda *et al.*, 2017) is ~90 % of the modern global ocean (Tréguer and De La Rocha, 2013) and was a major sink for oceanic DSi.

Bedded cherts are rocks composed of chert layers (Si-rich), interbedded with clay-rich shale partings (Si-poor), produced by differential compaction and diagenetic reactions of dissolution-precipitation usually forming opal-CT and later quartz (Isaacs, 1981; Tada, 1991). Radiolarian molds filled with nearly pure microquartz and/or chalcedony are found in the silica-rich matrix of cherts. The radiolarian molds that we measured are commonly spherical. Therefore, they could result from Spumellaria, which have regularly a spherical morphology, dwelling in a photic zone due to their symbiotic relation with photosynthetic algae (e.g. Swanberg and Anderson, 1985; Takahashi *et al.*, 2003).

Inuyama Sections

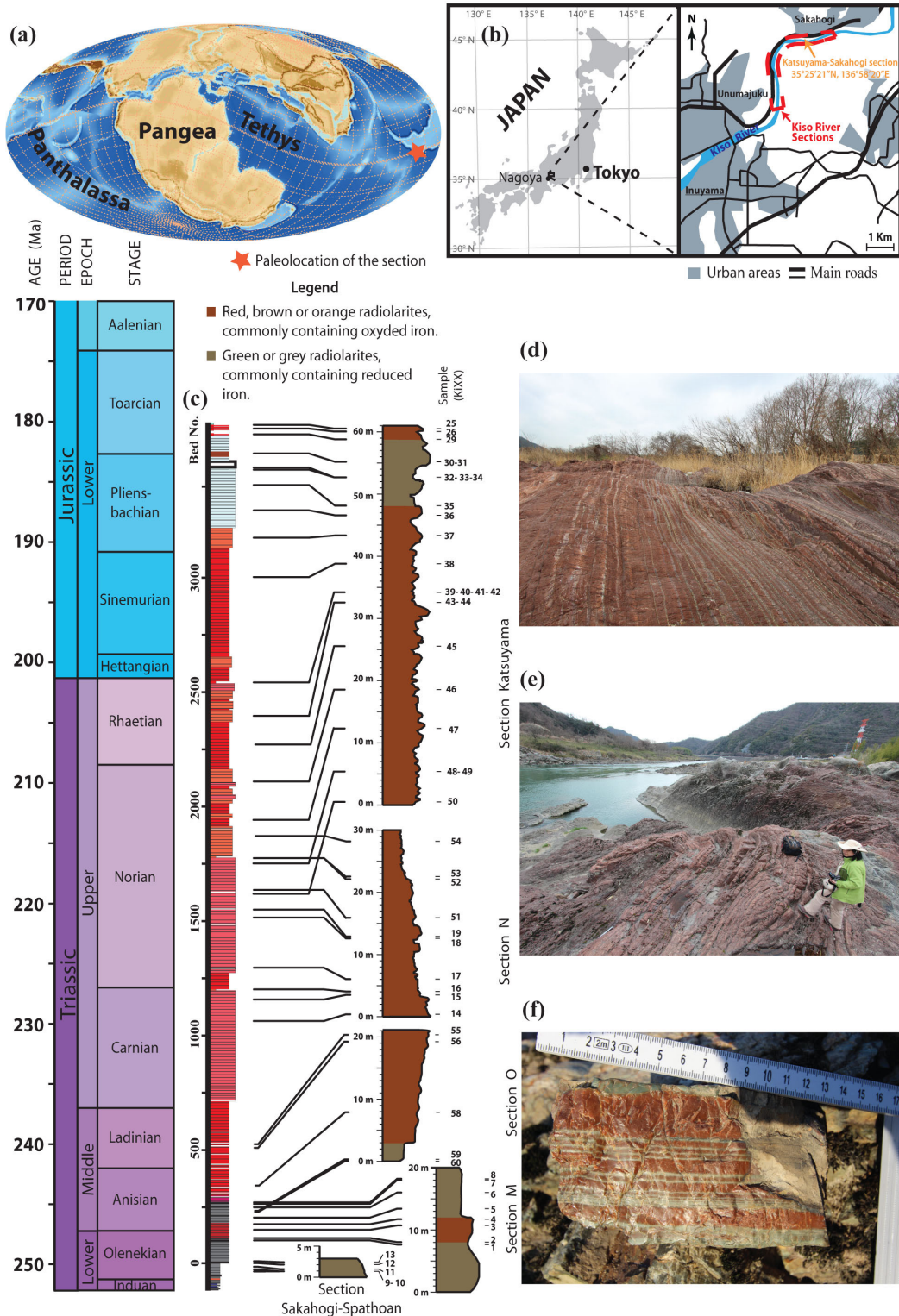


Fig. 1 Log of the Inuyama sections with their Triassic-Early Jurassic paleogeography on a map of Middle Jurassic (a) and their current location (b). The Paleomap (a) is from the Stampfli model developed at the University of Lausanne (Stampfli and Borel, 2002). The bed number log (c) is from Ikeda and Tada (2014). Additional information on the Kiso River sections can be found in Sugiyama (1997). These radiolarites are illustrated through (d) the nice parallel bedding for Late Triassic bedded chert ($35^{\circ}23'57''$ N, $136^{\circ}57'34''$ E), (e) outcropping of Rhaetian bedded chert along the Kiso River ($35^{\circ}25'21''$ N, $136^{\circ}58'16''$ E) and (f) millimetric laminations inside single Norian bed (Ki18).

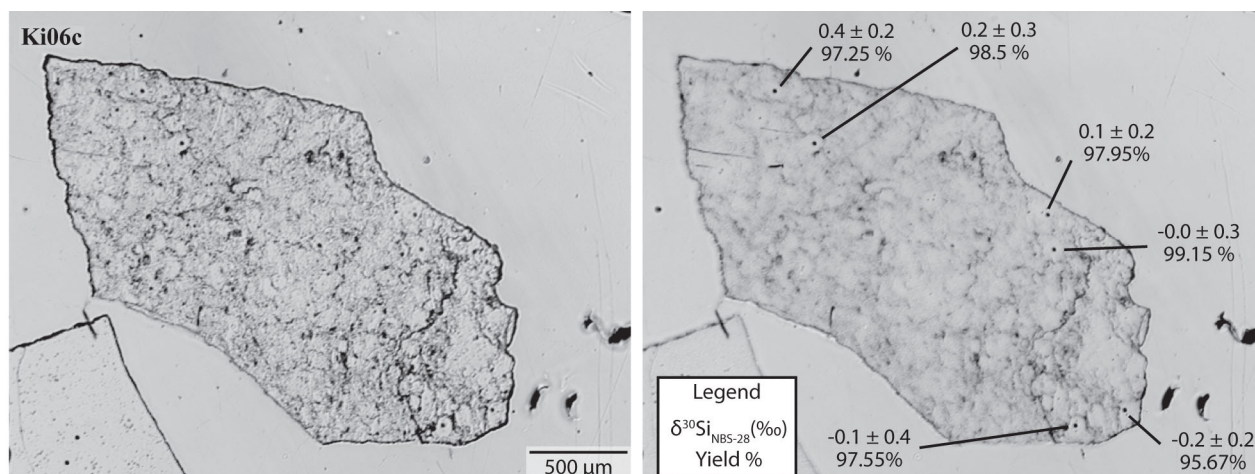


Fig. 2 $\delta^{30}\text{Si}_{\text{NBS-28}}$ measurements on sample Ki06c with analytical yield percent relative to the yield of the Paine Quartz Standard (UNIL_Q1). Image on the right are out of focus to better distinguish the analytical spots.

3. Methods

In total, 34 cherts were analysed for the Inuyama section. Sample holders consists of ten fragments of different samples mounted into epoxide around an internal standard. These fragments were previously polished into trapezoidal shapes and the presence of radiolarian molds was checked by optical methods.

The $\delta^{30}\text{Si}$ of micro-crystalline quartz precipitated inside radiolarian molds was measured by SIMS at University of Lausanne with a primary Cs^+ ion beam intensity of 2 nA, resulting in a $\sim 10 \mu\text{m}$ spot (cf. Seitz *et al.*, 2017), to avoid contamination from other sources of silicon in radiolarites (detrital/aeolian minerals). Secondary ions ^{30}Si and ^{28}Si were analyzed at 3000 MRP and collected on Faraday cups (FC) multi-collection mode. The resistances of the L'2 and H'2 FC were $10^{11} \Omega$ for the detection of ^{28}Si and ^{30}Si , respectively. FCs were calibrated in the beginning of each session, using the calibration routine. Mass calibration was performed at the beginning of each session and every 12 h. Samples were gold coated to dissipated charges. Each analysis consists of 20 cycles of 5 sec, and starts with a presputtering time of 30 sec to remove gold and stabilize the secondary ion emission. The standard deviation of each analysis is expressed as analytical standard deviation. The data have been obtained in 7 different sessions for $\delta^{30}\text{Si}$ measurements, over 7 months.

For each chert sample, we made 4-10 measurements within about 0.5 cm stratigraphic interval (Fig. 2). A quartz internal standard (UNIL_Q1; Paine Quartz; Seitz *et al.*, 2017 for $\delta^{18}\text{O}$ and method; $\delta^{30}\text{Si}_{\text{NBS-28}} = -0.13 \pm 0.02 \text{ ‰}$ (2SD)) was analysed every 6-10 measurements for instrumental drift correction and calibration.

We subsequently controlled by optical methods that the ion beam actually hit the radiolarian molds for each measurement. In addition, data were postprocessed using the analytical yield and the analytical deviation of each

measurement. The analytical yield depends on the nature of the analysed material (mineral species and matrix effect) and on the topography of the analysed surface which modifies the incident angle of the primary ion beam. In addition to instrumental instabilities, the high analytical deviation can also indicate heterogeneity and the analyse of a mixture of silica, clays minerals and/or oxides. Regarding these considerations, the analytical yield and deviation are objective parameters to decide if a measurement must be rejected.

The drift correction was realized using a least square regression line weighted for incertitude (σ^2). For the calibration, we calculated the least square $\delta^{30}\text{Si}$ -mean (\bar{x}) and standard deviation (σ_i) for the internal standard also weighted for incertitude (Equation 1 and 2) to keep consistent data processing with the least square drift correction. The calibrated $\delta^{30}\text{Si}$ for samples ($\delta^{30}\text{Si}_{\text{NBS-28}} \text{ Spl}$) depend on each sample measurement ($\delta^{30}\text{Si}_{\text{NBS-28}} \text{ Spl}_{\text{measured}}$) and are proportional to the measured least square $\delta^{30}\text{Si}$ -mean and the true $\delta^{30}\text{Si}_{\text{NBS-28}}$ from the internal standard ($\delta^{30}\text{Si}_{\text{NBS-28}} \text{ Std}_{\text{measured}}$ and $\delta^{30}\text{Si}_{\text{NBS-28}} \text{ Std}$, respectively) (Equation 3). The errors on the calibrated $\delta^{30}\text{Si}$ ($\sigma(\delta^{30}\text{Si}_{\text{NBS-28}} \text{ Spl})$) were obtained by error propagation (Equation 4). The weighted means and standard deviations (Table 2 and Appendix Tables A1 and A2) were then calculated for each sample following equation 1 and 2. Raw, drift corrected and calibrated data are indicated in appendix tables. The $\delta^{30}\text{Si}$ -data were then filtered with a 10 Ma moving windows average with a step of 5 Ma and compared with estimation of the BSi burial rates in the Inuyama area (Ikeda *et al.*, 2017).

Equation 1

$$\bar{x} = \sum \left(\frac{1}{\sigma_i^2} \times x_i \right) / \left(\frac{1}{\sigma_i^2} \right)$$

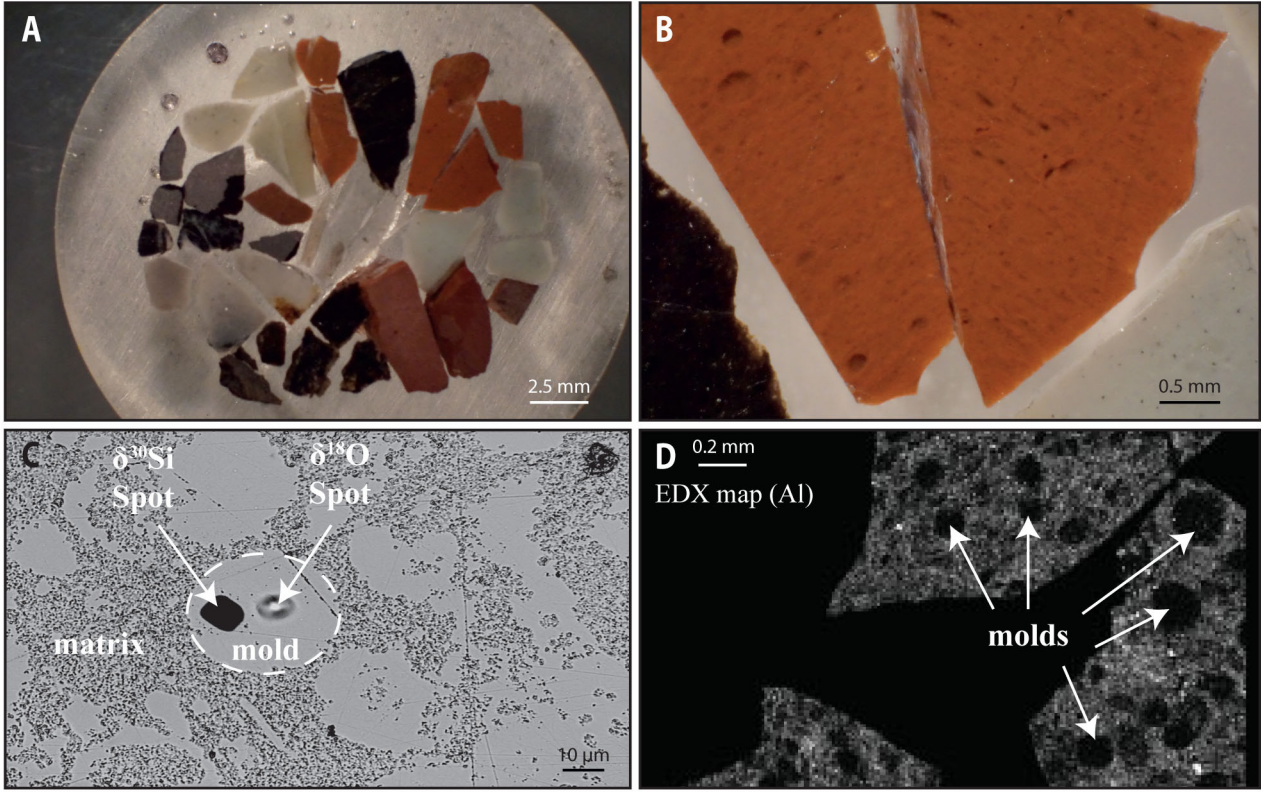


Fig. 3 Illustration of the analyzed materials. A) epoxy sample mount (Br7) including several fragments of about 10 samples. B) Zoom on a sample in this mount with binocular. C) Image of two spots left by a SIMS analysis ($\delta^{18}\text{O}$ and $\delta^{30}\text{Si}$) in radiolarian silica (radiolarian molds) on a gold coated sample mount. The $\delta^{18}\text{O}$ -spot on the right is covered by a new gold coating. The difference of polishing between the radiolarian molds of nearly pure microcrystalline quartz and the matrix is well illustrated on this image. D) SEM imaging of the aluminium distribution in the sample Ki08c (EDX map). Radiolarian molds are aluminium-free on this image.

Equation 2

$$\hat{\sigma}_i = \sqrt{\frac{\sum \left(\frac{1}{\sigma_i^2} \times (x_i - \bar{x})^2 \right)}{\sum \left(\frac{1}{\sigma_i^2} \right) \times \frac{N}{N-1}}}$$

Equation 3

$$\begin{aligned} \delta^{30}\text{Si}_{\text{NBS-28 Spl}} &= \left(\left(1 + \frac{\delta^{30}\text{Si Spl}_{\text{measured}}}{1000} \right) \right. \\ &\quad \left. / \frac{(1 + \delta^{30}\text{Si Std}_{\text{measured}} / 1000)}{(1 + \delta^{30}\text{Si}_{\text{NBS-28 Std}} / 1000)} - 1 \right) \times 1000 \end{aligned}$$

Equation 4

$$\sigma(\delta^{30}\text{Si}_{\text{NBS-28 Spl}}) = \sqrt{\left(\frac{\partial F}{\partial V_1} \times dV_1 \right)^2 + \left(\frac{\partial F}{\partial V_2} \times dV_2 \right)^2 + \left(\frac{\partial F}{\partial V_3} \times dV_3 \right)^2}$$

$$\text{With } F = \delta^{30}\text{Si}_{\text{NBS-28 Spl}}, V_1 = \delta^{30}\text{Si Spl}_{\text{measured}}, V_2 =$$

$$\delta^{30}\text{Si Std}_{\text{measured}} \text{ and } V_3 = \delta^{30}\text{Si}_{\text{NBS-28 Std}}$$

For the SIMS analyses, it is common to use 2SD, which make sense considering the high accuracy of the method or when measuring very homogenous samples. 2SD was thus also used to discuss the UNIL-Q1 $\delta^{30}\text{Si}$. We used 1SD for the LS-mean of samples and for their moving average, following usage in palaeoceanography, such as $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ in low magnesium calcium shells through time (e.g. Veizer *et al.*, 1999).

We also use scanning electron microscopy-energy dispersive X-ray spectrometry (SEM-EDS) in Lausanne University to map the elemental distribution in chert (Fig. 3).

4. Results

All the analytical dataset of our samples and a standard is presented in Appendix Tables, and is summarised in Tables 1 and 2. The means and standard deviations (2SD) of the raw $\delta^{30}\text{Si}$ from UNIL-Q1 range from -41.77 ‰ to -45.6 ‰ and from 0.33 ‰ to 0.69 ‰, respectively (Table 1). The drift correction only slightly reduced the standard deviations (0.29 ‰ to 0.68 ‰; 2SD). Calibrating data using the LS-means of standard, the arithmetic means and standard deviation (2SD) of the $\delta^{30}\text{Si}_{\text{NBS-28}}$ values

Table 1 Raw, drift corrected and calibrated $\delta^{30}\text{Si}$ means and standard deviations (2 SD) for the UNIL-Q1 standard (in ‰) between our different sessions. Instrumental fractionation is calculated based on the least square mean and standard deviation of the drift corrected $\delta^{30}\text{Si}$. The $\delta^{30}\text{Si}_{\text{NBS-28}}$ mean and standard deviation (2 SD) of calibrated data are given for all standard measured through a session and for the average $\delta^{30}\text{Si}_{\text{NBS-28}}$ of the different standard clusters. Reproducibility of the $\delta^{30}\text{Si}_{\text{NBS-28}}$ based on the average $\delta^{30}\text{Si}_{\text{NBS-28}}$ of the different standard clusters is much better than on all standard measured ($< 0.37\text{‰}$, 2SD vs $< 0.71\text{‰}$).

Session	Number of bracketing standard	Raw $\delta^{30}\text{Si}$ (‰)		drift corrected $\delta^{30}\text{Si}$ (‰)			Fractionation based on LS-mean			Calibrated $\delta^{30}\text{Si}_{\text{NBS-28}}$ (‰)				Reproducibility of $\delta^{30}\text{Si}_{\text{NBS-28}}$ (‰)			
		All standard		All standard			LS-mean			All standard		Standard clusters		All standard		Standard clusters	
		mean	2 SD	mean	2 SD	LS-mean	LS-STD (1SD)	LS-STD (1SD)	mean	2SD	mean	2SD	mean	2SD	2SD	2SD	
Br2	36	-42.67	0.38	-42.67	0.38	-42.65	0.19	1.0444	-0.15	0.40	-0.15	0.21	0.40	0.40	0.21	0.21	
Br3	28	-42.84	0.33	-42.84	0.31	-42.84	0.16	1.0446	-0.13	0.33	-0.13	0.15	0.33	0.33	0.15	0.15	
Br4	16	-45.60	0.49	-45.60	0.39	-45.61	0.19	1.0477	-0.12	0.40	-0.12	0.36	0.40	0.40	0.36	0.36	
Br4v2	18	-43.08	0.69	-43.08	0.68	-43.12	0.37	1.0449	-0.09	0.71	-0.09	0.02	0.71	0.71	0.03	0.03	
Br6	40	-45.36	0.67	-45.36	0.66	-45.36	0.31	1.0474	-0.15	0.69	-0.15	0.20	0.69	0.69	0.21	0.21	
Br7	40	-42.44	0.35	-42.44	0.33	-42.44	0.17	1.0442	-0.13	0.35	-0.13	0.25	0.35	0.35	0.25	0.25	
Br1	24	-41.77	0.37	-41.77	0.29	-41.76	0.12	1.0434	-0.14	0.30	-0.14	0.13	0.30	0.30	0.13	0.13	
Br5	44	-41.79	0.39	-41.79	0.39	-41.80	0.19	1.0435	-0.12	0.41	-0.12	0.17	0.41	0.41	0.17	0.17	
mean		-43.19		-43.20		-43.20			-0.13		-0.13						
2SD		2.97		2.98		2.98			0.04		0.04						

from our standards vary from -0.09‰ to -0.15‰ and from 0.30‰ to 0.71‰ , respectively. The average of the arithmetic $\delta^{30}\text{Si}_{\text{NBS-28}}$ means obtained during different sessions is thus $-0.13 \pm 0.04\text{‰}$ (2SD) which is relatively similar to the bulk UNIL-Q1 $\delta^{30}\text{Si}_{\text{NBS-28}}$ ($-0.13 \pm 0.02\text{‰}$; 2SD) and below the reproducibility of all $\delta^{30}\text{Si}_{\text{NBS-28}}$ from standard during a single session. We removed 6 on 189 measurements of samples which had yield deviating more than 10 % from the yield of the quartz standard or their analytical deviation (2SD) exceeding 0.37‰ .

LS-mean $\delta^{30}\text{Si}$ of measured radiolarian molds ranges from -0.3‰ to 2‰ (Table 2). $\delta^{30}\text{Si}$ -means have an inter-sample range of up to 2.3‰ , which is higher than their standard deviation (Table 2). The intra-sample standard deviations (1SD) of the $\delta^{30}\text{Si}$ -means varies between 0.1‰ and 0.75‰ . The $\delta^{30}\text{Si}$ -means from our samples range from -0.3‰ to 0.8‰ during the Early Triassic, from -0.3‰ to 1‰ during the Middle Triassic, from -0.3‰ to 1.5‰ during the Late Triassic, and from 0.5‰ to 2‰ during the Early Jurassic (Fig. 4). The low $\delta^{30}\text{Si}$ -values from 10-Myr moving windows average are overall associated with high BSi burial rates (Ikeda *et al.*, 2017). A mapping of the aluminium content by SEM-EDS shows that clay minerals are concentrated in the matrix (Fig. 3).

5. Discussion

5.1 $\delta^{30}\text{Si}$ of radiolarian molds and diagenesis

SIMS-measured $\delta^{30}\text{Si}$ for Mesozoic radiolarian molds ranges from -0.3‰ to 2‰ (Fig. 4), which overlap with the range of Cenozoic radiolarian tests (e.g. Fontorbe *et al.*, 2016; Fig. 5), potentially supporting that the Mesozoic radiolarian molds preserve the original values to some extent.

Radiolarian skeletons are originally composed of biogenic opal (opal-A), which is the most soluble silica phase (Walther and Helgeson 1977; Fournier and Rowe, 1977; Gunnarsson and Arnórsson, 2000). During the phase transitions from opal-A to opal-CT and quartz, silicon isotope of radiolarian molds might have changed by contamination of pore water DSi. However, migration of Si from the layers with low Si content (shale bed) to layers with high Si content (chert bed) allows us to ignore interbed migration of Si (Tada, 1991). The lower solubility of quartz ($<1000\text{ ppm}$) than opal-A ($<2000\text{ ppm}$) and opal-CT ($<1500\text{ ppm}$) (e.g. Gunnarsson and Arnórsson, 2000) further implicates the negligible effect of aeolian/detrital quartz dissolution on DSi of pore water. Additionally, clay mineral diagenesis occurs at higher temperature (80 °C ; Chamley, 1989; Fagel 2007) than opal-CT transition (65 °C ; Matheney and Knauth, 1993), which segregated biosiliceous sediments (Tada, 1991). Considering mass balance in the chert-dominant bedded chert succession with minor-clay component in the Inuyama area (Sugiyama, 1997), we can thus assume that the bulk $\delta^{30}\text{Si}$ of radiolarian molds in cherts would be equal to that of former opal-A.

Table 2 List of samples with their age, their $\delta^{30}\text{Si}_{\text{NBS-28}}$ Least square mean (LS-mean) and their $\delta^{30}\text{Si}$ least square standard deviation (LS-std; 1SD). The $\delta^{30}\text{Si}_{\text{NBS-28}}$ was averaged with a 10 Ma windows moving average (5 Ma step) and compared with BSi.

Sample	Age (Ma)	Number of measurements	Results		Curves	
			$\delta^{30}\text{Si}_{\text{NBS-28}}$ (‰)		$\delta^{30}\text{Si}_{\text{NBS-28}}$ (‰)	Biosilica burial rate (g cm ⁻² Kyr ⁻¹)
			This study		This study	Ikeda <i>et al.</i> , 2017
			LS-mean	LS-std (1SD)	10 Ma moving average	10 Ma smooth
Ki20	174.00	5	2.0	0.7	1.42	
Ki22c1	178.00	8	1.1	0.1	1.24	
Ki22c2	178.00	3	0.8	0.2	1.24	
Ki21	178.00	9	1.2	0.3	1.24	
Ki27	180.99	6	0.5	0.6	1.26	0.26
Ki24	182.00	6	1.9	0.2	1.31	0.26
Ki32	184.20	10	2.0	0.4	1.43	0.26
Ki34	184.37	10	1.0	0.3	1.43	0.26
Ki35	185.62	9	2.0	0.2	1.45	0.25
Ki38	193.31	10	0.8	0.4	0.96	0.19
Ki40	201.50	5	0.8	0.5	1.22	0.23
Ki39	201.50	5	1.4	0.3	1.22	0.23
Ki41	201.50	1	1.1		1.22	0.23
Ki42	201.50	10	1.4	0.3	1.22	0.23
Ki44	204.82	9	1.4	0.1	1.22	0.25
ki43	204.82	6	1.2	0.3	1.22	0.25
Ki46	210.27	10	-0.3	0.6	-0.09	0.27
Ki54	214.40	6	0.0	0.5	0.40	0.26
ki48	217.09	4	1.0	0.4	0.71	0.25
Ki51	219.00	6	1.1	0.3	0.94	0.25
Ki15	228.00	10	1.1	0.3	1.10	0.21
Ki57	241.00	9	1.0	0.4	0.52	0.36
Ki58	243.48	6	0.6	0.5	0.39	0.36
Ki08	244.90	10	-0.2	0.3	0.32	0.35
Ki07	245.00	6	0.9	0.2	0.32	0.35
Ki06	245.25	6	0.0	0.3	0.31	0.35
Ki06s	245.25	14	-0.1	0.4	0.31	0.35
Ki05	246.20	6	-0.3	0.2	0.30	0.33
Ki04	246.60	9	0.7	0.3	0.30	0.33
Ki03	247.20	10	0.2	0.6	0.29	0.32
Ki02	247.80	7	0.4	0.4	0.28	0.31
Ki01	248.00	10	0.3	0.5	0.28	0.31
Ki10	250.20	6	-0.2	0.3	0.25	
Ki09	250.30	9	0.7	0.3	0.25	

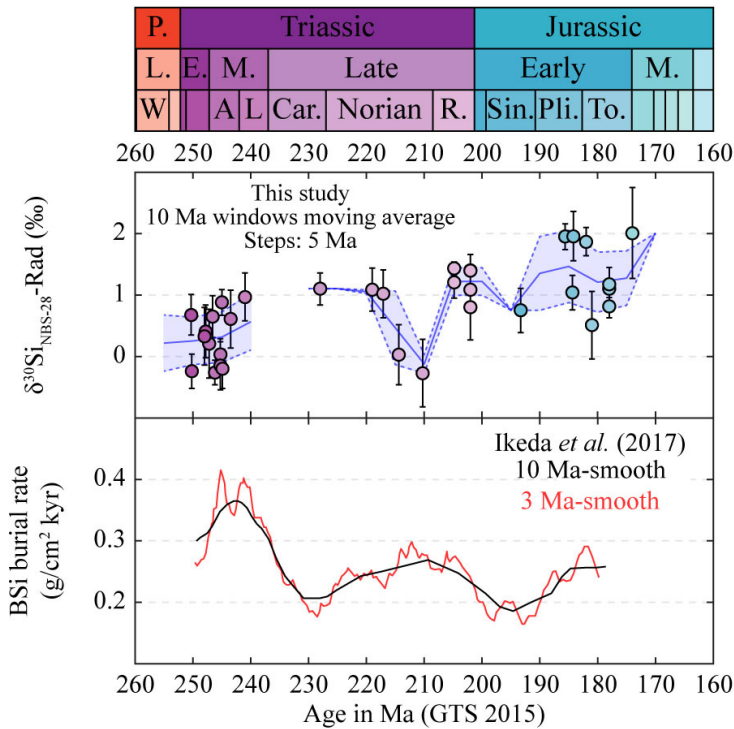


Fig. 4 Evolution of $\delta^{30}\text{Si}_{\text{NBS-28}}$ through time from radiolarian silica in the Inuyama Area (this study). Our results are compared with estimations of the BSi burial rate (Ikeda *et al.*, 2017). The trend and variation of $\delta^{30}\text{Si}$ can be correlated ($R = -0.73$) with the trend and/or variations observed in the BSi burial rate. The geological timescale (GTS 2015) used for this figure is the timescale of the international commission of stratigraphy (Cohen *et al.* 2013). The color filling inside markers corresponds to the color of the geological stage. The boundaries of curves, when plotted, are equivalent to 1SD. The moving average for radiolarian silica is realized using a 10 Ma windows and 5 Ma steps. The $\delta^{30}\text{Si}_{\text{NBS-28}}$ error bars correspond to the least square standard deviation of samples presented in Table 2.

The relatively large internal $\delta^{30}\text{Si}$ -scattering up to 0.8 ‰ (1SD) for each sample suggests that kyr-scale $\delta^{30}\text{Si}$ heterogeneity within each chert sample still exist after the diagenesis. Micrometric isotopic variations have been previously observed even in Precambrian cherts, supporting that the $\delta^{30}\text{Si}$ is not homogenised through time in cherts (Marin-Carbonne *et al.*, 2011, 2012). Even fragments of cherts included as enclave in tonalitic intrusions (>700 °C) or metamorphosed in amphibolite facies seem to preserve their $\delta^{30}\text{Si}$ (André *et al.*, 2006). Therefore, it is reasonable to assess $\delta^{30}\text{Si}$ records of radiolarian molds in Mesozoic bedded chert as those in Mesozoic radiolarian tests, despite of diagenetic homogenization to some extent.

5.2 Evolution of radiolarian $\delta^{30}\text{Si}$

The increasing trend of radiolarian $\delta^{30}\text{Si}$ through the Triassic might be interpreted as an increase of the radiolarian productivity resulting in a higher biogenic fractionation (e.g. De La Rocha *et al.* 1998), despite of large scattering and complex fractionation of $\delta^{30}\text{Si}$ (Fig. 4). However, this conventional interpretation contradicts with low $\delta^{30}\text{Si}$ mainly associated with higher BSi burial rates (Ikeda *et al.*, 2017) (Fig. 5). Upwelling of isotopically-light DSi might have affected the observed negative correlation between $\delta^{30}\text{Si}$ and BSi flux in equatorial Panthalassa. Regarding the radiolarian BSi as major sink of DSi in the Mesozoic ocean before the post-Cretaceous rise of diatoms (Ikeda *et al.*, 2017), however, radiolarian $\delta^{30}\text{Si}$ could have reflected $\delta^{30}\text{Si}$ of oceanic DSi on timescale longer than residence time of oceanic DSi (~100 kyr: Ritterbush *et al.*, 2015).

$\delta^{30}\text{Si}$ of oceanic DSi is controlled by changes in $\delta^{30}\text{Si}$

values of sources and sinks (e.g. Frings *et al.*, 2016). Major source of oceanic DSi is river input, which $\delta^{30}\text{Si}$ currently varies from 0 ‰ to 4 ‰, mainly depending on diatom uptake and rock types of provenance (e.g. Frings *et al.*, 2016). However, before the rise of diatom, biogenic uptake in continent can be negligible. Small difference exists between continental felsic rocks ($\delta^{30}\text{Si} = -0.5$ to 0.5 ‰) and mantle-origin mafic rocks ($\delta^{30}\text{Si} = -1$ to 0 ‰) (Opfergelt and Delmelle, 2012). Up to 1 ‰ amplitudes of 10-Myr scale $\delta^{30}\text{Si}$ data can be explained by changes in felsic/mafic ratios, although our $\delta^{30}\text{Si}$ data is too low-resolution to discuss its <10-Myr scale dynamics.

On another hand, $\delta^{30}\text{Si}$ of siliceous sponges varies from -6 ‰ to -1 ‰ (Frings *et al.*, 2016), whereas that of radiolarians and diatoms varies from -1.1 ‰ to 1.7 ‰ (Abelmann *et al.*, 2015; Fontorbe *et al.*, 2016) and from -1 ‰ to 3 ‰ (e.g. Frings *et al.*, 2016), respectively. Changes in the relative contribution of sponge BSi deposition might be a candidate to explain $\delta^{30}\text{Si}$ variations, despite of lack of evidence of massive sponge deposition, except for some biotic events after Carnian Pluvial Event, Norian Manicouagan impact, and the end-Triassic extinction (Thibodeau *et al.*, 2016; Onoue *et al.*, 2016; Shi *et al.*, 2017). However, there are no significant $\delta^{30}\text{Si}$ variations across the end-Triassic extinction, implying negligible effect of sponge deposition on Si cycle at this event (Fig. 4).

On the other hand, 10-Myr scale BSi burial flux also correlates with calculated global silicate weathering rate, which potentially linked with changes in weathering of highly-weatherable volcanic rocks with lighter silicon isotope (Ikeda *et al.*, 2017). This idea is consistent with

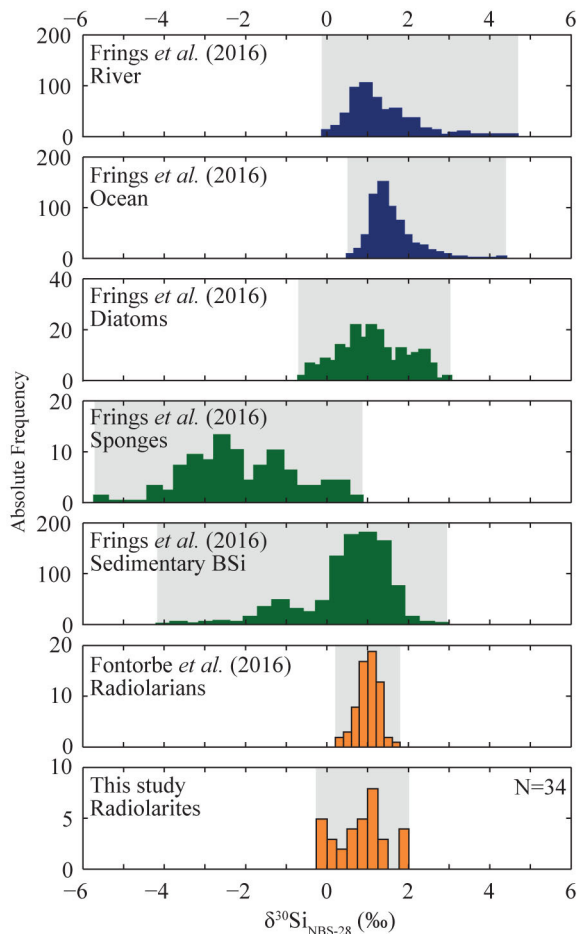


Fig. 5 $\delta^{30}\text{Si}_{\text{NBS-28}}$ -distribution in rivers, oceans, diatoms, sponges and sedimentary biogenic silica from Frings *et al.* (2016) compared with $\delta^{30}\text{Si}_{\text{NBS-28}}$ -distribution for Cenozoic radiolarians (Fontorbe *et al.*, 2016) and for the Triassic to Jurassic radiolarites from the Inuyama Area (this study). The relative similar range between Cenozoic radiolarians and Mesozoic radiolarites supports that the Mesozoic radiolarian molds preserve the original values.

the overall negative correlation between radiolarian $\delta^{30}\text{Si}$ and the BSi burial flux in the Inuyama area, despite of large scattering and low-resolution $\delta^{30}\text{Si}$ records (Fig. 4). Further high-resolution works are necessary to improve our understanding of $\delta^{30}\text{Si}$ cycle of radiolarian molds and unravelling some radiolarian crisis through geologic events, in response to bolide impact, massive volcanism, oceanic acidification, and oceanic anoxic events.

6. Conclusion and perspective

We measured $\delta^{30}\text{Si}$ of the Mesozoic radiolarian molds in Inuyama chert by SIMS. Range of $\delta^{30}\text{Si}$ between -0.3 and 2 ‰ is consistent with that of modern and Cenozoic radiolarian tests. Relatively large $\delta^{30}\text{Si}$ up to 0.8 ‰ (1SD) in intra-chert bed supports that $\delta^{30}\text{Si}$ of the Mesozoic radiolarian molds is not perfectly homogenized in a chert

bed during the diagenesis, and potentially record of kyr-scale changes in radiolarian $\delta^{30}\text{Si}$. 10-Myr scale trend of $\delta^{30}\text{Si}$ of the Mesozoic radiolarian molds from 250 Ma to 180 Ma is overall out-of-phase relation with BSi burial flux. This relation contradicts with interpretation of $\delta^{30}\text{Si}$ as a productivity proxy, despite of low-resolution and scattered $\delta^{30}\text{Si}$ records. Further high-resolution analysis will allow a better understanding of the past silica cycle, opening the possibility of accurate estimations of the past oceanic silica cycle and the contribution of past radiolarian productivity.

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Appendix

Table A1 Least square drift corrected and calibrated $\delta^{30}\text{Si}$ measurements (‰) for each sample. The LS-mean and standard deviation for each sample is indicated in bold font. All $\delta^{30}\text{Si}$ are relative to NBS-28 and all standard deviations are given as 2σ . Analytical standard deviation is sometime higher than 0.37 (2SD) due the use of error propagation during calibration.

Sample	Ki20	Ki22e2	Ki22e1	Ki21	Ki27	Ki24	Ki32	Ki34	Ki35	Ki38	Ki41	Ki39	
age	174.00	178.00	178.00	178.00	180.99	182.00	184.20	184.37	185.62	193.31	201.50	201.50	
LS-mean ($\delta^{30}\text{Si}_{\text{NBS-28}}$) and LS-SD (2 SD) (‰)	2.01 ± 1.47	0.82 ± 0.37	1.10 ± 0.29	1.18 ± 0.54	0.51 ± 1.10	1.87 ± 0.47	1.96 ± 0.79	1.04 ± 0.56	1.95 ± 0.42	0.75 ± 0.72	1.09 ± 0.20	1.40 ± 0.53	
Number of measurements	5	6	9	9	6	6	10	10	9	10	2	5	
Accepted measurements	5	5	8	9	6	6	10	10	9	10	1	4	
Rejected measurements	0	1	1	0	0	0	0	0	0	0	1	1	
$\delta^{30}\text{Si}_{\text{NBS-28}}$ and analytical standard deviation (2SD) of all accepted measurements (‰)	1	3.00 ± 0.31	0.79 ± 0.28	0.92 ± 0.28	1.17 ± 0.29	0.37 ± 0.25	2.02 ± 0.26	1.37 ± 0.20	0.58 ± 0.29	1.96 ± 0.26	1.39 ± 0.19	1.09 ± 0.20	1.68 ± 0.27
	2	2.13 ± 0.26	0.66 ± 0.27	0.93 ± 0.39	1.08 ± 0.21	1.16 ± 0.24	1.75 ± 0.20	1.92 ± 0.28	1.25 ± 0.21	2.22 ± 0.23	0.73 ± 0.33		1.31 ± 0.22
	3	2.52 ± 0.27	1.03 ± 0.33	1.08 ± 0.21	0.67 ± 0.28	0.96 ± 0.21	1.87 ± 0.27	1.74 ± 0.16	1.30 ± 0.25	2.06 ± 0.30	0.84 ± 0.34		1.62 ± 0.15
	4	1.16 ± 0.17		1.11 ± 0.28	1.52 ± 0.29	0.19 ± 0.19	2.16 ± 0.21	2.02 ± 0.24	0.85 ± 0.28	1.53 ± 0.19	1.14 ± 0.30		1.03 ± 0.17
	5	1.56 ± 0.19		1.18 ± 0.34	1.15 ± 0.24	-0.39 ± 0.33	1.92 ± 0.18	1.82 ± 0.19	1.18 ± 0.22	2.06 ± 0.24	0.64 ± 0.31		1.41 ± 0.24
	6			0.97 ± 0.34	1.06 ± 0.27	0.56 ± 0.48	1.49 ± 0.22	2.19 ± 0.33	0.99 ± 0.28	2.08 ± 0.26	0.64 ± 0.17		
	7			1.31 ± 0.25	1.44 ± 0.29			2.25 ± 0.24	1.19 ± 0.23	1.93 ± 0.30	0.74 ± 0.25		
	8			1.24 ± 0.24	1.03 ± 0.20			2.84 ± 0.24	1.36 ± 0.22	1.81 ± 0.29	0.49 ± 0.22		
	9				1.53 ± 0.26			1.73 ± 0.22	0.60 ± 0.25	2.03 ± 0.30	0.05 ± 0.37		
	10							1.86 ± 0.22	0.99 ± 0.29		0.63 ± 0.36		
	11												
	12												
	13												
	14												
Session	Br6	Br2	Br2	Br2	Br1	Br3	Br7	Br7	Br2	Br5	Br4v2	Br4	
Sample order during the session	5	3	2	4	3	5	7	9	9	2	3	5	

Sample	Ki40	Ki42	Ki43	Ki44	Ki46	Ki54	Ki48	Ki51	Ki15	Ki57	Ki58	Ki08	
age	201.50	201.50	204.82	204.82	210.27	214.4	217.09	219.00	228.00	241.00	243.48	244.90	
LS-mean ($\delta^{30}\text{Si}_{\text{NBS-28}}$) and LS-SD (2 SD) (‰)	0.80 ± 1.06	1.40 ± 0.52	1.21 ± 0.52	1.43 ± 0.22	-0.27 ± 1.10	0.03 ± 0.97	1.02 ± 0.78	1.09 ± 0.69	1.10 ± 0.52	0.97 ± 0.78	0.61 ± 0.93	-0.20 ± 0.64	
Number of measurements	5	10	6	9	10	6	5	6	10	10	6	10	
Accepted measurements	5	10	6	9	10	6	4	6	10	9	6	10	
Rejected measurements	0	0	0	0	0	0	1	0	0	1	0	0	
$\delta^{30}\text{Si}_{\text{NBS-28}}$ and analytical standard deviation (2SD) of all accepted measurements (‰)	1	-0.11 ± 0.18	1.47 ± 0.28	0.82 ± 0.24	1.44 ± 0.24	-0.57 ± 0.21	-0.12 ± 0.23	1.54 ± 0.22	0.51 ± 0.20	1.31 ± 0.35	0.65 ± 0.19	0.85 ± 0.23	0.27 ± 0.23
	2	0.88 ± 0.21	1.59 ± 0.29	1.03 ± 0.17	1.41 ± 0.31	0.83 ± 0.25	0.15 ± 0.33	0.63 ± 0.15	1.20 ± 0.19	1.31 ± 0.31	1.46 ± 0.26	0.91 ± 0.28	0.15 ± 0.20
	3	1.21 ± 0.20	1.38 ± 0.24	1.17 ± 0.26	1.30 ± 0.29	-1.06 ± 0.22	-0.67 ± 0.24	1.08 ± 0.32	1.36 ± 0.29	0.98 ± 0.29	1.23 ± 0.24	0.92 ± 0.24	-0.03 ± 0.23
	4	1.09 ± 0.17	1.37 ± 0.19	1.51 ± 0.23	1.26 ± 0.25	-0.62 ± 0.27	0.75 ± 0.21	0.91 ± 0.16	0.93 ± 0.25	1.14 ± 0.28	0.66 ± 0.25	0.90 ± 0.23	0.00 ± 0.20
	5	0.97 ± 0.18	0.76 ± 0.20	1.32 ± 0.27	1.43 ± 0.19	-0.01 ± 0.27	0.25 ± 0.27		1.45 ± 0.18	0.88 ± 0.24	1.54 ± 0.21	0.15 ± 0.26	-0.69 ± 0.18
	6		1.38 ± 0.25	1.40 ± 0.20	1.45 ± 0.22	-0.24 ± 0.19	-0.16 ± 0.17		1.09 ± 0.24	1.04 ± 0.24	1.04 ± 0.23	-0.15 ± 0.27	-0.17 ± 0.27
	7		1.56 ± 0.25		1.63 ± 0.18	-0.02 ± 0.20				1.55 ± 0.20	0.45 ± 0.27		-0.34 ± 0.28
	8		1.56 ± 0.35		1.41 ± 0.26	0.05 ± 0.25				0.99 ± 0.18	0.67 ± 0.23		-0.21 ± 0.32
	9		1.67 ± 0.21		1.52 ± 0.21	-0.01 ± 0.26				1.25 ± 0.22	0.98 ± 0.29		-0.29 ± 0.26
	10		1.36 ± 0.23			-0.90 ± 0.20				0.67 ± 0.22			-0.64 ± 0.20
	11												
	12												
	13												
	14												
Session	Br4	Br7	Br6	Br2	Br7	Br1	Br6	Br3	Br2	Br7	Br3	Br7	
Sample order during the session	1	3	4	8	5	5	2	4	5	1	7	4	

Sample	Ki07	Ki06	Ki06s	Ki05	Ki04	Ki3s	Ki03	Ki02	Ki01	Ki10	Ki09	Total
age	245.00	245.25	245.25	246.20	246.60	247.20	247.20	247.80	248.00	250.2	250.30	
LS-mean ($\delta^{30}\text{Si}_{\text{NBS-28}}$) and LS-SD (2 SD) (‰)	0.88 ± 0.42	0.03 ± 0.53	-0.15 ± 0.78	-0.26 ± 0.40	0.65 ± 0.67		0.21 ± 1.11	0.41 ± 0.85	0.33 ± 0.94	-0.24 ± 0.57	0.68 ± 0.65	
Number of measurements	6	6	14	6	9	8	10	7	10	6	9	266
Accepted measurements	6	6	14	6	9	0	10	7	10	6	9	252
Rejected measurements	0	0	0	0	0	8	0	0	0	0	0	14
$\delta^{30}\text{Si}_{\text{NBS-28}}$ and analytical standard deviation (2SD) of all accepted measurements (‰)	1	1.11 ± 0.24	0.43 ± 0.19	-0.03 ± 0.39	-0.09 ± 0.49	0.75 ± 0.29		0.31 ± 0.26	0.86 ± 0.23	0.63 ± 0.27	-0.38 ± 0.26	0.33 ± 0.26
	2	0.96 ± 0.31	0.16 ± 0.28	0.02 ± 0.32	-0.37 ± 0.30	0.83 ± 0.19		0.17 ± 0.28	-0.16 ± 0.30	0.79 ± 0.13	-0.42 ± 0.25	0.80 ± 0.29
	3	0.79 ± 0.16	0.07 ± 0.20	0.16 ± 0.22	-0.46 ± 0.25	1.14 ± 0.23		0.73 ± 0.27	0.77 ± 0.22	0.71 ± 0.19	-0.02 ± 0.25	0.97 ± 0.37
	4	1.14 ± 0.29	-0.10 ± 0.32	-0.19 ± 0.20	0.01 ± 0.25	0.59 ± 0.23		0.84 ± 0.33	0.52 ± 0.29	0.44 ± 0.21	-0.52 ± 0.28	0.26 ± 0.31
	5	0.70 ± 0.25	-0.23 ± 0.18	0.11 ± 0.25	-0.39 ± 0.23	0.26 ± 0.24		0.08 ± 0.17	0.72 ± 0.25	-0.27 ± 0.30	0.34 ± 0.42	1.03 ± 0.26
	6	0.64 ± 0.26	-0.23 ± 0.36	0.24 ± 0.24	-0.14 ± 0.33	0.19 ± 0.16		-0.23 ± 0.18	0.23 ± 0.19	0.77 ± 0.21	-0.20 ± 0.30	0.35 ± 0.26
	7			-0.24 ± 0.40		0.39 ± 0.28		0.73 ± 0.27	-0.16 ± 0.23	0.51 ± 0.19		0.72 ± 0.18
	8			0.22 ± 0.47		0.83 ± 0.25		0.46 ± 0.34		-0.48 ± 0.16		0.52 ± 0.28
	9			-0.23 ± 0.43		1.02 ± 0.34		0.43 ± 0.18		0.12 ± 0.18		1.15 ± 0.27
	10			0.14 ± 0.37				-1.00 ± 0.21		-0.14 ± 0.31		
	11			-0.65 ± 0.20								
	12			-1.06 ± 0.20								
	13			-0.09 ± 0.23								
	14			-0.03 ± 0.29								
Session	Br3	Br3	Br5	Br1	Br2	Br2	Br7	Br3	Br7	Br1	Br2	
Sample order during the session	3	6	8	2	7	1	2	8	6	1	6	

Table A2 Continued.

SIMS ANALYSES										Mount: BR2			Analyser: $\delta^{34}\text{S}$			Date: 03.09.2015		
Beam	$\text{H}^{21}\text{T}2$ CFS	$\text{H}^{21}\text{T}2$ CFS	$\text{L}2$ CFS	$\text{L}2$ CFS	$\text{H}^{21}\text{C}8\text{H}^{21}\text{C}6\text{O}4$ STD	$\text{H}^{21}\text{C}8\text{H}^{21}\text{C}6\text{O}4$ STD	$\text{H}^{21}\text{C}8\text{H}^{21}\text{C}6\text{O}4$ STD	Time	Yield CFS/nA	mean STD	Measurements $\delta^{34}\text{S}$ STD	mean STD	Diff. correction $\delta^{34}\text{S}$ STD	Calibration $\delta^{34}\text{S}$ STD	Value $\delta^{34}\text{S}$ STD	Comment		
430S1_030915 BR2_Panne617	2.28	3.2698E-02	7.0136E+07	1.1794E-01	2.2941E+06	1.0669E+01	2.2925E+06	2025	3.0830E+07	3.0799E+07	-42.67	-42.59	-42.59	-42.59	-0.07	Standard		
430S1_030915 BR2_Panne618	2.28	3.2688E-02	7.0037E+07	1.1699E-01	2.2925E+06	1.0701E+01	2.2925E+06	2028	3.0830E+07	1.5919E+05	-42.71	0.36	-42.71	0.36	0.23	Standard		
430S1_030915 BR2_Panne619	2.28	3.2698E-02	7.0373E+07	1.1615E-01	2.2967E+06	1.0721E+01	2.2967E+06	2032	3.0857E+07	3.0857E+07	-42.41	0.23	-42.41	0.23	0.25	Standard		
430S1_030915 BR2_Panne620	2.28	3.2693E-02	7.1574E+07	1.0877E-01	2.2866E+06	9.9959E+02	2.2866E+06	2035	3.0703E+07	3.0703E+07	-42.57	0.31	-42.57	-0.04	0.25	Standard		
430S1_030915 K15_rad@0	2.28	3.2716E-02	7.0976E+07	1.0497E-01	2.2759E+06	9.4929E+02	2.2759E+06	2038	3.0575E+07	3.0575E+07	-41.90	0.22	-41.90	0.22	0.65	K15		
430S1_030915 K15_rad@1	2.28	3.2705E-02	7.1304E+07	1.0225E-01	2.3231E+06	1.0837E+01	2.3231E+06	2042	3.1157E+07	3.0274E+07	-42.23	0.26	-41.89	-42.23	0.34	K15		
430S1_030915 K15_rad@2	2.28	3.2720E-02	7.1457E+07	1.0382E-01	2.2049E+06	7.7665E+02	2.2049E+06	2045	2.8696E+07	1.6946E+06	-41.78	0.29	-41.78	0.29	0.78	K15		
430S1_030915 K15_rad@3	2.28	3.2725E-02	7.1869E+07	1.1842E-01	2.1134E+06	1.1842E+01	2.1134E+06	2048	2.8592E+07	2.8592E+07	-41.62	0.37	-41.62	0.37	0.95	K15		
430S1_030915 K15_rad@4	2.28	3.2702E-02	7.0892E+07	1.2066E-01	2.3185E+06	1.2108E+01	2.3185E+06	2051	3.1112E+07	3.1112E+07	-42.29	0.31	-42.29	0.25	0.25	K15		
430S1_030915 K15_rad@5	2.28	3.2727E-02	7.1531E+07	1.1528E-01	2.2158E+06	1.1386E+01	2.2158E+06	2055	2.9729E+07	2.9729E+07	-41.56	0.26	-41.56	0.26	1.01	K15		
430S1_030915 K15_rad@6	2.28	3.2705E-02	7.0346E+07	1.1938E-01	2.3019E+06	1.1386E+01	2.3019E+06	2058	3.0867E+07	3.0867E+07	-42.21	0.26	-42.20	0.26	0.34	K15		
430S1_030915 K15_rad@7	2.28	3.2717E-02	7.1153E+07	1.1938E-01	2.2649E+06	1.0710E+01	2.2649E+06	2101	3.0376E+07	3.0376E+07	-41.85	0.18	-41.85	0.18	0.70	K15		
430S1_030915 K15_rad@8	2.28	3.2711E-02	7.1406E+07	1.1807E-01	2.2993E+06	1.1191E+01	2.2993E+06	2108	3.0846E+07	3.0846E+07	-42.04	0.28	-42.04	0.28	0.51	K15		
430S1_030915 K15_rad@9	2.28	3.2732E-02	7.1587E+07	1.1557E-01	2.2530E+06	1.1332E+01	2.2530E+06	2108	3.0184E+07	3.0184E+07	-41.44	0.27	-41.44	0.27	1.14	K15		
430S1_030915 BR2_Panne621	2.28	3.2681E-02	7.1908E+07	1.0591E-01	2.3502E+06	9.7210E+02	2.3502E+06	2111	3.1578E+07	3.1513E+07	-42.92	0.22	-42.79	-42.92	0.13	Standard		
430S1_030915 BR2_Panne622	2.28	3.2691E-02	7.1814E+07	1.0667E-01	2.3477E+06	9.1274E+02	2.3477E+06	2114	3.1505E+07	1.0345E+05	-42.63	0.28	0.24	-42.79	0.16	Standard		
430S1_030915 BR2_Panne623	2.28	3.2686E-02	7.1704E+07	1.0805E-01	2.3440E+06	9.7910E+02	2.3440E+06	2118	3.1452E+07	3.1452E+07	-42.78	0.23	-42.78	0.23	0.22	Standard		
430S1_030915 BR2_Panne624	2.28	3.2684E-02	7.1842E+07	1.0674E-01	2.3484E+06	9.4787E+02	2.3484E+06	2121	3.1571E+07	3.1571E+07	-42.83	0.28	-42.83	0.28	0.24	Standard		
430S1_030915 K44_rad@1	2.28	3.2718E-02	7.1400E+07	1.0978E-01	2.1963E+06	9.6489E+02	2.1963E+06	2124	2.9444E+07	3.0239E+07	-41.83	0.29	-41.82	0.29	0.74	K15		
430S1_030915 K44_rad@2	2.28	3.2721E-02	7.1928E+07	1.1916E-01	2.2605E+06	1.2609E+01	2.2605E+06	2128	3.0302E+07	1.1552E+06	-41.75	0.19	-41.75	0.19	0.81	K15		
430S1_030915 K44_rad@3	2.28	3.2731E-02	7.1254E+07	1.0528E-01	2.2172E+06	8.8027E+02	2.2172E+06	2131	2.9706E+07	2.9706E+07	-41.45	0.23	-41.45	0.23	1.13	K15		
430S1_030915 K44_rad@4	2.28	3.2713E-02	7.1513E+07	1.1598E-01	2.2495E+06	1.1341E+01	2.2495E+06	2134	3.0172E+07	3.0172E+07	-41.98	0.23	-41.97	0.23	0.58	K15		
430S1_030915 K44_rad@5	2.28	3.2702E-02	7.1848E+07	1.1848E-01	2.2870E+06	1.6041E+01	2.2870E+06	2138	3.0691E+07	3.0691E+07	-42.30	0.24	-42.29	0.24	0.23	K15		
430S1_030915 K44_rad@6	2.28	3.2700E-02	7.0270E+07	1.1848E-01	2.2870E+06	1.6041E+01	2.2870E+06	2141	3.0818E+07	3.0818E+07	-42.36	0.16	-42.35	0.16	0.18	K15		
430S1_030915 K44_rad@7	2.28	3.2706E-02	7.0889E+07	1.1848E-01	2.3180E+06	8.9241E+02	2.3180E+06	2144	3.0818E+07	3.0818E+07	-42.17	0.28	-42.17	0.28	0.38	K15		
430S1_030915 K44_rad@8	2.28	3.2721E-02	7.1376E+07	1.1848E-01	2.2859E+06	9.5888E+02	2.2859E+06	2144	3.0652E+07	3.0652E+07	-42.17	0.28	-42.17	0.28	0.38	K15		
430S1_030915 K44_rad@9	2.28	3.2721E-02	7.1281E+07	1.1848E-01	2.2829E+06	8.6294E+02	2.2829E+06	2147	3.0591E+07	3.0591E+07	-41.74	0.25	-41.74	0.25	0.82	K15		
430S1_030915 K44_rad@0	2.28	3.2727E-02	7.1676E+07	1.0680E-01	2.2040E+06	9.4344E+02	2.2040E+06	2151	2.9507E+07	2.9507E+07	-41.57	0.34	-41.56	0.34	1.01	K15		
430S1_030915 K44_rad@1	2.28	3.2741E-02	7.1908E+07	1.0680E-01	2.2544E+06	8.7651E+02	2.2544E+06	2154	3.0182E+07	3.0182E+07	-41.15	0.24	-41.15	0.24	0.26	K15		
430S1_030915 BR2_Panne625	2.28	3.2683E-02	7.1420E+07	1.0691E-01	2.3450E+06	8.8623E+02	2.3450E+06	2157	3.1453E+07	3.1416E+07	-42.87	0.29	-42.69	-42.68	-0.35	Standard		
430S1_030915 BR2_Panne626	2.28	3.2687E-02	7.1687E+07	1.0528E-01	2.3453E+06	1.0246E+01	2.3453E+06	2201	3.1438E+07	3.1438E+07	-42.73	0.18	-42.73	0.18	0.32	Standard		
430S1_030915 BR2_Panne627	2.28	3.2688E-02	7.1556E+07	1.0851E-01	2.3397E+06	1.0437E+01	2.3397E+06	2204	3.1370E+07	3.1370E+07	-42.43	0.29	-42.42	0.29	0.11	Standard		
430S1_030915 BR2_Panne628	2.28	3.2688E-02	7.1697E+07	1.0921E-01	2.3435E+06	8.8593E+02	2.3435E+06	2207	3.1401E+07	3.1401E+07	-42.72	0.14	-42.71	0.14	0.21	Standard		
430S1_030915 K44_rad@2	2.28	3.2740E-02	7.1251E+07	1.1251E-01	2.2910E+06	1.1307E+01	2.2910E+06	2210	3.0655E+07	3.0551E+07	-41.19	0.31	-41.18	-41.19	0.31	K15		
430S1_030915 K44_rad@3	2.28	3.2736E-02	7.1225E+07	1.1251E-01	2.3320E+06	7.928E+02	2.3320E+06	2214	3.1193E+07	3.0769E+05	-41.30	0.29	-41.29	0.29	0.25	K15		
430S1_030915 K44_rad@4	2.28	3.2735E-02	7.0045E+07	1.2242E-01	2.2931E+06	1.2302E+01	2.2931E+06	2217	3.0685E+07	3.0685E+07	-41.34	0.25	-41.33	0.25	1.25	K15		
430S1_030915 K44_rad@5	2.28	3.2740E-02	7.0955E+07	1.2491E-01	2.2906E+06	1.2556E+01	2.2906E+06	2220	3.0641E+07	3.0641E+07	-41.18	0.19	-41.17	0.19	1.42	K15		
430S1_030915 K44_rad@6	2.28	3.2741E-02	7.0689E+07	1.2902E-01	2.2505E+06	1.1969E+01	2.2505E+06	2224	3.0079E+07	3.0079E+07	-41.15	0.22	-41.14	0.22	1.44	K15		
430S1_030915 K44_rad@7	2.28	3.2747E-02	7.0431E+07	1.0452E-01	2.3063E+06	1.0159E+01	2.3063E+06	2227	3.0847E+07	3.0847E+07	-40.98	0.18	-40.97	0.18	1.62	K15		
430S1_030915 K44_rad@8	2.28	3.2740E-02	7.0208E+07	1.0631E-01	2.2976E+06	8.6094E+02	2.2976E+06	2230	3.0739E+07	3.0739E+07	-41.20	0.26	-41.19	0.26	0.24	K15		
430S1_030915 K44_rad@9	2.29	3.2743E-02	7.1388E+07	1.1388E-01	2.2401E+06	1.3814E+01	2.2401E+06	2233	2.9940E+07	2.9940E+07	-41.09	0.21	-41.08	0.21	1.51	K15		
430S1_030915 K15_rad@1	2.28	3.2758E-02	7.1289E+07	1.2380E-01	2.2638E+06	1.1696E+01	2.2638E+06	2237	3.0249E+07	3.0249E+07	-40.66	0.26	-40.66	0.26	0.24	K15		
430S1_030915 BR2_Panne629	2.28	3.2680E-02	7.1696E+07	1.0696E-01	2.3444E+06	9.1310E+02	2.3444E+06	2240	3.1391E+07	3.1377E+07	-42.94	0.25	-42.75	-42.75	-0.42	Standard		
430S1_030915 BR2_Panne630	2.28	3.2687E-02	7.1770E+07	1.0328E-01	2.3458E+06	1.0042E+01	2.3458E+06	2244	3.1429E+07	3.1429E+07	-42.73	0.27	-42.72	0.27	0.23	Standard		
430S1_030915 BR2_Panne631	2.28	3.2687E-02	7.1711E+07	1.0695E-01	2.3443E+06	9.6941E+02	2.3443E+06	2247	3.1408E+07	3.1408E+07	-42.75	0.31	-42.74	0.31	0.22	Standard		
430S1_030915 BR2_Panne632	2.28	3.2691E-02	7.1428E+07	1.0129E-01	2.3351E+06	9.5676E+02	2.3351E+06	2250	3.1282E+07	3.1282E+07	-42.63	0.20	-42.62	0.20	0.22	Standard		
430S1_030915 K15_rad@2	2.28	3.2766E-02	7.1402E+07	1.0508E-01	2.2527E+06	1.0399E+01	2.2527E+06	2253	3.0152E+07	2.9876E+07	-40.42	0.23	-40.41	0.23	0.23	K15		
430S1_030915 K15_rad@3	2.28	3.2761E-02	7.1499E+07	1.0747E-01	2.2596E+06	1.6210E+01	2.2596E+06	2257	3.0233E+07	1.0492E+06	-40.58	0.30	-40.56	0.30	2.05	K15		
430S1_030915 K15_rad@4	2.28	3.2744E-02	7.1516E+07	1.2516E-01	2.2429E+06	1.2349E+01	2.2429E+06	2300	3.0049E+07	3.0049E+07	-41.08	0.19	-41.07	0.19	1.52	K15		
430S1_030915 K15_rad@5	2.28	3.2761E-02	7.1268E+07															

Table A2 Continued.

SIMS ANALYSES										Mount: BR4-2			Standard: UNIL-Q1 (Paine)			Analyse: $\delta^{30}\text{Si}$			Value: -0.13 ± 0.02 (NBS-28, 2e)			Date: 03.09.2015	
Beam	nA	$\text{H}^2\text{L}12$ (30S/28Si)		$\text{L}12$ (30S/1Coef)		H^2 (28Si/Coef)		H^2 (28Si/Coef)		Time	Yield		Drift correction		Calibration		Comment						
		CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD		mean	2SD	$\delta^{30}\text{Si}$	2SD	$\delta^{30}\text{Si}_{\text{NBS-28}}$	2SD							
40S1_040915 BR4_Paineg01	2.33	3.2677E-02	1.1111E-02	8.4303E-07	1.4409E-01	2.7567E+06	1.3101E-01	3.6210E+07	3.5641E+07	0	3.6210E+07	3.5641E+07	-43.04	0.22	-43.01		Setting Standard						
40S1_040915 BR4_Paineg02	2.34	3.2691E-02	7.3868E-03	8.1704E-07	1.1866E-01	2.6711E+06	1.1937E-01	3.4901E+07	2.6702E+06	1	3.4901E+07	2.6702E+06	-42.63	0.15	0.78		Setting Standard						
40S1_040915 BR4_Paineg03	2.34	3.2678E-02	9.2309E-03	8.2091E-07	1.5683E-01	2.6826E+06	1.5333E-01	3.3121E+07	2.6702E+06	2	3.3121E+07	2.6702E+06	-43.04	0.18			Setting Standard						
40S1_040915 BR4_Paineg04	2.34	3.2693E-02	1.1562E-02	8.1632E-07	1.1800E-01	2.6909E+06	1.1802E-01	3.4904E+07	2.6702E+06	3	3.4904E+07	2.6702E+06	-42.56	0.23			Setting Standard						
40S1_040915 BR4_Paineg05	2.34	3.2676E-02	1.4900E-02	7.4776E-07	1.6906E-01	2.4436E+06	1.6931E-01	3.1937E+07	2.6702E+06	4	3.1937E+07	2.6702E+06	-43.07	0.30			Setting Standard						
40S1_040915 BR4_Paineg06	2.34	3.2659E-02	1.2601E-02	8.7057E-07	1.1632E-01	2.8433E+06	1.1725E-01	3.7203E+07	2.6702E+06	5	3.7203E+07	2.6702E+06	-43.56	0.25			Setting Standard						
40S1_040915 BR4_Paineg07	2.34	3.2679E-02	7.0994E-03	8.3225E-07	1.5934E-01	2.7219E+06	1.4642E-01	3.5099E+07	2.6702E+06	6	3.5099E+07	2.6702E+06	-42.99	0.14			Setting Standard						
40S1_040915 BR4_Paineg08	2.34	3.2680E-02	7.7632E-03	8.0416E-07	1.5914E-01	2.6282E+06	1.2099E-01	3.4298E+07	2.6702E+06	7	3.4298E+07	2.6702E+06	-42.95	0.16			Setting Standard						
40S1_040915 BR4_Paineg09	2.34	3.2679E-02	1.2674E-02	8.3395E-07	1.2314E-01	2.7236E+06	1.3337E-01	3.5578E+07	2.6702E+06	8	3.5578E+07	2.6702E+06	-42.97	0.25			Setting Standard						
40S1_040915 BR4_Paineg10	2.34	3.2644E-02	1.3700E-02	8.6099E-07	1.9825E-01	2.8106E+06	1.9444E-01	3.6740E+07	2.6702E+06	9	3.6740E+07	2.6702E+06	-43.99	0.27			Setting Standard						
40S1_040915 BR4_Paineg11	2.35	3.2679E-02	7.9642E-03	8.5974E-07	1.6549E-01	2.8095E+06	1.6595E-01	3.6059E+07	2.6702E+06	10	3.6059E+07	2.6702E+06	-42.96	0.16			Setting Standard						
40S1_040915 BR4_Paineg12	2.34	3.2657E-02	1.4912E-02	8.9971E-07	1.9070E-01	2.9089E+06	1.7753E-01	3.8079E+07	2.6702E+06	11	3.8079E+07	2.6702E+06	-43.62	0.30			Setting Standard						
40S1_040915 BR4_Paineg13	2.35	3.2684E-02	9.0542E-03	8.2865E-07	1.2373E-01	2.7085E+06	1.2837E-01	3.5333E+07	2.6702E+06	12	3.5333E+07	2.6702E+06	-42.83	0.18			Setting Standard						
40S1_040915 BR4_Paineg14	2.35	3.2686E-02	8.3286E-03	8.4266E-07	9.1881E-02	2.7544E+06	9.0038E-02	3.5931E+07	2.6702E+06	13	3.5931E+07	2.6702E+06	-42.78	0.17			Setting Standard						
40S1_040915 BR4_Paineg15	2.35	3.2694E-02	9.2168E-03	8.3456E-07	1.3326E-01	2.7285E+06	1.1490E-01	3.5899E+07	2.6702E+06	14	3.5899E+07	2.6702E+06	-42.53	0.18			Setting Standard						
40S1_040915 BR4_Paineg16	2.35	3.2677E-02	1.5172E-02	8.5898E-07	1.3879E-01	2.8047E+06	1.5102E-01	3.6618E+07	2.6702E+06	15	3.6618E+07	2.6702E+06	-43.03	0.30			Setting Standard						
40S1_040915 BR4_Paineg17	2.35	3.2689E-02	9.5313E-03	8.2581E-07	1.1661E-01	2.6993E+06	1.1507E-01	3.5216E+07	2.6702E+06	16	3.5216E+07	2.6702E+06	-42.69	0.19			Setting Standard						
40S1_040915 BR4_Paineg18	2.34	3.2679E-02	9.5364E-03	8.5453E-07	1.4588E-01	2.7923E+06	1.4702E-01	3.6583E+07	3.6642E+07	17	3.6583E+07	3.6642E+07	-42.98	0.19	-43.12	0.07	0.40	-0.08	Standard				
40S1_040915 BR4_Paineg19	2.35	3.2681E-02	1.0937E-02	8.6477E-07	1.3502E-01	2.8244E+06	1.5749E-01	3.6868E+07	1.3429E+06	18	3.6868E+07	1.3429E+06	-43.52	0.22	0.71	-0.49	0.40	0.74	Standard				
40S1_040915 BR4_Paineg20	2.35	3.2665E-02	8.3444E-03	8.7634E-07	1.7846E-01	2.8648E+06	1.6951E-01	3.7599E+07	2.6702E+06	19	3.7599E+07	2.6702E+06	-43.38	0.17	-43.33	0.17	-0.35	0.40	Standard				
40S1_040915 BR4_Paineg21	2.35	3.2692E-02	1.0079E-02	8.3962E-07	8.1061E-02	2.7449E+06	7.9766E-02	3.5757E+07	2.6702E+06	20	3.5757E+07	2.6702E+06	-42.61	0.20	-42.56	0.20	0.45	0.40	Standard				
40S1_040915 BR4_Paineg22	2.35	3.2678E-02	9.5927E-03	8.4321E-07	1.3506E-01	2.7554E+06	1.5549E-01	3.5888E+07	3.6211E+07	31	3.5888E+07	3.6211E+07	-43.02	0.19	-43.08	-0.01	0.40	-0.09	Standard				
40S1_040915 BR4_Paineg23	2.35	3.2684E-02	8.4275E-03	8.4676E-07	1.1396E-01	2.7693E+06	1.0541E-01	3.6004E+07	1.8631E+06	32	3.6004E+07	1.8631E+06	-42.84	0.17	0.76	0.17	0.40	0.79	Standard				
40S1_040915 BR4_Paineg24	2.35	3.2687E-02	1.0581E-02	8.5300E-07	1.6291E-01	2.7800E+06	1.6437E-01	3.6262E+07	2.6702E+06	33	3.6262E+07	2.6702E+06	-42.74	0.21	0.80	0.28	0.40	0.40	Standard				
40S1_040915 BR4_Paineg25	2.35	3.2659E-02	1.1782E-02	8.9236E-07	1.2595E-01	2.9145E+06	1.2747E-01	3.7931E+07	2.6702E+06	34	3.7931E+07	2.6702E+06	-43.55	0.24	-43.55	0.24	-0.58	0.41	Standard				
40S1_040915 BR4_Paineg26	2.35	3.2651E-02	7.1539E-03	8.8410E-07	1.8994E-01	2.8868E+06	1.9097E-01	3.7664E+07	2.6702E+06	35	3.7664E+07	2.6702E+06	-43.81	0.14	-43.81	0.14	-0.85	0.39	Standard				
40S1_040915 BR4_Paineg27	2.35	3.2683E-02	1.7097E-02	8.3575E-07	1.3612E-01	2.7317E+06	1.3510E-01	3.5514E+07	2.6702E+06	36	3.5514E+07	2.6702E+06	-42.86	0.34	-42.86	0.34	0.15	0.42	Standard				
40S1_040915 BR4_Paineg28	2.35	3.2659E-02	1.1533E-02	8.6054E-07	1.5429E-01	2.8105E+06	1.5835E-01	3.6589E+07	2.6702E+06	37	3.6589E+07	2.6702E+06	-43.55	0.23	-43.55	0.23	-0.59	0.40	Standard				
40S1_040915 BR4_Paineg29	2.35	3.2686E-02	1.3047E-02	8.3492E-07	1.5429E-01	2.7291E+06	1.0150E-01	3.5498E+07	2.6702E+06	38	3.5498E+07	2.6702E+06	-42.76	0.26	-42.77	0.26	0.24	0.41	Standard				
40S1_040915 BR4_Paineg30	2.35	3.2682E-02	1.1241E-02	8.2751E-07	1.2243E-01	2.7060E+06	1.1355E-01	3.5183E+07	2.6702E+06	39	3.5183E+07	2.6702E+06	-42.89	0.22	-42.90	0.22	0.10	0.40	Standard				
40S1_040915 BR4_Paineg31	2.35	3.2686E-02	1.0208E-02	8.3726E-07	1.3262E-01	2.7367E+06	1.1319E-01	3.5841E+07	2.6702E+06	40	3.5841E+07	2.6702E+06	-42.78	0.20	-42.79	0.20	0.22	0.40	Standard				
40S1_040915 K41_rnd01	2.36	3.2720E-02	1.0150E-02	7.7345E-07	1.3637E-01	2.5306E+06	1.3666E-01	3.3841E+07	3.3700E+07	49	3.3841E+07	3.3700E+07	-41.76	0.20	-42.01	-1.80	1.25	0.40	K41				
40S1_040915 K41_rnd02	2.35	3.2729E-02	1.8727E-02	7.7343E-07	6.1451E-02	2.5317E+06	6.4292E-02	3.3853E+07	3.0611E+06	50	3.3853E+07	3.0611E+06	-41.51	0.37	1.32	1.64	1.72	0.40	K41				
40S1_040915 BR4_Paineg32	2.35	3.2681E-02	9.9304E-03	8.2620E-07	1.4261E-01	2.7002E+06	1.4438E-01	3.4993E+07	3.5677E+07	51	3.4993E+07	3.5677E+07	-42.90	0.20	-43.04	-0.20	0.05	0.40	Standard				
40S1_040915 BR4_Paineg33	2.36	3.2673E-02	1.0927E-02	8.4631E-07	1.0688E-01	2.7661E+06	1.0690E-01	3.4935E+07	7.4899E+05	52	3.4935E+07	7.4899E+05	-43.16	0.22	0.19	-0.21	0.40	0.20	Standard				
40S1_040915 BR4_Paineg34	2.36	3.2678E-02	1.1001E-02	8.3915E-07	1.2060E-01	2.7420E+06	1.2043E-01	3.5810E+07	2.6702E+06	53	3.5810E+07	2.6702E+06	-43.01	0.22	-43.06	-0.07	0.40	0.40	Standard				
40S1_040915 BR4_Paineg35	2.36	3.2675E-02	1.1106E-02	8.4421E-07	9.5202E-02	2.7584E+06	9.2224E-02	3.5830E+07	2.6702E+06	54	3.5830E+07	2.6702E+06	-43.15	0.22	-43.15	-0.16	0.40	0.40	Standard				

Table A2 Continued.

SIMS ANALYSES										Mount: BR6				Standard: UNIL-QI (Paint)				Analyse $\delta^{34}\text{S}$				Value = -0.13 ± 0.02 (NBS-28, 2c)				Date: 03.09.2015	
Beam nA	H ²¹ /2 (308/281)		H ²² (288/287)		H ²³ (288/287)		Time		Yield		mean		mean		Drift correction		Calibration		Comment								
	CPS	2SD	CPS	2SD	CPS	2SD	2SD	CPS/nA	2SD	$\delta^{34}\text{S}$	2SD	$\delta^{34}\text{S}$	2SD	$\delta^{34}\text{S}$	2SD	$\delta^{34}\text{S}$	2SD										
2.85	3.2600E-02	1.0308E-02	9.8177E+07	8.3010E-02	3.2010E+06	7.4661E+02	1510	3.4493E+07	3.4162E+07	-45.29	0.21	-45.10	0.24						Setting Standard								
2.85	3.2608E-02	8.4777E-03	9.8585E+07	1.3462E-01	3.2148E+06	1.3181E+02	1513	3.4531E+07	1.5657E+06	44.97	0.17	0.24							Setting Standard								
2.86	3.2611E-02	1.1732E-02	9.6759E+07	7.9006E-02	3.1569E+06	9.1071E+02	1517	3.3885E+07		44.91	0.23								Setting Standard								
2.86	3.2603E-02	9.1721E-03	9.8461E+07	7.8345E-02	3.2103E+06	7.9922E+02	1520	3.4484E+07		45.21	0.19								Setting Standard								
2.86	3.2609E-02	1.1222E-02	9.9075E+07	1.1920E-01	3.2501E+06	1.1222E+02	1523	3.4875E+07		45.04	0.22								Setting Standard								
2.86	3.2608E-02	7.9083E-03	9.3462E+07	3.1953E-01	3.0474E+06	1.1779E+01	1526	3.2703E+07		45.06	0.16								Setting Standard								
2.85	3.2598E-02	9.7979E-03	9.3629E+07	1.8384E-01	3.0349E+06	1.6920E-01	1538	3.2819E+07		45.36	0.20	-45.23	-0.10	0.34	0.01				Standard								
2.86	3.2606E-02	1.0866E-02	9.6410E+07	1.3948E-01	3.1457E+06	1.2793E-01	1551	3.3690E+07		45.11	0.22	-45.09	0.22	-0.15	0.34	0.22			duplicate								
2.86	3.2600E-02	1.1102E-02	1.0067E+08	1.5093E-01	3.2838E+06	1.3535E-01	1554	3.3151E+07		45.29	0.22	-45.27	0.22	-0.03	0.34				Standard								
2.86	3.2600E-02	1.3121E-02	9.5810E+07	2.2866E-01	3.1253E+06	1.1242E-01	1558	3.3468E+07		45.29	0.26	-45.27	0.26	-0.04	0.35				Standard								
2.87	3.2592E-02	9.9742E-03	9.9673E+07	2.0366E-01	3.2516E+06	1.9003E-01	1621	3.3778E+07		45.52	0.20	-45.34	-0.20	-0.34	-0.10				Standard								
2.87	3.2602E-02	9.6423E-03	9.7008E+07	1.2312E-01	3.1624E+06	1.1213E-01	1624	3.3824E+07		45.24	0.19	0.29	0.02	0.34	0.30				Standard								
2.87	3.2595E-02	1.0272E-02	9.7065E+07	1.3189E-01	3.1640E+06	1.3189E-01	1627	3.3856E+07		45.45	0.21	-45.43	0.21	-0.20	0.34				Standard								
2.87	3.2604E-02	1.0983E-02	9.6249E+07	1.1832E-01	3.1381E+06	1.1709E-01	1630	3.3557E+07		45.17	0.13	-45.16	0.13	0.09	0.33				Standard								
2.87	3.2645E-02	1.4083E-02	9.1341E+07	7.4821E-02	2.9820E+06	7.3635E-02	1634	3.1866E+07		43.71	0.22	-44.19	-0.39	0.22	0.34				K648								
2.87	3.2624E-02	7.7468E-03	9.1638E+07	1.8806E-01	2.9898E+06	1.8353E-01	1637	3.1974E+07		44.58	0.15	0.73	0.44	0.70	0.33				K648								
2.87	3.2639E-02	1.6062E-02	9.1689E+07	1.4856E-01	2.9926E+06	1.4161E-01	1640	3.1989E+07		44.15	0.32	-44.13	0.32	1.15	0.36				K648								
2.87	3.2634E-02	7.9662E-03	8.3345E-02	5.4489E-02	2.7198E+06	5.3755E-02	1643	2.9085E+07		44.31	0.16	-44.29	0.16	0.99	0.33				K648								
2.87	3.2623E-02	4.9264E-02	4.9264E-02	-7.6538E-03	2.7198E+06	4.1698E-02	1647	1.7166E+07		-9558.22									Yield								
2.87	3.2597E-02	9.3871E-03	9.8644E+07	1.9022E-01	3.2258E+06	1.7664E-01	1650	3.4495E+07		45.38	0.19	-45.38	0.19	-45.37	-0.14	0.34	-0.13		Standard								
2.87	3.2597E-02	1.1024E-02	1.0222E+08	2.2235E-01	3.3633E+06	2.1936E-01	1653	3.3976E+07		45.91	0.22	0.77	-45.90	0.22	0.66	-0.29	0.34	0.70	Standard								
2.87	3.2609E-02	1.1687E-02	9.6092E+07	7.9924E-02	3.1335E+06	9.3768E-02	1657	3.3491E+07		45.02	0.23	-45.01	0.23	-0.64	0.34				Standard								
2.87	3.2603E-02	9.7144E-03	9.7703E+07	6.0574E-02	3.1853E+06	9.2870E-02	1700	3.4055E+07		45.20	0.20	-45.19	0.20	0.05	0.34				Standard								
2.87	3.2609E-02	1.1020E-02	9.6472E+07	1.1339E-01	3.1459E+06	1.1780E-01	1726	3.3601E+07		45.04	0.29	-45.40	0.29	-0.36	-0.17				Standard								
2.87	3.2587E-02	1.0200E-02	1.0016E+08	1.6617E-01	3.2639E+06	1.6556E-01	1729	3.4868E+07		45.66	0.22	0.56	-45.65	0.22	0.49	-0.43	0.34	0.51	Standard								
2.87	3.2589E-02	9.7018E-03	9.5567E+07	1.1007E-01	3.1163E+06	9.9056E-02	1733	3.3299E+07		45.60	0.19	-45.59	0.19	-0.37	0.34				Standard								
2.87	3.2598E-02	1.4101E-02	9.6797E+07	1.1054E-01	3.1554E+06	1.0929E-01	1736	3.3697E+07		45.34	0.28	-45.34	0.28	-0.11	0.35				Standard								
2.87	3.2631E-02	1.1986E-02	9.1157E+07	1.9092E-01	2.9743E+06	1.8759E-01	1739	3.1737E+07		44.40	0.24	-44.02	0.24	0.89	0.35				K643								
2.87	3.2632E-02	8.4439E-03	9.5318E+07	1.2148E-01	3.1111E+06	1.1630E-01	1742	3.1695E+07		44.19	0.17	0.49	-44.19	0.17	1.10	0.34			K643								
2.87	3.2642E-02	1.2908E-02	9.0341E+07	1.5164E-01	2.9489E+06	1.4815E-01	1746	3.1432E+07		44.06	0.26	-44.06	0.26	1.24	0.35				K643								
2.87	3.2653E-02	1.1504E-02	8.6235E+07	1.8871E-01	2.8158E+06	1.8121E-01	1749	3.0004E+07		43.73	0.23	-43.73	0.23	1.58	0.34				K643								
2.88	3.2647E-02	1.3404E-02	9.2920E-02	2.9897E+06	4.5444E-02	1754	3.1802E+07		43.91	0.27	-43.91	0.27	1.39	0.35				K643									
2.88	3.2650E-02	1.0058E-02	9.1817E+07	1.2289E-01	2.9997E+06	1.2973E-01	1757	3.1930E+07		43.84	0.20	-43.83	0.20	-0.43	0.34				K643								
2.87	3.2616E-02	1.1356E-02	9.0220E+07	1.6610E-01	2.9424E+06	1.6369E-01	1800	3.1410E+07		44.83	0.23	-45.41	0.23	-0.43	0.34	-0.18			Standard								
2.88	3.2564E-02	1.1604E-02	8.8765E+07	3.3713E-01	2.8908E+06	3.1633E-01	1804	3.0896E+07		46.33	0.23	1.35	-46.33	0.23	1.17	-1.15	0.34	1.22	Standard								
2.87	3.2610E-02	1.1234E-02	9.6270E+07	1.2863E-01	3.1195E+06	1.0698E-01	1807	3.3527E+07		45.00	0.22	-44.99	0.22	0.25	0.34				Standard								
2.87	3.2584E-02	1.6640E-02	9.3402E+07	3.2807E-01	2.8508E+06	3.1861E-01	1810	3.2607E+07		45.48	0.33	-45.48	0.33	-0.25	0.36				Standard								
2.87	3.2702E-02	1.3311E-02	7.2862E+07	1.1339E-01	2.3846E+06	1.3011E-01	1814	2.3570E+07		42.31	0.31	-42.31	0.31	3.07	0.36				K620								
2.87	3.2674E-02	1.3194E-02	7.5096E+07	9.7037E-02	2.4536E+06	9.7077E-02	1817	2.6136E+07		43.14	0.26	-43.14	0.26	2.20	0.35				K620								
2.87	3.2686E-02	1.3627E-02	6.8071E+07	1.0425E-01	2.2250E+06	1.0136E-01	1820	2.3702E+07		42.77	0.27	-42.77	0.27	2.59	0.35				K620								
2.88	3.2642E-02	8.6278E-03	8.4333E+07	1.7955E-01	2.7520E+06	1.8100E-01	1823	2.9319E+07		44.07	0.17	-44.07	0.17	1.62	0.34				K620								
2.87	3.2655E-02	9.5962E-03	7.9744E+07	1.5535E-01	2.6037E+06	1.5697E-01	1827	2.7745E+07		43.69	0.19	-43.69	0.19	1.62	0.34				K620								
2.87	3.2601E-02	1.0628E-02	1.0252E+08	1.3549E-01	3.2866E+06	1.3637E-01	1830	3.4895E+07		45.26	0.21	-45.54	0.21	-0.43	0.34	-0.32			Standard								
2.88	3.2590E-02	8.2813E-03	9.1912E+07	1.8276E-01	3.2328E+06	1.8266E-01	1833	3.4498E+07		45.59	0.17	0.80	-45.59	0.17	0.69	-0.37	0.33	0.72	Standard								
2.87	3.2573E-02	9.0740E-03	1.0625E+08	1.6989E-01	3.4634E+06	1.6183E-01	1836	3.6978E+07		46.08	0.18	-46.08	0.18	-0.89	0.34				Standard								
2.87	3.2603E-02	1.0416E-02	9.9135E+07	1.2055E-01	3.2339E+06	1.1218E-01	1840	3.4488E+07		45.21	0.21	-45.22	0.21	0.02	0.34				Standard								
2.87	3.2604E-02	1.0537E-02	9.9222E+07	1.3592E-01	3.2372E+06	1.2450E-01	1903	3.4537E+07		45.17	0.21	-45.38	0.21	-0.45	0.34	-0.16			Standard								
2.88	3.2605E-02	1.0568E-02	9.8211E+07	9.2243E-02	3.2024E+06	8.9789E-02	1906	3.4146E+07		45.16	0.21	0.55	-45.16	0.21	0.47	0.08	0.34	0.50	Standard								
2.87	3.2585E-02	8.6040E-03	1.0499E+08	1.3089E-01	3.4232E+06	1.2243E-01	1909	3.6533E+07		45.73	0.17	-45.74	0.17	-0.53	0.34				Standard								
2.87	3.2594E-02	1.0260E-02	9.9013E+07	3.7047E-01	3.2274E+06	3.1727E-01	1912	3.4451E+07		45.46	0.21	-45.47	0.21	-0.25	0.34				Standard								
2.88	3.2595E-02	7.1932E-03	9.7245E+07	3.1926E-01	3.1697E+06	3.1870E-01	1935	3.3816E+07		45.44	0.14	-45.44	0.14	-0.46	0.33	-0.23			Standard								
2.88	3.2567E-02	1.2160E-02	1.0084E+08	3.1331E-01	3.2843E+06	3																					

Table A2 Continued.

SIMS ANALYSES										Mount: BR7				Standard: UNIL-Q1 (Paine)				Analyser: δ ³¹ Si				Value: -0.13±0.02 (NBS-28, 2e)				Date: 03.09.2015	
Beam	ID	H12U12 (30S/28Si)		L12 (30S/16O)		H12 (30S/Coef)		CPS		Time	Yield		mean		Drift correction		Calibration		Comment								
		2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS		2SD	CPS	2SD	int	2SD	δ ³¹ Si	2SD	δ ³¹ Si		2SD							
40S1_04015	BR7_Paine_test@1	2.35	3.2698E-02	8.5542E-03	7.8302E-07	1.8324E-01	1.8324E-01	13.49	3.1338E+07	13.49	3.1338E+07	42.42	0.17	-42.55					Setting Standard								
40S1_04015	BR7_Paine_test@2	2.36	3.2695E-02	8.3876E-03	7.8302E-07	1.2329E-01	1.2329E-01	13.53	3.1979E+07	13.53	3.1979E+07	42.50	0.30	0.47					Setting Standard								
40S1_04015	BR7_Paine_test@3	2.36	3.2695E-02	8.1036E-03	7.8302E-07	1.3377E-01	1.3377E-01	13.56	3.3396E+07	13.56	3.3396E+07	42.45	0.23						Setting Standard								
40S1_04015	BR7_Paine_test@4	2.36	3.2695E-02	8.0159E-03	7.8302E-07	1.2891E-01	1.2891E-01	13.59	3.3978E+07	13.59	3.3978E+07	42.51	0.23						Setting Standard								
40S1_04015	BR7_Paine_test@5	2.36	3.2695E-02	8.1602E-03	7.8302E-07	1.2891E-01	1.2891E-01	14.03	3.3394E+07	14.03	3.3394E+07	42.48	0.20						Setting Standard								
40S1_04015	BR7_Paine_test@6	2.36	3.2695E-02	9.6415E-03	8.0167E-07	1.2827E-01	1.2827E-01	14.06	3.3940E+07	14.06	3.3940E+07	42.78	0.19						Setting Standard								
40S1_04015	BR7_Paine_test@7	2.36	3.2702E-02	1.0227E-02	7.3960E-07	3.9608E-01	3.9608E-01	14.09	3.0636E+07	14.09	3.0636E+07	42.30	0.20						Setting Standard								
40S1_04015	BR7_Paine_test@8	2.36	3.2678E-02	1.1543E-02	7.5879E-07	2.7848E-01	2.7848E-01	14.13	3.2104E+07	14.13	3.2104E+07	43.01	0.23						Setting Standard								
40S1_04015	BR7_Paine@1	2.36	3.2700E-02	1.6150E-02	7.7564E-07	1.7097E-01	1.7097E-01	14.17	3.2867E+07	14.17	3.2867E+07	3.3035E+07	0.32	-42.37	-42.43	-0.13	-0.24	-0.15	Standard								
40S1_04015	BR7_Paine@2	2.37	3.2700E-02	7.3254E-03	7.9038E-07	1.5356E-01	1.5356E-01	14.20	3.3411E+07	14.20	3.3411E+07	42.38	0.15	0.04	-42.46	0.15	0.03	-0.15	0.19	0.03							
40S1_04015	BR7_Paine@3	2.36	3.2700E-02	1.3863E-02	7.7696E-07	1.4544E-01	1.4544E-01	14.23	3.2865E+07	14.23	3.2865E+07	42.37	0.28		-42.45	0.28		-0.14	0.22	Standard							
40S1_04015	BR7_Paine@4	2.36	3.2699E-02	1.1811E-02	7.8019E-07	1.4630E-01	1.4630E-01	14.27	3.2997E+07	14.27	3.2997E+07	42.40	0.24		-42.47	0.24		-0.17	0.21	Standard							
40S1_04015	K57_rad@1	2.36	3.2726E-02	9.7275E-03	7.7400E-07	1.3478E-01	1.3478E-01	14.30	3.2735E+07	14.30	3.2735E+07	41.61	0.19	-41.14	-41.68	0.19	0.65	0.20	0.20	K57							
40S1_04015	K57_rad@2	2.36	3.2752E-02	1.3111E-02	6.9962E-07	1.7990E-01	1.7990E-01	14.33	2.2916E+06	14.33	2.2916E+06	40.83	0.26	1.28	-40.91	0.26	1.46	0.22	0.21	K57							
40S1_04015	K57_rad@3	2.36	3.2744E-02	1.2240E-02	7.4255E-07	1.3508E-01	1.3508E-01	14.37	3.1399E+07	14.37	3.1399E+07	41.06	0.24		-41.13	0.24	1.23	0.21	0.21	K57							
40S1_04015	K57_rad@4	2.37	3.2726E-02	1.2459E-02	7.5864E-07	1.5457E-01	1.5457E-01	14.40	2.4845E+06	14.40	2.4845E+06	41.60	0.25		-41.68	0.25	0.66	0.21	0.21	K57							
40S1_04015	K57_rad@5	2.37	3.2755E-02	1.0355E-02	7.3148E-07	1.8980E-01	1.8980E-01	14.43	3.2859E+06	14.43	3.2859E+06	40.76	0.21		-40.83	0.21	1.54	0.20	0.20	K57							
40S1_04015	K57_rad@6	2.37	3.2738E-02	1.1352E-02	7.3338E-07	1.6555E-01	1.6555E-01	14.47	3.1031E+06	14.47	3.1031E+06	41.24	0.23		-41.31	0.23	1.04	0.21	0.21	K57							
40S1_04015	K57_rad@7	2.36	3.2719E-02	1.1335E-02	7.6744E-07	1.2646E-01	1.2646E-01	14.50	3.2468E+07	14.50	3.2468E+07	41.81	0.27		-41.88	0.27	0.45	0.22	0.22	K57							
40S1_04015	K57_rad@8	2.37	3.2728E-02	1.1628E-02	7.5307E-07	9.0513E-02	9.0513E-02	14.53	3.1789E+07	14.53	3.1789E+07	41.60	0.23		-41.66	0.23	0.67	0.21	0.21	K57							
40S1_04015	K57_rad@9	2.37	3.2793E-02	1.3895E-02	6.9961E-07	1.0132E-01	1.0132E-01	14.56	2.9555E+07	14.56	2.9555E+07	39.63	0.28		-41.38	0.29	0.98	0.23	0.23	K57							
40S1_04015	K57_rad@10	2.37	3.2736E-02	1.4276E-02	7.4133E-07	2.5698E-01	2.5698E-01	15.00	2.2499E+06	15.00	2.2499E+06	41.31	0.29		-41.38	0.29	0.98	0.23	0.23	K57							
40S1_04015	K3c_rad@1	2.37	3.2714E-02	1.2811E-02	7.7656E-07	1.5262E-01	1.5262E-01	15.03	3.2821E+07	15.03	3.2821E+07	41.95	0.26		-42.01	0.26	0.31	0.22	0.22	K03C							
40S1_04015	BR7_Paine@5	2.36	3.2701E-02	1.1867E-02	7.9516E-07	1.3463E-01	1.3463E-01	15.07	2.6004E+06	15.07	2.6004E+06	3.3415E+07	0.24	-42.39	-42.39	-0.09	0.21	-0.08	Standard								
40S1_04015	BR7_Paine@6	2.37	3.2702E-02	1.0781E-02	7.7973E-07	1.3932E-01	1.3932E-01	15.10	2.5497E+06	15.10	2.5497E+06	42.29	0.22	0.17	-42.35	0.22	0.14	-0.05	0.20	0.15							
40S1_04015	BR7_Paine@7	2.37	3.2704E-02	1.2501E-02	7.9432E-07	1.5085E-01	1.5085E-01	15.13	2.5977E+06	15.13	2.5977E+06	42.24	0.25		-42.31	0.25	0.01	0.21	0.21	Standard							
40S1_04015	BR7_Paine@8	2.37	3.2697E-02	1.2610E-02	7.9350E-07	1.5563E-01	1.5563E-01	15.16	3.3526E+07	15.16	3.3526E+07	42.44	0.25		-42.50	0.25	-0.20	0.22	0.22	Standard							
40S1_04015	K3c_rad@2	2.37	3.2709E-02	1.4099E-02	7.8436E-07	1.6892E-01	1.6892E-01	15.20	3.3137E+07	15.20	3.3137E+07	42.09	0.28		-42.15	0.28	0.17	0.22	0.22	K03C							
40S1_04015	K3c_rad@3	2.37	3.2727E-02	1.3517E-02	7.3853E-07	2.5015E-01	2.5015E-01	15.23	2.4170E+06	15.23	2.4170E+06	41.56	0.27	1.06	-41.61	0.27	0.73	0.22	0.22	K03C							
40S1_04015	K3c_rad@4	2.37	3.2731E-02	1.6375E-02	7.4904E-07	1.0266E-01	1.0266E-01	15.26	3.1633E+07	15.26	3.1633E+07	41.45	0.33		-41.50	0.33	0.84	0.24	0.24	K03C							
40S1_04015	K3c_rad@5	2.37	3.2706E-02	8.2757E-03	7.8671E-07	1.3618E-01	1.3618E-01	15.30	3.2325E+07	15.30	3.2325E+07	42.18	0.17		-42.23	0.17	0.08	0.19	0.19	K03C							
40S1_04015	K3c_rad@6	2.37	3.2696E-02	9.1298E-03	7.9008E-07	1.5155E-01	1.5155E-01	15.33	2.5833E+06	15.33	2.5833E+06	42.47	0.18		-42.53	0.18	-0.23	0.20	0.20	K03C							
40S1_04015	K3c_rad@7	2.37	3.2727E-02	1.3385E-02	7.6718E-07	1.4830E-01	1.4830E-01	15.36	3.2085E+07	15.36	3.2085E+07	41.56	0.27		-41.61	0.27	0.73	0.22	0.22	K03C							
40S1_04015	K3c_rad@8	2.37	3.2718E-02	1.7101E-02	7.7635E-07	1.3033E-01	1.3033E-01	15.40	3.2416E+06	15.40	3.2416E+06	41.82	0.34		-41.87	0.34	0.46	0.24	0.24	K03C							
40S1_04015	K3c_rad@9	2.37	3.2717E-02	8.9040E-03	7.6749E-07	1.5763E-01	1.5763E-01	15.43	3.2430E+07	15.43	3.2430E+07	41.85	0.18		-41.90	0.18	0.43	0.20	0.20	K03C							
40S1_04015	K3c_rad@10	2.37	3.2671E-02	1.0620E-02	7.9130E-07	1.4193E-01	1.4193E-01	15.46	3.2821E+07	15.46	3.2821E+07	43.22	0.21		-43.27	0.21	-1.00	0.20	0.20	K03C							
40S1_04015	K52_rad@1	2.37	3.2723E-02	1.4032E-02	7.2386E-07	2.0918E-01	2.0918E-01	15.50	3.1659E+07	15.50	3.1659E+07	40.85	0.28		-43.09	0.28	1.87	0.22	0.22	K52							
40S1_04015	BR7_Paine@9	2.37	3.2696E-02	1.1891E-02	8.0320E-07	1.3071E-01	1.3071E-01	15.53	3.3926E+07	15.53	3.3926E+07	42.48	0.24	-42.60	-42.53	0.24	-42.64	-0.22	0.21	-0.34							
40S1_04015	BR7_Paine@10	2.37	3.2693E-02	1.3240E-02	7.9106E-07	1.3460E-01	1.3460E-01	15.56	3.3397E+07	15.56	3.3397E+07	42.56	0.26	0.21	-42.61	0.26	0.18	-0.31	0.22	0.19							
40S1_04015	BR7_Paine@11	2.37	3.2687E-02	1.0073E-02	7.9791E-07	1.3400E-01	1.3400E-01	16.00	3.3703E+07	16.00	3.3703E+07	42.73	0.20		-42.78	0.20	-0.49	0.20	0.20	Standard							
40S1_04015	BR7_Paine@12	2.37	3.2691E-02	1.1246E-02	7.8429E-07	1.3907E-01	1.3907E-01	16.03	3.3118E+07	16.03	3.3118E+07	42.61	0.22		-42.65	0.22	-0.36	0.21	0.21	Standard							
40S1_04015	K42_rad@1	2.37	3.2755E-02	1.4653E-02	7.2759E-07	1.1304E-01	1.1304E-01	16.06	3.0726E+07	16.06	3.0726E+07	40.75	0.29	-40.92	-40.79	0.29	1.59	0.23	0.23	K42							
40S1_04015	K42_rad@2	2.37	3.2748E-02	1.2128E-02	7.3407E-07	1.4083E-01	1.4083E-01	16.10	3.0995E+07	16.10	3.0995E+07	40.95	0.24	0.48	-40.99	0.24	1.38	0.21	0.21	K42							
40S1_04015	K42_rad@3	2.37	3.2748E-02	9.4508E-03	7.6768E-07	1.5635E-01	1.5635E-01	16.13	3.2390E+07	16.13	3.2390E+07	40.96	0.19		-41.00	0.19	1.37	0.20	0.20	K42							
40S1_04015	K42_rad@4	2.37	3.2728E-02	1.0159E-02	7.7803E-07	1.4340E-01	1.4340E-01	16.16	2.8516E+06	16.16	2.8516E+06	41.54	0.20		-41.58	0.20	0.76	0.20	0.20	K42							
40S1_04015	K42_rad@5	2.37	3.2748E-02	1.2327E-02	7.7803E-07	1.4448E-01																					

Table A2 Continued.

Beam nr1	SIMS ANALYSES			Mount: BR7			Standard: UNIL-QI (Paine)			Analyse: $\delta^{34}\text{S}$			Date: 03.09.2015		
	H2172 (RMS/28S)	H2172 (CPS)	$\delta^{34}\text{S}$	L2 (RMS/Coeff)	L2 (CPS)	$\delta^{34}\text{S}$	H2 (RMS/Coeff)	H2 (CPS)	$\delta^{34}\text{S}$	Yield CPS/nA	mean $\delta^{34}\text{S}$	mean $\delta^{34}\text{S}$	Drift correction $\delta^{34}\text{S}$	Calibration $\delta^{34}\text{S}_{\text{NIST}}$	Comment
430S1_040915_K4c_rad@2	2.37	3.2729E-02	1.3325E-02	7.6127E-07	1.9181E-01	2.4913E-06	1.8649E-01	17.39	3.2063E+07	3.2878E+07	-41.52	-42.56	-41.52	0.83	K04C
430S1_040915_K4c_rad@3	2.37	3.2667E-02	1.3328E-01	7.8304E-07	1.3228E-01	2.5580E+06	1.2756E-01	17.43	3.2974E+07	3.4332E+07	-43.33	-42.33	-1.06	0.21	K04C
430S1_040915_K4c_rad@4	2.37	3.2681E-02	1.3359E-02	7.8335E-07	1.4863E-01	2.5601E+06	1.4785E-01	17.46	3.3005E+07	6.9141E+05	-42.91	1.06	-43.33	-1.06	0.21
430S1_040915_K4c_rad@5	2.37	3.2701E-02	1.3378E-02	7.9001E-07	1.3996E-01	2.5743E+06	1.3648E-01	17.49	3.3148E+07	42.32	0.27	-42.91	0.27	-0.62	0.22
430S1_040915_K4c_rad@6	2.37	3.2694E-02	1.3408E-02	7.9715E-07	1.4777E-01	2.5827E+06	1.4388E-01	17.53	3.3264E+07	42.34	0.27	-42.91	-0.01	0.22	K04C
430S1_040915_K4c_rad@7	2.38	3.2701E-02	1.3431E-02	7.9888E-07	1.5143E-01	2.5472E+06	1.5421E-01	17.56	3.2794E+07	42.34	0.26	-42.33	-0.24	0.20	K04C
430S1_040915_K4c_rad@8	2.38	3.2703E-02	1.3456E-02	7.7486E-07	1.4084E-01	2.6635E+06	1.6635E-01	17.59	3.2620E+07	42.34	0.25	-42.33	0.05	0.21	K04C
430S1_040915_K4c_rad@9	2.37	3.2701E-02	1.3472E-02	7.7737E-07	1.3778E-01	2.5436E+06	1.3778E-01	18.03	3.2766E+07	42.32	0.26	-42.33	-0.01	0.22	K04C
430S1_040915_K4c_rad@10	2.38	3.2672E-02	1.3518E-02	7.8574E-07	1.4518E-01	2.5670E+06	1.4134E-01	18.06	3.3075E+07	43.17	0.20	-43.17	-0.90	0.20	K04C
430S1_040915_BR7_Paine@21	2.37	3.2709E-02	1.3471E-02	7.8733E-07	1.4037E-01	2.4499E+06	1.1294E-02	18.09	3.1531E+07	3.3433E+07	-41.71	0.27	-41.71	0.63	0.22
430S1_040915_BR7_Paine@22	2.38	3.2703E-02	1.3545E-02	7.9802E-07	1.3314E-01	2.6127E+06	1.3371E-01	18.13	3.1665E+07	3.3433E+07	-42.10	0.21	-42.32	0.23	0.20
430S1_040915_BR7_Paine@23	2.38	3.2698E-02	1.3505E-02	7.8776E-07	1.3508E-01	2.5759E+06	1.3108E-01	18.19	3.1632E+07	6.0414E+05	-42.38	0.31	0.33	0.05	0.23
430S1_040915_BR7_Paine@24	2.38	3.2693E-02	1.3523E-03	8.0299E-07	1.2220E-01	2.6225E+06	1.2037E-01	18.23	3.1733E+07	42.36	0.18	-42.36	-0.07	0.20	Standard
430S1_040915_K1R_rad@2	2.38	3.2727E-02	1.3680E-03	7.1271E-07	1.3266E-02	2.4325E+06	7.9017E-02	18.26	3.0004E+07	3.2318E+07	-41.57	0.13	-42.03	-0.16	0.19
430S1_040915_K1R_rad@3	2.38	3.2725E-02	1.3709E-03	7.6277E-07	1.4098E-01	2.4963E+06	1.4784E-01	18.29	3.2094E+07	2.1061E+06	-41.64	0.19	-41.56	0.71	0.20
430S1_040915_K1R_rad@4	2.38	3.2716E-02	1.3750E-02	7.6657E-07	1.3781E-01	2.4881E+06	1.3670E-01	18.33	3.1998E+07	4.189	0.22	-41.89	0.22	0.44	0.20
430S1_040915_K1R_rad@5	2.38	3.2692E-02	1.3520E-02	7.8438E-07	1.5438E-01	2.5664E+06	1.3945E-01	18.36	3.3025E+07	42.59	0.30	-42.57	0.30	-0.27	0.23
430S1_040915_K1R_rad@6	2.38	3.2726E-02	1.3710E-01	7.5802E-07	1.6705E-01	2.4807E+06	1.6308E-01	18.39	3.1874E+07	41.59	0.21	-41.57	0.21	0.20	K01
430S1_040915_K1R_rad@7	2.38	3.2718E-02	1.3743E-02	7.8066E-07	1.2434E-02	2.5522E+06	8.3304E-02	18.43	3.2817E+07	41.84	0.19	-41.82	0.19	0.51	0.20
430S1_040915_K1R_rad@8	2.38	3.2685E-02	1.3687E-02	7.9685E-07	1.2023E-01	2.6047E+06	1.2536E-01	18.46	3.3523E+07	42.79	0.16	-42.77	-0.48	0.19	K01
430S1_040915_K1R_rad@9	2.38	3.2705E-02	1.3733E-03	7.8350E-07	1.2523E-01	2.5623E+06	1.2536E-01	18.49	3.2982E+07	42.21	0.18	-42.19	0.18	-0.12	0.20
430S1_040915_K1R_rad@10	2.38	3.2696E-02	1.3568E-02	7.9151E-07	1.4851E-01	2.5900E+06	1.4061E-01	18.53	3.3326E+07	42.47	0.31	-42.44	-0.14	0.23	K01
430S1_040915_K32_rad@1	2.38	3.2746E-02	1.4596E-02	7.6427E-07	1.5956E-01	2.5029E+06	1.5584E-01	18.56	3.2172E+07	41.03	0.20	-41.00	0.20	1.37	K32
430S1_040915_BR7_Paine@25	2.37	3.2705E-02	1.3691E-02	8.0291E-07	1.4897E-01	2.6256E+06	1.4693E-01	18.59	3.3811E+07	3.3515E+07	-42.22	0.21	-42.32	0.12	0.00
430S1_040915_BR7_Paine@26	2.38	3.2698E-02	1.3691E-02	7.8955E-07	1.4315E-01	2.5815E+06	1.4135E-01	19.03	3.3211E+07	5.7907E+05	-42.42	0.21	-42.39	-0.08	0.21
430S1_040915_BR7_Paine@27	2.38	3.2697E-02	1.3651E-02	8.0114E-07	1.3865E-01	2.6211E+06	1.2694E-01	19.06	3.3707E+07	42.44	0.25	-42.41	0.25	-0.11	0.21
430S1_040915_BR7_Paine@28	2.38	3.2702E-02	1.3578E-02	7.9193E-07	1.4346E-01	2.5915E+06	1.3255E-01	19.09	3.3331E+07	42.30	0.27	-42.27	0.27	0.04	0.22
430S1_040915_K32_rad@2	2.38	3.2763E-02	1.4153E-02	7.2852E-07	1.8988E-02	2.3856E+06	8.0526E-02	19.13	3.0657E+07	3.0858E+07	-40.59	0.28	-40.46	1.92	0.23
430S1_040915_K32_rad@3	2.38	3.2758E-02	1.4130E-03	7.4174E-07	1.8626E-01	2.4297E+06	1.8333E-01	19.16	3.1221E+07	1.9450E+06	-40.67	0.16	-40.64	0.16	1.74
430S1_040915_K32_rad@4	2.38	3.2767E-02	1.4007E-02	7.4106E-07	1.9446E-01	2.4280E+06	1.9323E-01	19.19	3.1175E+07	40.41	0.24	-40.38	0.24	2.02	0.21
430S1_040915_K32_rad@5	2.38	3.2760E-02	1.4522E-03	7.1487E-07	1.8862E-01	2.3590E+06	1.8708E-01	19.23	3.0072E+07	-40.60	0.19	-40.57	0.19	1.82	0.20
430S1_040915_K32_rad@6	2.38	3.2772E-02	1.4343E-02	7.1276E-07	2.0414E-01	2.3559E+06	1.9833E-01	19.26	2.9999E+07	-40.25	0.33	-40.21	0.33	2.19	0.24
430S1_040915_K32_rad@7	2.38	3.2774E-02	1.1828E-02	6.9465E-07	2.1179E-01	2.2734E+06	2.0952E-01	19.29	2.9184E+07	-40.19	0.24	-40.15	0.24	2.25	0.21
430S1_040915_K32_rad@8	2.38	3.2793E-02	1.1851E-02	7.2993E-07	2.2250E-01	2.3937E+06	2.2439E-01	19.33	3.0718E+07	-39.63	0.24	-39.59	0.24	2.84	0.21
430S1_040915_K32_rad@9	2.38	3.2757E-02	1.1899E-02	7.4943E-07	1.0623E-01	2.4551E+06	1.0541E-01	19.36	3.1515E+07	-40.69	0.22	-40.65	0.22	1.73	0.21
430S1_040915_K32_rad@10	2.38	3.2761E-02	1.0789E-02	7.6603E-07	1.6177E-01	2.5097E+06	1.5683E-01	19.39	3.2237E+07	-40.58	0.22	-40.53	0.22	1.86	0.21
430S1_040915_BR7_Paine@29	2.38	3.2699E-02	1.0052E-02	7.8953E-07	1.2690E-01	2.5832E+06	1.1595E-01	19.46	3.3243E+07	3.3342E+07	-42.39	0.20	-42.39	-0.03	0.20
430S1_040915_BR7_Paine@30	2.38	3.2705E-02	1.0676E-02	8.0284E-07	1.3436E-01	2.6270E+06	1.2416E-01	19.49	3.3775E+07	7.9618E+05	-42.21	0.21	-42.16	0.16	0.20
430S1_040915_BR7_Paine@31	2.38	3.2692E-02	1.1636E-02	7.9674E-07	1.5080E-01	2.6064E+06	1.3727E-01	19.53	3.3509E+07	42.58	0.23	-42.54	0.23	-0.32	
430S1_040915_BR7_Paine@32	2.38	3.2693E-02	1.4567E-02	7.8079E-07	1.4935E-01	2.5544E+06	1.3539E-01	19.56	3.2842E+07	-42.38	0.29	-42.53	-0.23	0.23	
430S1_040915_K34_rad@1	2.38	3.2718E-02	1.4395E-02	7.2962E-07	1.7414E-01	2.3870E+06	3.7193E-01	20.29	3.0681E+07	-41.82	0.29	-41.76	0.29	0.58	0.23
430S1_040915_BR7_Paine@33	2.38	3.2699E-02	1.1880E-02	7.9544E-07	1.5090E-01	2.6099E+06	1.4877E-01	20.33	3.4431E+07	3.3038E+07	-42.32	0.24	-42.32	-0.01	0.21
430S1_040915_BR7_Paine@34	2.38	3.2697E-02	1.1580E-02	7.7549E-07	1.5769E-01	2.5538E+06	1.5533E-01	20.36	3.5825E+07	1.7158E+06	-42.39	0.23	0.06	-0.08	0.21
430S1_040915_BR7_Paine@35	2.38	3.2697E-02	1.0881E-02	8.0952E-07	1.1373E-01	2.6487E+06	1.0839E-01	20.39	3.4029E+07	42.45	0.22	-42.39	0.22	0.08	0.21
430S1_040915_BR7_Paine@36	2.38	3.2698E-02	1.4187E-02	7.6419E-07	1.8788E-01	2.4685E+06	1.8414E-01	20.43	3.2109E+07	-42.43	0.28	-42.36	-0.05	0.22	
430S1_040915_K34S_rad@2	2.38	3.2740E-02	1.0323E-02	7.8865E-07	1.3088E-01	2.5321E+06	1.3395E-01	20.46	3.1923E+07	3.3045E+07	-41.18	0.21	-41.40	0.21	1.25
430S1_040915_K34S_rad@3	2.38	3.2742E-02	1.2278E-02	7.9381E-07	1.4501E-01	2.6056E+06	1.4477E-01	20.49	3.3440E+07	1.8242E+06	-41.13	0.25	-41.06	0.25	1.30
430S1_040915_K34S_rad@4	2.38	3.2738E-02	1.3991E-02	7.7557E-07	1.7056E-01	2.5400E+06	1.4536E-01	20.53	3.2598E+07	-41.56	0.28	-41.49	0.28	1.85	
430S1_040915_K34S_rad@5	2.38	3.2738E-02	1.1103E-02	7.9996E-07	1.3153E-01	2.6190E+06	1.2928E-01	20.56	3.3603E+07	-41.25	0.22	-41.18	0.22	1.18	
430S1_040915_K34S_rad@6	2.38	3.2732E-02	1.3780E-02	7.8167E-07	1.6749E-01	2.5387E+06	1.6541E-01	20.59	3.2824E+07	-41.24	0.28	-41.17	0.28	0.99	
430S1_040915_K34S_rad@7	2.38	3.2738E-02	1.1524E-02	7.7029E-07	1.7029E-01	2.6177E+06	1.6226E-01	21.03	3.3614E+07	-41.44	0.23	-41.37	0.23	1.19	
430S1_040915_K34S_rad@8	2.38	3.2744E-02	1.0931E-02	7.8033E-07	1.4801E-01	2.5567E+06	1.4371E-01	21.06	3.2790E+07	-41.08	0.22	-41.01	0.22	1.36	
430S1_040915_K34S_rad@9	2.38	3.2719E-02	1.2330E-02	8.0386E-07											

Table A2 Continued.

SIMS ANALYSES										Mount: BR5				Standard: UNIL-Q1 (Paine)				Analyse: $\delta^{30}\text{Si}$				Value: -0.13 ± 0.02 (NBS-28, 2e)				Date: 03.09.2015			
Beam	nA	H ₂ /O ₂ (60S/28S)		L ₂ (60S/12C)		H ₂ (60S/28S)		CPS		Time	Yield	mean		Drift correction		Calibration		Comment											
		CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD			2SD Int	2SD Int	2SD Int	2SD Int	$\delta^{30}\text{Si}$ (avg)	2SD												
40S1_03016_PaneBR5@1	2.39	3.2730E-02	1.2599E-02	8.9573E-07	1.0093E-01	2.9130E-06	1.0717E-01	1617	3.7463E+07	16.17	3.7463E+07	-41.50	0.25	-41.84	-41.50	0.25	-41.85	0.18	-0.18	Setting Standard									
40S1_03016_PaneBR5@2	2.39	3.2721E-02	1.0458E-02	8.9823E-07	1.0732E-01	2.9390E-06	1.0732E-01	1621	3.7592E+07	16.21	3.7592E+07	-41.76	0.21	-41.76	-41.76	0.21	0.37	0.09	0.22	Setting Standard									
40S1_03016_PaneBR5@3	2.39	3.2725E-02	1.2325E-02	8.9977E-07	1.0489E-01	2.9438E-06	1.0862E-01	1624	3.7605E+07	16.24	3.7605E+07	-41.62	0.25	-41.63	-41.63	0.25	0.05	0.23	0.29	Setting Standard									
40S1_03016_PaneBR5@4	2.39	3.2711E-02	1.2128E-02	8.9726E-07	9.7604E-02	2.9159E-06	1.0627E-01	1627	3.7230E+07	16.27	3.7230E+07	-42.03	0.42	-42.04	-42.04	0.42	-0.38	0.29	0.29	Setting Standard									
40S1_03016_PaneBR5@5	2.39	3.2716E-02	1.1808E-02	8.9233E-07	1.1680E-01	2.9191E-06	1.1611E-01	1630	3.7355E+07	16.30	3.7355E+07	-41.90	0.44	-41.91	-41.91	0.44	-0.24	0.29	0.24	Setting Standard									
40S1_03016_PaneBR5@6	2.40	3.2712E-02	1.3882E-02	8.9644E-07	9.7443E-02	2.9324E-06	1.0325E-01	1634	3.7386E+07	16.34	3.7386E+07	-42.02	0.28	-42.03	-42.03	0.28	-0.37	0.24	0.24	Setting Standard									
40S1_03016_PaneBR5@7	2.40	3.2715E-02	1.4699E-02	8.9665E-07	9.7678E-02	2.9335E-06	1.0510E-01	1637	3.7371E+07	16.37	3.7371E+07	-41.92	0.29	-41.93	-41.93	0.29	-0.26	0.24	0.24	Setting Standard									
40S1_03016_PaneBR5@8	2.40	3.2715E-02	1.4699E-02	8.9665E-07	9.7678E-02	2.9335E-06	1.1201E-01	1640	3.7355E+07	16.40	3.7355E+07	-42.00	0.20	-42.01	-42.01	0.20	-0.34	0.22	0.22	Setting Standard									
40S1_03016_PaneBR5@9	2.40	3.2722E-02	1.1383E-02	8.9731E-07	9.3143E-02	2.9363E-06	9.5543E-02	1732	3.7433E+07	17.32	3.7433E+07	-41.72	0.23	-41.73	-41.73	0.23	-41.79	-0.06	0.22	-0.12	Standard								
40S1_03016_PaneBR5@10	2.40	3.2707E-02	1.6588E-02	9.0586E-07	1.1606E-01	2.9623E-06	1.0919E-01	1733	3.7812E+07	17.33	3.7812E+07	-42.15	0.33	-42.15	-42.15	0.33	0.42	-0.50	0.25	0.44	Standard								
40S1_03016_PaneBR5@11	2.39	3.2726E-02	1.1308E-02	8.9312E-07	9.5622E-02	2.9228E-06	8.8626E-02	1739	3.7316E+07	17.39	3.7316E+07	-41.60	0.23	-41.60	-41.60	0.23	0.08	0.22	0.22	Standard									
40S1_03016_PaneBR5@12	2.39	3.2723E-02	1.1822E-02	8.9801E-07	9.9524E-02	2.9094E-06	1.0737E-01	1742	3.7276E+07	17.42	3.7276E+07	-41.68	0.24	-41.69	-41.69	0.24	-0.11	0.23	0.23	Standard									
40S1_03016_KB86@1	2.38	3.2767E-02	9.6572E-03	8.7800E-07	1.1483E-01	2.8639E-06	1.1743E-01	1745	3.6678E+07	17.45	3.6678E+07	-40.40	0.19	-41.03	-40.40	0.19	1.33	-0.00	0.22	K338									
40S1_03016_KB86@2	2.39	3.2745E-02	1.6488E-02	8.8321E-07	1.0313E-01	2.8895E-06	1.1671E-01	1748	3.7001E+07	17.48	3.7001E+07	-41.04	0.33	-41.04	-41.04	0.33	0.66	0.25	0.25	K338									
40S1_03016_KB86@3	2.39	3.2749E-02	1.6855E-02	8.7888E-07	1.2609E-01	2.8783E-06	1.1716E-01	1752	3.6789E+07	17.52	3.6789E+07	-40.92	0.34	-40.93	-40.93	0.34	0.78	0.26	0.26	K338									
40S1_03016_KB86@4	2.39	3.2759E-02	1.4869E-02	8.6769E-07	9.4572E-02	2.8429E-06	8.9309E-02	1755	3.6233E+07	17.55	3.6233E+07	-40.63	0.30	-40.64	-40.64	0.30	1.08	0.24	0.24	K338									
40S1_03016_KB86@5	2.40	3.2742E-02	1.5259E-02	8.7890E-07	9.5538E-02	2.8782E-06	1.0338E-01	1758	3.6660E+07	17.58	3.6660E+07	-41.12	0.31	-41.13	-41.13	0.31	0.57	0.25	0.25	K338									
40S1_03016_KB86@6	2.39	3.2743E-02	8.3122E-03	8.821E-07	1.2460E-01	2.8877E-06	1.2989E-01	1801	3.6855E+07	18.01	3.6855E+07	-41.12	0.17	-41.12	-41.12	0.17	0.58	0.21	0.21	K338									
40S1_03016_KB86@7	2.39	3.2746E-02	1.2357E-02	8.7202E-07	1.2460E-01	2.8553E-06	1.1565E-01	1805	3.6523E+07	18.05	3.6523E+07	-41.02	0.25	-41.02	-41.02	0.25	0.69	0.23	0.23	K338									
40S1_03016_KB86@8	2.39	3.2738E-02	1.0977E-02	8.8044E-07	1.1828E-01	2.8824E-06	1.1828E-01	1808	3.6892E+07	18.08	3.6892E+07	-41.26	0.22	-41.26	-41.26	0.22	0.43	0.22	0.22	K338									
40S1_03016_KB86@9	2.39	3.2723E-02	1.8451E-02	8.8567E-07	1.0884E-01	2.8966E-06	1.1093E-01	1811	3.7007E+07	18.11	3.7007E+07	-41.68	0.37	-41.69	-41.69	0.37	-0.01	0.27	0.27	K338									
40S1_03016_KB86@10	2.40	3.2742E-02	1.8200E-02	8.8034E-07	1.0545E-01	2.8807E-06	1.1634E-01	1814	3.6744E+07	18.14	3.6744E+07	-41.13	0.36	-41.13	-41.13	0.36	0.57	0.27	0.27	K338									
40S1_03016_PaneBR5@13	2.39	3.2723E-02	2.506E-02	8.9681E-07	8.3345E-02	2.9340E-06	8.0731E-02	1818	3.7470E+07	18.18	3.7470E+07	-41.69	0.50	-41.70	-41.70	0.50	-41.92	-0.02	0.32	-0.26	Standard								
40S1_03016_PaneBR5@14	2.39	3.2709E-02	9.5730E-03	8.9544E-07	9.9818E-02	2.9297E-06	1.0433E-01	1821	3.7506E+07	18.21	3.7506E+07	-42.09	0.19	-42.10	-42.10	0.19	0.30	-0.44	0.22	0.32	Standard								
40S1_03016_PaneBR5@15	2.39	3.2712E-02	1.6328E-02	8.9727E-07	1.0251E-01	2.9355E-06	1.0840E-01	1824	3.7582E+07	18.24	3.7582E+07	-42.01	0.33	-42.01	-42.01	0.33	-0.35	0.25	0.25	Standard									
40S1_03016_PaneBR5@16	2.40	3.2717E-02	1.4308E-02	8.9624E-07	8.3673E-02	2.9239E-06	8.7203E-02	1827	3.7414E+07	18.27	3.7414E+07	-41.87	0.29	-41.88	-41.88	0.29	-0.21	0.24	0.24	Standard									
40S1_03016_PaneBR5@17	2.39	3.2714E-02	2.0163E-02	8.9743E-07	8.7382E-02	2.9369E-06	8.7169E-02	1910	3.7581E+07	19.10	3.7581E+07	-41.96	0.40	-41.68	-41.96	0.40	-41.68	-0.20	0.28	-0.01	Standard								
40S1_03016_PaneBR5@18	2.39	3.2733E-02	1.3335E-02	8.9291E-07	9.2416E-02	2.9227E-06	9.7886E-02	1913	3.7429E+07	19.13	3.7429E+07	-42.39	0.27	-42.47	-42.39	0.27	0.41	-0.30	0.24	0.42	Standard								
40S1_03016_PaneBR5@19	2.39	3.2721E-02	9.3495E-03	8.9482E-07	1.0740E-01	2.9281E-06	1.1181E-01	1916	3.7562E+07	19.16	3.7562E+07	-41.74	0.19	-41.75	-41.75	0.19	-0.07	0.22	0.22	Standard									
40S1_03016_PaneBR5@20	2.38	3.2725E-02	1.0092E-02	8.967E-07	9.9850E-02	2.9442E-06	1.0295E-01	1919	3.7688E+07	19.19	3.7688E+07	-41.64	0.20	-41.64	-41.64	0.20	0.04	0.22	0.22	Standard									
40S1_03016_PaneBR5@21	2.38	3.2713E-02	8.0536E-03	8.9526E-07	1.0469E-01	2.9286E-06	1.0780E-01	1919	3.7688E+07	19.19	3.7688E+07	-41.74	0.20	-41.74	-41.74	0.20	-0.33	0.21	-0.06	Standard									
40S1_03016_PaneBR5@22	2.38	3.2719E-02	1.0759E-02	1.0102E-01	2.9333E-02	2.9333E-06	1.0780E-01	2015	3.7641E+07	20.15	3.7641E+07	-41.80	0.22	-41.80	-41.80	0.22	0.49	-0.13	0.22	0.51	Standard								
40S1_03016_PaneBR5@23	2.39	3.2735E-02	1.8993E-02	8.9235E-07	8.2768E-02	2.9202E-06	8.2768E-02	2018	3.7402E+07	20.18	3.7402E+07	-41.34	0.38	-41.34	-41.34	0.38	0.35	0.27	0.27	Standard									
40S1_03016_PaneBR5@24	2.38	3.2719E-02	1.8026E-02	8.9237E-07	8.9843E-02	2.9197E-06	8.9744E-02	2021	3.7493E+07	20.21	3.7493E+07	-41.82	0.36	-41.82	-41.82	0.36	-0.15	0.26	0.26	Standard									
40S1_03016_PaneBR5@25	2.37	3.2725E-02	9.4497E-03	8.9020E-07	9.3873E-02	2.9132E-06	9.4630E-02	2057	3.7497E+07	20.57	3.7497E+07	-41.62	0.19	-41.62	-41.62	0.19	-41.68	0.05	0.22	0.00	Standard								
40S1_03016_PaneBR5@26	2.37	3.2714E-02	9.4758E-03	8.9256E-07	1.2561E-01	2.9126E-06	1.2561E-01	2100	3.7552E+07	21.00	3.7552E+07	-41.94	0.40	-41.94	-41.94	0.40	-0.28	0.22	0.36	Standard									
40S1_03016_PaneBR5@27	2.38	3.2724E-02	1.1021E-02	8.9223E-07	9.4406E-02	2.9231E-06	9.7635E-02	2104	3.7547E+07	21.04	3.7547E+07	-41.67	0.22	-41.67	-41.67	0.22	0.01	0.22	0.22	Standard									
40S1_03016_PaneBR5@28	2.38	3.2731E-02	2.3450E-02	8.9232E-07	8.0476E-02	2.9228E-06	8.5623E-02	2107	3.7496E+07	21.07	3.7496E+07	-41.47	0.47	-41.47	-41.47	0.47	0.22	0.30	0.30	Standard									
40S1_03016_PaneBR5@29	2.38	3.2722E-02	1.7666E-02	8.9386E-07	1.0125E-01	2.9250E-06	1.0752E-01	2159	3.7602E+07	21.59	3.7602E+07	-41.71	0.35	-41.73	-41.71	0.35	-41.73	-0.03	0.26	-0.06	Standard								
40S1_03016_PaneBR5@30	2.39	3.2717E-02	1.7077E-02	8.9044E-07	7.6190E-02	2.9124E-06	6.6389E-02	2202	3.7318E+07	22.02	3.7318E+07	-41.86	0.34	-41.86	-41.86	0.34	0.19	-0.19	0.26	0.20	Standard								
40S1_03016_PaneBR5@31	2.38	3.2720E-02	1.9733E-02	8.9071E-07	8.3952E-02	2.9145E-06	8.9301E-02	2205	3.7436E+07	22.05	3.7436E+07	-41.77	0.44	-41.76	-41.76	0.44	-0.09	0.29	0.29	Standard									
40S1_03016_PaneBR5@32	2.37	3.2726E-02	1.5505E-02	8.8837E-07	8.3440E-02	2.9078E-06	8.6851E-02	2209	3.7411E+07	22.09	3.7411E+07	-41.59	0.31	-41.59	-41.59														

二次イオン質量分析法 (SIMS) を用いた
中部日本犬山地域中生代層状チャート中の放散虫殻 Si 同位体分析

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要 旨

全球シリカ循環は長期的気候システムの重要な要素だが、その制御要因は古環境指標の制約に乏しいため、不確実性が大きい。本論では、二次イオン質量分析計 (SIMS) によって測定された犬山地域の中生代チャートに含まれる放散虫化石のシリカ変動 ($\delta^{30}\text{Si}$) を報告する。測定の結果、放散虫殻 $\delta^{30}\text{Si}$ は $-0.3 \sim 2 \text{ ‰}$ で、現在及び新生代の放散虫殻の値と調和的であった。さらに、予察的な $\delta^{30}\text{Si}$ 変動は低解像度にもかかわらず、1,000 万年スケールでは生物起源シリカ (BSi) 埋没速度と逆相関し、従来の古生産性プロキシとしての $\delta^{30}\text{Si}$ の解釈に矛盾する結果となった。この時間スケールでは BSi 埋没速度は風化速度に依存するため、風化しやすく低 $\delta^{30}\text{Si}$ の苦鉄質岩の風化速度変化によって、この逆相関は説明されるかもしれない。さらに高解像度で $\delta^{30}\text{Si}$ 記録を測定することで、過去のシリカ循環をより深く理解できると期待される。

Oxygen isotope analysis of Mesozoic radiolarites using SIMS

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Abstract: The oxygen isotope ($\delta^{18}\text{O}$) analysis of carbonate fossils is widely applied for palaeoceanographic analysis, whereas that of siliceous fossils is only limited partly due to technical constraints and uncertain fractionation factors. Here we used a secondary ion mass spectrometer (SIMS) for $\delta^{18}\text{O}$ of radiolarian silica, precipitated inside radiolarian molds in Mesozoic radiolarites from Japan, Italy, Switzerland and Romania in order to examine its potential for palaeoceanographic proxy. 507 measurements of the isotopic oxygen signature relative to the Vienna Standard Mean Ocean Water ($\delta^{18}\text{O}_{\text{VSMOW}}$) of 53 chert samples range between 19.8 to 35.3 ‰ overlapping with that of modern and Cenozoic radiolarian tests in the equatorial Pacific. Relatively large intra-chert variability supports that $\delta^{18}\text{O}$ of the Mesozoic radiolarian tests are not perfectly homogenized within a chert bed during the diagenetic segregation. The temporal changes in the $\delta^{18}\text{O}$ values of radiolarians ($\delta^{18}\text{O}_{\text{radiolarians}}$) show an Early-Middle Triassic slight positive excursion, a Late Triassic high plateau, an Early Jurassic negative excursion with up to 8 ‰, a Middle Jurassic slight positive excursion, and a few light values for the Cretaceous despite of their low resolution. A comparison of $\delta^{18}\text{O}$ between radiolarian molds, conodont apatite, and the low magnesium calcium shells show overall similar secular variations during the Triassic, but different trends was observed during the Early Jurassic. Because our data is low-resolution, further cross check of $\delta^{18}\text{O}_{\text{radiolarians}}$ is necessary to use as a proxy for paleoceanography.

Keywords: $\delta^{18}\text{O}$, Mesozoic, radiolarites, radiolarians, SIMS

1. Introduction

One of the most widely used palaeoceanographic techniques is the oxygen isotope ($\delta^{18}\text{O}$) analysis of carbonate shells which reflects past environmental changes, such as temperature, ice sheet volume, and precipitation/evaporation ratio (e.g., Emiliani, 1955; Shackleton and Kennett, 1975). However, a significant caveat in the paleoceanographic analyses using carbonate shells is their scarcity or complete absence in some sediments for large sections of the globe and deep past, due to dissolution below carbonate compensation depth (CCD) and/or carbonate organism evolution in pelagic ocean after the Late Triassic. The most easily available archive of seawater $\delta^{18}\text{O}$ for such sediments is biogenic silica (BSi), such as diatoms, sponges, and radiolarians (e.g. Jaffrés *et al.*, 2007). At least, $\delta^{18}\text{O}$ values of opal

have been recognized as a potential proxy of past seawater temperature and isotopic composition as referred for diatoms (Labeyrie, 1974; Mikkelsen *et al.*, 1978; Juillet-Leclerc and Labeyrie, 1987; Shemesh *et al.*, 1992, 2001; Schmidt *et al.*, 1997; Swann *et al.*, 2008; Swann and Leng, 2009; Maier *et al.*, 2013), even if this is still debated. In Southern Ocean cores, $\delta^{18}\text{O}$ values of diatoms and radiolarians show similar patterns with similar values from 43 ‰ to 45 ‰ for at least last 30 ky (cf. Abelmann *et al.*, 2015). Therefore, $\delta^{18}\text{O}_{\text{radiolarians}}$ might also be a potential proxy for paleoceanography.

Radiolarians dominated as BSi producers in Paleozoic and Mesozoic open ocean (Hein and Parrish, 1987), whereas siliceous sponges were and are today largely restricted to marginal settings, and diatoms became quantitatively important in the Cenozoic (Racki and Cordey, 2000; Kidder and Erwin, 2001). Radiolarites were deposited in

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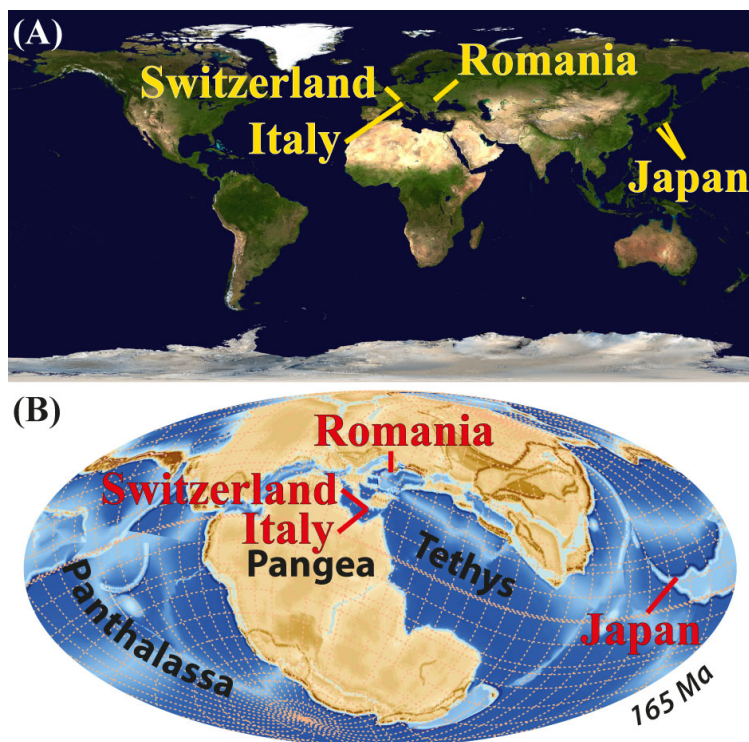


Fig. 1 Location of the studied sections (A) and their paleogeography during the Middle Jurassic (B). The Paleomap is from the Stampfli model developed at the University of Lausanne (Stampfli and Borel, 2002).

a broad low-latitude belt, and radiolarian-bearing siliceous mudstones also dominated in mid-latitudes (Baumgartner, 2013). Therefore, radiolarian $\delta^{18}\text{O}$ provides potentially important information for paleoceanography of the Paleozoic and Mesozoic.

Some evidence supports that radiolarian $\delta^{18}\text{O}$ may reflect environmental change (Knauth and Epstein, 1976; Wu *et al.*, 1997; Abelmann *et al.*, 2015). Although diagenetic isotopic fractionation of radiolarian test is still debated, even for Cenozoic (e.g. Fontorbe *et al.*, 2016), that for early Mesozoic bedded chert might be more simple system due to diagenetic segregation; This segregation results from the migration of silica from layers with low-Si content to layers with high-Si content during the transformation from opal-A to opal-CT (Isaacs, 1981; Tada, 1991). Thus, the cm-scale alternations of chert and shale might have limited migration of BSi within a chert-shale couplet. Contrary to this advantage, such diagenetic processes make extraction and picking of radiolarian molds without strong chemical procedures quite hard. To solve this disadvantage, secondary ion mass spectrometer (SIMS) is a powerful technique to measure the Mesozoic radiolarian molds on $\sim 10\ \mu\text{m}$ scale. Here we used SIMS for the Mesozoic radiolarian $\delta^{18}\text{O}$ to examine their paleoceanographic and diagenetic imprints.

2. Materials

We collected 55 chert samples from Mesozoic radiolarites in Japan, Italy, Switzerland and Romania (Fig. 1). Only fresh cherts were sampled to avoid alteration. The

Triassic to Early Jurassic Panthalassan bedded cherts are distributed in the Inuyama Area from Japan (Kiso River sections; Nakaseko and Nishimura, 1979; Yao *et al.*, 1980; Mizutani and Koike, 1982; Hori, 1988; Sugiyama, 1997; Yao and Kuwahara, 1997; Carter and Hori, 2005; Ikeda and Tada, 2014). These bedded cherts consist of several tectonic imbricates outcropping along the Kiso River which were formed during the Jurassic accretion (Kimura and Hori, 1993). The Cretaceous Panthalassan radiolarites come from the Goshikikahama section (Okamura and Uto, 1982; Kodama *et al.*, 1983). We sampled Middle Jurassic Tethyan radiolarites from the Sogno section (Gaetani and Poliani, 1978; Baumgartner *et al.*, 1980; Kocher, 1981; Baumgartner, 1984; Baumgartner *et al.*, 1995; Ikeda *et al.*, 2016). Additional material (Pi01, Pe01, Ro01 and Ca01) are a Middle Jurassic radiolarite sample from the southern part of Switzerland ($45^{\circ}54'12''\text{N}$, $8^{\circ}59'55''\text{E}$), a Late Permian radiolarian chert from Neo section, Japan ($35^{\circ}41'39''\text{N}$, $136^{\circ}39'25''\text{E}$), a Middle Jurassic radiolarian chert from the Rarau synclinal of the Carpathian Mountain along a road going to Lagu Rosu, Romania ($46^{\circ}47'29''\text{N}$, $25^{\circ}47'32''\text{E}$; Dumitrica, 1995) and a Berriasian diagenetic chert nodule from the Capriolo section in the Lombardian basin, Italy ($45^{\circ}38'40''\text{N}$, $9^{\circ}57'32''\text{E}$; Weissert *et al.*, 1979; Channell *et al.*, 1987; Lini *et al.*, 1992; Föllmi *et al.*, 2012), respectively.

Bedded chert is composed of silica-rich chert layers interbedded with silica-poorer shale partings (Davis, 1918; Tada, 1991; Hori *et al.*, 1993). In radiolarian cherts, radiolarian silica can be easily distinguished from a stained matrix enclosing radiolarian molds often filled with nearly

Table 1 Raw, drift corrected, calibrated for our $\delta^{18}\text{O}_{\text{VSMOW}}$ least square mean and standard deviations on the UNIL-Q1 standard (in ‰; 2SD). The identical raw and drift corrected $\delta^{18}\text{O}_{\text{VSMOW}}$ suggests that there was no major instrumental drift during our sessions. Overall, the reproductivity of the UNIL-Q1 on the 269 measurements after calibration is 0.3 ‰ (2SD). The raw $\delta^{18}\text{O}_{\text{VSMOW}}$ of the PRIM123 session for Ki08C, Pi01 and Ca01 is in average 2.5 ‰ higher than the other session due to instrumental fractionation differences.

Session	Number of measurements	Raw $\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)		Drift corrected $\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)		Calibrated $\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)	
		LSmean	LSstd	LSmean	LSstd	LSmean	LSstd
Prim123	48	6.5	0.3	6.5	0.3	9.8	0.3
Br2	47	4.4	0.5	4.4	0.5	9.8	0.5
Br3	31	4.5	0.4	4.4	0.4	9.8	0.3
Br4	52	3.8	0.3	3.8	0.3	9.8	0.3
Br6	28	3.7	0.3	3.7	0.3	9.8	0.3
Br7	63	3.9	0.3	3.9	0.3	9.8	0.3
All UNIL-Q1 measurements	269					9.8	0.3

pure microquartz and/or chalcedony (Bôle *et al.*, 2020). Most of these microquartz are early diagenetic products from opal-A precipitating in the pore space produced by radiolarian skeletons which are the target of this study.

3. Methods

Secondary ion mass spectrometer (SIMS) can achieve accurate in situ analyses of very small sample amount. The aim of in situ analyses in the infills radiolarian molds by SIMS is to measure the oxygen derived from their microquartz without contamination from clays or aeolian/detrital minerals present in the matrix. Microquartz infills in radiolarian molds, which are likely an early silica precipitation derived from biosilica, were analysed with a CAMECA IMS 1280HR at the University of Lausanne (e.g. Seitz *et al.*, 2016).

About 10 subsamples were mounted in each 1-inch micron-polished epoxy sample holder (Epopit resin) around an internal standard. An amorphous glass standard from the National Institute of Standard and Technology (NIST-610) was used for the initial three mounts with samples Ki08C, Pi01, and Ca01, to control instrumental drift in absence of a proper quartz standard. The quartz internal standard UNIL-Q1 (Seitz *et al.*, 2016) have been used for all mounts, including initial three samples reanalysed, to correct instrumental drift and mass fractionation. As instrumental mass fractionation on SIMS depend on minerals species (Marin *et al.*, 2010), data from chert samples are calibrated with UNIL-Q1. Before SIMS measurements, radiolarian molds and their microcrystalline quartz infill was carefully examined on each subsample on mount by optical and/or scanning electrons microscopy (SEM).

The $\delta^{18}\text{O}$ from 532 points of radiolarian molds was measured with a primary Cs⁺ ion beam intensity of ~2 nA, resulting in a beam diameter of ~10 μm . Electrical charges were compensated using an electron flood gun,

with normal incidence and the conductivity of the sample surface was assured by a gold coating connected to electrical ground. ^{16}O and ^{18}O secondary ions, accelerated at 10 kV, were analyzed at a mass resolving power of 3000 and collected on Faraday cups (FC) multi-collection mode. The resistances of the L'2 and H'2 FC were set to $10^{10} \Omega$ and $10^{11} \Omega$ for the detection of ^{16}O and ^{18}O , respectively. FC were calibrated in the beginning of each session, using the calibration routine. Mass calibration was performed at the beginning of each session and every 12 h.

Each analysis took less than 4 minutes and consists of 20 cycles of 5 seconds starting with a presputtering time of 30 seconds to remove gold, stabilized the secondary ion emission allowing automatic centring of the secondary ion beam. This setting allowed an average reproducibility better than 0.3 ‰ (2 standard deviation, 2SD) on UNIL-Q1 (Seitz *et al.*, 2016; Table 1) at the beginning each session, and analytical standard deviation for each analysis lower than 0.3 ‰ (2SD). The analytical standard deviation expressed here is the standard deviation of each data from different analytical cycles. A minimum set of 4 analyses of UNIL-Q1 quartz standard (9.18 ± 0.14 ‰ (VSMOW); Seitz *et al.*, 2016), inserted in each mount, has been measured every 6-10 analyses for monitoring the instrument stability, and for accurately correcting the instrumental mass fractionation, as it can slightly differ from mount to mount. The data have been obtained in 10 different sessions for $\delta^{18}\text{O}$ measurements, over 11 months. The variation of the UNIL-Q1 quartz standard over the entire sessions is < 0.3‰ (2SD after the drift correction (Table 1).

The target locations were controlled by optical methods and SEM for each measurement to check that radiolarian molds were effectively hit. Analytical yield and standard deviation of each measurement were also used to check the validation. The analytical yield is the quantity of elements measured relative to the intensity of primary ion beam (cps/nA), and here we used relative analytical yield normalized

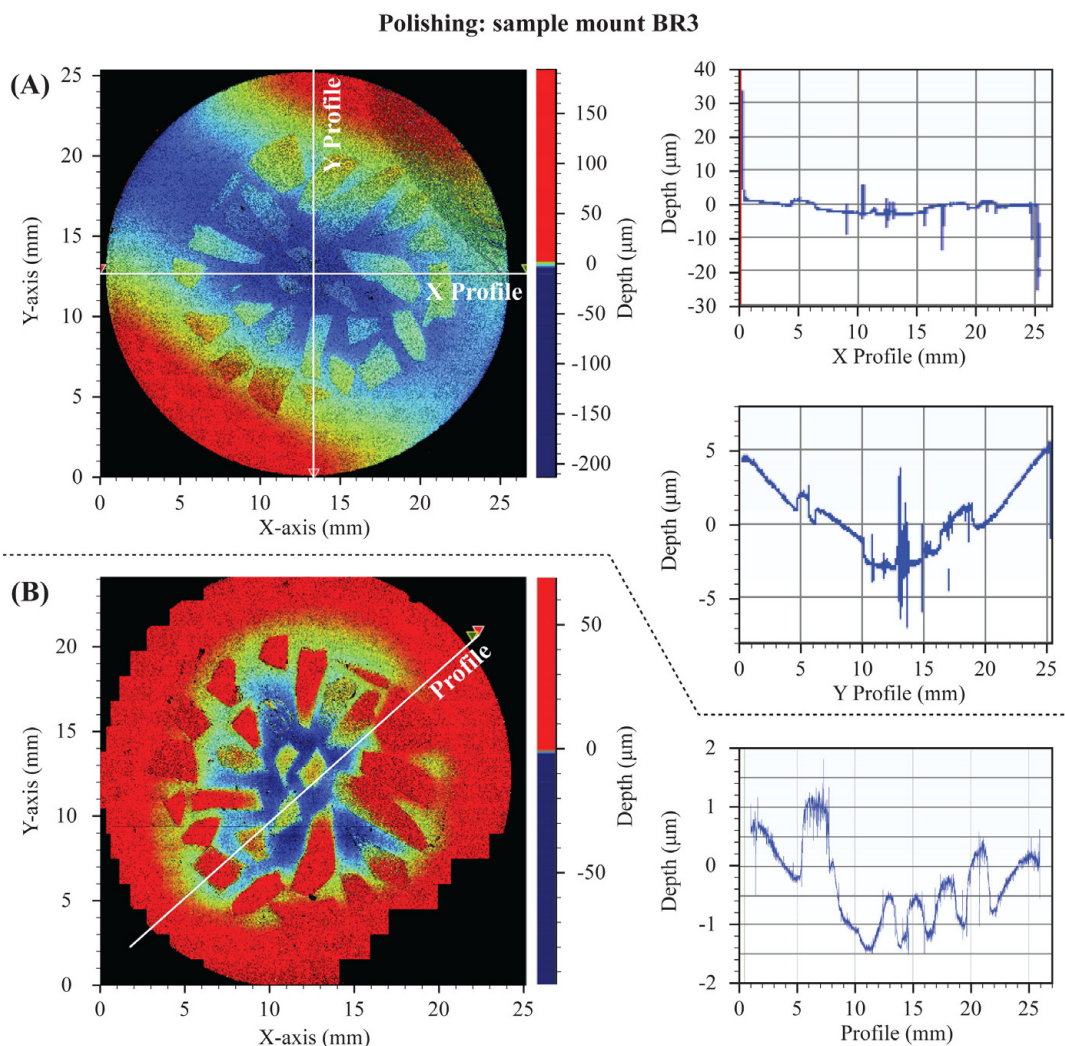


Fig.2 Topography of the sample mount BR3 after polishing. (A) Between the lower part of the rim and the centre of the sample mount, there is a difference of about 7.5 μm in depth. This difference was due to an internal tension triggered by a screw which initially hold the sample holder during the polishing. (B) After a second polishing without the screw, the topography of the rim is better and the difference between the rim and the center is reduced to 0.25 μm . The profiles on the right part are represented by white lines on the surface of the sample mount.

by that of internal standard UNIL-Q1. Different analytical yields result from the nature of the analysed material (mineral species and matrix effect) and from the topography of the analysed surface, which modifies the incident angle of the primary ion beams and thus the energy per surface of the primary ion beam. Therefore, the planarity of our sample mounts after polishing up to 0.25 μm was checked using white light profilometer (Brucker: Countour GT; Fig. 2) to have elevation differences of less than 1.5 μm between the standard and the samples. High analytical standard deviation depends on isotopic heterogeneities in the analysed minerals, but also on modifications of the instrumental mass fractionation which could be triggered by changes of the analytical conditions (topography, beam and analyser stability) or by the analyses of a mixture of silica, clay minerals and oxides. Overall, analytical yield

similar to the one of our UNIL-Q1 standard (~ 7.9 cps/nA, but depends on sessions) and analytical standard deviation lower than 0.35 ‰ (2SD) are objective criterions to decide if a measurement should be accepted.

The drift correction was realized using a least square regression line weighted for incertitude (σ_i^2). For the calibration, the weighted $\delta^{18}\text{O}$ -mean (\bar{x}) and standard deviation ($\hat{\sigma}_i$) for the internal standard was calculated also using the incertitude (Equation 1 and 2) to keep consistent data processing with the least square drift correction. The calibrated $\delta^{18}\text{O}$ for samples ($\delta^{18}\text{O}_{VSMOW} Spl$), depend on each sample measurement ($\delta^{18}\text{O}_{Spl_{measured}}$) and are proportional to the measured least square $\delta^{18}\text{O}$ -mean and the true $\delta^{18}\text{O}_{VSMOW}$ from the internal standard ($\delta^{18}\text{O}_{Std_{measured}}$ and $\delta^{18}\text{O}_{VSMOW} Std$, respectively) (Equation 3). The errors on the calibrated $\delta^{18}\text{O}$ ($\sigma(\delta^{18}\text{O}_{VSMOW} Spl)$) were

obtained by error propagation (Equation 4). The weighted means and standard deviations (Table 2 and Appendix Table A1) were then calculated for each sample following equation 1 and 2. Conversion from VSMOW to VPDB was realized using the Equation 5 (Kim *et al.*, 2015). For comparison, $\delta^{18}\text{O}$ values relative to the Vienna Pee Dee Belemnite standard for low magnesium calcium shells ($\delta^{18}\text{O}_{\text{VPDB LMC}}$) from Grossman (2012) were digitalized using PlotDigitizer (2.6.8).

Equation 1

$$\dot{x} = \sum \left(\frac{1}{\sigma_i^2} \times x_i \right) / \left(\frac{1}{\sigma_i^2} \right)$$

Equation 2

$$\dot{\sigma}_i = \sqrt{\sum \left(\frac{1}{\sigma_i^2} \times (x_i - \dot{x})^2 \right) / \sum \left(\frac{1}{\sigma_i^2} \right) \times \frac{N}{N-1}}$$

Equation 3

$$\begin{aligned} \delta^{18}\text{O}_{\text{VSMOW Spl}} &= x_i \\ &= \left(\left(\left(1 + \frac{\delta^{18}\text{O Spl}_{\text{measured}}}{1000} \right) / \left(1 + \frac{\delta^{18}\text{O Std}_{\text{measured}}}{1000} \right) \right) \right. \\ &\quad \left. - 1 \right) \times 1000 \end{aligned}$$

Equation 4

$$\begin{aligned} \sigma(\delta^{18}\text{O}_{\text{VSMOW Spl}}) &= \sqrt{\left(\frac{\partial F}{\partial V_1} \times dV_1 \right)^2 + \left(\frac{\partial F}{\partial V_2} \times dV_2 \right)^2 + \left(\frac{\partial F}{\partial V_3} \times dV_3 \right)^2} \\ \text{With } F &= \delta^{18}\text{O}_{\text{VSMOW Spl}}, V_1 = \delta^{18}\text{O Spl}_{\text{measured}}, V_2 = \\ &\delta^{18}\text{O Std}_{\text{measured}} \text{ and } V_3 = \delta^{18}\text{O}_{\text{VSMOW Std}} \end{aligned}$$

Equation 5

$$\delta^{18}\text{O}_{\text{VSMOW}} = 1.0392 \times \delta^{18}\text{O}_{\text{VPDB}} + 30.92 \text{ ‰}$$

4. Results

Here we present some raw data from the SIMS to illustrate the measurements of single radiolarian molds and their validation. Raw results from repetitive measurements on some samples are also described to highlight reproducibility of our analyses. Finally, we present the $\delta^{18}\text{O}_{\text{VSMOW}}$ trends from our measurements through the Mesozoic.

4.1 Quality 234 check of SIMS analyses

We validate the quality of the SIMS measurement by post-checking of analysed spots by optical and/or SEM, and relative analytical yield. Illustrating the necessity of the pre- and post-checking, some spherical zones on the

Late Triassic Inuyama chert sample Ki48 (Fig. 3A and B) are difficult to be identified as radiolarian molds or epoxy, and have been measured with very low analytical yield (~3 % of the UNIL-Q1 yield) and very light $\delta^{18}\text{O}_{\text{RAW}}$ (uncalibrated $\delta^{18}\text{O}$) (<-30 ‰) (Appendix Table A2). The normal analytical yield and $\delta^{18}\text{O}_{\text{RAW}}$ of this sample (e.g. Fig. 3C) are ~97.8 % and ~25 ‰, respectively, with calibrated $\delta^{18}\text{O}_{\text{VSMOW}}$ of ~31.3 ‰.

The measurements of mixings of epoxy and radiolarian molds for some analytical spots have low analytical yields and light $\delta^{18}\text{O}_{\text{RAW}}$ -values. Such value can be found in an Early Triassic siliceous mudstone sample Ki12s from Inuyama (e.g. Fig. 3E). The detection of potential analytical sites in this sample is particularly difficult due to scarcity of radiolarian molds. Alternatively, we also commonly observed quartz patches, which could be the result from radiolarian silica precipitation, aleatory distributed that we tried to analyse (Fig. 3D). The post checking of these patches was nearly impossible due to their small size (smaller or equal to the 10 μm beam size) and we exclude them because they have usually low analytical yield (<90 %). The measurement of one of these patches has however a relatively good yield (94.9 %) but an uncommon light $\delta^{18}\text{O}_{\text{RAW}}$ (2.8 ‰), corresponding to a calibrated $\delta^{18}\text{O}_{\text{VSMOW}}$ of 8.7 ‰ (Fig. 3D).

High analytical $\delta^{18}\text{O}$ -yields were measured in the Middle Jurassic Sogno sample So29 (~130.2 ‰). SEM-EDX analyses demonstrate that the molds in this sample were filled with carbonates, and that the matrix is more siliceous (Fig. 4). This observation contrasts with the radiolarian molds being commonly more siliceous than the matrix, which contains clays (Fig. 5 and Fig. 6).

4.2 Reproducibility of SIMS measurements for radiolarian $\delta^{18}\text{O}$

Three chert samples (Pi01, Ki08C, Ca01) have been initially measured in detail relative to the NIST-610 standard, and then together relative to UNIL-Q1 standard (Table 3). NIST-610 was used to correct instrumental drift, but not to calibrate $\delta^{18}\text{O}$ measurements because the instrument fractionation depends on matrix and structure of the material analysed, mainly chert in this study. Overall, reproducibility on the UNIL-Q1 standard was 0.3 ‰ for $\delta^{18}\text{O}$ over all SIMS sessions (Table 1).

In detail, $\delta^{18}\text{O}$ -measurements on Pi01 (a Middle Jurassic radiolarite sample in the southern Switzerland) with NIST-610 were done during two sessions. Uncalibrated $\delta^{18}\text{O}$ -results for Pi01 after drift correction for the first and second sequences are $28.2 \pm 1.5 \text{ ‰}$ (10 points, 2SD) and $28.4 \pm 5.1 \text{ ‰}$ (27 points, 2SD) with NIST-610 values at $10.4 \pm 0.2 \text{ ‰}$ (5 points, 2SD) and $10.8 \pm 0.4 \text{ ‰}$ (16 points, 2SD), respectively. $\delta^{18}\text{O}$ -measurements on Ki08C (a Middle Triassic chert in Inuyama) and Ca01 (a diagenetic chert lacking radiolarian molds in the Cretaceous Maiolica Formation of the Lombardian basin) were measured relative to NIST-610 during single session. Uncalibrated $\delta^{18}\text{O}$ -results for Ki08C after drift correction

Table 2 List of samples with their age, the number of measurements, their $\delta^{18}\text{O}$ least square mean (LS-mean) and their $\delta^{18}\text{O}$ least square standard deviation (LS std; 1SD). The $\delta^{18}\text{O}$ was average with a 10 Ma windows moving average (5 Ma step) and compared with $\delta^{18}\text{O}$ from low magnesium carbonate shells for tropical and temperate regions (Grossman, 2012).

Sample	Age (Ma)	Number of measurements	Results		Curves			
			$\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)		$\delta^{18}\text{O}_{\text{VSMOW}}$ Radiolarian molds (‰)		$\delta^{18}\text{O}_{\text{VPDB}}$ LMC shells (‰)	
			This study		This study		Grossman, 2012	
			LS-mean	LS-std	Panthalassa	Tethys	Tropical	Temperate
10 Ma moving average					4 Ma moving average			
Pe01	253.15	10	24.0	0.2	27.21		-3.0	
Ki09	250.30	9	29.3	0.4	28.15			
Ki01	248.00	10	30.9	1.4	28.57			
Ki02	247.80	6	30.8	1.0	28.60			
Ki03s	247.20	4	20.9	2.3	28.69			
Ki03s	247.20	5	23.9	1.1	28.69			
Ki03	247.20	10	30.4	0.7	28.69			
Ki04	246.60	9	31.3	0.9	28.79			
Ki06	245.25	6	31.9	1.2	29.00		-4.0	
Ki07	245.00	6	29.3	1.8	29.04		-4.1	
Ki08	244.90	10	30.6	1.0	29.06		-4.1	
Ki08	244.90	86	30.8	0.6	29.06		-4.1	
Ki58C	243.48	6	31.2	0.7	29.24		-4.3	
Ki57	241.00	9	26.8	1.8	29.56		-3.8	
Ki15	228.00	10	31.5	0.7	31.49		-2.5	
Ki51	219.00	6	31.9	0.3	31.35		-2.1	
Ki48	217.09	3	30.6	0.9	31.52		-2.1	
Ki46	210.27	10	32.6	1.6	32.55		-1.8	
Ki44	204.82	9	31.6	0.4	31.48		-1.7	
Ki43s	204.82	6	31.3	0.4	31.48		-1.7	
Ki39	201.50	9	31.4	0.5	31.48			
Ki40	201.50	10	31.7	0.6	31.48			
Ki41	201.50	10	31.4	0.8	31.48			
Ki42	201.50	10	31.4	0.5	31.48			
Ki35	185.62	9	30.0	0.4	28.92		-0.8	
Ki34s	184.37	1	30.8	0.0	28.66		-1.0	
Ki34s	184.37	8	25.8	1.4	28.66		-1.0	
Ki32	184.20	10	30.2	0.9	28.62		-1.1	
Ki24	182.00	6	30.8	0.7	28.22		-1.9	
Ki25	180.66	8	25.2	2.5	27.97		-2.4	
Ki22c1	178.00	9	27.0	1.2	26.90		-2.7	
Ki22c2	178.00	6	26.1	1.5	26.90		-2.7	
Ki21	178.00	10	27.2	1.1	26.90		-2.7	
Ki20	174.00	5	21.9	1.9	24.76		-2.3	
So07	169.50	6	32.5	0.8		33.53	0.3	
So08	169.50	7	34.0	0.4		33.53	0.3	
So09	169.30	6	34.3	0.4		33.54	0.3	
So12	168.27	6	35.0	0.5		33.63	0.3	
So13	168.13	9	34.3	0.4		33.64	0.3	
So15	167.54	6	34.7	0.5		33.69	0.4	
So17	166.89	9	32.7	2.6		33.75	0.5	
So21	166.09	6	32.3	1.0		33.81	0.5	
Ro01	166.00	6	30.9	0.8		33.82	0.5	
So23	165.10	6	34.0	0.6		33.90	0.4	
So25	164.31	9	35.3	0.3		34.06	0.3	
Pi01	163.00	32	34.6	0.4		34.35	-0.1	
So26	162.65	6	35.1	0.3		34.43	-0.2	
So28	161.14	8	34.7	0.7		34.76	-0.4	
So31	159.97	9	35.2	0.3		35.02	-0.5	
Ca01	139.00	10	34.3	0.2		34.36	-0.3	
Go68	131.70	6	26.5	0.6	26.49		-0.3	
Go73	97.95	9	28.5	1.0	24.09		-2.6	
Go74	97.59	10	19.8	0.7	24.09		-2.8	

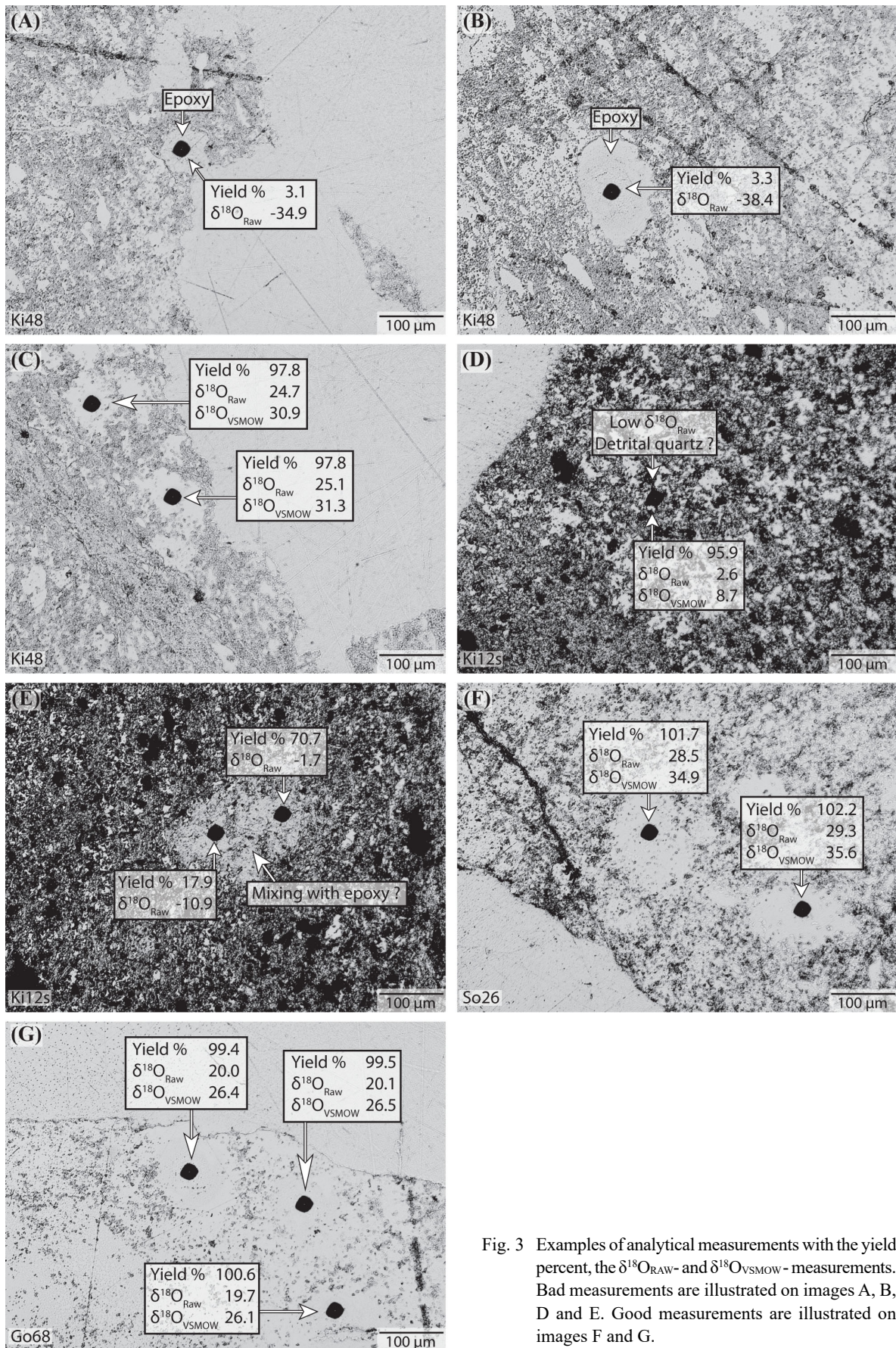


Fig. 3 Examples of analytical measurements with the yield percent, the $\delta^{18}\text{O}_{\text{RAW}}$ - and $\delta^{18}\text{O}_{\text{VSMOW}}$ -measurements. Bad measurements are illustrated on images A, B, D and E. Good measurements are illustrated on images F and G.

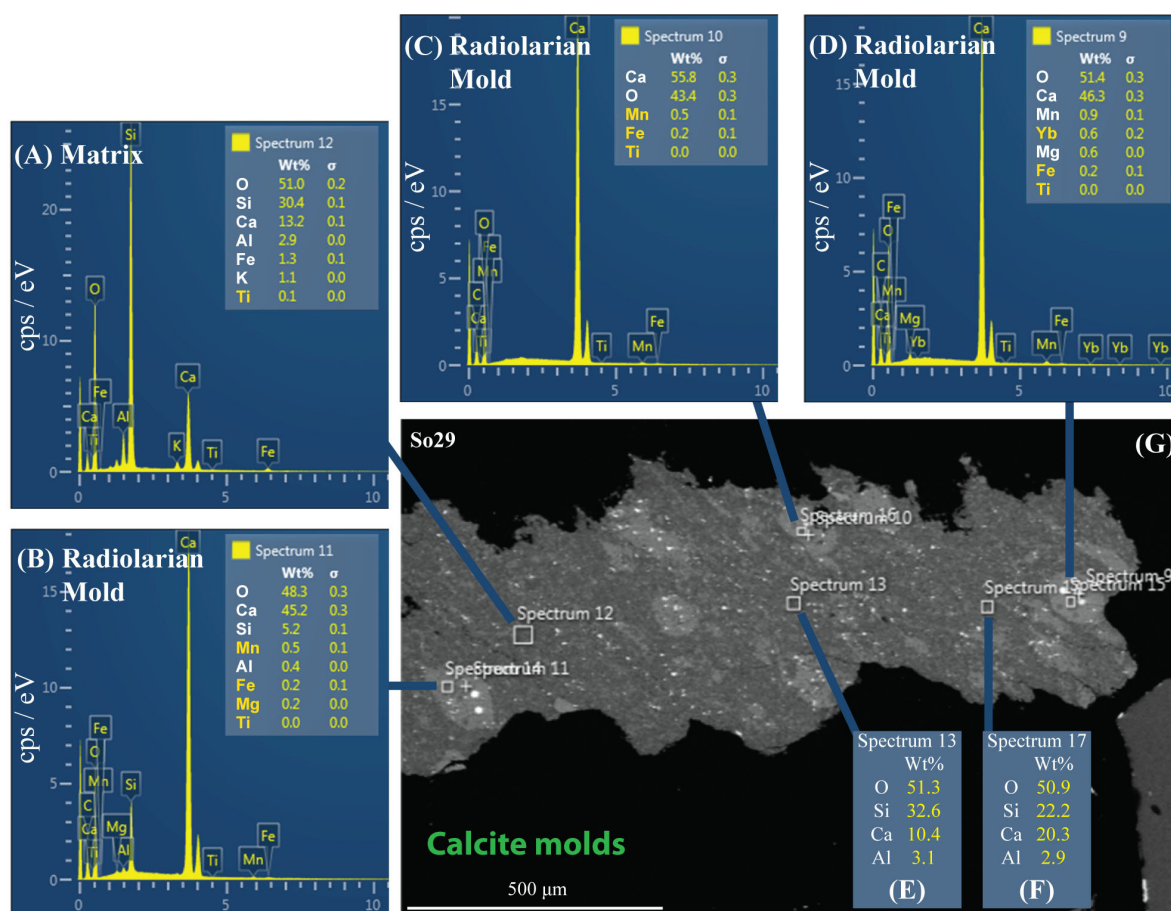


Fig. 4 EDX analyses of sample So29 with SEM. The EDX spectrums prove that the molds were filled with calcites in this sample (B, C and D) during diagenesis. The $\delta^{18}\text{O}$ - and $\delta^{30}\text{Si}$ -yields were respectively about 130.2 % and 11 % of the UNIL-Q1 quartz standard yield. The matrix is more siliceous but also included some carbonates (A, E, F). BSE image is show in (G).

are 25.4 ± 1.8 ‰ (27 points, 2SD) for radiolarians, 23.6 ± 1.2 ‰ (3 points, 2SD) for the matrix and 10.5 ± 0.4 ‰ (16 points, 2SD) for NIST-610. Uncalibrated $\delta^{18}\text{O}$ -results for Ca01 after drift correction are 29.0 ± 1.6 ‰ (11 points, 2SD) for sample and 10.6 ± 0.2 ‰ (8 points, 2SD) for NIST-610.

The samples Pi01, Ca01 and Ki08C were then measured together during a single session relative to UNIL-Q1 (PRIM123) to calibrate oxygen isotopes correcting instrumental mass fractionation due to the mineralogy of samples (chemical composition and crystal structure). Uncalibrated $\delta^{18}\text{O}$ -results after drift correction for this UNIL-Q1 session are 6.5 ± 0.3 ‰ (48 points, 2SD) for UNIL-Q1, 31.2 ± 0.7 ‰ (32 points, 2SD) for Pi01, 27.4 ± 1.2 ‰ (86 points, 2SD) for Ki08C, and 30.9 ± 0.4 ‰ (10 points, 2SD) for Ca01, respectively. Measurements of the base, middle and top of samples Ki08C are 27.3 ± 0.6 ‰ (4 points, 2SD), 27.9 ± 1 ‰ (28 points, 2SD) and 27.0 ± 1 ‰ (15 points, 2SD), respectively, with the average value of 27.4 ± 1.2 ‰ (86 points, 1SD).

Between the NIST-610 and the UNIL-Q1 sessions, we observed an $\delta^{18}\text{O}$ -offset of $\sim 2.4 \pm 1.2$ ‰ for the uncalibrated $\delta^{18}\text{O}$ of the samples Pi01, Ki08C and Ca01.

However, the $\delta^{18}\text{O}$ -results for Ki08C are systematically lighter of about 3.3 ± 0.9 ‰ (2SD) than the two other samples (Pi01 and Ca01). Moreover, it is interesting to note that UNIL-Q1 has an uncalibrated $\delta^{18}\text{O}$ of 6.5 ± 0.3 ‰ during the PRIM123 session whereas UNIL-Q1 has an average value of 4.0 ± 0.7 ‰ during the other session, and that correspond relatively well with the $\sim 2.4 \pm 1.2$ ‰ offset detected (Table 2).

4.3 Mesozoic $\delta^{18}\text{O}_{\text{radiolarians}}$ trends and fluctuations

Here, we present the $\delta^{18}\text{O}_{\text{VSMOW}}$ trends after having checked and calibrated our results (Fig. 7 and Fig. 8). The Early Triassic to Early Jurassic Panthalassan samples from the Inuyama area have least square means (LS-means) $\delta^{18}\text{O}_{\text{VSMOW}}$ range from 20.9 ‰ to 32.6 ‰. The Early to Middle Triassic LS-means $\delta^{18}\text{O}_{\text{VSMOW}}$ range from 20.9 ‰ to 31.9 ‰ with a slightly increasing trend. The Late Triassic LS-means $\delta^{18}\text{O}_{\text{VSMOW}}$ range has relatively low variation between 30.6 ‰ to 32.6 ‰. The Early Jurassic LS-means $\delta^{18}\text{O}_{\text{VSMOW}}$ range from 21.9 ‰ to 30.8 ‰ with a decreasing trend.

In the Middle Jurassic Lombardian basin of the Tethys, the LS-means $\delta^{18}\text{O}_{\text{VSMOW}}$ values are relatively high from

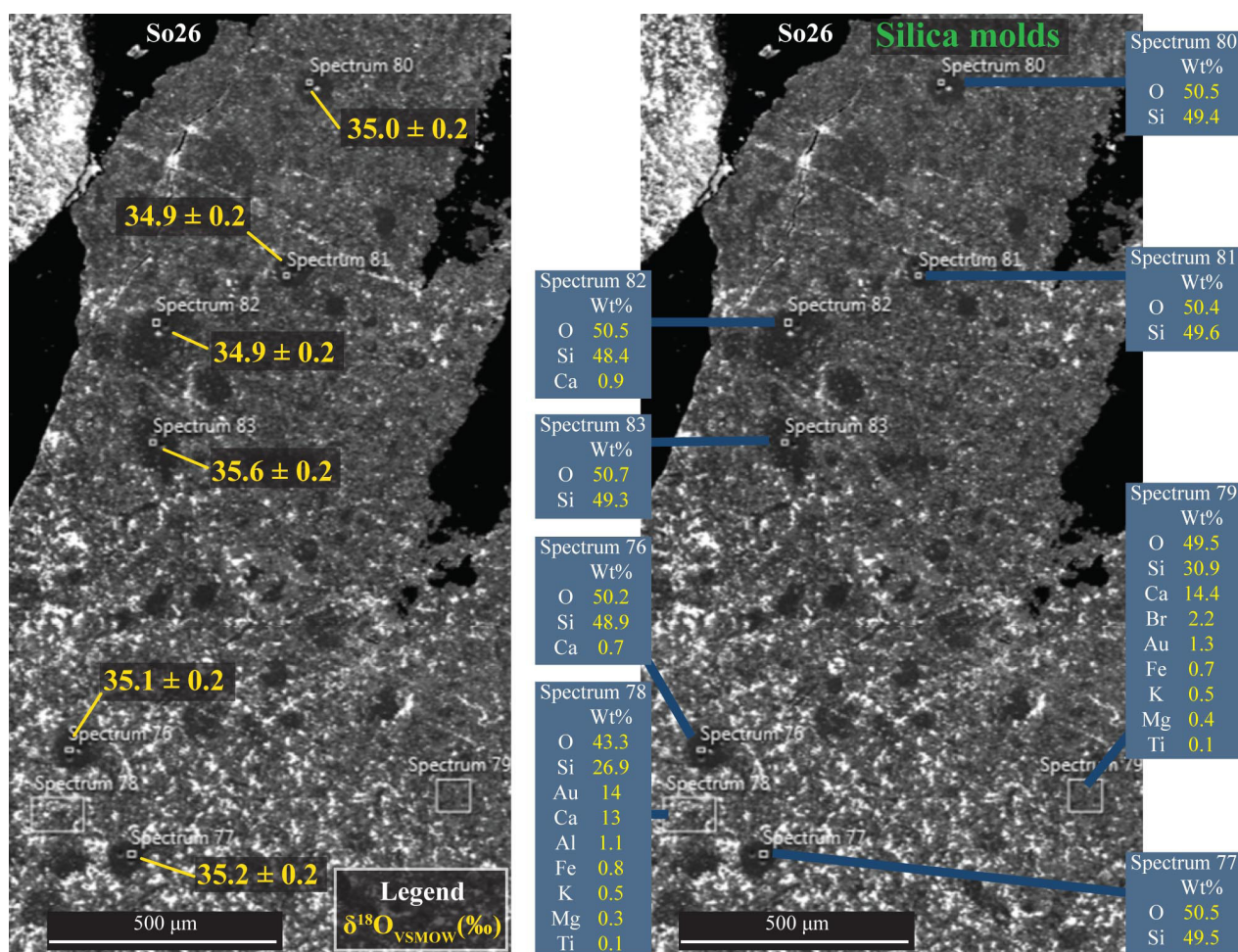


Fig. 5 Comparison between $\delta^{18}\text{O}_{\text{VSMOW}}$ and the chemical composition estimated by SEM-EDX spectrum in sample So26.

32.3 ‰ to 35.3 ‰. Our Middle Jurassic radiolarite sample from southern Switzerland (Pi01) has a LS-mean of 34.6 ‰, which is coherent with the other $\delta^{18}\text{O}$ range from the Sogno section in north Italy. A Berriasian diagenetic chert (Ca01) from the Lombardian basin has similar LS-means of 34.3 ‰. However, our Middle Jurassic Romanian sample (Ro01) has lighter $\delta^{18}\text{O}_{\text{VSMOW}}$ (30.9 ‰).

For the Cretaceous, the $\delta^{18}\text{O}_{\text{VSMOW}}$ values of the Goshikigahama samples are relatively light. A Hauterivian sample has a LS-mean 26.5 ‰, and two other Cenomanian samples have LS-means of 28.5 ‰ and 19.8 ‰.

5. Discussion

5.1 Quality of the SIMS analysis

To evaluate quality of SIMS results, we examined analytical yields relative to UNIL-Q1 yield, in addition to optical check. We rejected analytical spots with low yield values, generally <90 ‰, potentially the measurement of a mixing with epoxy or other minerals. Even if the spot analysed of the Early Triassic Inuyama shale sample Ki12s on Fig. 3D was discarded after optical control, it is interesting to note that it has one of

the best analytical yield (94.9 ‰) of the spots measured in Ki12s. In this sample, radiolarian molds are scarce, but quartz patches are common. The quartz patches are potentially precipitated radiolarian silica, but could be a diagenetic product. Its $\delta^{18}\text{O}_{\text{RAW}}$ is very light (2.8 ‰), corresponding to a calibrated $\delta^{18}\text{O}_{\text{VSMOW}}$ of 8.7 ‰. Such light value is common for igneous quartz (e.g. Seitz *et al.*, 2016), metamorphic quartz or eventually diagenetic quartz (cf. Bindeman, 2008), but not for radiolarian molds (e.g. Viswanathan and Mahabaleswar, 2014). We thus discarded this measurement.

The chemical composition of some excluded molds was also checked by SEM-EDX. SEM-EDX shows calcium as the first major element of the radiolarian molds in chert sample So29 from Sogno section (Fig. 4), and too high analytical $\delta^{18}\text{O}$ -yields (~130.2 ‰) were measured in this sample. Therefore, the chemical composition of these radiolarian molds is necessary to be checked before interpretation of isotopic data. In total, 53 samples and 507 measurements passed our checking from the initial 55 samples and 532 measurements.

Calibration of our measurements requires to know the nature of our samples (mineralogy and chemical

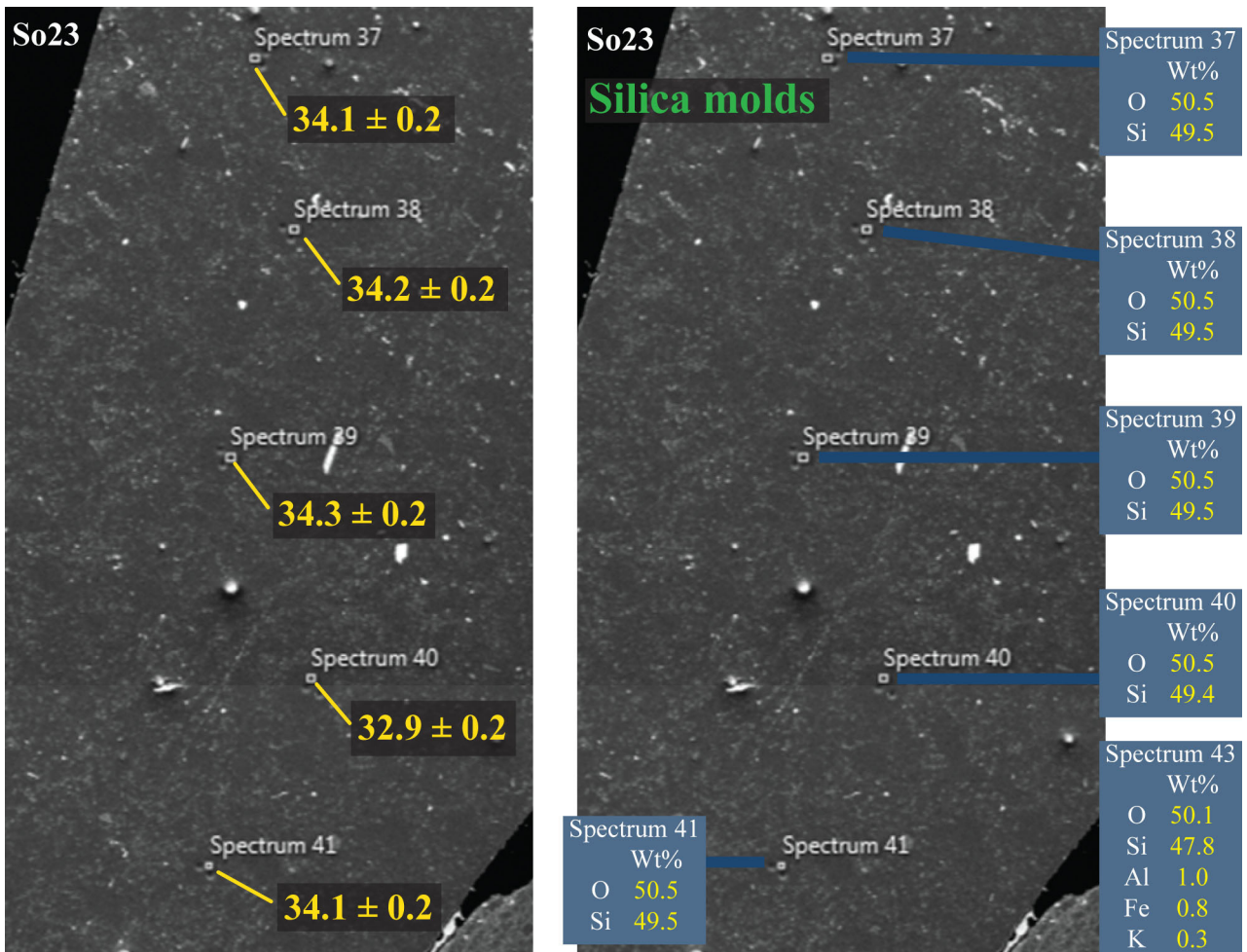


Fig. 6 Comparison between $\delta^{18}\text{O}_{\text{VSMOW}}$ and the chemical composition estimated by SEM-EDX spectrum in sample So23.

composition), because it affects the instrumental fractionation. If we correct the offset of $2.4 \pm 1.2 \text{ ‰}$ between the NIST-610 glass standard and the UNIL-Q1 quartz standard using the initial chert samples (Pi01, Ki08C, and Ca01) as references, values of the UNIL-Q1 standards should be about 6.4 ‰ lighter than the NIST standard. This difference is in contradiction with absolute value of UNIL-Q1 standard ($9.81 \pm 0.14 \text{ ‰}$, VSMOW, 2SD) and of NIST-610 standard (10.79 ‰ , VSMOW following GEOREM database from Jochum *et al.*, 2005), and might be explained by different instrumental fractionations between quartz and silica glass (amorphous material). Similar instrumental fractionations having already been inferred between microquartz, filling radiolarian molds, and quartz (Marin *et al.*, 2010), it is likely that the instrumental fractionation is different with silica glass. Higher uncalibrated $\delta^{18}\text{O}$ values for radiolarian molds ($25.4 \pm 1.8 \text{ ‰}$, 2SD) than for the matrix ($23.6 \pm 1.2 \text{ ‰}$, 2SD) in sample Ki08C might also be the result of a different instrumental fractionation between pure quartz in radiolarian molds and the mineral mixture in the matrix (matrix effect).

The $\delta^{18}\text{O}$ -values relative to VSMOW calibrated

by UNIL-Q1 are $34.55 \pm 0.4 \text{ ‰}$ (32 points, 1SD) for Pi01, $30.75 \pm 0.61 \text{ ‰}$ (86 points, 1SD) for Ki08C and $34.33 \pm 0.23 \text{ ‰}$ (10 points, 1SD) for Ca01. These intra-sample and inter-sample $\delta^{18}\text{O}$ -variability in radiolarite molds are larger than instrumental resolution ($<0.3 \text{ ‰}$, 2SD), indicating that our measurements represent some signatures in addition to instrumental uncertainties.

Our $\delta^{18}\text{O}_{\text{VSMOW}}$ ranges from LS-means (19.8 ‰ to 35.3 ‰ ; Fig. 8) is also consistent with the 18 ‰ to 38 ‰ of bulk $\delta^{18}\text{O}_{\text{VSMOW}}$ of Phanerozoic cherts by Knauth (1973). This further supports the validity of our SIMS $\delta^{18}\text{O}$ measurements and our calibration routine.

5.2 SIMS $\delta^{18}\text{O}$ from radiolarian molds and diagenetic effects

We here discuss some of the factors influencing the $\delta^{18}\text{O}_{\text{radiolarians}}$ based on our data and comparison with other dataset. Numbers of factors could have influenced on $\delta^{18}\text{O}_{\text{radiolarians}}$ in Mesozoic cherts, including such as: 1) temperature and $\delta^{18}\text{O}$ of seawater for oxygen in original radiolarian opal-A as referred for diatoms (e.g. Juillet-Leclerc and Labeyrie, 1987; Brandriss *et al.*, 1998), 2) Vital effects (Swann *et al.* 2007), 3) Dissolution and

Table 3 Comparison of our SIMS $\delta^{18}\text{O}$ least square means and standard deviations (in ‰; 2SD) for the samples Pi01, Ki08C and Ca01 between their NIST-610 and UNIL-Q1 sessions. Raw $\delta^{18}\text{O}_{\text{VSMOW}}$ of our samples are slightly shifted between the NIST-610 and UNIL-Q1 sessions due to changes of the analytical parameters but they conserve a similar pattern (~2.5 ‰). Least square standard deviations do not change between raw and drift corrected $\delta^{18}\text{O}_{\text{VSMOW}}$ indicating that there were no major instrumental drifts during our sessions.

Sample	Points	Yield (CPS/nA)			Raw $\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)		Drift corrected $\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)		Calibrated $\delta^{18}\text{O}_{\text{VSMOW}}$ (‰)	
		Mean	Std	Yield % relative to NIST-610	LS-Mean	LS-Std	LS-Mean	LS-Std	LS-Mean	LS-Std

NIST Sessions	NIST-610	5	9.37E+08	4.55E+06		10.5	0.2	10.4	0.2
	Pi01	10	9.65E+08	6.18E+06	102.9%	28.2	1.5	28.2	1.5
	NIST	16	9.28E+08	3.43E+07		10.8	0.4	10.8	0.4
	Pi01	27	9.49E+08	3.29E+07	102.2%	28.4	5.1	28.4	5.1
	NIST-610	16	9.04E+08	7.96E+06		10.5	0.4	10.5	0.4
	Ki08C	27	9.17E+08	1.65E+07	101.5%	25.4	1.8	25.4	1.8
	Matrix	3	8.96E+08	1.49E+07	99.1%	23.6	1.2	23.6	1.2
	NIST-610	8	9.41E+08	1.65E+07		10.6	0.2	10.6	0.1
	Ca01	11	9.52E+08	6.01E+06	101.1%	29.0	1.6	29.0	1.6

UNIL-Q1 Session	UNIL-Q1	48
	Ki08C base	4
	Ki08C med	28
	Ki08C top	15
	Total Ki08C	86
	Pi01	32
	Ca01	10

6.5	0.3	6.5	0.3	9.8	0.3
27.3	0.6	27.3	0.6	30.7	0.6
27.9	1.0	27.9	1.0	31.2	1.0
27.0	1.0	27.0	1.0	30.4	1.0
27.4	1.2	27.4	1.2	30.8	1.2
31.2	0.7	31.2	0.7	34.6	0.8
30.9	0.4	30.9	0.4	34.3	0.5

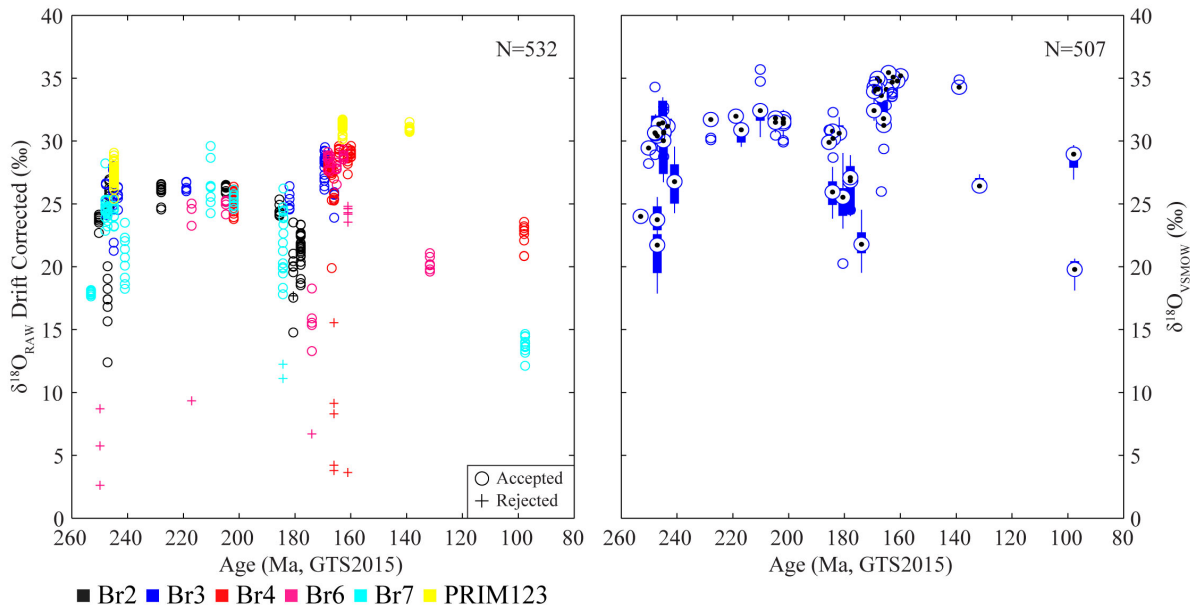


Fig. 7 $\delta^{18}\text{O}$ -measurements in function of sample mounts on left and boxplot for each sample on the right. For the measurements, each color corresponds to a sample mount showing that samples of similar age tend to converge toward similar $\delta^{18}\text{O}$. Cross markers correspond to analyses with lower analytical yields, uncommon internal analytical errors or which did not pass the microscopic checking. Based on these parameters, these values are not considered as valid and excluded for the boxplot diagrams and the average values of samples. Br2 to 7 and PRIM123 are the session names.

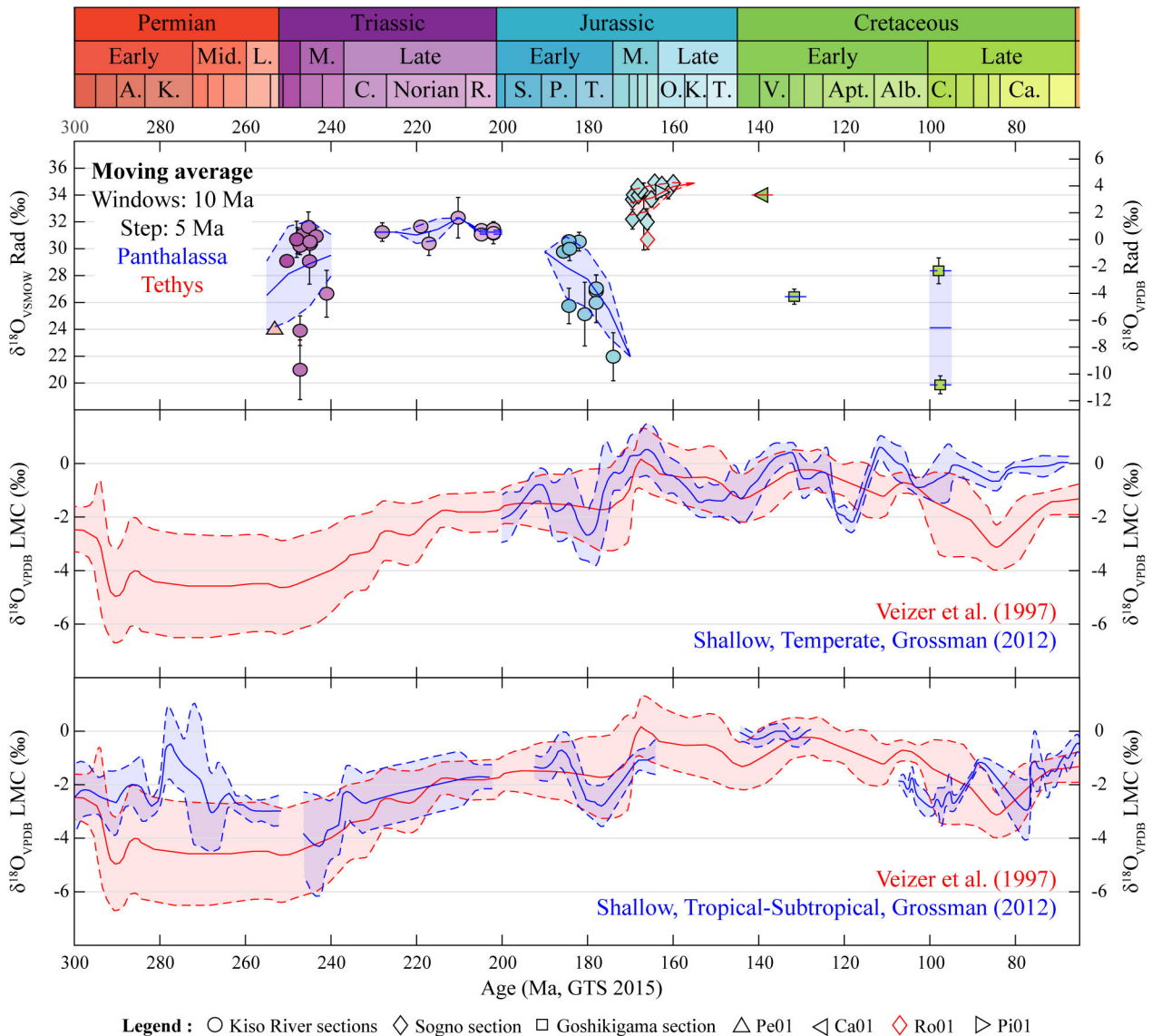


Fig. 8 Comparison between $\delta^{18}\text{O}$ from radiolarian silica ($\delta^{18}\text{O}$ Rad; this study) and digitalized curves for $\delta^{18}\text{O}$ from low magnesium calcite shells ($\delta^{18}\text{O}$ LMC; Veizer *et al.*, 1999 and Grossman, 2012), expressed either relative to the Vienna Standard Mean Ocean Water ($\delta^{18}\text{O}_{\text{VSMOW}}$) or to the Vienna Pee Dee Belemnite standard ($\delta^{18}\text{O}_{\text{VPDB}}$). The minimum values of $\delta^{18}\text{O}$ measured in radiolarian silica correspond to light values in the LMC curves during the Early Triassic, the Toarcian and eventually during the Early Cretaceous. In addition, the Middle Jurassic acme is also recorded with heavier $\delta^{18}\text{O}$ in the radiolarites. The geological timescale (GTS 2015) used for this figure is the timescale of the international commission of stratigraphy in 2015 based on Cohen *et al.* (2013). The color filling inside markers corresponds to the color of the geological stage. The boundaries of curves, when plotted, are equivalent to 1SD. The moving averages for radiolarian silica are realized separately for Panthalassa and Tethys using both a 10 Ma windows and 5 Ma steps. The $\delta^{18}\text{O}$ error bars correspond to the least square standard deviation of samples presented in Table 2.

dehydroxylation during settling and early diagenesis preferentially releasing light ^{16}O as for diatom frustules (see Schmidt *et al.* 2001; Moschen *et al.*, 2006; Swann and Leng, 2009), 4) Diagenetic temperature and isotopic composition of sediment pore water (e.g. Matheny and Knauth, 1993). In addition to these factors, we can not exclude that $\delta^{18}\text{O}$ is also influenced by 5) degree of diagenetic migration of opal-A during the phase transitions

from opal-A to quartz via opal-CT, which is mainly controlled by the solubilities of the different phases (e.g. Gunnarsson *et al.*, 2000) and 6) partially by kinetic isotopic fractionation associated with silica tetrahedrons, which at least occur for $\delta^{30}\text{Si}$ during silica precipitation under 200 °C (see He and Liu, 2015; Pollington *et al.*, 2016). Factors influencing the kinetic of the phase transitions from opal-A to quartz via opal-CT, such as temperature

(Ernst and Calvert, 1969; Dralus *et al.*, 2011) and pH of pore water associated with the presence of accessory minerals (c.f. Kastner *et al.*, 1977; Isaacs, 1981; Hinman, 1998) might thus also have some influences.

In the Cenozoic unconsolidated sediments, $\delta^{18}\text{O}_{\text{radiolarians}}$ ranges from 21 ‰ to 35 ‰ in equatorial Pacific (Wu *et al.*, 1997) and from 42 ‰ to 45 ‰ in Southern Ocean (e.g. Abelmann *et al.*, 2015), possibly related with different oceanographic setting. The former $\delta^{18}\text{O}_{\text{radiolarians}}$ from equatorial Pacific are within the range of our measurements on low latitude Mesozoic deep-sea bedded cherts (19.8 ‰ to 35.3 ‰). Contrary to ~6 ‰ changes in Cenozoic $\delta^{18}\text{O}_{\text{benthic foraminifera}}$ (Zachos *et al.*, 2008), different dwelling depth of radiolarians might be attributed to ~7 ‰ (Katz *et al.*, 2010; Xu *et al.*, 2012; Völpel *et al.*, 2017). Other factor is partial silica dissolution and isotopic enrichment of up to 6.8 ‰ for diatom (Moschen *et al.*, 2006). Because dissolution rate mainly depends on Si undersaturation and pH, fluctuation of $\delta^{18}\text{O}$ values amplified the temperature effect. Although contamination of siliciclastics with light $\delta^{18}\text{O}$ values, ranging from 10 to 20 ‰ (Eiler, 2001), cannot be rejected, similar ~10 ‰ large fluctuations in pelagic equatorial (paleo-) Pacific $\delta^{18}\text{O}_{\text{radiolarians}}$ during the Cenozoic and Mesozoic might imply that diagenetic $\delta^{18}\text{O}$ shift is insignificant relative to unconsolidated biosiliceous sediments, and that $\delta^{18}\text{O}$ from radiolarian molds still preserve an environmental signature after diagenesis.

The cause of spatial variations of $\delta^{18}\text{O}_{\text{radiolarians}}$ needs to be examined. Our equatorial paleo-Pacific $\delta^{18}\text{O}_{\text{radiolarians}}$ from Triassic-Early Jurassic Inuyama and Cretaceous Goshikigahama sections (~20 to 32 ‰) is slightly lighter than the low-middle latitude $\delta^{18}\text{O}_{\text{radiolarians}}$ from the Middle Jurassic Tethys regions (~30 to 35 ‰), although our data cannot reject the possibility of age difference (Fig. 8). In the Tethys region, oppositely, a higher latitude $\delta^{18}\text{O}_{\text{radiolarians}}$ from Romania is slightly lighter than other data from Italy (Fig. 8). Further spatio-temporal comparison will be needed to understand the nature of $\delta^{18}\text{O}_{\text{radiolarians}}$ in the past.

The intra-chert variability of $\delta^{18}\text{O}$ (0.16 to 2.49 ‰) larger than analytical errors could be related with different degree of diagenesis and/or original variations. Such $\delta^{18}\text{O}$ microvariations in cherts have also described from Precambrian chert (Marin *et al.*, 2010). During the transformation from opal-A to opal-CT, the diagenetic migration of opal-A from layers with low-Si content to layers with high-Si content could have homogenized $\delta^{18}\text{O}$ within a chert (Isaacs, 1981; Tada, 1991). Subsequent transformation from opal-CT to quartz could have also homogenized through similar mechanism. Nevertheless, significant $\delta^{18}\text{O}$ -variability in each chert suggests that such diagenetic segregation could not have perfectly homogenized $\delta^{18}\text{O}$ within each chert bed, and original variability might be larger.

If original, the intra-sample variations within cherts (0.16 to 2.49 ‰) corresponds to temperature differences of 2.3 to 40.9 °C, using equation from Brandriss *et*

al. (1998) with an initial seawater $\delta^{18}\text{O}$ of about 0 ‰. A temperature difference of 40.9 °C is certainly too high to reflect original $\delta^{18}\text{O}$ from radiolarian skeletons, suggesting possible changes in degree of diagenesis and/or paleoceanographic condition, such as global ice volume, precipitation/evaporation, and/or upwelling of deep-water. Because most of the Mesozoic is ice-free in polar region (e.g. Frakes *et al.*, 1992) and because Triassic-Jurassic Inuyama samples were deposited under equatorial pelagic Panthalassa, changes in upwelling of deep-water, dwelling depth of radiolarians, partial silica dissolution, and diagenesis might be likely cause of the intra-sample variations. The sedimentary rhythms of bedded chert are hypothesized to be linked with periodic changes in upwelling intensity and associated radiolarian productivity based on the systematic changes in cosmic spherule content between chert and mudstone (e.g. Hori *et al.*, 1993). The average duration of a chert-shale couplet is ~20 kyr based on radiolarian and conodont biostratigraphic age model in the early Mesozoic bedded chert sequence in the Inuyama area, Japan, which is consistent with a dominant periodicity of a precession cycle (Hori *et al.*, 1993; Ikeda *et al.*, 2010; Ikeda and Tada, 2014). Therefore, <20 kyr-scale oscillation in radiolarian $\delta^{18}\text{O}$ might be preserved in bedded chert to some extent after diagenesis. Further detailed works are necessary to understand this issue.

5.3 Mesozoic $\delta^{18}\text{O}$ trends of radiolarian silica and other proxies

$\delta^{18}\text{O}$ of low magnesium carbonate shells (LMC shells) are widely used for paleo- $\delta^{18}\text{O}$ of seawater during the Paleozoic and Mesozoic because of their less diagenetic overprint (e.g. Veizer *et al.* 1999; Grossman, 2012). The Mesozoic $\delta^{18}\text{O}$ of LMC shells from Tethys ocean shows a ~2 ‰ positive excursion during the Early-Middle Triassic, a relatively stable plateau during the Late Triassic, a large variation with <10-Myr variations during the Early Jurassic, a slight positive excursion during the Middle Jurassic (Fig. 8; Veizer *et al.* 1999; Grossman, 2012).

Although our SIMS measurement of $\delta^{18}\text{O}_{\text{radiolarians}}$ is low-resolution, our LS-mean data also shows a positive excursion during the Early-Middle Triassic, and relatively stable plateau during the Late Triassic, but up to ~8 ‰ negative excursion is not present in $\delta^{18}\text{O}$ curves of LMC shells (Fig. 8). Subsequent ~2 ‰ positive excursion during the Middle Jurassic seems to be also similar (Fig. 8). Although <10 Myr scale variability in our data is also large, similar repetitive measurements on sample Pi01 (uncalibrated LS-means of 28.2 ± 1.5 ‰ and 28.4 ± 5.1 ‰ from 10 and 27 measurements, respectively (2SD)) and Ki08C (27.3 ± 0.6 ‰, 27.9 ± 1.0 ‰, and 27.0 ± 1.0 ‰ from 4, 28, and 15 measurements (2SD)) imply that heterogeneities are well distributed in each chert and that chert could be relatively homogenous at bigger scale. Similar $\delta^{18}\text{O}$ trends of LMC shells and radiolarians might imply similar factors controlling these $\delta^{18}\text{O}$ records during

the Triassic and the Middle Jurassic (Fig. 8).

The cause of the ~8 ‰ negative excursion of radiolarian $\delta^{18}\text{O}$ during the Early Jurassic is unclear. Because our data is radiolarian silica of equatorial Panthalassa, various factors can be related, such as differences in paleoceanographic setting, temperature and pH of water column and sediment pore water for early dissolution, diagenetic processes. Considering large scattering of LMC shells records (e.g. Veizer *et al.* 1999; Grossman, 2012), further high-resolution and multi-proxy works are needed to examine this issue.

The $\delta^{18}\text{O}$ of conodont apatite ($\delta^{18}\text{O}_{\text{conodont}}$) is recently used for paleoceanographic analysis, and also shows similar positive excursion during the Early-Middle Triassic, and relatively stable plateau during the Late Triassic (e.g. Trotter *et al.*, 2015). Unfortunately, $\delta^{18}\text{O}_{\text{conodont}}$ cannot be applied for post-Triassic successions due to the complete extinction of conodont at the end-Triassic extinction (e.g. Clark, 1983). The $\delta^{18}\text{O}_{\text{conodont}}$ could be useful to compare with $\delta^{18}\text{O}_{\text{radiolarians}}$ with high-resolution because conodonts co-occur with radiolarians in chert and other siliceous sediments. Further crosscheck with other $\delta^{18}\text{O}$ -signatures, such as conodont, have to be done to validate radiolarian silica as a paleoceanographic proxy.

6. Concluding remarks

Here we report the in situ $\delta^{18}\text{O}_{\text{VPDB}}$ -values from Mesozoic radiolarian molds show range between 19.8 to 35.8 ‰, which is consistent with that of modern and Cenozoic radiolarian tests from deep-sea core of equatorial Pacific (Wu *et al.*, 1997). Relatively large variability of $\delta^{18}\text{O}$ in intra-chert bed could support $\delta^{18}\text{O}$ of the Mesozoic radiolarian tests are not perfectly homogenized in a chert bed during the diagenetic segregation and through time.

The temporal changes in the $\delta^{18}\text{O}_{\text{VSMOW}}$ values from Mesozoic radiolarian molds show an Early-Middle Triassic slight positive excursion, a Late Triassic high plateau, an Early Jurassic negative excursion with up to 8 ‰, a Middle Jurassic slight positive excursion, and few light values for the Cretaceous. Although the Early Jurassic negative excursion is not consistent with the $\delta^{18}\text{O}$ trend of less-diagenetic low-Mg calcite shells in shallow marine Tethys, similar $\delta^{18}\text{O}$ trends among radiolarians, LMC shells, and conodonts during the Triassic and the Middle Jurassic imply potential preservation of an environmental component even after diagenesis of biogenic silica. Further crosscheck with other $\delta^{18}\text{O}$ -signatures have to be done to validate radiolarian silica as a paleoceanographic proxy.

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Appendix

Table A1 Least square drift corrected and calibrated $\delta^{18}\text{O}$ measurements for each sample and for each accepted analyses. All $\delta^{18}\text{O}$ are relative to VPDB and all standard deviations are given as 2SD.

Sample	Pe01	Ki09	Ki12s	Ki01	Ki02	Ki03s	Ki03s	Ki03	Ki04	Ki06	Ki07	
age	253.15	250.30	249.9	248	247.80	247.20	247.20	247.20	246.60	245.25	245.00	
LS-mean ($\delta^{18}\text{O}_{\text{SSMOW}}$) and LS-Standard deviation (2SD) (%)	23.99 ± 0.33	29.27 ± 0.88		30.93 ± 2.81	30.83 ± 2.01	20.92 ± 4.61	23.92 ± 2.29	30.44 ± 1.38	31.27 ± 1.8	31.88 ± 2.32	29.25 ± 3.55	
Number of measurements	10	9	6	10	6	4	5	10	9	6	6	
Accepted measurements	10	9	0	10	6	4	5	10	9	6	6	
Rejected measurements			6									
$\delta^{18}\text{O}_{\text{SSMOW}}$ and analytical standard deviation (2SD) of all accepted measurements (%)	1	24.04 ± 0.18	29.2 ± 0.31		28.91 ± 0.2	29.59 ± 0.21	22.31 ± 0.32	24.55 ± 0.25	30.24 ± 0.2	30.04 ± 0.12	31.42 ± 0.13	30.39 ± 0.25
	2	23.86 ± 0.22	29.45 ± 0.19		30.53 ± 0.18	32.13 ± 0.19	22.9 ± 0.22	22.92 ± 0.19	30.87 ± 0.19	30.19 ± 0.19	30.81 ± 0.19	31.03 ± 0.24
	3	23.97 ± 0.2	29.48 ± 0.16		30.59 ± 0.13	30.58 ± 0.16	21.15 ± 0.2	22.9 ± 0.22	30.5 ± 0.13	30.5 ± 0.27	31.49 ± 0.23	27.39 ± 0.25
	4	23.99 ± 0.22	29.35 ± 0.18		31.29 ± 0.13	30.18 ± 0.15	17.87 ± 0.14	23.75 ± 0.24	31.37 ± 0.2	31.36 ± 0.17	33.2 ± 0.19	29.99 ± 0.24
	5	23.67 ± 0.25	29.54 ± 0.19		31.38 ± 0.26	30.57 ± 0.2		25.55 ± 0.21	30.09 ± 0.33	32.5 ± 0.21	30.93 ± 0.17	26.73 ± 0.21
	6	24.15 ± 0.19	29.68 ± 0.18		30.97 ± 0.25	32.02 ± 0.22			30.31 ± 0.18	31.15 ± 0.18	33.48 ± 0.17	30.08 ± 0.23
	7	23.79 ± 0.21	29.47 ± 0.19		34.3 ± 0.17				31.4 ± 0.26	31.6 ± 0.25		
	8	24.15 ± 0.24	28.21 ± 0.2		30.45 ± 0.23				29.7 ± 0.23	31.8 ± 0.18		
	9	24.07 ± 0.21	29.05 ± 0.17		29.96 ± 0.2				30.78 ± 0.17	32.36 ± 0.16		
	10	24.18 ± 0.23			30.72 ± 0.2				29.26 ± 0.13			
	11											
	12											
	13											
	14											
	15											
	16											
	17											
	18											
	19											
	20											
Session	Br7	Br2	Br6	Br7	Br3	Br2	Br2	Br7	Br2	Br3	Br3	
Sample order during the session	8	6	7	11	8	1	2	5	8	6	3	

Sample	Ki08	Ki08				Ki58C	Ki57	Ki15	Ki51	Ki48		
age	244.9	244.9				243.48	241.00	228.00	219.00	217.09		
LS-mean ($\delta^{18}\text{O}_{\text{SSMOW}}$) and LS-Standard deviation (2SD) (%)	30.58 ± 2.05	30.75 ± 1.21				31.16 ± 1.32	26.76 ± 3.62	31.49 ± 1.43	31.91 ± 0.68	30.61 ± 1.83		
Number of measurements	10	86				6	10	10	6	6		
Accepted measurements	10	86				6	9	10	6	3		
Rejected measurements							1			3		
$\delta^{18}\text{O}_{\text{SSMOW}}$ and analytical standard deviation (2SD) of all accepted measurements (%)	1	30.05 ± 0.23	30.9 ± 0.2	32.43 ± 0.17	30.12 ± 0.1	30.81 ± 0.18	30.34 ± 0.12	31.78 ± 0.2	24.27 ± 0.25	31.74 ± 0.18	31.76 ± 0.22	29.55 ± 0.21
	2	30.58 ± 0.18	30.64 ± 0.18	31.25 ± 0.23	29.77 ± 0.17	30.4 ± 0.16	30.01 ± 0.12	31.22 ± 0.24	26.78 ± 0.24	31.47 ± 0.2	32.23 ± 0.25	31.33 ± 0.19
	3	32.72 ± 0.22	30.89 ± 0.15	30.79 ± 0.22	30.11 ± 0.2	30.78 ± 0.1	30 ± 0.17	30 ± 0.24	27.45 ± 0.18	30.25 ± 0.2	31.49 ± 0.16	30.9 ± 0.18
	4	29.25 ± 0.23	30.25 ± 0.17	30.68 ± 0.17	30.3 ± 0.15	30.91 ± 0.11	30.9 ± 0.21	31.84 ± 0.22	26.13 ± 0.22	32.13 ± 0.26	32.2 ± 0.24	
	5	30.92 ± 0.31	31.73 ± 0.14	30.68 ± 0.16	30.09 ± 0.27	30.36 ± 0.12	30.54 ± 0.14	31.13 ± 0.17	29.56 ± 0.18	31.92 ± 0.16	31.63 ± 0.2	
	6	29.3 ± 0.2	31.02 ± 0.14	31.04 ± 0.16	30.35 ± 0.17	30.94 ± 0.15	30.26 ± 0.16	30.95 ± 0.21	28.09 ± 0.19	31.69 ± 0.25	32.25 ± 0.17	
	7	30.7 ± 0.19	30.31 ± 0.16	30.91 ± 0.2	29.84 ± 0.17	30.63 ± 0.14			24.64 ± 0.23	32.07 ± 0.25		
	8	31.35 ± 0.21	31.01 ± 0.2	30.95 ± 0.18	30.25 ± 0.1	30.08 ± 0.17			28.36 ± 0.26	31.88 ± 0.2		
	9	30.9 ± 0.21	30.96 ± 0.19	31.24 ± 0.12	30.49 ± 0.1	30.15 ± 0.12			25.16 ± 0.19	30.09 ± 0.26		
	10	30.08 ± 0.26	31.3 ± 0.22	30.52 ± 0.18	31.19 ± 0.18	30.23 ± 0.16				31.61 ± 0.24		
	11		31.22 ± 0.21	30.66 ± 0.17	30.64 ± 0.2	31 ± 0.18						
	12		31.16 ± 0.16	32.22 ± 0.14	30.32 ± 0.15	30.64 ± 0.14						
	13		30.83 ± 0.14	30.51 ± 0.17	30.54 ± 0.21	30.8 ± 0.14						
	14		31.48 ± 0.11	31.11 ± 0.23	30.34 ± 0.15	28.77 ± 0.16						
	15		31.25 ± 0.14	31.44 ± 0.19	29.51 ± 0.13	30.61 ± 0.16						
	16		31.31 ± 0.16	31.17 ± 0.14	31.93 ± 0.15	30.64 ± 0.14						
	17		31.74 ± 0.19	30.4 ± 0.18	30.88 ± 0.2	31 ± 0.12						
	18		31.7 ± 0.18	29.8 ± 0.17	31.79 ± 0.19	30.63 ± 0.17						
	19		31.81 ± 0.21	30.9 ± 0.18	30.42 ± 0.18	31.22 ± 0.16						
	20		31.77 ± 0.15	30.24 ± 0.16	30.58 ± 0.19	30.76 ± 0.15						
Session	Br7	PRIM123				Br3	Br7	Br2	Br3	Br6		
Sample order during the session	3	1				7	6	7	4	2		

Table A1 Continued.

Sample	Ki46	Ki43s	Ki44	Ki39	Ki40	Ki41	Ki42	Ki35	Ki34s	Ki34s	Ki32	
age	210.27	204.82	204.82	201.50	201.50	201.50	201.5	185.62	184.37	184.37	184.20	
LS-mean ($\delta^{18}\text{O}_{\text{SSMOW}}$) and LS-Standard deviation (2SD) (%)	32.59 ± 3.13	31.3 ± 0.89	31.65 ± 0.86	31.41 ± 1.05	31.72 ± 1.17	31.38 ± 1.57	31.42 ± 0.94	29.98 ± 0.81	25.82 ± 2.74	30.79 ± 0.3	30.21 ± 1.83	
Number of measurements	10	6	9	9	10	10	10	9	9	1	10	
Accepted measurements	10	6	9	9	10	10	10	9	8	1	10	
Rejected measurements									1			
$\delta^{18}\text{O}_{\text{SSMOW}}$ and analytical standard deviation (2SD) of all accepted measurements (%)	1	31.63 ± 0.22	31.47 ± 0.25	31.44 ± 0.19	30.79 ± 0.15	31.88 ± 0.17	31.46 ± 0.25	32.2 ± 0.19	29.95 ± 0.25	23.83 ± 0.12	30.79 ± 0.3	29.98 ± 0.17
	2	30.32 ± 0.19	31.57 ± 0.14	31.95 ± 0.2	31.51 ± 0.12	31.46 ± 0.24	32.49 ± 0.16	31.2 ± 0.18	30.47 ± 0.28	27.29 ± 0.21		30.43 ± 0.22
	3	32.43 ± 0.21	31.25 ± 0.22	32 ± 0.22	32.08 ± 0.24	32.28 ± 0.16	31.33 ± 0.18	30.66 ± 0.25	29.88 ± 0.17	25.5 ± 0.28		29.41 ± 0.16
	4	32.51 ± 0.15	31.52 ± 0.22	32.06 ± 0.24	32.16 ± 0.19	31.79 ± 0.23	31.84 ± 0.15	30.94 ± 0.18	29.72 ± 0.2	26 ± 0.21		30.55 ± 0.2
	5	34.75 ± 0.22	30.46 ± 0.18	30.69 ± 0.22	30.52 ± 0.17	32.12 ± 0.21	29.91 ± 0.19	30.88 ± 0.2	29.59 ± 0.26	27.94 ± 0.29		28.69 ± 0.25
	6	32.41 ± 0.14	31.56 ± 0.19	31.91 ± 0.19	31.39 ± 0.14	31.71 ± 0.26	30.34 ± 0.18	31.85 ± 0.19	29.65 ± 0.14	24.35 ± 0.26		29.96 ± 0.18
	7	32.47 ± 0.18		31.83 ± 0.19	31.69 ± 0.28	31.79 ± 0.2	31.28 ± 0.2	31.51 ± 0.19	29.92 ± 0.22	26.31 ± 0.19		30.44 ± 0.24
	8	31.2 ± 0.24		31.57 ± 0.19	31.35 ± 0.15	31.94 ± 0.14	31.9 ± 0.19	31.57 ± 0.15	29.9 ± 0.18	25.9 ± 0.2		32.3 ± 0.21
	9	32.38 ± 0.2		31.4 ± 0.18	31.55 ± 0.14	30.06 ± 0.24	31.14 ± 0.2	31.73 ± 0.22	30.89 ± 0.29			29.92 ± 0.21
	10	35.7 ± 0.18				31.76 ± 0.16	31.96 ± 0.19	31.56 ± 0.16				30.45 ± 0.24
	11											
	12											
	13											
	14											
	15											
	16											
	17											
	18											
	19											
	20											
Session	Br7	Br6	Br2	Br4	Br4	Br4	Br7	Br2	Br7	Br7	Br7	
Sample order during the session	10	4	9	3	6	8	4	11	2	2	9	

Sample	Ki24	Ki25	Ki22c2	Ki22c1	Ki21	Ki20	So07	So08	So09	So12	So13	
age	182.00	180.66	178.00	178.00	178.00	174.00	169.50	169.50	169.30	168.27	168.13	
LS-mean ($\delta^{18}\text{O}_{\text{SSMOW}}$) and LS-Standard deviation (2SD) (%)	30.76 ± 1.45	25.19 ± 4.92	26.07 ± 3.07	26.97 ± 2.44	27.16 ± 2.11	21.92 ± 3.73	32.47 ± 1.52	33.99 ± 0.82	34.35 ± 0.83	34.96 ± 0.9	34.31 ± 0.83	
Number of measurements	6	9	6	9	10	6	6	7	6	6	9	
Accepted measurements	6	8	6	9	10	5	6	7	6	6	9	
Rejected measurements		1				1						
$\delta^{18}\text{O}_{\text{SSMOW}}$ and analytical standard deviation (2SD) of all accepted measurements (%)	1	31.9 ± 0.23	23.04 ± 0.28	24.03 ± 0.25	26.42 ± 0.22	26.33 ± 0.21	19.52 ± 0.21	32.8 ± 0.29	34.01 ± 0.14	34.21 ± 0.22	34.93 ± 0.16	33.99 ± 0.26
	2	30.88 ± 0.22	25.56 ± 0.18	26.76 ± 0.31	27.1 ± 0.22	28.02 ± 0.19	21.61 ± 0.26	32.34 ± 0.14	34.74 ± 0.16	33.99 ± 0.15	35.19 ± 0.22	34.36 ± 0.19
	3	30.36 ± 0.2	25.96 ± 0.24	24.21 ± 0.22	28.86 ± 0.25	27.14 ± 0.14	21.79 ± 0.28	33.81 ± 0.22	33.89 ± 0.21	34.68 ± 0.31	34.01 ± 0.28	33.74 ± 0.24
	4	30.03 ± 0.26	26.56 ± 0.17	26.93 ± 0.2	26.33 ± 0.2	26.64 ± 0.2	22.14 ± 0.24	32.52 ± 0.18	33.95 ± 0.18	34.38 ± 0.14	35.05 ± 0.15	34.83 ± 0.18
	5	30.11 ± 0.17	20.26 ± 0.3	27.4 ± 0.26	24.53 ± 0.19	27.95 ± 0.19	24.53 ± 0.21	31.62 ± 0.18	34.15 ± 0.18	33.96 ± 0.14	35.44 ± 0.15	34.04 ± 0.13
	6	31.25 ± 0.19	29.06 ± 0.28	27.12 ± 0.21	27.36 ± 0.19	27.04 ± 0.22		31.89 ± 0.21	33.4 ± 0.17	35.02 ± 0.2	34.9 ± 0.19	34.13 ± 0.27
	7		25.07 ± 0.17		28.21 ± 0.24	28.88 ± 0.2			33.76 ± 0.22			33.96 ± 0.2
	8		25.51 ± 0.13		26.91 ± 0.26	25.91 ± 0.17						34.9 ± 0.12
	9				27.22 ± 0.13	28.12 ± 0.21						34.53 ± 0.21
	10					25.67 ± 0.2						
	11											
	12											
	13											
	14											
	15											
	16											
	17											
	18											
	19											
	20											
Session	Br3	Br2	Br2	Br2	Br2	Br6	Br3	Br3	Br3	Br6	Br4	
Sample order during the session	5	10	4	3	5	5	2	1	10	9	4	

Table A1 Continued.

Sample	So15	So17	So21	Ro01	So23	So25	P01	So26	So28	So29	
age	167.54	166.89	166.09	166.00	165.10	164.31	163.00	162.65	161.14	161.06	
LS-mean ($\delta^{18}\text{O}_{\text{VSMOW}}$) and LS-Standard deviation (2SD) (%)	34.68 ± 1.05	32.69 ± 5.18	32.27 ± 2.04	30.92 ± 1.54	34.02 ± 1.18	35.3 ± 0.65	34.55 ± 0.8	35.12 ± 0.55	34.71 ± 1.37		
Number of measurements	6	9	11	6	6	9	32	6	9	6	
Accepted measurements	6	9	6	6	6	9	32	6	8	0	
Rejected measurements			5						1	6	
$\delta^{18}\text{O}_{\text{VSMOW}}$ and analytical standard deviation (2SD) of all accepted measurements (%)	1	34.58 ± 0.18	33.73 ± 0.15	31.96 ± 0.25	31.2 ± 0.16	34.61 ± 0.2	35.74 ± 0.21	34.74 ± 0.15	34.69 ± 0.17	35.03 ± 0.2	35.38 ± 0.22
	2	34.99 ± 0.19	33.93 ± 0.28	31.49 ± 0.17	30.85 ± 0.13	34.06 ± 0.16	35.5 ± 0.25	33.54 ± 0.15	34.92 ± 0.1	34.85 ± 0.18	34.52 ± 0.2
	3	35.18 ± 0.2	33.44 ± 0.23	33.75 ± 0.16	29.38 ± 0.21	32.85 ± 0.2	35.23 ± 0.2	33.53 ± 0.2	34.56 ± 0.18	34.89 ± 0.2	34.81 ± 0.24
	4	33.92 ± 0.26	25.99 ± 0.21	31.39 ± 0.26	31.29 ± 0.22	34.32 ± 0.17	34.88 ± 0.19	34.73 ± 0.1	33.75 ± 0.14	35.62 ± 0.22	35.15 ± 0.17
	5	34.12 ± 0.21	31.42 ± 0.2	31.62 ± 0.2	31.42 ± 0.24	34.15 ± 0.16	34.91 ± 0.27	34.75 ± 0.16	34.55 ± 0.17	35.15 ± 0.22	34.75 ± 0.17
	6	35.15 ± 0.22	33.59 ± 0.13	33.15 ± 0.24	31.39 ± 0.21	34.07 ± 0.17	35.45 ± 0.17	34.59 ± 0.14	34.74 ± 0.1	35.22 ± 0.18	34.33 ± 0.19
	7		34.01 ± 0.17				34.89 ± 0.18	34.66 ± 0.12	34.48 ± 0.1		33.49 ± 0.13
	8		34.26 ± 0.24				35.5 ± 0.18	34.88 ± 0.13	35 ± 0.17		35.56 ± 0.2
	9		33.62 ± 0.17				35.55 ± 0.2	34.71 ± 0.2	34.33 ± 0.14		
	10							34.77 ± 0.15	34.62 ± 0.17		
	11							35.09 ± 0.22	34.84 ± 0.2		
	12							34.92 ± 0.15	34.73 ± 0.11		
	13							34.47 ± 0.16			
	14							34.62 ± 0.1			
	15							33.87 ± 0.13			
	16							33.89 ± 0.2			
	17							34.68 ± 0.13			
	18							34.72 ± 0.17			
	19							34.8 ± 0.16			
	20							34.43 ± 0.16			
Session	Br6	Br4	Br4	Br3	Br6	Br4	PRIM123	Br6	Br4	Br6	
Sample order during the session	1	2	10	9	3	7	2	8	1	6	

Sample	So31	Ca01	Go68	Go73	Go74	
age	159.97	139.00	131.70	97.95	97.59	
LS-mean ($\delta^{18}\text{O}_{\text{VSMOW}}$) and LS-Standard deviation (2SD) (%)	35.25 ± 0.57	34.33 ± 0.45	26.53 ± 1.16	28.52 ± 1.99	19.76 ± 1.4	
Number of measurements	9		6	9	10	
Accepted measurements	9		6	9	10	
Rejected measurements						
$\delta^{18}\text{O}_{\text{VSMOW}}$ and analytical standard deviation (2SD) of all accepted measurements (%)	1	35.42 ± 0.22	34.27 ± 0.14	27.06 ± 0.26	29.68 ± 0.2	18.12 ± 0.29
	2	35.09 ± 0.19	34.3 ± 0.12	25.89 ± 0.19	26.95 ± 0.18	19.61 ± 0.22
	3	35.76 ± 0.15	34.89 ± 0.17	26.39 ± 0.17	28.71 ± 0.14	20.49 ± 0.23
	4	34.87 ± 0.22	34.18 ± 0.17	26.47 ± 0.26	28.2 ± 0.15	19.36 ± 0.13
	5	34.9 ± 0.24	34.1 ± 0.13	26.06 ± 0.18	28.96 ± 0.18	19.16 ± 0.15
	6	35.1 ± 0.19	34.13 ± 0.13	27.36 ± 0.17	26.96 ± 0.17	20.03 ± 0.14
	7	35.19 ± 0.18	34.36 ± 0.19		29.33 ± 0.19	20.45 ± 0.19
	8	35.42 ± 0.16	34.28 ± 0.17		29 ± 0.22	19.89 ± 0.11
	9	35.32 ± 0.24	34.51 ± 0.2		29.19 ± 0.23	20.65 ± 0.22
	10		34.4 ± 0.19			19.7 ± 0.17
	11					
	12					
	13					
	14					
	15					
	16					
	17					
	18					
	19					
	20					
Session	Br4	PRIM123	Br6	Br4	Br7	
Sample order during the session	9	3	10	5	1	

Table A2 All the analytical dataset of the samples and standards. This data set include raw information such as the beam intensity (nA) during the analyses, counts per second on the L2 and H2 FC for ¹⁶O and ¹⁸O, respectively, time at which was realized (HH:MM). The raw isotopic ratios (H2/L2 (¹⁸O/¹⁶O)) is computed from the counts per second on the L2 and H2 FC and transformed into raw δ¹⁸O using the regular formula (δ¹⁸O = [(¹⁸O / ¹⁶O)sp / (¹⁸O / ¹⁶O)ref] - 1) with VSMOW as reference. For each standard and sample cluster, we calculated mean and standard deviation for raw δ¹⁸O. Raw δ¹⁸O were drifted corrected and calibrated using the procedure described in methods. The analytical yield (cps/nA) is calculated from the counts per second on the L2 FC and the beam intensity.

Beam id	SIMS ANALYSES										Mount: P01	Standard: UNIL-Q1 (Paine)			Analyse: δ ¹⁸ O			Value: 9.81±0.14 (VSMOW, 2σ)			Date: 06.12.2013	Comment
	H2/L2 (16O)	H2/L2 (18O)	12C (16C/12C)	12C (18C/12C)	H2 (16O/16O)	H2 (18O/16O)	Time	Yield	Yield	Yield		Mean	SD	Drift correction (‰)	Calibration (‰)	SD						
4180_NIST_051213_Ra@21	1.76	2.03E+03	6.74E+03	1.65E+09	1.69E+01	3.39E+06	1.71E+01	18.40	93703723.6	10.53	0.07	10.50	10.50	0.07	0.07	0.07	Standard					
4180_NIST_051213_Ra@22	1.75	2.03E+03	6.74E+03	1.65E+09	2.22E+01	3.29E+06	2.24E+01	18.46	92975404.2	10.44	0.09	10.42	10.42	0.09	0.09	0.09	Standard					
4180_NIST_051213_Ra@23	1.76	2.03E+03	8.94E+03	1.65E+09	2.19E+01	3.34E+06	2.21E+01	18.52	94029392.6	10.55	0.09	10.52	10.52	0.09	0.09	0.09	Standard					
4180_NIST_051213_Ra@24	1.75	2.03E+03	8.15E+03	1.65E+09	1.78E+01	3.34E+06	1.77E+01	18.57	94074702.5	10.26	0.08	10.25	10.25	0.08	0.08	0.08	Standard					
4180_NIST_051213_Ra@25	1.73	2.06E+03	8.40E+03	1.69E+09	1.88E+01	3.48E+06	1.87E+01	19.03	97583338.6	10.25	0.08	10.24	10.24	0.08	0.08	0.08	PrOI					
4180_NIST_051213_Ra@26	1.74	2.06E+03	8.93E+03	1.68E+09	1.88E+01	3.49E+06	1.86E+01	19.09	970639709.1	28.52	0.09	28.51	28.51	0.09	0.09	0.09	matrix					
4180_NIST_051213_Ra@27	1.74	2.06E+03	7.66E+03	1.68E+09	1.79E+01	3.46E+06	1.79E+01	19.14	965642159.9	28.34	0.07	28.34	28.33	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@28	1.73	2.06E+03	9.30E+03	1.68E+09	1.79E+01	3.47E+06	1.79E+01	19.20	96271924.0	28.39	0.09	28.39	28.39	0.09	0.09	0.09	PrOI					
4180_NIST_051213_Ra@29	1.74	2.06E+03	6.68E+03	1.68E+09	1.47E+01	3.47E+06	1.48E+01	19.25	958772118.2	28.61	0.08	28.61	28.62	0.08	0.08	0.08	PrOI					
4180_NIST_051213_Ra@30	1.72	2.06E+03	8.09E+03	1.68E+09	1.78E+01	3.47E+06	1.78E+01	19.31	964038027.7	26.02	0.06	26.03	26.03	0.06	0.06	0.06	matrix					
4180_NIST_051213_Ra@31	1.73	2.06E+03	7.24E+03	1.68E+09	1.78E+01	3.47E+06	1.78E+01	19.37	964038027.7	26.02	0.06	26.03	26.03	0.06	0.06	0.06	matrix					
4180_NIST_051213_Ra@32	1.71	2.06E+03	7.24E+03	1.68E+09	1.78E+01	3.47E+06	1.78E+01	19.48	97184331.5	28.58	0.07	28.58	28.60	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@33	1.73	2.06E+03	7.24E+03	1.68E+09	1.78E+01	3.47E+06	1.78E+01	19.54	96927526.0	28.13	0.08	28.13	28.16	0.08	0.08	0.08	PrOI					
4180_NIST_051213_Ra@34	1.74	2.03E+03	9.51E+03	1.64E+09	1.78E+01	3.32E+06	1.89E+01	19.59	939541026.8	10.44	0.09	10.44	10.48	0.09	0.09	0.09	Standard					
4180_NIST_051213_Ra@35	1.74	-5.54E+02	6.84E+03	8.08E+08	2.00E+01	1.61E+06	2.07E+01	20.04	465282628.7	-28621.71	0.00					Standard						
4180_NIST_051213_Ra@36	1.97	2.03E+03	9.82E+03	1.91E+09	2.92E+01	3.86E+06	2.93E+01	10.08	966977620.6	10.66	0.10	10.66	10.67	0.10	0.10	0.10	Standard					
4180_NIST_051213_Ra@37	1.99	2.03E+03	4.46E+03	1.90E+09	2.65E+01	3.86E+06	2.65E+01	10.13	958449755.8	10.62	0.04	10.62	10.62	0.04	0.04	0.04	Standard					
4180_NIST_051213_Ra@38	1.99	2.03E+03	6.66E+03	1.91E+09	2.79E+01	3.86E+06	2.77E+01	10.19	95849040.1	10.90	0.07	10.90	10.90	0.07	0.07	0.07	Standard					
4180_NIST_051213_Ra@39	1.99	2.03E+03	4.97E+03	1.89E+09	2.14E+01	3.82E+06	2.17E+01	10.24	946796831.9	10.94	0.05	10.94	10.94	0.05	0.05	0.05	Standard					
4180_NIST_051213_Ra@40	1.99	2.03E+03	8.31E+03	1.89E+09	2.01E+01	3.82E+06	2.01E+01	10.20	945772952.2	10.94	0.08	10.93	10.93	0.08	0.08	0.08	Standard					
4180_NIST_051213_Ra@41	1.99	2.03E+03	1.01E+04	1.95E+09	2.92E+01	3.81E+06	1.98E+01	10.35	942748487.2	11.09	0.10	11.09	11.07	0.10	0.10	0.10	Standard					
4180_NIST_051213_Ra@42	1.99	2.03E+03	6.50E+03	1.88E+09	2.05E+01	3.82E+06	2.08E+01	10.41	949252425.5	11.04	0.06	11.04	11.02	0.06	0.06	0.06	Standard					
4180_NIST_051213_Ra@43	1.99	2.03E+03	5.83E+03	1.88E+09	3.38E+01	3.82E+06	3.28E+01	10.57	949252425.5	10.13	0.09	10.13	10.12	0.09	0.09	0.09	Standard					
4180_NIST_051213_Ra@44	1.99	2.03E+03	8.73E+03	1.90E+09	1.79E+01	3.92E+06	1.79E+01	10.57	945750963.4	10.94	0.08	10.94	10.94	0.08	0.08	0.08	matrix					
4180_NIST_051213_Ra@45	1.99	2.06E+03	8.73E+03	1.90E+09	1.79E+01	3.92E+06	1.79E+01	10.57	945750963.4	28.87	0.08	28.87	28.85	0.08	0.08	0.08	matrix					
4180_NIST_051213_Ra@46	1.98	2.06E+03	7.34E+03	1.91E+09	1.52E+01	3.94E+06	1.52E+01	11.03	96402824.2	28.66	0.07	28.66	28.65	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@47	1.98	2.06E+03	6.87E+03	1.89E+09	1.17E+01	3.90E+06	1.19E+01	11.09	956311609.9	28.74	0.07	28.74	28.73	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@48	1.97	2.06E+03	7.44E+03	1.90E+09	1.51E+01	3.92E+06	1.54E+01	11.14	964120635.3	28.51	0.07	28.51	28.50	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@49	1.97	2.06E+03	8.10E+03	1.87E+09	1.10E+01	3.86E+06	1.09E+01	11.20	948241294.4	28.67	0.08	28.67	28.66	0.08	0.08	0.08	PrOI					
4180_NIST_051213_Ra@50	1.97	2.06E+03	4.89E+03	1.89E+09	1.23E+01	3.90E+06	1.23E+01	11.25	957616444.5	28.53	0.05	28.53	28.52	0.05	0.05	0.05	PrOI					
4180_NIST_051213_Ra@51	1.97	2.03E+03	8.08E+03	1.82E+09	1.67E+01	3.68E+06	1.67E+01	11.31	923048359.4	10.79	0.08	10.78	10.78	0.08	0.08	0.08	Standard					
4180_NIST_051213_Ra@52	1.96	2.03E+03	6.45E+03	1.81E+09	1.18E+01	3.67E+06	1.18E+01	11.36	920637037.3	10.77	0.06	10.77	10.77	0.06	0.06	0.06	Standard					
4180_NIST_051213_Ra@53	1.96	2.03E+03	1.03E+04	1.81E+09	1.59E+01	3.67E+06	1.58E+01	11.42	925487784.3	10.81	0.10	10.80	10.80	0.10	0.10	0.10	Standard					
4180_NIST_051213_Ra@54	1.95	2.06E+03	6.81E+03	1.82E+09	1.42E+01	3.68E+06	1.43E+01	11.47	929622947.4	10.85	0.07	10.84	10.84	0.07	0.07	0.07	Standard					
4180_NIST_051213_Ra@55	1.95	2.06E+03	4.82E+03	1.81E+09	1.60E+01	3.68E+06	1.62E+01	11.53	93272381.9	28.80	0.04	28.80	28.80	0.04	0.04	0.04	PrOI					
4180_NIST_051213_Ra@56	1.95	2.06E+03	6.32E+03	1.81E+09	6.96E+02	3.72E+06	6.90E+02	11.58	925183065.1	28.54	0.06	28.54	28.54	0.06	0.06	0.06	PrOI					
4180_NIST_051213_Ra@57	1.95	2.06E+03	6.32E+03	1.81E+09	6.96E+02	3.72E+06	6.90E+02	11.58	925183065.1	27.47	0.07	27.47	27.47	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@58	1.94	2.06E+03	7.07E+03	1.80E+09	8.37E+02	3.71E+06	8.49E+02	12.09	93071680.5	28.48	0.07	28.48	28.48	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@59	1.94	2.06E+03	8.55E+03	1.79E+09	7.33E+02	3.69E+06	7.08E+02	12.15	922863978.0	27.99	0.08	27.99	27.99	0.08	0.08	0.08	PrOI					
4180_NIST_051213_Ra@60	1.93	2.06E+03	5.83E+03	1.87E+09	9.65E+02	3.85E+06	8.70E+02	12.21	96985903.1	25.64	0.06	25.64	25.64	0.06	0.06	0.06	PrOI					
4180_NIST_051213_Ra@61	1.93	2.06E+03	5.99E+03	1.88E+09	1.39E+01	3.86E+06	1.39E+01	12.26	973256706.8	22.40	0.06	22.40	22.40	0.06	0.06	0.06	PrOI					
4180_NIST_051213_Ra@62	1.93	2.05E+03	1.32E+04	2.03E+09	3.54E+01	4.16E+06	3.44E+01	12.32	1052497065.5	22.41	0.13	22.41	22.41	0.13	0.13	0.13	PrOI					
4180_NIST_051213_Ra@63	1.92	2.06E+03	8.04E+03	1.89E+09	3.20E+02	3.88E+06	3.50E+02	12.37	98512842.5	25.51	0.08	25.51	25.51	0.08	0.08	0.08	PrOI					
4180_NIST_051213_Ra@64	1.92	2.06E+03	7.11E+03	1.81E+09	7.41E+02	3.71E+06	7.50E+02	12.43	94294176.3	24.94	0.07	24.94	24.95	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@65	1.92	2.03E+03	7.31E+03	1.73E+09	1.17E+01	3.50E+06	1.10E+01	12.48	901636040.5	11.01	0.07	11.01	11.01	0.07	0.07	0.07	Standard					
4180_NIST_051213_Ra@66	1.92	2.03E+03	7.52E+03	1.75E+09	1.22E+01	3.56E+06	1.21E+01	12.54	916158385.4	10.78	0.07	10.78	10.78	0.07	0.07	0.07	Standard					
4180_NIST_051213_Ra@67	1.90	2.03E+03	6.83E+03	1.75E+09	1.42E+01	3.54E+06	1.45E+01	12.59	91675062.4	10.90	0.07	10.91	10.91	0.07	0.07	0.07	Standard					
4180_NIST_051213_Ra@68	1.90	2.03E+03	9.97E+03	1.71E+09	1.15E+01	3.47E+06	1.17E+01	13.05	909093011.1	10.74	0.08	10.75	10.75	0.08	0.08	0.08	Standard					
4180_NIST_051213_Ra@69	1.90	2.07E+03	6.08E+03	1.76E+09	7.42E+02	3.48E+06	7.48E+02	13.11	929756446.0	31.64	0.06	31.64	31.64	0.06	0.06	0.06	PrOI					
4180_NIST_051213_Ra@70	1.88	2.07E+03	6.79E+03	1.70E+09	8.30E+02	3.53E+06	8.85E+02	13.16	903009555.7	34.29	0.07	34.30	34.30	0.07	0.07	0.07	PrOI					
4180_NIST_051213_Ra@71																						

Table A2. Continued.

Beam	SIMS ANALYSES				Mount: PRIM123	Standard: UNIL-Q1 (Paine)	Analyse: $\delta^{18}O$				Value: 9.81±0.14 (VSMOW, 2 σ)				Date: 26.06.2014	Comment
	$^{18}O/^{16}O$	CPS	$^{17}O/^{16}O$	SDP			Yield	CPS	SDP	mean	SDP	DBE correction (%)	Calibration (%)	$\delta^{18}O_{VSMOW}$		
d180_primi23_paine@48	1754				18:07	6.69	0.17	6.59	0.23	6.66	0.18	9.94	0.20	Standard		
d180_primi23_paine@49					18:12	6.73	0.18	6.70	0.18	6.73	0.18	9.88	0.20	Standard		
d180_primi23_paine@50					18:17	6.76	0.18	6.73	0.18	6.76	0.18	10.02	0.20	Standard		
d180_primi23_paine@51					18:21	6.49	0.12	6.51	0.21	6.46	0.12	9.75	0.19	Standard		
d180_primi23_paine@52					18:26	6.54	0.21	6.51	0.21	6.51	0.21	9.80	0.21	Standard		
d180_primi23_paine@53					18:31	6.62	0.14	6.59	0.14	6.59	0.14	9.88	0.19	Standard		
d180_primi23_paine@54					18:35	6.55	0.16	6.52	0.16	6.52	0.16	9.81	0.20	Standard		
d180_primi23_paine@55					18:38	6.65	0.13	6.62	0.13	6.62	0.13	9.91	0.19	Standard		
d180_primi23_paine@56					18:40	6.54	0.16	6.51	0.16	6.51	0.16	9.85	0.20	Standard		
d180_primi23_paine@57					18:43	6.57	0.17	6.54	0.17	6.54	0.17	9.85	0.20	Standard		
d180_primi23_paine@58					18:50	6.43	0.17	6.40	0.17	6.40	0.17	9.68	0.20	Standard		
d180_primi23_paine@59					18:54	6.41	0.20	6.38	0.20	6.38	0.20	9.66	0.21	Standard		
d180_primi23_paine@60					18:59	6.52	0.17	6.50	0.17	6.50	0.17	9.78	0.20	Standard		
d180_primi23_f4c1@1					19:04	27.57	0.20	27.49	0.20	27.54	0.20	30.90	0.21	base-KIOBC		
d180_primi23_f4c1@2					19:09	27.31	0.18	1.00	27.28	0.18	30.64	0.20	base-KIOBC			
d180_primi23_f4c1@3					19:13	27.56	0.15	27.54	0.15	27.54	0.15	30.89	0.20	base-KIOBC		
d180_primi23_f4c1@4					19:18	26.92	0.17	26.90	0.17	26.90	0.17	30.25	0.20	base-KIOBC		
d180_primi23_f1c2@1					19:23	28.40	0.14	28.40	0.14	28.37	0.14	31.73	0.20	middle-KIOBC		
d180_primi23_f1c2@2					19:27	27.69	0.14	27.67	0.14	27.67	0.14	31.02	0.20	middle-KIOBC		
d180_primi23_f1c2@3					19:32	26.99	0.16	26.96	0.16	26.96	0.16	30.31	0.20	middle-KIOBC		
d180_primi23_paine@61					19:37	6.47	0.17	6.53	0.17	6.44	0.17	9.73	0.20	Standard		
d180_primi23_paine@62					19:42	6.52	0.14	6.51	0.14	6.51	0.14	9.79	0.19	Standard		
d180_primi23_paine@63					19:46	6.48	0.14	6.45	0.14	6.45	0.14	9.79	0.19	Standard		
d180_primi23_f1c2@4					19:51	6.68	0.23	6.66	0.23	6.66	0.23	9.95	0.21	Standard		
d180_primi23_f2c2@1					19:56	27.68	0.20	28.08	0.20	27.66	0.20	31.01	0.21	middle-KIOBC		
d180_primi23_f2c2@2					20:01	27.63	0.19	0.82	27.61	0.19	30.96	0.21	middle-KIOBC			
d180_primi23_f2c2@3					20:05	27.97	0.22	27.94	0.22	27.94	0.22	31.30	0.21	middle-KIOBC		
d180_primi23_f2c2@4					20:10	27.89	0.21	27.86	0.21	27.86	0.21	31.22	0.21	middle-KIOBC		
d180_primi23_f2c2@5					20:15	27.83	0.16	27.81	0.16	27.81	0.16	31.16	0.20	middle-KIOBC		
d180_primi23_f2c2@6					20:20	27.49	0.14	27.47	0.14	27.47	0.14	30.83	0.20	middle-KIOBC		
d180_primi23_f2c2@7					20:24	28.14	0.11	28.12	0.11	28.12	0.11	31.48	0.20	middle-KIOBC		
d180_primi23_f3c2@1					20:29	27.92	0.14	27.90	0.14	27.90	0.14	31.25	0.20	middle-KIOBC		
d180_primi23_f3c2@2					20:34	27.98	0.16	27.96	0.16	27.96	0.16	31.31	0.20	middle-KIOBC		
d180_primi23_f3c2@3					20:38	28.40	0.19	28.38	0.19	28.38	0.19	31.74	0.21	middle-KIOBC		
d180_primi23_f3c2@4					20:43	28.36	0.18	28.34	0.18	28.34	0.18	31.70	0.20	middle-KIOBC		
d180_primi23_f4c2@1					20:48	28.47	0.21	28.45	0.21	28.45	0.21	31.81	0.21	middle-KIOBC		
d180_primi23_f4c2@2					20:53	28.43	0.15	28.41	0.15	28.41	0.15	31.77	0.20	middle-KIOBC		
d180_primi23_f4c2@3					20:57	29.09	0.17	29.07	0.17	29.07	0.17	32.43	0.20	middle-KIOBC		
d180_primi23_f4c2@4					21:02	27.92	0.23	27.90	0.23	27.90	0.23	31.25	0.22	middle-KIOBC		
d180_primi23_paine@65					21:07	6.36	0.12	6.35	0.12	6.35	0.12	9.63	0.19	Standard		
d180_primi23_paine@66					21:12	6.67	0.14	6.56	0.14	6.65	0.14	9.84	0.19	Standard		
d180_primi23_paine@67					21:16	6.27	0.12	6.25	0.12	6.25	0.12	9.54	0.19	Standard		
d180_primi23_paine@68					21:21	6.87	0.17	6.86	0.17	6.86	0.17	10.14	0.20	Standard		
d180_primi23_f4c2@5					21:26	27.45	0.22	27.62	0.22	27.44	0.22	30.79	0.21	middle-KIOBC		
d180_primi23_f4c2@6					21:31	27.34	0.17	27.32	0.17	27.32	0.17	30.68	0.20	middle-KIOBC		
d180_primi23_f4c2@7					21:35	27.70	0.16	27.69	0.16	27.69	0.16	31.04	0.20	middle-KIOBC		
d180_primi23_f4c2@8					21:40	27.70	0.16	27.69	0.16	27.69	0.16	31.04	0.20	middle-KIOBC		
d180_primi23_f5c1@1					21:45	27.57	0.20	27.55	0.20	27.55	0.20	30.91	0.21	middle-KIOBC		
d180_primi23_f5c1@2					21:49	27.61	0.18	27.60	0.18	27.60	0.18	30.95	0.20	middle-KIOBC		
d180_primi23_f5c1@3					21:54	27.90	0.12	27.88	0.12	27.88	0.12	31.24	0.19	middle-KIOBC		
d180_primi23_f5c1@4					21:59	27.18	0.18	27.17	0.18	27.17	0.18	30.52	0.21	middle-KIOBC		
d180_primi23_f5c1@5					22:03	27.32	0.17	27.31	0.17	27.31	0.17	30.66	0.20	middle-KIOBC		
d180_primi23_f5c1@6					22:08	28.88	0.14	28.87	0.14	28.87	0.14	32.22	0.20	middle-KIOBC		
d180_primi23_f1c3@1					22:13	27.17	0.17	27.16	0.17	27.16	0.17	30.51	0.20	pop-KIOBC		
d180_primi23_f2c3@1					22:18	27.77	0.23	27.76	0.23	27.76	0.23	31.11	0.22	pop-KIOBC		
d180_primi23_f2c3@2					22:22	28.10	0.19	28.08	0.19	28.08	0.19	31.44	0.21	pop-KIOBC		
d180_primi23_f2c3@3					22:27	27.83	0.14	27.82	0.14	27.82	0.14	31.17	0.20	pop-KIOBC		
d180_primi23_f2c3@4					22:32	27.06	0.18	27.05	0.18	27.05	0.18	30.40	0.20	pop-KIOBC		
d180_primi23_paine@69					22:37	6.73	0.12	6.45	0.12	6.72	0.12	10.00	0.19	Standard		
d180_primi23_paine@70					22:41	6.53	0.20	6.60	0.20	6.52	0.20	9.81	0.20	Standard		
d180_primi23_paine@71					22:46	6.50	0.14	6.49	0.14	6.49	0.14	9.78	0.19	Standard		
d180_primi23_paine@72					22:51	6.02	0.19	6.01	0.19	6.01	0.19	9.30	0.20	Standard		
d180_primi23_f3c3@1					22:56	26.46	0.17	26.95	0.17	26.45	0.17	29.80	0.20	pop-KIOBC		
d180_primi23_f3c3@2					23:00	27.56	0.18	0.78	27.55	0.18	30.90	0.20	pop-KIOBC			
d180_primi23_f3c3@3					23:05	26.90	0.16	26.89	0.16	26.89	0.16	30.24	0.20	pop-KIOBC		
d180_primi23_f4c3@1					23:10	26.77	0.10	26.76	0.10	26.76	0.10	30.12	0.19	pop-KIOBC		
d180_primi23_f4c3@2					23:15	26.43	0.17	26.42	0.17	26.42	0.17	29.77	0.20	pop-KIOBC		
d180_primi23_f4c3@3					23:19	26.77	0.20	26.76	0.20	26.76	0.20	30.11	0.21	pop-KIOBC		
d180_primi23_f4c3@4					23:24	26.95	0.15	26.94	0.15	26.94	0.15	30.30	0.20	pop-KIOBC		
d180_primi23_f5c3@1					23:29	26.74	0.27	26.74	0.27	26.74	0.27	30.09	0.23	pop-KIOBC		
d180_primi23_f5c3@2					23:34	27.00	0.17	27.00	0.17	27.00	0.17	30.35	0.20	pop-KIOBC		

Table A2 Continued.

Beam #d	SIMS ANALYSES				Mount: PRIM23				Standard: UNIL-QI (Panne)				Analyse: $\delta^{18}O$				Value: 9.81±0.14 (VSMOW, 2σ)				Date: 26.06.2014	
	$H^2(1^0C/0^0)$		$L^2(1^0C/0^0)$		$H^2(1^0C/0^0)$		$L^2(1^0C/0^0)$		Yield		Measurements (%)		Drift correction (%)		Calibration (%)		Comments					
	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS/aA	mean	2SD	$\delta^{18}O$	2SD	$\delta^{18}O_{VSMOW}$	2SD	$\delta^{18}O_{VSMOW}$	2SD					
d180_jrmi23_1r5c@3											26.50	0.17	26.49	0.17	29.84	0.20		top Kf8C				
d180_jrmi23_1r6c@1											26.91	0.10	26.90	0.10	30.25	0.19		Kf8C				
d180_jrmi23_1r6c@2											27.14	0.10	27.14	0.10	30.49	0.19		Kf8C				
d180_jrmi23_1r6c@3											27.84	0.18	27.84	0.18	31.19	0.21		Kf8C				
d180_jrmi23_1r6c@4											27.29	0.20	27.29	0.20	30.64	0.21		Kf8C				
d180_jrmi23_1r6c@5											26.97	0.15	26.97	0.15	30.72	0.20		Kf8C				
d180_jrmi23_panne@74											6.44	0.11	6.44	0.11	9.72	0.19		Standard				
d180_jrmi23_panne@75											6.43	0.13	6.43	0.13	9.71	0.19		Standard				
d180_jrmi23_panne@76											6.44	0.13	6.44	0.13	9.73	0.19		Standard				
d180_jrmi23_panne@77											6.41	0.10	6.40	0.10	9.69	0.19		Standard				
d180_jrmi23_1r2c@1											27.19	0.21	27.37	0.21	30.54	0.21		Kf8C				
d180_jrmi23_1r2c@2											26.99	0.15	26.99	0.15	30.34	0.20		Kf8C				
d180_jrmi23_1r2c@3											26.16	0.13	26.16	0.13	29.51	0.19		Kf8C				
d180_jrmi23_1r3c@1											28.58	0.15	28.58	0.15	31.95	0.20		Kf8C				
d180_jrmi23_1r3c@2											27.52	0.20	27.52	0.20	30.88	0.21		Kf8C				
d180_jrmi23_1r3c@3											28.44	0.19	28.44	0.19	31.79	0.21		Kf8C				
d180_jrmi23_1r3c@4											27.07	0.18	27.07	0.18	30.42	0.20		Kf8C				
d180_jrmi23_1r3c@5											27.23	0.19	27.23	0.19	30.58	0.21		Kf8C				
d180_jrmi23_1r4c@1											27.46	0.18	27.46	0.18	30.81	0.20		Kf8C				
d180_jrmi23_1r4c@2											27.05	0.16	27.05	0.16	30.40	0.20		Kf8C				
d180_jrmi23_1r4c@3											27.42	0.10	27.42	0.10	30.78	0.19		Kf8C				
d180_jrmi23_1r4c@4											27.56	0.11	27.56	0.11	30.91	0.19		Kf8C				
d180_jrmi23_1r5c@1											27.00	0.12	27.01	0.12	30.36	0.19		Kf8C				
d180_jrmi23_1r5c@2											27.59	0.15	27.59	0.15	30.94	0.20		Kf8C				
d180_jrmi23_1r5c@3											27.27	0.14	27.27	0.14	30.63	0.20		Kf8C				
d180_jrmi23_panne@78											6.77	0.15	6.77	0.15	10.06	0.19		Standard				
d180_jrmi23_panne@79											6.79	0.16	6.80	0.16	10.08	0.20		Standard				
d180_jrmi23_panne@80											6.33	0.24	6.33	0.24	9.61	0.22		Standard				
d180_jrmi23_panne@81											6.66	0.11	6.66	0.11	9.95	0.19		Standard				
d180_jrmi23_1r5c@1											26.73	0.17	26.73	0.17	30.08	0.20		Kf8C				
d180_jrmi23_1r5c@2											26.79	0.12	26.79	0.12	30.15	0.19		Kf8C				
d180_jrmi23_1r5c@3											26.88	0.16	26.88	0.16	30.23	0.20		Kf8C				
d180_jrmi23_1r5c@4											27.64	0.18	27.65	0.18	31.00	0.20		Kf8C				
d180_jrmi23_1r5c@5											27.28	0.14	27.29	0.14	30.64	0.20		Kf8C				
d180_jrmi23_1r5c@6											27.44	0.14	27.45	0.14	30.80	0.20		Kf8C				
d180_jrmi23_1r5c@7											25.41	0.16	25.42	0.16	28.77	0.20		Kf8C				
d180_jrmi23_1r5c@8											27.25	0.16	27.26	0.16	30.61	0.20		Kf8C				
d180_jrmi23_1r5c@9											27.28	0.14	27.28	0.14	30.64	0.20		Kf8C				
d180_jrmi23_1r5c@10											27.64	0.12	27.64	0.12	31.00	0.19		Kf8C				
d180_jrmi23_1r5c@11											27.27	0.17	27.28	0.17	30.63	0.20		Kf8C				
d180_jrmi23_1r5c@12											27.85	0.16	27.86	0.16	31.22	0.20		Kf8C				
d180_jrmi23_1r5c@13											27.40	0.15	27.41	0.15	30.76	0.20		Kf8C				
d180_jrmi23_1r5c@14											26.98	0.12	26.99	0.12	30.34	0.19		Kf8C				
d180_jrmi23_1r5c@15											26.65	0.12	26.66	0.12	30.01	0.19		Kf8C				
d180_jrmi23_panne@82											6.66	0.18	6.67	0.18	9.96	0.20		Standard				
d180_jrmi23_panne@83											6.09	0.18	6.10	0.18	9.39	0.20		Standard				
d180_jrmi23_panne@84											6.64	0.13	6.65	0.13	9.94	0.19		Standard				
d180_jrmi23_panne@85											6.57	0.14	6.58	0.14	9.87	0.19		Standard				
d180_jrmi23_1r4c@5											26.63	0.17	27.06	0.17	30.00	0.20		Kf8C				
d180_jrmi23_1r4c@6											27.53	0.21	27.55	0.21	30.90	0.21		Kf8C				
d180_jrmi23_1r4c@7											27.17	0.14	27.18	0.14	30.54	0.20		Kf8C				
d180_jrmi23_1r4c@8											26.89	0.16	26.90	0.16	30.26	0.20		Kf8C				
d180_jrmi23_1r4c@9											31.36	0.15	31.16	0.15	34.74	0.20		Pf01				
d180_jrmi23_1r4c@10											30.17	0.15	30.18	0.15	33.54	0.20		Pf01				
d180_jrmi23_1r4c@11											30.15	0.20	30.16	0.20	33.53	0.21		Pf01				
d180_jrmi23_1r4c@12											31.35	0.10	31.36	0.10	34.73	0.19		Pf01				
d180_jrmi23_1r4c@13											31.37	0.16	31.39	0.16	34.75	0.20		Pf01				
d180_jrmi23_1r4c@14											31.21	0.14	31.23	0.14	34.59	0.20		Pf01				
d180_jrmi23_1r4c@15											31.28	0.12	31.29	0.12	34.66	0.19		Pf01				
d180_jrmi23_1r4c@16											31.49	0.13	31.51	0.13	34.88	0.19		Pf01				
d180_jrmi23_1r4c@17											31.32	0.20	31.34	0.20	34.71	0.21		Pf01				
d180_jrmi23_1r4c@18											31.38	0.15	31.40	0.15	34.77	0.20		Pf01				
d180_jrmi23_1r4c@19											31.71	0.22	31.72	0.22	35.09	0.21		Pf01				

Table A2 Continued.

Beam #4	SIMS ANALYSES			Mount: PRIM123			Standard: UNIL-Q1 (Paine)			Analyse: $\delta^{18}O$			Value: 981±0.14 (VSMOW, 2 σ)			Date: 26/06/2014	Comment	
	$H2.L12$ ($^{18}O/^{16}O$)	$L2$ ($^{18}O/^{16}O$)	$H2.L12$ ($^{18}O/^{16}O$)	CPS	Yield	Time	CPS	Yield	Time	CPS	Yield	Time	CPS	Yield	Time			CPS
	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD	CPS	2SD
d18O_prim123_paine@86	4.37		6.37	0.22	6.55	6.39	0.22	9.67	0.21	9.67	0.21	9.67	0.21	9.67	0.21	9.67	0.21	Standard
d18O_prim123_paine@87	4.42		6.64	0.12	0.28	6.66	0.12	6.66	0.12	6.66	0.12	6.66	0.12	6.66	0.12	6.66	0.12	Standard
d18O_prim123_paine@88	4.46		6.52	0.21	6.54	6.54	0.21	6.54	0.21	6.54	0.21	6.54	0.21	6.54	0.21	6.54	0.21	Standard
d18O_prim123_paine@89	4.51		6.67	0.19	6.69	6.69	0.19	6.69	0.19	6.69	0.19	6.69	0.19	6.69	0.19	6.69	0.19	Standard
d18O_prim123_f4c6@1	4.56		31.54	0.15	31.12	31.56	0.15	31.56	0.15	31.56	0.15	31.56	0.15	31.56	0.15	31.56	0.15	P001
d18O_prim123_f4c6@2	5.01		31.09	0.16	0.75	31.11	0.16	31.11	0.16	31.11	0.16	31.11	0.16	31.11	0.16	31.11	0.16	P001
d18O_prim123_f4c6@3	5.05		31.24	0.10	31.26	31.26	0.10	31.26	0.10	31.26	0.10	31.26	0.10	31.26	0.10	31.26	0.10	P001
d18O_prim123_f5c6@1	5.10		30.49	0.13	30.50	30.50	0.13	30.50	0.13	30.50	0.13	30.50	0.13	30.50	0.13	30.50	0.13	P001
d18O_prim123_f5c6@2	5.15		30.50	0.20	30.52	30.52	0.20	30.52	0.20	30.52	0.20	30.52	0.20	30.52	0.20	30.52	0.20	P001
d18O_prim123_f5c6@3	5.20		31.29	0.13	31.31	31.31	0.13	31.31	0.13	31.31	0.13	31.31	0.13	31.31	0.13	31.31	0.13	P001
d18O_prim123_f1c7@1	5.25		31.34	0.17	31.36	31.36	0.17	31.36	0.17	31.36	0.17	31.36	0.17	31.36	0.17	31.36	0.17	P001
d18O_prim123_f1c7@2	5.29		31.41	0.16	31.43	31.43	0.16	31.43	0.16	31.43	0.16	31.43	0.16	31.43	0.16	31.43	0.16	P001
d18O_prim123_f1c7@3	5.34		31.04	0.16	31.07	31.07	0.16	31.07	0.16	31.07	0.16	31.07	0.16	31.07	0.16	31.07	0.16	P001
d18O_prim123_f1c7@4	5.39		31.31	0.17	31.33	31.33	0.17	31.33	0.17	31.33	0.17	31.33	0.17	31.33	0.17	31.33	0.17	P001
d18O_prim123_f2c7@1	5.44		31.53	0.10	31.55	31.55	0.10	31.55	0.10	31.55	0.10	31.55	0.10	31.55	0.10	31.55	0.10	P001
d18O_prim123_f2c7@2	5.48		31.17	0.18	31.19	31.19	0.18	31.19	0.18	31.19	0.18	31.19	0.18	31.19	0.18	31.19	0.18	P001
d18O_prim123_f3c7@1	5.53		30.36	0.14	30.38	30.38	0.14	30.38	0.14	30.38	0.14	30.38	0.14	30.38	0.14	30.38	0.14	P001
d18O_prim123_f3c7@2	5.58		31.16	0.17	31.18	31.18	0.17	31.18	0.17	31.18	0.17	31.18	0.17	31.18	0.17	31.18	0.17	P001
d18O_prim123_f4c7@1	6.03		31.35	0.10	31.38	31.38	0.10	31.38	0.10	31.38	0.10	31.38	0.10	31.38	0.10	31.38	0.10	P001
d18O_prim123_paine@90	6.08		6.78	0.24	6.50	6.81	0.24	10.09	0.22	10.09	0.22	10.09	0.22	10.09	0.22	10.09	0.22	Standard
d18O_prim123_paine@91	6.12		6.21	0.15	0.51	6.23	0.15	6.23	0.15	6.23	0.15	6.23	0.15	6.23	0.15	6.23	0.15	Standard
d18O_prim123_paine@92	6.17		6.62	0.14	6.64	6.64	0.14	6.64	0.14	6.64	0.14	6.64	0.14	6.64	0.14	6.64	0.14	Standard
d18O_prim123_paine@93	6.22		6.38	0.14	6.41	6.41	0.14	6.41	0.14	6.41	0.14	6.41	0.14	6.41	0.14	6.41	0.14	Standard
d18O_prim123_f4c7@2	6.27		31.09	0.10	31.07	31.12	0.10	31.12	0.10	31.12	0.10	31.12	0.10	31.12	0.10	31.12	0.10	P001
d18O_prim123_f5c7@1	6.31		31.61	0.17	0.36	31.64	0.17	31.64	0.17	31.64	0.17	31.64	0.17	31.64	0.17	31.64	0.17	P001
d18O_prim123_f5c7@2	6.36		30.94	0.14	30.96	30.96	0.14	30.96	0.14	30.96	0.14	30.96	0.14	30.96	0.14	30.96	0.14	P001
d18O_prim123_f5c7@3	6.41		31.23	0.17	31.26	31.26	0.17	31.26	0.17	31.26	0.17	31.26	0.17	31.26	0.17	31.26	0.17	P001
d18O_prim123_f5c7@4	6.46		31.45	0.20	31.48	31.48	0.20	31.48	0.20	31.48	0.20	31.48	0.20	31.48	0.20	31.48	0.20	P001
d18O_prim123_f5c7@5	6.50		31.34	0.11	31.37	31.37	0.11	31.37	0.11	31.37	0.11	31.37	0.11	31.37	0.11	31.37	0.11	P001
d18O_prim123_f1c8@1	6.55		30.88	0.14	30.91	30.91	0.14	30.91	0.14	30.91	0.14	30.91	0.14	30.91	0.14	30.91	0.14	C601
d18O_prim123_f1c8@2	7.00		30.91	0.12	30.93	30.93	0.12	30.93	0.12	30.93	0.12	30.93	0.12	30.93	0.12	30.93	0.12	C601
d18O_prim123_f2c8@1	7.05		31.49	0.17	31.52	31.52	0.17	31.52	0.17	31.52	0.17	31.52	0.17	31.52	0.17	31.52	0.17	C601
d18O_prim123_f2c8@2	7.09		30.79	0.17	30.82	30.82	0.17	30.82	0.17	30.82	0.17	30.82	0.17	30.82	0.17	30.82	0.17	C601
d18O_prim123_f3c8@1	7.14		30.71	0.13	30.73	30.73	0.13	30.73	0.13	30.73	0.13	30.73	0.13	30.73	0.13	30.73	0.13	C601
d18O_prim123_f3c8@2	7.19		30.74	0.13	30.77	30.77	0.13	30.77	0.13	30.77	0.13	30.77	0.13	30.77	0.13	30.77	0.13	C601
d18O_prim123_f4c8@1	7.24		30.96	0.19	30.99	30.99	0.19	30.99	0.19	30.99	0.19	30.99	0.19	30.99	0.19	30.99	0.19	C601
d18O_prim123_f4c8@2	7.28		30.88	0.17	30.91	30.91	0.17	30.91	0.17	30.91	0.17	30.91	0.17	30.91	0.17	30.91	0.17	C601
d18O_prim123_f5c8@1	7.33		31.11	0.20	31.14	31.14	0.20	31.14	0.20	31.14	0.20	31.14	0.20	31.14	0.20	31.14	0.20	C601
d18O_prim123_f5c8@2	7.38		31.00	0.19	31.03	31.03	0.19	31.03	0.19	31.03	0.19	31.03	0.19	31.03	0.19	31.03	0.19	C601
d18O_prim123_paine@94	7.43		6.55	0.17	6.48	6.59	0.17	6.48	0.17	6.48	0.17	6.48	0.17	6.48	0.17	6.48	0.17	Standard
d18O_prim123_paine@95	7.48		6.45	0.19	0.12	6.48	0.19	6.48	0.19	6.48	0.19	6.48	0.19	6.48	0.19	6.48	0.19	Standard
d18O_prim123_paine@96	7.52		6.45	0.13	6.48	6.48	0.13	6.48	0.13	6.48	0.13	6.48	0.13	6.48	0.13	6.48	0.13	Standard

Table A2 Continued.

SIMS ANALYSES										Mount: BR2				Analyse: $\delta^{18}O$				Value: 9.81±0.14 (VSMOW, 2 σ)				Date: 14.04.2015	
Beam	$^{17}O/^{16}O$	$^{18}O/^{16}O$	$^{17}O/^{16}O$ (Cath)	$^{18}O/^{16}O$ (Cath)	$^{17}O/^{16}O$ (Cath)	$^{18}O/^{16}O$ (Cath)	Time	Yield	Measurements (%)	Drift correction (%)	Calibration (%)	$\delta^{18}O_{VSMOW}$	$\delta^{18}O_{VSMOW}$	$\delta^{18}O_{VSMOW}$	$\delta^{18}O_{VSMOW}$	Comment							
id	GS	GS	GS	GS	GS	GS		atom	SD	SD	SD	SD	SD	SD	SD								
2.00	2.0252E+03	8.2597E+03	1.5308E+09	1.0875E+02	3.0359E+06	1.3709E+02	23.56	7.547E+08	3.65	0.17	4.18	3.65	0.17	9.08	0.27	Standard							
2.00	2.0414E+03	1.4626E+09	1.4626E+09	4.3975E+02	2.9458E+06	5.4077E+02	0.01	7.3000E+08	4.42	0.27	0.76	4.42	0.27	9.84	0.29	Standard							
2.00	2.0364E+03	1.4850E+09	1.4850E+09	2.2870E+02	2.9900E+06	2.2158E+02	0.06	7.4219E+08	4.17	0.19	0.19	4.17	0.19	9.58	0.28	Standard							
2.00	2.0422E+03	1.1873E+02	1.4833E+09	3.8832E+02	2.9919E+06	3.9387E+02	0.10	7.4366E+08	4.49	0.24	0.24	4.49	0.24	9.91	0.28	Standard							
2.00	2.0521E+03	1.5478E+02	1.5408E+09	6.3102E+02	3.0344E+06	6.7789E+02	0.15	7.8210E+08	23.68	0.31	23.75	23.68	0.31	29.20	0.31	K609							
1.99	2.0522E+03	9.4119E+03	1.5709E+09	2.9077E+02	3.2385E+06	3.4585E+02	0.20	7.9252E+08	23.93	0.19	0.87	23.93	0.19	29.45	0.28	K609							
1.99	2.0532E+03	7.8983E+03	1.5668E+09	2.2060E+02	3.2173E+06	2.2523E+02	0.24	7.8751E+08	23.95	0.16	0.16	23.95	0.16	29.48	0.27	K609							
1.99	2.0530E+03	8.9644E+03	1.5339E+09	2.3064E+02	3.1902E+06	3.0209E+02	0.29	7.8061E+08	23.83	0.18	0.18	23.83	0.18	29.35	0.28	K609							
1.99	2.0534E+03	9.2730E+03	1.5748E+09	1.7150E+02	3.2381E+06	1.7185E+02	0.34	7.8796E+08	24.02	0.19	0.19	24.02	0.19	29.54	0.28	K609							
2.00	2.0538E+03	8.7920E+03	1.5326E+09	2.0553E+02	3.1470E+06	2.2561E+02	0.38	7.6632E+08	24.16	0.18	0.18	24.16	0.18	29.68	0.28	K609							
2.00	2.0531E+03	1.5408E+09	1.5408E+09	1.3827E+02	3.1623E+06	1.5197E+02	0.43	7.7121E+08	23.95	0.19	0.19	23.95	0.19	29.47	0.28	K609							
2.00	2.0507E+03	9.8889E+03	1.4999E+09	1.1862E+02	3.0799E+06	1.8338E+02	0.48	7.5146E+08	23.70	0.20	0.20	23.70	0.20	28.21	0.28	K609							
2.00	2.0524E+03	8.3424E+03	1.5178E+09	2.3201E+02	3.1148E+06	2.3700E+02	0.52	7.6045E+08	23.53	0.17	0.17	23.53	0.17	29.05	0.28	K609							
1.00515	br2_mine@25	8.2597E+03	1.5088E+09	1.0837E+02	3.0359E+06	1.3709E+02	0.57	7.547E+08	3.65	0.17	4.18	3.65	0.17	9.06	0.27	Standard							
1.00515	br2_mine@26	1.4626E+09	1.4626E+09	4.3975E+02	2.9458E+06	5.4077E+02	1.02	7.3000E+08	4.42	0.27	0.76	4.42	0.27	9.84	0.29	Standard							
1.00515	br2_mine@27	1.4850E+09	1.4850E+09	2.2870E+02	2.9900E+06	2.2158E+02	1.06	7.4219E+08	4.17	0.19	0.19	4.17	0.19	9.58	0.28	Standard							
1.00515	br2_mine@28	1.1873E+02	1.4833E+09	3.8832E+02	2.9919E+06	3.9387E+02	1.11	7.4366E+08	4.49	0.24	0.24	4.49	0.24	9.91	0.28	Standard							
1.99	2.0544E+03	6.2092E+03	1.5709E+09	1.5287E+02	3.2268E+06	1.2770E+02	1.16	7.8745E+08	24.52	0.12	25.75	24.52	0.12	30.04	0.27	K610							
2.00	2.0547E+03	9.7278E+03	1.5733E+09	2.5268E+02	3.2329E+06	2.5208E+02	1.20	7.8586E+08	24.66	0.19	1.78	24.66	0.19	30.19	0.28	K610							
2.00	2.0551E+03	1.3332E+02	1.5745E+09	2.9982E+02	3.2349E+06	2.8401E+02	1.25	7.8791E+08	24.98	0.27	0.27	24.98	0.27	30.50	0.30	K610							
2.00	2.0570E+03	6.6872E+03	1.5756E+09	1.3244E+02	3.2088E+06	1.0544E+02	1.29	7.7890E+08	25.83	0.17	0.17	25.83	0.17	31.36	0.28	K610							
2.00	2.0591E+03	1.0546E+02	1.5500E+09	1.9507E+02	3.1927E+06	2.4505E+02	1.34	7.7631E+08	26.96	0.21	0.21	26.96	0.21	32.50	0.28	K610							
2.00	2.0566E+03	8.8795E+03	1.5845E+09	2.3793E+02	3.2383E+06	2.0786E+02	1.39	7.9390E+08	25.63	0.18	0.18	25.63	0.18	31.15	0.28	K610							
2.00	2.0575E+03	1.2316E+02	1.5632E+09	1.8248E+02	3.2163E+06	2.0853E+02	1.43	7.8180E+08	26.07	0.25	0.25	26.07	0.25	31.60	0.29	K610							
1.99	2.0579E+03	9.2116E+03	1.5852E+09	1.6525E+02	3.2462E+06	2.0276E+02	1.48	7.8111E+08	26.27	0.18	0.18	26.27	0.18	31.80	0.28	K610							
1.99	2.0590E+03	8.0113E+03	1.5713E+09	1.2521E+02	3.2325E+06	1.3137E+02	1.53	7.8934E+08	26.83	0.16	0.16	26.83	0.16	32.36	0.28	K610							
1.99	2.0438E+03	1.2262E+02	1.4803E+09	2.0883E+02	2.9809E+06	2.1863E+02	1.57	7.4390E+08	4.18	0.25	4.53	4.18	0.25	9.60	0.29	Standard							
1.99	2.0441E+03	1.0785E+02	1.4823E+09	4.4082E+02	2.9915E+06	4.8745E+02	2.02	7.4540E+08	4.45	0.22	0.56	4.45	0.22	9.86	0.28	Standard							
2.00	2.0445E+03	1.3614E+02	1.4826E+09	4.3495E+02	2.9873E+06	4.4552E+02	2.07	7.4126E+08	4.63	0.27	4.63	4.63	0.27	10.05	0.29	Standard							
2.00	2.0499E+03	1.8551E+02	1.4837E+09	6.7282E+02	2.9893E+06	6.6640E+02	2.11	7.4474E+08	4.85	0.21	0.21	4.85	0.21	10.27	0.28	Standard							
2.00	2.0571E+03	9.5420E+03	1.5571E+09	2.1809E+02	3.2037E+06	2.2621E+02	2.16	7.7975E+08	25.91	0.19	26.12	25.90	0.19	31.44	0.28	K614							
2.00	2.0582E+03	1.0153E+02	1.5408E+09	2.4376E+02	3.1705E+06	2.6485E+02	2.21	7.7190E+08	26.42	0.20	0.87	26.41	0.20	31.95	0.28	K614							
2.00	2.0583E+03	1.0939E+02	1.5424E+09	2.5605E+02	3.1745E+06	3.1688E+02	2.25	7.7170E+08	26.46	0.22	0.22	26.46	0.22	32.00	0.29	K614							
2.01	2.0584E+03	1.1965E+02	1.5462E+09	2.6002E+02	3.1827E+06	3.2304E+02	2.30	7.6922E+08	26.53	0.24	0.24	26.53	0.24	32.06	0.29	K614							
2.01	2.0587E+03	1.9386E+02	1.5377E+09	3.3796E+02	3.1608E+06	2.0001E+02	2.35	7.6648E+08	25.17	0.22	0.22	25.16	0.22	30.69	0.29	K614							
2.00	2.0581E+03	9.5489E+03	1.5643E+09	2.2990E+02	3.2198E+06	2.0001E+02	2.39	7.8088E+08	26.38	0.19	0.19	26.38	0.19	31.91	0.28	K614							
2.00	2.0579E+03	9.5134E+03	1.5516E+09	2.2132E+02	3.1923E+06	2.3971E+02	2.44	7.7574E+08	26.30	0.19	0.19	26.30	0.19	31.83	0.28	K614							
2.01	2.0574E+03	1.9481E+02	1.5498E+09	2.5269E+02	3.1876E+06	2.1038E+02	2.48	7.7035E+08	26.04	0.19	0.19	26.03	0.19	31.57	0.28	K614							
2.01	2.0571E+03	1.5798E+03	1.5798E+09	2.0792E+02	3.2046E+06	2.4797E+02	2.53	7.7403E+08	25.87	0.18	0.18	25.87	0.18	31.40	0.28	K614							
2.01	2.0338E+03	1.0547E+02	1.4448E+09	3.9720E+02	2.9898E+06	4.3025E+02	2.58	7.3905E+08	4.29	0.21	4.48	4.29	0.21	9.70	0.28	Standard							
2.00	2.0422E+03	8.1319E+03	1.4902E+09	5.5718E+02	3.0077E+06	5.7874E+02	3.02	7.4328E+08	4.49	0.16	0.27	4.49	0.16	9.91	0.27	Standard							
2.01	2.0441E+03	1.0499E+02	1.4800E+09	5.8475E+02	2.9932E+06	5.9384E+02	3.07	7.4053E+08	4.60	0.21	4.59	4.60	0.21	10.01	0.28	Standard							
2.01	2.0432E+03	1.2257E+02	1.4821E+09	2.6190E+02	2.9849E+06	3.3507E+02	3.12	7.3852E+08	4.56	0.25	4.55	4.55	0.25	9.97	0.29	Standard							
2.01	2.0404E+03	1.4050E+02	1.5284E+09	2.5807E+02	3.1185E+06	2.3655E+02	3.16	7.6098E+08	17.55	0.28	19.42	17.55	0.28	23.04	0.30	K625up							
2.01	2.0454E+03	9.1025E+03	1.5208E+09	3.1637E+02	3.1230E+06	3.3011E+02	3.21	7.5890E+08	20.06	0.18	4.97	20.06	0.18	23.56	0.28	K625up							
2.01	2.0462E+03	1.2212E+02	1.5325E+09	1.6666E+02	3.1361E+06	1.5799E+02	3.26	7.6390E+08	20.46	0.24	0.24	20.46	0.24	25.96	0.29	K625up							
2.01	2.0474E+03	8.5655E+03	1.5157E+09	1.9028E+02	3.1073E+06	1.9533E+02	3.30	7.5298E+08	21.06	0.17	0.17	21.05	0.17	26.56	0.28	K625up							
2.02	2.0489E+03	1.5043E+02	1.4608E+09	2.8121E+02	2.9900E+06	3.7696E+02	3.35	7.2828E+08	14.79	0.30	0.30	14.79	0.30	20.26	0.30	K625up							
2.02	2.0524E+03	1.3965E+02	1.5273E+09	1.9213E+02	3.1361E+06	3.1977E+02	3.40	7.5328E+08	25.54	0.28	0.28	25.54	0.28	29.06	0.30	K625up							
2.02	2.0407E+03	1.5395E+02	1.5262E+09	3.1763E+02	3.1145E+06	3.4770E+02	3.44	7.5625E+08	17.70	0.31	0.31	17.70	0.31	25.07	0.28	K625up							
2.01	2.0445E+03	8.4551E+03	1.4997E+09	2.8287E+02	3.0655E+06	2.4565E+02	3.49	7.4628E+08	19.58	0.17	0.17	19.57	0.17	25.07	0.28	K625up							
2.01	2.0453E+03	6.3718E+03	1.4909E+09	3.8126E+02	3.0627E+06	3.6883E+02	3.54	7.4372E+08	20.02	0.13	20.01	20.01	0.13	25.51	0.27	K625up							
2.01	2.0424E+03	1.3017E+02	1.4842E+09	6.8412E+02	2.9895E+06	7.3530E+02	3.58	7.3850E+08	4.50	0.26	4.42	4.50	0.26	9.91	0.29	Standard							
2.01	2.0399E+03	1.3256E+02	1.4713E+09	3.8917E+02	2.9631E+06	3.8151E+02	4.03	7.3009E+08	4.34	0.27	0.14	4.34	0.27	9.75	0.29	Standard							
2.01	2.0440E+03	1.2609E+02	1.4901E+09	3.6778E+02	3.0011E+06	4.0231E+02	4.08	7.4098E+08	4.38	0.25	4.38	4.38	0.25	9.80	0.29	Standard							
2.01	2.0414E+03	1.1095E+02	1.4848E+09	4.7488E+02	2.9898E+06	4.5383E+02	4.12	7.3989E+08	4.46	0.22	4.46	4.46	0.22	9.87	0.28	Standard							
2.01	2.0542E+03	1.2385E+02	1.5052E+09	2.8608E+02	3.0917E+06	3.9707E+02	4.17	7.4808E+08	24.43	0.25	24.48	24.42	0.25	29.95	0.29	K635							
2.01	2.0552E+03	1.3942E+02	1.4681E+09	5.0650E+02	3.0169E+06	4.9550E+02	4.22	7.2920E+08	24.94	0.28	0.84	24.94	0.28	30.47	0.30	K635							
2.01	2.0540E+03	8.7273E+03	1.5458E+09	3.2119E+02	2.9878E+06																		

Table A2 Continued.

SIMS ANALYSES			Mount: BR2			Standard: UNIL-QI (Piano)			Analyse: $\delta^{18}\text{O}$			Value: 9.81 ± 0.14 (VSMOW, 2 σ)			Date: 14.04.2015		
Beam	H_2^{21}O (%)	H_2^{22}O (%)	L^{17} (°C/Coef)	L^{17} (°C/Coef)	CPS	IP^{17} (°C/Coef)	CPS	Yield	Measurements (%)	Diff correction (%)	Calibration (%)	Comment					
μm	ΔD	ΔD	ΔD	ΔD	ΔD	ΔD	ΔD	$\text{CPS}/\mu\text{m}^2$	$\delta^{18}\text{O}$	$\delta^{18}\text{O}$	$\delta^{18}\text{O}_{\text{norm}}$						
									mean	SD	SD						
140515_bz_psm6@41	2.01	2.0139E-03	1.2666E-02	1.4916E-09	4.4111E-02	4.8006E-02	3.0041E-06	7.4051E+08	4.36	0.25	4.36	9.77	Standard				
140515_bz_psm6@42	2.02	2.0143E-03	7.2291E-03	1.4967E-09	5.8138E-02	6.0334E-02	3.0150E-06	7.4266E+08	4.54	0.15	4.53	9.95	Standard				
140515_bz_psm6@43	2.01	2.0136E-03	9.7661E-03	1.4941E-09	4.9058E-02	5.0592E-02	3.0085E-06	7.4173E+08	4.19	0.20	4.19	9.60	Standard				
140515_bz_psm6@44	2.01	2.0143E-03	1.0264E-02	1.4966E-09	5.7832E-02	3.0146E-06	5.8963E-02	7.4592E+08	4.52	0.21	4.52	9.93	Standard				
140515_bz_psm6@45	2.00	2.0142E-03	7.4181E-03	1.4883E-09	4.0642E-02	3.7716E-02	2.9978E-06	7.4327E+08	4.47	0.15	4.47	9.88	Standard				
140515_bz_psm6@46	2.00	2.0146E-03	1.4133E-02	1.4684E-09	4.9321E-02	5.5406E-02	2.9581E-06	7.3316E+08	4.68	0.28	4.68	10.10	Standard				
140515_bz_psm6@47	2.01	2.0145E-03	1.0599E-02	1.4818E-09	3.0994E-02	3.8462E-02	2.9841E-06	7.3796E+08	4.62	0.21	4.62	10.03	Standard				

Table A2 Continued.

Beam n/d	SIMS ANALYSES				Mount: BR3				Standard: UNL-Q1 (Paine)				Analyse: $\delta^{18}O$				Date: 14.04.2015	
	$^{112}O/^{16}O$ CPS	$^{12}C/^{16}O$ CPS	$^{13}C/^{16}O$ CPS	$^{12}C/^{16}O$ SD	$^{112}C/^{16}O$ CPS	$^{13}C/^{16}O$ CPS	Time	Yield	Measurments (%)	Drift correction (%)	Calibration (%)	Comment						
d180_mdolo10515_B3a_p7@01	1.99	2.0579E-03	9.8242E-03	1.5945E-09	4.1554E-02	3.2810E-06	4.7304E-02	14.15	8.0094E-08	8.0172E+08	26.26		0.20	28.41	26.30	0.20	31.78	0.22
d180_mdolo10515_B3a_p7@02	1.99	2.0576E-03	1.1881E-02	1.8755E-09	3.6403E-02	3.2602E-06	4.3088E-02	14.19	7.9653E-08	7.1286E+06	25.71	0.24	1.70	25.75	0.24	31.22	0.23	K158
d180_mdolo10515_B3a_p7@03	1.99	2.0543E-03	1.1860E-02	1.6048E-09	3.8438E-02	3.2939E-06	4.3818E-02	14.24	8.0583E-08		24.48	0.24		24.53	0.24	30.00	0.23	K158
d180_mdolo10515_B3a_p7@04	2.00	2.0588E-03	1.0815E-02	1.5997E-09	3.3924E-02	3.2919E-06	3.5451E-02	14.29	8.0115E-08		26.31	0.22		26.36	0.22	31.84	0.22	K158
d180_mdolo10515_B3a_p7@05	2.00	2.0585E-03	8.6944E-03	1.6991E-09	4.7592E-02	3.3093E-06	4.7838E-02	14.33	8.0595E-08		25.60	0.17		25.66	0.17	31.13	0.21	K158
d180_mdolo10515_B3a_p7@06	2.00	2.0552E-03	1.0553E-02	1.5985E-09	5.5373E-02	3.2832E-06	6.2954E-02	14.38	8.0013E-08		25.42	0.21		25.47	0.21	30.95	0.22	K158
d180_mdolo10515_B3a_p8@01	1.99	2.0553E-03	1.0402E-02	1.6098E-09	5.4207E-02	3.2834E-06	5.5081E-02	14.42	8.0478E-08	8.0602E+08	24.07	0.21	22.30	24.13	0.21	29.59	0.22	K102c
d180_mdolo10515_B3a_p8@02	1.99	2.0585E-03	9.4521E-03	1.6209E-09	5.2628E-02	3.3369E-06	5.0725E-02	14.47	8.1461E-08	9.0089E+06	26.59	0.19	16.02	26.65	0.19	32.13	0.22	K102c
d180_mdolo10515_B3a_p8@03	2.00	2.0546E-03	8.1389E-03	1.6088E-09	5.9228E-02	3.3054E-06	6.0502E-02	14.52	8.0398E-08		25.04	0.16		25.11	0.16	30.58	0.21	K102c
d180_mdolo10515_B3a_p8@04	2.00	2.0546E-03	7.3647E-03	1.6120E-09	5.1208E-02	3.3113E-06	2.5613E-02	14.56	8.0713E-08		24.64	0.15		24.71	0.15	30.18	0.21	K102c
d180_mdolo10515_B3a_p8@05	2.00	2.0544E-03	9.8535E-03	1.6038E-09	4.9651E-02	3.2946E-06	5.1795E-02	15.01	8.0277E-08		25.03	0.20		25.10	0.20	30.57	0.22	K102c
d180_mdolo10515_B3a_p8@06	2.00	2.0583E-03	1.0083E-02	1.6032E-09	5.6043E-02	3.2988E-06	5.8913E-02	15.05	8.0279E-08		26.47	0.22		26.54	0.22	32.02	0.22	K102c
d180_mdolo10515_B3a_p9@01	1.99	2.0137E-03	1.1746E-02	1.6216E-09	4.8692E-02	3.2607E-06	5.8164E-02	15.10	8.1317E-08	8.0763E+08	4.25	0.23	4.31	4.33	0.23	9.69	0.22	Standard
d180_mdolo10515_B3a_p9@02	1.99	2.0138E-03	1.4682E-02	1.6079E-09	5.4959E-02	3.2377E-06	5.7121E-02	15.15	8.0594E-08	8.4847E+06	4.19	0.29	0.23	4.27	0.29	9.63	0.24	Standard
d180_mdolo10515_B3a_p9@03	1.99	2.0139E-03	1.4136E-02	1.6090E-09	5.4740E-02	3.2405E-06	6.1151E-02	15.19	8.0771E-08		4.34	0.28		4.43	0.28	9.79	0.24	Standard
d180_mdolo10515_B3a_p9@04	2.00	2.0141E-03	1.0659E-02	1.6023E-09	3.9530E-02	3.2268E-06	4.0008E-02	15.24	8.0288E-08		4.45	0.21		4.54	0.21	9.91	0.22	Standard
d180_mdolo10515_B3a_p9@05	2.00	2.0566E-03	7.8349E-03	1.6095E-09	3.6382E-02	3.2927E-06	3.8441E-02	15.28	8.0379E-08	8.0457E+08	25.63	0.16	25.81	25.73	0.16	31.20	0.21	R-01
d180_mdolo10515_B3a_p9@06	2.00	2.0589E-03	6.5617E-03	1.6152E-09	5.5207E-02	3.3240E-06	5.7953E-02	15.33	8.0843E-08	4.3833E+06	25.27	0.13	2.85	25.37	0.13	30.85	0.20	R-01
d180_mdolo10515_B3a_p9@07	1.99	2.0529E-03	1.0474E-02	1.5905E-09	4.4644E-02	3.2741E-06	4.8653E-02	15.38	8.0251E-08		23.81	0.21		23.91	0.21	29.38	0.22	R-01
d180_mdolo10515_B3a_p9@08	2.00	2.0588E-03	1.1286E-02	1.6068E-09	5.4170E-02	3.3040E-06	5.5845E-02	15.42	8.0499E-08		25.71	0.22		25.82	0.22	31.29	0.22	R-01
d180_mdolo10515_B3a_p9@09	2.00	2.0570E-03	1.1767E-02	1.6041E-09	2.9637E-02	3.2939E-06	2.9767E-02	15.47	8.0266E-08		25.84	0.24		25.95	0.24	31.42	0.23	R-01
d180_mdolo10515_B3a_p9@10	2.00	2.0569E-03	1.0355E-02	1.6144E-09	5.5648E-02	3.3209E-06	5.8142E-02	15.51	8.0532E-08		25.80	0.21		25.91	0.21	31.39	0.22	R-01
d180_mdolo10515_B3a_p9@11	2.00	2.0602E-03	1.0984E-02	1.6014E-09	6.1231E-02	3.3039E-06	6.9828E-02	15.56	8.0096E-08	8.0481E+08	28.60	0.22	25.27	28.72	0.22	34.21	0.22	S-09
d180_mdolo10515_B3a_p9@12	1.99	2.0621E-03	7.3478E-03	1.6060E-09	3.8774E-02	3.3115E-06	3.9605E-02	16.01	8.0583E-08	8.1086E+06	28.38	0.15	18.44	28.50	0.15	33.99	0.21	S-09
d180_mdolo10515_B3a_p9@13	1.99	2.0635E-03	1.5496E-02	1.6192E-09	6.4650E-02	3.3411E-06	6.6000E-02	16.05	8.1212E-08		29.06	0.31		29.18	0.31	34.68	0.25	S-09
d180_mdolo10515_B3a_p9@14	2.00	2.0629E-03	6.7724E-03	1.6048E-09	6.3745E-02	3.3077E-06	6.5363E-02	16.10	8.0192E-08		28.76	0.14		28.89	0.14	34.38	0.21	S-09
d180_mdolo10515_B3a_p9@15	2.00	2.0620E-03	7.1795E-03	1.6086E-09	4.5577E-02	3.3171E-06	4.4597E-02	16.14	8.0428E-08		28.34	0.14		28.47	0.14	33.96	0.21	S-09
d180_mdolo10515_B3a_p9@16	2.01	2.0641E-03	9.8248E-03	1.6107E-09	5.9519E-02	3.3278E-06	5.7385E-02	16.19	8.0260E-08		29.39	0.20		29.52	0.20	35.02	0.22	S-09
d180_mdolo10515_B3a_p9@17	2.00	2.0140E-03	1.0113E-02	1.6228E-09	3.3302E-02	3.2637E-06	3.2709E-02	16.24	8.1001E-08	8.1053E+08	4.38	0.20	4.32	4.32	0.20	9.88	0.21	Standard
d180_mdolo10515_B3a_p9@18	2.00	2.0138E-03	7.3749E-03	1.6226E-09	6.7910E-02	3.2677E-06	6.5738E-02	16.28	8.1055E-08	4.1039E+06	4.29	0.15	0.21	4.43	0.15	9.79	0.20	Standard
d180_mdolo10515_B3a_p9@19	2.01	2.0139E-03	7.8518E-03	1.6245E-09	6.8367E-02	3.2714E-06	6.5860E-02	16.33	8.0845E-08		4.36	0.16		4.51	0.16	9.87	0.21	Standard
d180_mdolo10515_B3a_p9@20	2.00	2.0137E-03	9.6395E-03	1.6249E-09	5.7790E-02	3.2731E-06	5.6906E-02	16.37	8.1335E-08		4.24	0.19		4.39	0.19	9.75	0.21	Standard
d180_mdolo10515_B3a_p9@21	2.00	2.0142E-03	1.0498E-02	1.6114E-09	5.7144E-02	3.2475E-06	5.6851E-02	16.42	8.0747E-08		4.48	0.32		4.63	0.32	9.99	0.25	Standard
d180_mdolo10515_B3a_p9@22	2.00	2.0138E-03	1.0493E-02	1.6223E-09	3.9454E-02	3.2671E-06	3.7203E-02	16.47	8.1240E-08		4.31	0.21		4.46	0.21	9.82	0.22	Standard
d180_mdolo10515_B3a_p9@23	2.00	2.0135E-03	1.6753E-02	1.6213E-09	2.1055E-02	3.2649E-06	2.3084E-02	16.52	8.1090E-08		4.15	0.34		4.31	0.34	9.67	0.25	Standard

Table A2 Continued.

SIMS ANALYSES		Mount: BR4		Standard: UNL-QI (Piano)		Analyse: $\delta^{18}O$		Date: 15/04/2015		
Beam #1	$172 (^{\circ}C, vac)$ cps	$172 (^{\circ}C, vac)$ cps	Time	Yield cps/na	Measurements (%) $\delta^{18}O$	mean $\delta^{18}O$	Drift correction (%) $\delta^{18}O$	Calibration (%) $\delta^{18}O_{VSMOW}$	Comment	
2.04	2.0523E+03	1.6492E+09	4.713E+02	4.417E+02	3.743E+06	23.59	22.46	23.58	29.68	0.18
2.04	2.0470E+03	8.981E+03	7.103E+02	7.740E+02	3.374E+06	8.906E+08	8.031E+08	8.031E+08	29.68	0.18
2.04	2.0470E+03	9.895E+03	7.404E+02	7.383E+02	3.426E+06	8.030E+08	8.094E+06	8.094E+06	26.95	0.17
2.03	2.0506E+03	6.779E+03	5.264E+02	5.402E+02	3.358E+06	20.87	1.98	20.86	28.71	0.16
2.03	2.0495E+03	1.6380E+09	6.003E+02	6.327E+02	3.385E+06	22.62	0.14	22.61	28.20	0.17
2.04	2.0511E+03	8.989E+03	5.819E+02	6.029E+02	3.324E+06	22.87	0.18	22.86	28.96	0.17
2.04	2.0471E+03	8.684E+03	6.099E+02	5.748E+02	3.454E+06	20.88	0.17	20.88	26.96	0.17
2.04	2.0518E+03	9.362E+03	6.140E+02	6.359E+02	3.356E+06	23.24	0.19	23.23	29.33	0.17
2.04	2.0511E+03	1.0948E+02	4.660E+02	4.378E+06	3.378E+06	22.90	0.22	22.90	29.00	0.18
2.03	2.0515E+03	1.6306E+09	6.640E+02	6.857E+02	3.344E+06	23.16	0.23	23.09	29.19	0.18
2.03	2.0128E+03	6.816E+03	7.202E+02	7.597E+02	3.353E+06	3.78	0.14	3.78	9.76	0.16
2.03	2.0131E+03	1.0038E+02	7.317E+02	7.272E+02	3.332E+06	3.94	0.20	3.94	9.93	0.18
2.04	2.0126E+03	8.963E+03	6.202E+02	6.390E+02	3.359E+06	3.67	0.18	3.67	9.65	0.18
2.03	2.0127E+03	1.7129E+09	3.931E+01	3.956E+01	3.395E+06	3.72	0.21	3.72	9.70	0.18
2.08	2.0509E+03	8.6119E+03	6.125E+02	6.049E+02	3.393E+06	23.38	0.17	23.56	31.88	0.17
2.03	2.0506E+03	1.1792E+02	6.880E+09	7.228E+02	3.408E+06	7.105E+02	25.35	0.24	31.46	0.19
2.03	2.0577E+03	7.751E+03	6.378E+02	6.749E+02	3.404E+06	26.16	0.16	26.16	32.28	0.17
2.04	2.0587E+03	1.1418E+02	6.170E+09	6.438E+06	3.438E+06	25.67	0.23	25.68	31.79	0.19
2.06	2.0573E+03	1.0457E+02	6.438E+09	6.019E+02	3.381E+06	26.00	0.21	26.00	32.12	0.18
2.04	2.0585E+03	1.3100E+02	6.476E+09	4.084E+02	3.387E+06	25.59	0.26	25.59	31.71	0.20
2.04	2.0567E+03	9.9386E+03	6.640E+09	6.976E+02	3.385E+06	25.67	0.20	25.68	31.79	0.18
2.04	2.0570E+03	7.239E+03	4.422E+02	4.400E+02	3.372E+06	25.82	0.14	25.82	31.94	0.16
2.03	2.0532E+03	1.1217E+02	6.195E+09	5.974E+02	3.419E+06	23.94	0.24	23.95	30.06	0.19
2.04	2.0506E+03	8.1340E+03	6.658E+09	6.496E+02	3.424E+06	25.64	0.16	25.64	31.76	0.17
2.04	2.0131E+03	7.759E+03	5.530E+02	5.426E+02	3.384E+06	3.94	0.16	3.94	9.93	0.16
2.04	2.0129E+03	1.0935E+02	6.224E+02	6.860E+02	3.381E+06	3.86	0.22	3.87	9.86	0.18
2.03	2.0128E+03	1.674E+09	6.524E+02	6.510E+02	3.370E+06	3.77	0.20	3.78	9.77	0.18
2.03	2.0128E+03	9.539E+03	7.800E+02	7.800E+02	3.383E+06	3.78	0.19	3.79	9.78	0.17
2.04	2.0643E+03	1.0391E+02	5.327E+02	5.501E+02	3.476E+06	29.59	0.21	29.14	35.74	0.18
2.04	2.0643E+03	1.2728E+02	6.692E+09	7.402E+02	3.446E+06	29.36	0.25	29.36	35.50	0.20
2.04	2.0633E+03	9.800E+03	6.681E+02	6.681E+02	3.421E+06	29.08	0.20	29.09	35.23	0.18
2.03	2.0628E+03	9.429E+03	6.657E+09	6.919E+02	3.435E+06	28.73	0.19	28.74	34.88	0.18
2.03	2.0629E+03	1.3748E+02	6.839E+02	6.332E+02	3.455E+06	28.77	0.27	28.78	34.91	0.20
2.03	2.0639E+03	8.532E+03	5.667E+02	5.873E+02	3.449E+06	29.29	0.17	29.29	35.45	0.17
2.03	2.0628E+03	9.052E+03	6.831E+09	7.098E+02	3.471E+06	28.74	0.18	28.76	34.89	0.17
2.04	2.0640E+03	1.899E+03	6.097E+02	6.198E+02	3.432E+06	29.38	0.18	29.36	35.50	0.17
2.04	2.0641E+03	1.0231E+02	6.616E+09	6.783E+02	3.429E+06	29.39	0.20	29.41	35.55	0.17
2.04	2.0129E+03	1.0114E+02	5.710E+02	6.081E+02	3.369E+06	3.83	0.20	3.78	9.83	0.18
2.04	2.0126E+03	1.0055E+02	6.423E+02	6.483E+02	3.383E+06	3.71	0.20	3.73	9.71	0.18
2.03	2.0123E+03	1.2408E+02	5.984E+02	6.575E+02	3.375E+06	3.64	0.24	3.66	9.64	0.19
2.03	2.0131E+03	8.726E+03	7.380E+02	7.576E+02	3.382E+06	3.96	0.17	3.98	9.97	0.17
2.03	2.0506E+03	1.2459E+02	6.840E+09	6.423E+02	3.421E+06	23.33	0.25	23.33	29.35	0.25
2.03	2.0580E+03	7.899E+03	5.624E+02	5.508E+02	3.434E+06	26.37	0.16	26.37	32.49	0.17
2.03	2.0587E+03	9.160E+03	8.850E+02	8.839E+02	3.454E+06	25.30	0.18	25.32	31.33	0.17
2.03	2.0587E+03	7.726E+03	6.448E+02	6.013E+02	3.448E+06	25.70	0.15	25.72	31.84	0.17
2.03	2.0529E+03	9.530E+03	8.542E+02	8.668E+02	3.434E+06	23.78	0.19	23.81	29.91	0.18
2.03	2.0537E+03	9.0134E+03	5.272E+02	5.679E+02	3.387E+06	24.21	0.18	24.23	30.34	0.17
2.03	2.0556E+03	1.0172E+02	6.617E+09	7.309E+02	3.4159E+06	25.14	0.20	25.16	31.28	0.18
2.03	2.0569E+03	1.687E+09	6.906E+02	6.758E+02	3.438E+06	25.76	0.19	25.79	31.90	0.18
2.03	2.0553E+03	1.0145E+02	7.341E+02	7.482E+02	3.444E+06	25.00	0.20	25.02	31.14	0.18
2.03	2.0570E+03	9.690E+03	6.995E+02	7.450E+02	3.450E+06	25.82	0.19	25.85	31.96	0.18
2.03	2.0129E+03	7.291E+03	3.398E+02	3.398E+02	3.398E+06	3.86	0.15	3.73	9.88	0.16
2.04	2.0123E+03	1.310E+02	6.211E+02	6.211E+02	3.377E+06	3.53	0.26	3.56	9.54	0.19
2.03	2.0132E+03	6.478E+03	5.498E+02	5.318E+02	3.367E+06	4.00	0.13	4.03	10.02	0.16
2.03	2.0123E+03	9.891E+03	6.800E+02	7.173E+02	3.340E+06	3.55	0.20	3.58	9.56	0.18
2.03	2.0039E+03	1.1245E+02	5.660E+02	5.705E+02	3.448E+06	29.25	0.22	29.28	35.42	0.19
2.04	2.0033E+03	9.627E+03	6.517E+02	6.510E+02	3.482E+06	28.92	0.19	28.95	35.09	0.18
2.03	2.0045E+03	7.629E+03	4.943E+02	5.256E+02	3.433E+06	29.59	0.15	29.62	35.76	0.17
2.03	2.0028E+03	1.0990E+02	8.105E+02	8.205E+02	3.480E+06	28.70	0.22	28.74	34.87	0.17
2.03	2.0028E+03	1.3228E+02	7.345E+02	7.485E+02	3.454E+06	28.73	0.24	28.77	34.90	0.19
2.03	2.0032E+03	9.567E+03	6.650E+02	6.455E+02	3.455E+06	28.92	0.19	28.96	35.10	0.18
2.03	2.0034E+03	9.134E+03	7.814E+02	7.945E+02	3.454E+06	29.02	0.18	29.06	35.19	0.17
2.03	2.0036E+03	7.789E+03	7.466E+02	7.466E+02	3.457E+06	29.24	0.16	29.28	35.42	0.17
2.04	2.0036E+03	1.182E+02	6.421E+02	6.509E+02	3.429E+06	29.15	0.24	29.19	35.32	0.19

Table A2 Continued.

Beam no.	SIMS ANALYSES			Mount: BR4			Standard: UNIL-QI (Paine)			Analyse: $\delta^{18}O$			Value: 9.81 ± 0.14 (VSMOW, σ)			Date: 18.04.2015	Comment
	$112U/12C$ (cps)	$17O/16O$ (cps)	$18O/16O$ (cps)	$17O/16O$ (cps)	$18O/16O$ (cps)	Time	Yield CPS/A	Measurements (%) $\delta^{18}O$	mean $\delta^{18}O$	Diff. correction $\delta^{18}O$	Calibration (%) $\delta^{18}O_{VSMOW}$						
2.03	2.0128E+03	9.4525E+03	1.6693E+09	5.6510E+02	3.3403E+06	5.2901E+02	3.22	8.2374E+08	8.2356E+08	3.80	0.19	3.75	3.83	0.19	9.82	0.17	Standard
2.02	2.0124E+03	1.1556E+02	1.6689E+09	6.9823E+02	3.3585E+06	6.9862E+02	3.26	8.2425E+08	1.0853E+06	3.58	0.23	0.26	3.62	0.23	9.60	0.18	Standard
2.03	2.0130E+03	6.5448E+03	1.6689E+09	6.9907E+02	3.3597E+06	6.1402E+02	3.31	8.2310E+08		3.89	0.13		3.93	0.13	9.91	0.16	Standard
2.03	2.0127E+03	1.1426E+02	1.6739E+09	5.5716E+02	3.3695E+06	5.4761E+02	3.35	8.2331E+08		3.73	0.23		3.77	0.23	9.76	0.18	Standard
2.03	2.0569E+03	1.2726E+02	1.6393E+09	7.8974E+02	3.3720E+06	8.5507E+02	3.40	8.0672E+08	8.1658E+08	25.80	0.25	17.93	25.84	0.25	31.96	0.20	Sc21
2.03	2.0569E+03	8.6852E+03	1.7088E+09	1.0990E+01	3.4948E+06	1.0932E+01	3.44	8.3935E+08	2.9905E+07	25.33	0.17	19.67	25.37	0.17	31.49	0.17	Sc21
2.04	2.0632E+03	1.4332E+02	1.6862E+09	6.6989E+02	3.4335E+06	6.7297E+02	3.49	8.2744E+08		15.52	0.29		27.62	0.16	33.75	0.17	Sc21
2.05	2.0605E+03	7.5222E+03	1.6728E+09	6.7246E+02	3.4448E+06	6.9840E+02	3.53	8.1767E+08		27.57	0.16						Sc21
2.04	2.0136E+03	8.8607E+03	1.6799E+09	6.1453E+02	3.3825E+06	5.9988E+02	3.58	8.2334E+08		4.19	0.18						Sc21
2.03	2.0128E+03	1.5543E+02	1.6625E+09	8.5399E+02	3.3446E+06	8.7025E+02	4.02	8.1888E+08		3.77	0.31						Sc21
2.03	2.0235E+03	9.8608E+03	1.6563E+09	1.5207E+01	3.3515E+06	1.5392E+01	4.07	8.1449E+08		9.10	0.20						Sc21
2.04	2.0458E+03	1.3207E+02	1.6498E+09	1.0724E+01	3.3911E+06	1.0481E+01	4.11	8.0846E+08		25.23	0.26		25.27	0.26	31.39	0.20	Sc21
2.04	2.0582E+03	9.9306E+03	1.6654E+09	1.0453E+01	3.4204E+06	1.0807E+01	4.16	8.1395E+08		25.45	0.20		25.50	0.20	31.62	0.18	Sc21
2.04	2.0593E+03	1.2063E+02	1.6390E+09	7.4606E+02	3.3751E+06	7.3852E+02	4.20	8.0425E+08		26.97	0.24		27.02	0.24	33.15	0.19	Sc21
2.03	2.0218E+03	1.4269E+02	1.6421E+09	6.7439E+02	3.3199E+06	7.0025E+02	4.25	8.0779E+08		8.27	0.29						Sc21
2.03	2.0122E+03	1.0628E+02	1.6738E+09	5.6269E+02	3.3687E+06	5.6469E+02	4.29	8.2387E+08	8.2506E+08	3.62	0.21	3.73	3.67	0.21	9.65	0.18	Standard
2.03	2.0127E+03	1.2121E+02	1.6810E+09	7.8859E+02	3.3833E+06	7.7703E+02	4.34	8.2699E+08	3.3265E+06	3.72	0.24	0.19	3.77	0.24	9.76	0.19	Standard
2.04	2.0123E+03	9.1256E+03	1.6765E+09	6.7367E+02	3.3736E+06	7.0797E+02	4.38	8.2308E+08		3.55	0.18		3.60	0.18	9.58	0.17	Standard
2.04	2.0128E+03	1.2353E+02	1.6848E+09	6.5180E+02	3.3916E+06	6.4496E+02	4.42	8.2769E+08		3.80	0.25		3.85	0.25	9.84	0.19	Standard
2.04	2.0129E+03	9.3673E+03	1.6881E+09	5.5338E+02	3.3976E+06	5.5380E+02	4.47	8.2695E+08		3.82	0.19		3.87	0.19	9.86	0.17	Standard
2.04	2.0127E+03	1.4054E+02	1.6846E+09	7.3290E+02	3.3906E+06	7.3790E+02	4.51	8.2616E+08		3.75	0.28		3.81	0.28	9.79	0.20	Standard
2.04	2.0127E+03	1.3301E+02	1.6849E+09	5.3908E+02	3.3916E+06	5.3300E+02	4.56	8.2679E+08		3.74	0.27		3.80	0.27	9.79	0.20	Standard
2.03	2.0128E+03	6.0876E+03	1.6806E+09	7.0343E+02	3.3830E+06	7.1467E+02	5.00	8.2698E+08		3.81	0.12		3.86	0.12	9.85	0.16	Standard

Table A2 Continued.

SIMS ANALYSES										Mount: BR6			Standard: UNIL-QI (Paino)			Analyse: $\delta^{18}O$			Value: 9.81 ± 0.14 (VSMOW, 2 σ)			Date: 15.04.2015
Beam	$^{12}O/^{16}O$ (V)	$^{17}O/^{16}O$ (V)	$^{18}O/^{16}O$ (V)	$^{17}O/^{16}O$ (C)	$^{18}O/^{16}O$ (C)	$^{17}O/^{16}O$ (C)	$^{18}O/^{16}O$ (C)	Time	Yield	Measurements	Mean	Diff. correction	Calibration	Comment								
z4	CFB	CFB	CFB	CFB	CFB	CFB	CFB	min	$\delta^{18}O$	$\delta^{18}O$	$\delta^{18}O$	$\delta^{18}O$	$\delta^{18}O$									
	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	SD	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$	$\pm 2SD$									
41802_radiol140515_Bre_p06/1	2.04	2.0623E-03	9.8599E-03	1.6298E-09	9.8028E-02	3.3491E-06	1.0313E-01	14:20	7.9946E-08	8.0161E-08	28.67	0.20	28.76	28.71	0.20	35.03	0.19	Sc26				
41802_radiol140515_Bre_p06/2	2.04	2.0623E-03	8.9823E-03	1.6240E-09	9.5544E-02	3.3491E-06	9.9399E-02	14:24	7.9640E-08	7.8550E-06	28.49	0.18	0.55	28.54	0.18	34.85	0.18	Sc26				
41802_radiol140515_Bre_p06/3	2.04	2.0624E-03	9.9467E-03	1.6371E-09	1.2179E-01	3.3763E-06	1.1920E-01	14:29	8.0410E-08	8.6612E-02	28.53	0.20	0.22	28.58	0.20	34.89	0.19	Sc26				
41802_radiol140515_Bre_p06/4	2.04	2.0638E-03	1.1112E-02	1.6444E-09	8.8016E-02	3.3938E-06	8.6612E-02	14:34	8.0772E-08	8.6612E-02	29.25	0.22	0.22	29.30	0.22	35.62	0.20	Sc26				
41802_radiol140515_Bre_p06/5	2.04	2.0629E-03	1.0914E-02	1.6333E-09	8.2866E-02	3.3693E-06	8.5715E-02	14:38	8.0904E-08	8.5715E-02	28.78	0.22	0.22	28.84	0.22	35.15	0.19	Sc26				
41802_radiol140515_Bre_p06/6	2.04	2.0620E-03	9.2247E-03	1.6370E-09	1.1487E-01	3.3774E-06	1.1415E-01	14:43	8.0165E-08	8.0165E-08	28.84	0.18	0.18	28.90	0.18	35.22	0.18	Sc26				
41802_radiol140515_Bre_p06/21	2.04	2.0128E-03	1.6277E-09	9.1939E-02	3.2760E-06	8.9803E-06	8.9803E-06	14:48	7.9848E-08	7.9743E-08	3.71	0.19	3.57	3.77	0.19	9.93	0.18	Standard				
41802_radiol140515_Bre_p06/22	2.04	2.0124E-03	1.6259E-02	1.6347E-09	1.0899E-01	3.2898E-06	1.0965E-01	14:52	8.0031E-08	5.3192E-06	3.60	0.21	0.37	3.67	0.21	9.83	0.19	Standard				
41802_radiol140515_Bre_p06/23	2.06	2.0118E-03	9.4764E-03	1.6371E-09	8.6866E-02	3.2827E-06	8.3940E-02	14:57	7.9389E-08	8.3940E-02	3.30	0.19	0.19	3.37	0.19	9.53	0.18	Standard				
41802_radiol140515_Bre_p06/24	2.05	2.0120E-03	6.1522E-03	1.6339E-09	8.7112E-02	3.2835E-06	8.7837E-02	15:02	7.9758E-08	8.7837E-02	3.68	0.12	0.12	3.75	0.12	9.91	0.17	Standard				
41802_radiol140515_Bre_p06/31	2.05	2.0624E-03	8.0223E-03	1.6290E-09	1.1387E-01	3.3400E-06	1.1390E-01	15:06	7.9511E-08	7.9483E-08	28.54	0.16	28.52	28.62	0.16	34.93	0.18	Sc12				
41802_radiol140515_Bre_p06/32	2.04	2.0629E-03	1.0816E-02	1.6147E-09	9.1531E-02	3.3311E-06	9.4773E-02	15:11	7.9120E-08	7.5463E-06	28.80	0.22	0.96	28.88	0.22	35.19	0.19	Sc12				
41802_radiol140515_Bre_p06/33	2.04	2.0606E-03	1.3828E-02	1.6191E-09	1.1003E-01	3.3516E-06	1.0988E-01	15:16	7.9519E-08	27.63	0.28	0.28	27.71	0.28	34.01	0.21	Sc12					
41802_radiol140515_Bre_p06/34	2.04	2.0627E-03	7.7437E-03	1.6132E-09	8.8813E-02	3.3273E-06	9.0414E-02	15:20	7.9173E-08	8.6612E-02	28.66	0.15	28.74	28.74	0.15	35.05	0.18	Sc12				
41802_radiol140515_Bre_p06/35	2.01	2.0634E-03	7.6818E-03	1.6122E-09	1.1141E-01	3.3248E-06	1.1283E-01	15:42	8.0173E-08	29.03	0.15	29.13	29.13	0.15	35.44	0.18	Sc12					
41802_radiol140515_Bre_p06/36	2.03	2.0623E-03	9.7393E-03	1.6116E-09	9.5404E-02	3.3238E-06	9.5484E-02	15:47	7.9402E-08	28.48	0.19	28.48	28.58	0.19	34.90	0.19	Sc12					
41802_radiol140515_Bre_p10/01	2.03	2.0467E-03	1.2809E-02	1.6129E-09	9.1036E-02	3.3010E-06	8.9446E-02	15:51	7.9351E-08	7.9023E-08	20.70	0.26	20.17	20.80	0.26	27.06	0.20	Gr68				
41802_radiol140515_Bre_p10/02	2.04	2.0444E-03	9.4878E-03	1.6057E-09	1.0884E-01	3.2788E-06	1.0660E-01	15:56	7.8477E-08	1.0667E-07	19.53	0.19	1.13	19.63	0.19	25.89	0.18	Gr68				
41802_radiol140515_Bre_p10/03	2.04	2.0433E-03	8.3673E-03	1.6034E-09	9.2713E-02	3.2792E-06	8.8639E-02	16:01	7.8609E-08	20.02	0.17	20.13	20.13	0.17	26.39	0.18	Gr68					
41802_radiol140515_Bre_p10/04	2.04	2.0455E-03	1.2785E-02	1.6043E-09	1.0073E-01	3.2815E-06	9.8914E-02	16:06	7.8626E-08	20.10	0.26	20.21	20.21	0.26	26.47	0.20	Gr68					
41802_radiol140515_Bre_p10/05	2.04	2.0447E-03	9.0250E-03	1.6268E-09	1.0858E-01	3.3243E-06	1.1120E-01	16:10	7.9533E-08	19.69	0.18	19.80	19.80	0.18	26.06	0.18	Gr68					
41802_radiol140515_Bre_p10/06	2.05	2.0471E-03	8.3413E-03	1.6303E-09	9.1741E-02	3.3374E-06	9.3787E-02	16:15	7.9533E-08	20.98	0.17	21.09	21.09	0.17	27.36	0.18	Gr68					
41802_radiol140515_Bre_p06/25	2.05	2.0118E-03	1.6203E-02	1.6288E-09	8.3739E-02	3.2771E-06	8.4823E-02	16:20	7.9446E-08	7.9513E-08	3.30	0.20	3.49	3.42	0.20	9.57	0.19	Standard				
41802_radiol140515_Bre_p06/26	2.04	2.0127E-03	8.2003E-03	1.6238E-09	1.0693E-01	3.2683E-06	1.0829E-01	16:24	7.9408E-08	2.3744E-06	3.75	0.16	0.37	3.87	0.16	10.03	0.18	Standard				
41802_radiol140515_Bre_p06/27	2.04	2.0121E-03	7.8594E-03	1.6228E-09	1.0413E-01	3.2651E-06	1.0365E-01	16:29	7.9520E-08	3.45	0.16	0.16	3.57	0.16	9.73	0.17	Standard					
41802_radiol140515_Bre_p06/28	2.05	2.0122E-03	8.3880E-03	1.6296E-09	9.4821E-02	3.2792E-06	9.4330E-02	16:34	7.9676E-08	3.49	0.17	3.62	3.62	0.17	9.77	0.18	Standard					

Table A2 Continued.

SIMS ANALYSES										Mount-BK7			Analyse: $\delta^{18}\text{O}$			Value: 98144.14 (VSMOW, 2 σ)			Date: 13.04.2015		
Beam	n_f	$11\text{T}12\text{L}2\text{ (}^{16}\text{O}^+$)	$11\text{T}12\text{L}2\text{ (}^{17}\text{O}^+$)	$11\text{T}12\text{L}2\text{ (}^{18}\text{O}^+$)	$12\text{L}2\text{ (}^{16}\text{O}^{\text{eff}})$	$12\text{L}2\text{ (}^{17}\text{O}^{\text{eff}})$	$12\text{L}2\text{ (}^{18}\text{O}^{\text{eff}})$	$12\text{L}2\text{ (}^{16}\text{O}^{\text{eff}})$	$12\text{L}2\text{ (}^{17}\text{O}^{\text{eff}})$	$12\text{L}2\text{ (}^{18}\text{O}^{\text{eff}})$	Time	Yield	Measurements (%)	Drift correction (%)	Calibration (%)	Comment					
		CPA	CPA	CPA	CPA	CPA	CPA	CPA	CPA	CPA	min	CPA	2SD	2SD	2SD						
130415_pane@1	2.04	2.0156E+03	7.8006E+03	1.0012E+01	3.1228E+06	1.0012E+01	3.1228E+06	1.0012E+01	3.1228E+06	1.0012E+01	12.46	7.6238E-08	5.21	0.16	5.26	Standard					
130415_pane@2	2.03	2.0151E+03	9.4127E+03	1.0082E+01	3.1871E+06	1.0082E+01	3.1871E+06	1.0082E+01	3.1871E+06	1.0082E+01	12.50	7.6238E-08	4.94	0.19	1.82	Standard					
130415_pane@3	2.03	2.0157E+03	1.2454E+02	7.0947E+02	3.2071E+06	7.0947E+02	3.2071E+06	7.0947E+02	3.2071E+06	7.0947E+02	12.54	7.8302E-08	5.22	0.25		Standard					
130415_pane@4	2.03	2.0157E+03	1.2454E+02	7.0947E+02	3.2071E+06	7.0947E+02	3.2071E+06	7.0947E+02	3.2071E+06	7.0947E+02	12.54	7.8302E-08	5.22	0.25		Standard					
130415_pane@5	2.03	2.0166E+03	7.6623E+03	1.3474E+01	3.3027E+06	1.3474E+01	3.3027E+06	1.3474E+01	3.3027E+06	1.3474E+01	13.03	7.4605E-08	5.22	0.14		Standard					
130415_pane@6	2.03	2.0162E+03	9.1504E+03	1.5383E+01	6.8801E+02	3.1095E+06	6.8801E+02	3.1095E+06	6.8801E+02	3.1095E+06	13.07	7.5718E-08	5.50	0.18		Standard					
130415_pane@7	2.03	2.0167E+03	1.0013E+02	1.5109E+01	2.7171E+02	3.0650E+06	2.7171E+02	3.0650E+06	2.7171E+02	3.0650E+06	13.11	7.4896E-08	5.71	0.22		Standard					
130415_pane@8	1.99	2.0170E+03	1.2067E+02	1.5135E+01	3.0277E+06	3.0277E+06	3.0277E+06	1.3241E+01	3.0277E+06	1.3241E+01	13.25	7.6155E-08	5.90	0.24		Standard					
130415_pane@9	1.99	2.0167E+03	1.3372E+02	1.5321E+01	8.0729E+02	3.0875E+06	8.0729E+02	3.0875E+06	8.0729E+02	3.0875E+06	13.29	7.6895E-08	5.73	0.27		Standard					
130415_pane@10	1.99	2.0163E+03	1.1646E+02	1.5304E+01	7.1198E+02	3.0875E+06	7.1198E+02	3.0875E+06	7.1198E+02	3.0875E+06	13.33	7.6621E-08	5.54	0.23		Standard					
130415_pane@11	1.99	2.0163E+03	1.1646E+02	1.5304E+01	7.1198E+02	3.0875E+06	7.1198E+02	3.0875E+06	7.1198E+02	3.0875E+06	13.33	7.6621E-08	5.54	0.23		Standard					
130415_pane@12	1.99	2.0171E+03	1.0453E+02	1.5319E+01	5.3360E+02	3.0889E+06	5.3360E+02	3.0889E+06	5.3360E+02	3.0889E+06	13.37	7.6908E-08	5.96	0.21		Standard					
130415_pane@13	2.00	2.0174E+03	1.0641E+02	1.5045E+01	8.9284E+02	3.0350E+06	8.9284E+02	3.0350E+06	8.9284E+02	3.0350E+06	14.02	7.5318E-08	6.10	0.21		Standard					
130415_pane@14	1.97	2.0180E+03	1.0980E+02	1.3095E+01	2.8525E+02	2.6426E+06	2.8525E+02	2.6426E+06	2.8525E+02	2.6426E+06	13.49	6.6623E-08	6.39	0.22		Standard					
130415_pane@15	1.97	2.0184E+03	8.4346E+03	1.3226E+01	1.8888E+01	2.6888E+06	1.3226E+01	2.6888E+06	1.3226E+01	2.6888E+06	14.13	6.7607E-08	6.56	0.17		Standard					
130415_pane@16	1.97	2.0179E+03	1.0718E+02	1.3408E+01	1.8888E+01	2.7025E+06	1.3408E+01	2.7025E+06	1.3408E+01	2.7025E+06	14.13	6.8276E-08	6.54	0.21		Standard					
130415_pane@17	1.96	2.0188E+03	1.3144E+02	1.3246E+01	1.8181E+01	2.6798E+06	1.3246E+01	2.6798E+06	1.3246E+01	2.6798E+06	14.21	6.7586E-08	6.79	0.28		Standard					
130415_pane@18	1.96	2.0188E+03	1.3144E+02	1.3246E+01	1.8181E+01	2.6798E+06	1.3246E+01	2.6798E+06	1.3246E+01	2.6798E+06	14.21	6.7586E-08	6.79	0.28		Standard					
130415_pane@19	1.93	2.0162E+03	1.1524E+02	1.3349E+01	4.1660E+02	2.6401E+06	4.1660E+02	2.6401E+06	4.1660E+02	2.6401E+06	14.41	6.9388E-08	5.49	0.23		Standard					
130415_pane@20	1.93	2.0162E+03	1.1524E+02	1.3349E+01	4.1660E+02	2.6401E+06	4.1660E+02	2.6401E+06	4.1660E+02	2.6401E+06	14.41	6.9388E-08	5.49	0.23		Standard					
130415_pane@21	1.94	2.0165E+03	1.3010E+02	1.3168E+01	2.7791E+02	2.6554E+06	2.7791E+02	2.6554E+06	2.7791E+02	2.6554E+06	14.46	6.8206E-08	5.65	0.26		Standard					
130415_pane@22	1.94	2.0165E+03	1.3010E+02	1.3168E+01	2.7791E+02	2.6554E+06	2.7791E+02	2.6554E+06	2.7791E+02	2.6554E+06	14.46	6.8206E-08	5.65	0.26		Standard					
130415_pane@23	1.94	2.0155E+03	1.5079E+02	1.3085E+01	2.0965E+02	2.6420E+06	2.0965E+02	2.6420E+06	2.0965E+02	2.6420E+06	14.50	6.7723E-08	5.22	0.33		Standard					
130415_pane@24	1.94	2.0155E+03	1.5079E+02	1.3085E+01	2.0965E+02	2.6420E+06	2.0965E+02	2.6420E+06	2.0965E+02	2.6420E+06	14.50	6.7723E-08	5.22	0.33		Standard					
130415_pane@25	1.94	2.0163E+03	1.4509E+02	1.3067E+01	1.6958E+02	2.6447E+06	1.6958E+02	2.6447E+06	1.6958E+02	2.6447E+06	14.54	6.7252E-08	5.15	0.30		Standard					
130415_pane@26	1.93	2.0159E+03	1.3068E+02	1.2995E+01	4.8062E+02	2.6194E+06	4.8062E+02	2.6194E+06	4.8062E+02	2.6194E+06	14.58	6.7459E-08	5.53	0.29		Standard					
130415_pane@27	1.93	2.0163E+03	1.5462E+02	1.3044E+01	2.8943E+02	2.6217E+06	2.8943E+02	2.6217E+06	2.8943E+02	2.6217E+06	15.07	6.7227E-08	5.54	0.26		Standard					
130415_pane@28	1.93	2.0158E+03	1.2187E+02	1.3071E+01	4.1994E+02	2.6318E+06	4.1994E+02	2.6318E+06	4.1994E+02	2.6318E+06	15.11	6.7624E-08	5.28	0.24		Standard					
130415_pane@29	1.93	2.0133E+03	1.6687E+02	1.4018E+01	5.7281E+02	3.0044E+06	5.7281E+02	3.0044E+06	5.7281E+02	3.0044E+06	15.17	7.7501E-08	4.02	0.33		Standard					
130415_pane@30	1.93	2.0129E+03	6.9970E+03	1.4888E+01	4.7419E+02	2.9972E+06	4.7419E+02	2.9972E+06	4.7419E+02	2.9972E+06	15.21	7.9004E-08	3.86	0.14		Standard					
130415_pane@31	1.93	2.0130E+03	8.8625E+03	1.5904E+01	8.4649E+02	3.0798E+06	8.4649E+02	3.0798E+06	8.4649E+02	3.0798E+06	15.25	7.9794E-08	3.86	0.14		Standard					
130415_pane@32	2.00	2.0128E+03	1.0751E+02	1.5833E+01	3.4511E+02	3.1485E+06	3.4511E+02	3.1485E+06	3.4511E+02	3.1485E+06	15.30	7.7717E-08	3.79	0.22		Standard					
130415_pane@33	2.02	2.0128E+03	9.1744E+03	1.5627E+01	3.3777E+02	3.1485E+06	3.3777E+02	3.1485E+06	3.3777E+02	3.1485E+06	15.34	7.7182E-08	3.69	0.18		Standard					
130415_pane@34	2.04	2.0130E+03	1.2846E+02	1.5871E+01	6.8061E+02	3.1965E+06	6.8061E+02	3.1965E+06	6.8061E+02	3.1965E+06	15.38	7.7530E-08	3.90	0.26		Standard					
130415_pane@35	2.05	2.0132E+03	1.2803E+02	1.5844E+01	6.4136E+02	3.2175E+06	6.4136E+02	3.2175E+06	6.4136E+02	3.2175E+06	15.42	7.8004E-08	3.97	0.26		Standard					
130415_pane@36	2.04	2.0129E+03	9.9295E+03	1.6026E+01	4.8911E+02	3.2321E+06	4.8911E+02	3.2321E+06	4.8911E+02	3.2321E+06	15.47	7.8405E-08	3.83	0.20		Standard					
130415_BK7_pane@1	2.04	2.0127E+03	1.0896E+02	1.5984E+01	7.6244E+02	3.1901E+06	7.6244E+02	3.1901E+06	7.6244E+02	3.1901E+06	16.56	7.6434E-08	3.76	0.22	9.64	0.20					
130415_BK7_pane@2	2.04	2.0133E+03	1.4286E+02	1.5709E+01	4.3023E+02	3.1629E+06	4.3023E+02	3.1629E+06	4.3023E+02	3.1629E+06	17.00	7.6624E-08	4.04	0.29	3.99	0.29					
130415_BK7_pane@3	2.05	2.0128E+03	1.0913E+02	1.5792E+01	3.8833E+02	3.1702E+06	3.8833E+02	3.1702E+06	3.8833E+02	3.1702E+06	17.04	7.6933E-08	3.79	0.22	3.74	0.22					
130415_BK7_pane@4	2.05	2.0132E+03	1.1310E+02	1.5805E+01	3.4133E+02	3.1854E+06	3.4133E+02	3.1854E+06	3.4133E+02	3.1854E+06	17.08	7.6933E-08	4.01	0.23	3.96	0.23					
130415_BK7_pane@5	2.06	2.0127E+03	1.3649E+02	1.5863E+01	3.4889E+02	3.1972E+06	3.4889E+02	3.1972E+06	3.4889E+02	3.1972E+06	17.13	7.6956E-08	3.72	0.27	3.67	0.27					
130415_BK7_pane@6	2.06	2.0129E+03	1.2716E+02	1.5639E+01	4.3366E+02	3.1726E+06	4.3366E+02	3.1726E+06	4.3366E+02	3.1726E+06	17.17	7.6908E-08	3.84	0.25	3.80	0.25					
130415_BK7_pane@7	2.06	2.0138E+03	1.2409E+02	1.5754E+01	3.7702E+02	3.1684E+06	3.7702E+02	3.1684E+06	3.7702E+02	3.1684E+06	17.21	7.6471E-08	4.30	0.25	4.26	0.25					
130415_BK7_pane@8	2.06	2.0131E+03	6.2288E+03	1.5904E+01	2.8036E+01	3.2015E+06	2.8036E+01	3.2015E+06	2.8036E+01	3.2015E+06	17.26	7.7191E-08	3.92	0.12	9.80	0.18					
130415_BK7_pane@9	2.06	2.0134E+03	1.2815E+02	1.5918E+01	4.2764E+02	3.2048E+06	4.2764E+02	3.2048E+06	4.2764E+02	3.2048E+06	17.30	7.7323E-08	4.08	0.26	4.04	0.26					
130415_BK7_pane@10	2.06	2.0135E+03	7.8612E+03	1.6008E+01	2.3403E+02	3.2232E+06	2.3403E+02	3.2232E+06	2.3403E+02	3.2232E+06	17.34	7.7731E-08	4.14	0.16	10.02	0.19					
130415_BK7_pane@11	2.06	2.0296E+03	1.4348E+02	1.9044E+01	2.3808E+01	3.9183E+06	2.3808E+01	3.9183E+06	2.3808E+01	3.9183E+06	17.38	9.3679E-08	12.18	0.29	13.80	12.14					
130415_BK7_pane@12	2.05	2.0326E+03	1.0811E+02	1.9391E+01	9.5799E+02	3.9350E+06	9.5799E+02	3.9350E+06	9.5799E+02	3.9350E+06	17.43	9.4235E-08	13.67	0.22	13.62	0.22					
130415_BK7_pane@13	2.05	2.0244E+03	1.1562E+02	1.9411E+01	9.3487E+02	3.8994E+06	9.3487E+02	3.8994E+06	9.3487E+02	3.8994E+06	17.47	9.3805E-08	14.54	0.23	14.50	0.23					
130415_BK7_pane@14	2.06	2.0321E+03	6.3894E+03	1.9441E+01	1.7510E+01	3.9690E+06	1.7510E+01	3.9690E+06	1.7510E+01	3.9690E+06	17.51	9.4881E-08	13.42	0.13	13.38	0.13					
130415_BK7_pane@15	2.06	2.0317E+03	7.5534E+03	1.8609E+01	1.5122E+01	3.7808E+06	1.4849E+01	3.7808E+06	1.4849E+01	3.7808E+06	17.56	9.0586E-08	13.22	0.15	13.18	0.15					
130415_BK7_pane@16	2.06	2.0344E+03	7.0346E+03	1.8754E+01	9.0756E+02	3.8380E+06	9.3422E+02	3.8380E+06	9.3422E+02	3.8380E+06	18.00	9.1533E-08	14.08	0.14	14.04	0.14					
130415_BK7_pane@17	2.06	2.0341E+03	9.5619E+03																		

Table A2 Continued.

SIMS ANALYSES		Montic BR7		Analyte: δ ¹⁸ O		Value: 981±0.14 (V-SMOW, 2σ)		Date: 13.04.2015							
Standard	Standard	Standard	Standard	Standard	Standard	Standard	Standard	Standard	Standard						
130415_BR7_psm0016	9.9864E-03	6.0341E-09	6.4520E-02	3.2482E+06	6.1022E-02	19.31	7.8939E+08	4.02	0.20	3.88	3.99	0.20	9.92	0.20	
130415_BR7_psm0017	6.1700E-03	1.6248E+09	5.2735E-02	3.2708E+06	5.2544E-02	19.35	7.8939E+08	3.84	0.12	0.17	3.81	0.12	9.74	0.18	
130415_BR7_psm0018	1.4825E-02	1.6259E+09	4.7767E-02	3.2728E+06	3.2728E-02	19.39	7.8980E+08	3.86	0.30	3.83	0.30	3.83	0.30	9.76	0.23
130415_BR7_psm0019	9.5201E-03	1.6287E+09	5.5131E-02	3.2785E+06	6.2063E-02	19.43	7.9131E+08	3.85	0.19	3.82	0.19	3.79	0.19	9.75	0.20
130415_BR7_psm0020	6.1603E-03	1.6290E+09	5.8555E-02	3.2789E+06	5.5126E-02	19.48	7.8920E+08	3.81	0.12	3.79	0.12	3.72	0.12	9.71	0.18
130415_BR7_psm0021	9.5340E-03	1.5811E+09	3.7022E-02	3.2435E+06	1.6316E-02	19.52	7.6161E+08	26.16	0.19	25.38	26.14	0.19	32.20	0.20	
130415_BR7_psm0022	9.1871E-03	1.5577E+09	4.7364E-02	3.2409E+06	4.7538E-02	19.56	7.5866E+08	25.17	0.18	0.95	25.15	0.18	31.20	0.20	
130415_BR7_psm0023	1.2620E-02	1.5505E+09	7.4777E-02	3.1858E+06	7.6874E-02	20.01	7.5302E+08	24.64	0.25	24.61	0.25	24.61	0.25	30.66	0.22
130415_BR7_psm0024	2.0851E-03	1.5065E+09	9.0569E-02	3.2275E+06	8.6779E-02	20.05	7.6412E+08	24.91	0.18	24.89	0.18	24.89	0.18	30.94	0.20
130415_BR7_psm0025	1.0185E-02	1.5665E+09	4.0333E-02	3.2192E+06	4.3887E-02	20.09	7.5909E+08	24.85	0.20	24.83	0.20	24.83	0.20	30.88	0.20
130415_BR7_psm0026	2.0570E-03	1.5485E+09	5.0414E-02	3.1858E+06	5.3561E-02	20.13	7.4882E+08	25.82	0.19	25.79	0.19	25.79	0.19	31.85	0.20
130415_BR7_psm0027	9.3402E-03	1.5546E+09	3.1660E-02	3.1955E+06	4.6993E-02	20.18	7.5352E+08	25.48	0.19	25.46	0.19	25.46	0.19	31.51	0.20
130415_BR7_psm0028	7.7441E-03	1.5385E+09	3.0539E-02	3.1745E+06	2.9888E-02	20.22	7.4667E+08	25.53	0.15	25.51	0.15	25.51	0.15	31.57	0.19
130415_BR7_psm0029	1.0821E-02	1.5721E+09	7.9837E-02	3.2488E+06	7.1183E-02	20.26	7.6331E+08	25.69	0.22	25.67	0.22	25.67	0.22	31.73	0.21
130415_BR7_psm0030	7.8836E-03	1.5536E+09	6.1468E-02	3.1946E+06	6.7371E-02	20.31	7.5520E+08	25.52	0.16	25.50	0.16	25.50	0.16	31.56	0.19
130415_BR7_psm0031	2.0133E-03	1.6395E+09	2.9975E-02	3.3044E+06	3.9715E-02	20.35	7.9542E+08	4.14	0.17	3.82	4.12	0.17	10.05	0.19	
130415_BR7_psm0032	1.0441E-02	1.6386E+09	6.2560E-02	3.2978E+06	6.3786E-02	20.39	7.9236E+08	3.72	0.21	0.41	3.70	0.21	9.62	0.20	
130415_BR7_psm0033	1.0813E-02	1.6413E+09	4.8296E-02	3.3039E+06	5.0713E-02	20.43	7.9121E+08	3.89	0.22	3.87	0.22	3.87	0.22	9.80	0.20
130415_BR7_psm0034	7.8606E-03	1.6408E+09	4.0344E-02	3.3023E+06	3.9687E-02	20.48	7.9402E+08	3.71	0.16	3.70	0.16	3.70	0.16	9.62	0.19
130415_BR7_psm0035	9.1706E-03	1.6200E+09	3.9576E-02	3.2685E+06	3.9093E-02	20.52	7.8483E+08	3.63	0.18	3.61	0.18	3.61	0.18	9.54	0.19
130415_BR7_psm0036	1.5711E-09	4.7466E-02	3.2261E+06	4.8092E-02	3.9715E-02	20.56	7.6286E+08	24.21	0.20	24.42	24.19	0.20	30.24	0.20	
130415_BR7_psm0037	9.9265E-03	1.5448E+09	6.5358E-02	3.1743E+06	6.3441E-02	21.01	7.4770E+08	24.83	0.19	1.36	24.81	0.19	30.87	0.20	
130415_BR7_psm0038	6.2825E-03	1.5344E+09	4.1491E-02	3.1516E+06	4.5519E-02	21.05	7.4088E+08	24.47	0.13	24.45	0.13	24.45	0.13	30.50	0.19
130415_BR7_psm0039	1.0020E-02	1.5451E+09	5.0174E-02	3.1956E+06	5.1811E-02	21.09	7.5043E+08	25.33	0.20	25.32	0.20	25.32	0.20	31.37	0.20
130415_BR7_psm0040	1.6695E-02	1.5465E+09	2.4248E-02	3.1794E+06	2.9871E-02	21.13	7.4871E+08	24.06	0.33	24.04	0.33	24.04	0.33	30.09	0.24
130415_BR7_psm0041	8.9509E-03	1.5508E+09	6.4003E-02	3.1855E+06	6.5277E-02	21.18	7.5303E+08	24.28	0.18	24.28	0.18	24.28	0.18	30.31	0.20
130415_BR7_psm0042	2.0861E-03	1.5577E+09	3.7181E-02	3.2055E+06	3.7255E-02	21.22	7.5309E+08	25.36	0.26	25.36	0.26	25.36	0.26	31.40	0.22
130415_BR7_psm0043	1.1876E-02	1.5531E+09	5.4693E-02	3.1880E+06	5.8873E-02	21.26	7.5171E+08	23.66	0.23	23.65	0.23	23.65	0.23	29.70	0.21
130415_BR7_psm0044	8.4085E-03	1.5575E+09	2.3984E-02	3.1966E+06	2.3769E-02	21.31	7.5168E+08	24.74	0.17	24.73	0.17	24.73	0.17	30.78	0.19
130415_BR7_psm0045	6.2589E-03	1.5955E+09	2.4792E-02	3.1921E+06	2.5309E-02	21.35	7.4879E+08	23.23	0.13	23.22	0.13	23.22	0.13	29.26	0.18
130415_BR7_psm0046	1.0319E-02	1.6243E+09	3.5448E-02	3.2699E+06	3.8232E-02	21.39	7.8960E+08	4.03	0.21	3.83	4.02	0.21	9.94	0.20	
130415_BR7_psm0047	7.6486E-03	1.6159E+09	5.8938E-02	3.2325E+06	6.0901E-02	21.44	7.8657E+08	3.70	0.15	0.25	3.69	0.15	9.62	0.19	
130415_BR7_psm0048	1.3055E-02	1.6212E+09	5.3563E-02	3.2639E+06	4.5590E-02	21.48	7.8777E+08	3.76	0.23	3.75	0.23	3.75	0.23	9.88	0.20
130415_BR7_psm0049	9.3745E-03	1.6257E+09	7.3638E-02	3.2731E+06	7.0575E-02	21.52	7.8888E+08	3.88	0.19	3.87	0.19	3.87	0.19	9.80	0.19
130415_BR7_psm0050	1.1082E-02	1.6261E+09	4.9226E-02	3.2729E+06	4.8042E-02	21.56	7.8965E+08	3.79	0.22	3.78	0.22	3.78	0.22	9.71	0.20
130415_BR7_psm0051	1.5701E-02	1.5701E+09	3.6428E-02	3.1383E+06	3.6716E-02	22.01	7.4791E+08	18.27	0.25	19.85	18.26	0.25	24.27	0.21	
130415_BR7_psm0052	1.1904E-02	1.5371E+09	4.1123E-02	3.1391E+06	4.5267E-02	22.05	7.4490E+08	20.75	0.24	6.32	20.75	0.24	26.78	0.22	
130415_BR7_psm0053	1.6670E-02	1.5830E+09	9.4649E-02	3.2132E+06	9.4160E-02	22.09	7.6971E+08	12.27	0.33	12.27	0.33	12.27	0.33	30.87	0.20
130415_BR7_psm0054	8.9288E-03	1.5427E+09	3.2754E-02	3.1951E+06	3.1420E-02	22.14	7.4833E+08	21.43	0.18	21.42	0.18	21.42	0.18	27.45	0.20
130415_BR7_psm0055	1.0884E-02	1.5673E+09	4.2129E-02	3.2071E+06	4.7898E-02	22.18	7.5758E+08	20.11	0.22	20.10	0.22	20.10	0.22	26.13	0.21
130415_BR7_psm0056	8.9179E-03	1.5794E+09	4.1225E-02	3.2485E+06	4.7880E-02	22.22	7.6548E+08	23.52	0.18	23.52	0.18	23.52	0.18	29.56	0.20
130415_BR7_psm0057	9.4479E-03	1.5644E+09	6.2714E-02	3.2084E+06	5.8429E-02	22.27	7.5623E+08	22.06	0.19	22.05	0.19	22.05	0.19	28.09	0.20
130415_BR7_psm0058	1.1356E-02	1.5588E+09	5.7982E-02	3.1778E+06	6.1339E-02	22.31	7.5380E+08	18.63	0.23	18.63	0.23	18.63	0.23	24.64	0.21
130415_BR7_psm0059	1.3074E-02	1.5618E+09	3.3990E-02	3.2016E+06	3.8277E-02	22.35	7.5467E+08	22.33	0.26	22.33	0.26	22.33	0.26	28.36	0.22
130415_BR7_psm0060	9.3345E-03	1.5385E+09	3.3097E-02	3.1489E+06	3.1000E-02	22.39	7.4488E+08	19.15	0.19	19.15	0.19	19.15	0.19	25.16	0.20
130415_BR7_psm0061	6.2186E-03	1.6355E+09	4.3566E-02	3.2923E+06	4.0631E-02	22.44	7.9246E+08	3.73	0.25	3.91	3.73	0.25	9.65	0.21	
130415_BR7_psm0062	1.4755E-02	1.6394E+09	4.8020E-02	3.2939E+06	4.7674E-02	22.48	7.9321E+08	3.94	0.12	0.28	3.94	0.12	9.87	0.18	
130415_BR7_psm0063	1.3488E-02	1.6395E+09	4.4770E-02	3.3013E+06	4.8020E-02	22.52	7.9605E+08	4.00	0.27	4.01	0.27	4.01	0.27	9.93	0.22
130415_BR7_psm0064	1.0335E-02	1.6373E+09	5.2963E-02	3.2961E+06	5.3691E-02	22.57	7.9460E+08	4.07	0.21	4.07	0.21	4.07	0.21	10.00	0.20
130415_BR7_psm0065	1.1630E-02	1.6344E+09	5.1662E-02	3.2886E+06	4.9216E-02	23.01	7.9251E+08	3.81	0.23	3.81	0.23	3.81	0.23	9.74	0.21
130415_BR7_psm0066	6.2186E-03	1.5784E+09	3.1388E-02	3.2316E+06	3.3335E-02	23.05	7.6522E+08	17.82	0.12	18.89	17.82	0.12	23.83	0.18	
130415_BR7_psm0067	1.0654E-02	1.5944E+09	4.5810E-02	3.1694E+06	4.5206E-02	23.10	7.5072E+08	21.26	0.21	6.36	21.26	0.21	27.29	0.20	
130415_BR7_psm0068	1.4171E-02	1.5584E+09	5.5255E-02	3.1850E+06	4.8882E-02	23.14	7.6208E+08	19.48	0.28	19.48	0.28	19.48	0.28	25.50	0.23
130415_BR7_psm0069	1.0290E-02	1.5412E+09	3.5255E-02	3.1520E+06	3.9955E-02	23.18	7.5074E+08	19.98	0.21	19.98	0.21	19.98	0.21	26.00	0.20
130415_BR7_psm0070	1.4270E-02	1.5255E+09	5.5228E-02	3.1260E+06	4.4152E-02	23.23	7.4195E+08	21.91	0.29	21.91	0.29	21.91	0.29	27.94	0.23
130415_BR7_psm0071	1.2054E-02	1.5431E+09	5.5999E-02	3.1397E+06	4.4451E-02	23.27	7.4904E+08	11.12	0.24	11.12	0.24	11.12	0.24	24.35	0.22
130415_BR7_psm0072	1.2947E-02	1.5812E+09	3.1247E-02	3.2287E+06	3										

SIMS を用いた中生代放散虫岩の酸素同位体分析

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要 旨

化石殻の炭酸カルシウムの酸素同位体比 ($\delta^{18}\text{O}$) を用いた古海洋研究が広く用いられているが、珪質化石殻については分析の制約や同位体分別の不確定性等のため、古海洋研究への適用例は限られている。本論では、二次イオン質量分析計 (SIMS) によって測定した日本、イタリア、スイス、ルーマニアの中生代チャートに含まれる放散虫化石 $\delta^{18}\text{O}$ 変動の古海洋指標としての有用性について報告する。53 試料 507 点の測定の結果、放散虫殻 $\delta^{18}\text{O}$ は 19.8 ~ 35.3 ‰ で、現世及び新生代の放散虫殻の値と調和的であり、標準試料 UNIL-Q1 の繰り返し測定誤差 0.3 ‰ 以上に 1 チャート試料中の $\delta^{18}\text{O}$ 変化がみられる。このことから、続成作用 (セグリゲーション) の影響による均一化は完全ではなく、初生的な値が保存されている可能性を支持する。さらに、予察的な放散虫化石の $\delta^{18}\text{O}$ 記録は低解像度にもかかわらず、1,000 万年スケールではコノドントのアパタイトや低 Mg 炭酸塩殻に確認される前期-中期三畳紀の正のシフトや後期三畳紀の安定した高い値と調和的であるが、前期ジュラ紀のパンサラッサ海遠洋域における放散虫化石 $\delta^{18}\text{O}$ の約 8 ‰ の負のシフトはテチス海沿岸域の低 Mg 炭酸塩殻には確認されない。さらに高解像度で他指標と比較することで、放散虫化石の $\delta^{18}\text{O}$ 記録の古海洋学的意義をより深く理解できると期待される。

Radiolarian research by the Geological Survey of Japan, AIST, with bibliographic lists from 1950 to 2019

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ITO Tsuyoshi, NAKAE Satoshi and ITAKI Takuya (2020) Radiolarian research by the Geological Survey of Japan, AIST, with bibliographic lists from 1950 to 2019. *Bulletin of the Geological Survey of Japan*, vol. 71(4), p. 395–437, 7 figs, 6 tables, 1 appendix.

Abstract: The Geological Survey of Japan (GSJ), established in 1882, marked its 135th anniversary in 2017 and has issued numerous publications, such as geologic maps, research articles and newsletters, during its history. This article compiles previous GSJ publications related to radiolarian research for future reference. In the GSJ publications from 1950 to 2019, the term of RADIOLARIA in Japanese appears in 252 Geological Maps of the Quadrangle Series (1:50,000), in 21 Geological Maps of the Quadrangle Series (1:200,000), in 75 articles of the Bulletin of the Geological Survey of Japan, in 14 items in Chishitsu News, in 21 items in GSJ Chishitsu News and in seven articles in the Cruise Report. The GSJ publications related to radiolarian research increased during the 1980s, which is consistent with the commonly called Radiolarian Revolution.

Keywords: radiolaria, bibliography, compilation, Paleozoic, Mesozoic, Cenozoic, Japan

1. Introduction

The Geological Survey of Japan (GSJ), a Japanese public organization for geological survey, was established in 1882 under the Ministry of Agriculture and Commerce. In 2001, the National Institute of Advanced Industrial Science and Technology (AIST) was extensively restructured as an independent administrative agency to integrate 15 research institutes, including GSJ.

Since its establishment, GSJ has aimed to make geological maps of Japan and has published many geological maps on several scales (Fig. 1) (Kato *et al.*, 2011). In 1890, GSJ published a geological map of the Japanese Islands (1:3,000,000) for the first time (Fig. 2A). Geologic maps of the Japanese Islands have been often renewed. The most recently published geological map of the Japanese Islands (1:1,000,000) was published in 1992 as the 3rd Edition (Fig. 2B). Twenty quadrangular areas (1:500,000) cover the Japanese Islands (Fig. 3).

GSJ began publishing Geological Maps of the Quadrangle Series (1:50,000) in the 1950s and continued the publication thereafter (e.g. Saito, 2009; Miyazaki, 2018). Radiolarians are important index fossils used to make the geologic maps of the series, some of which contain descriptions of radiolarians. In addition to the

geological maps, GSJ has issued various publications, such as the Bulletin of the Geological Survey of Japan and GSJ Chishitsu News. Some of these publications also contain radiolarian information.

GSJ has also conducted marine surveys since the 1970s (e.g. Arai *et al.*, 2013). Some of their survey results have been presented via GSJ publications, such as in the Cruise Report. Description of radiolarians are included in many of these publications because they are marine protozoa that are generally included in ocean deposits around the Japanese Islands.

Geological Maps of the Quadrangle Series (1:50,000), the Bulletin of the Geological Survey of Japan, GSJ Chishitsu News and Chishitsu News can be downloaded as portable document format (PDF) files from the website of GSJ (Appendix). All documents are OCRed; thus, we searched the documents for the term RADIOLARIA in Japanese (=“放散虫”). The Cruise Report was also downloaded as a PDF file from the website; however, the files are not OCRed.

Here, we present a brief history of radiolarian research by GSJ via the compilation of previous publications. This paper aims to provide bibliographic lists related to radiolarians for future reference.

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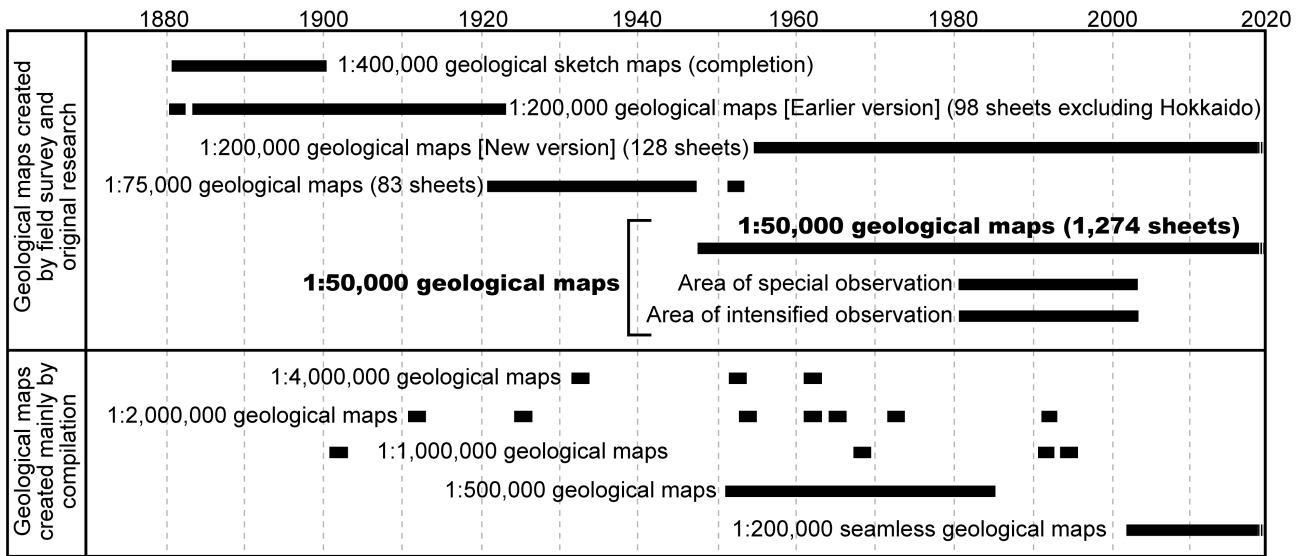


Fig. 1 History of the creation of geological maps by GSI (modified from Kato *et al.*, 2011).

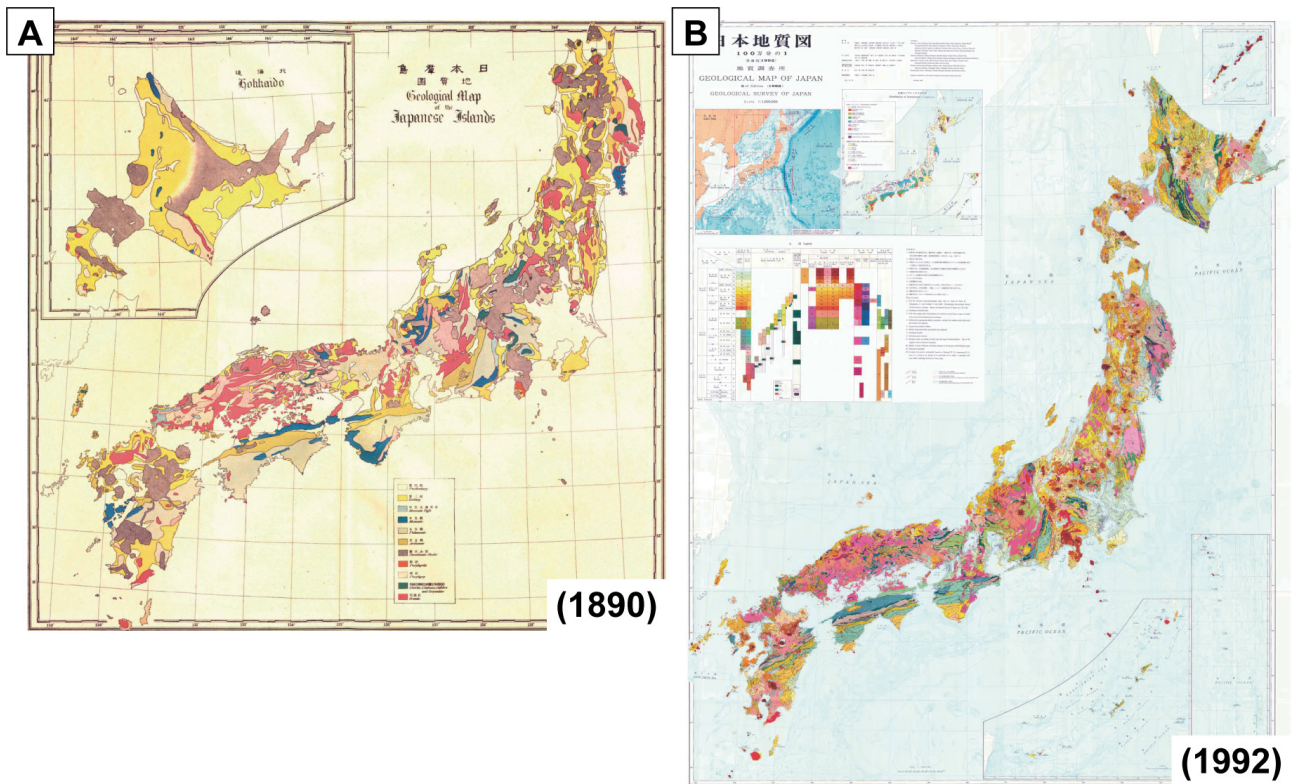


Fig. 2 Geological maps of GSI. A: First published geological map of the Japanese Islands (1:3,000,000) by GSI in 1890, as drawn by T. Harada (reprints from Kato *et al.*, 2011); B: Most recently published geological map of the Japanese Islands (1:1,000,000, 3rd Edition) by GSI in 1992.

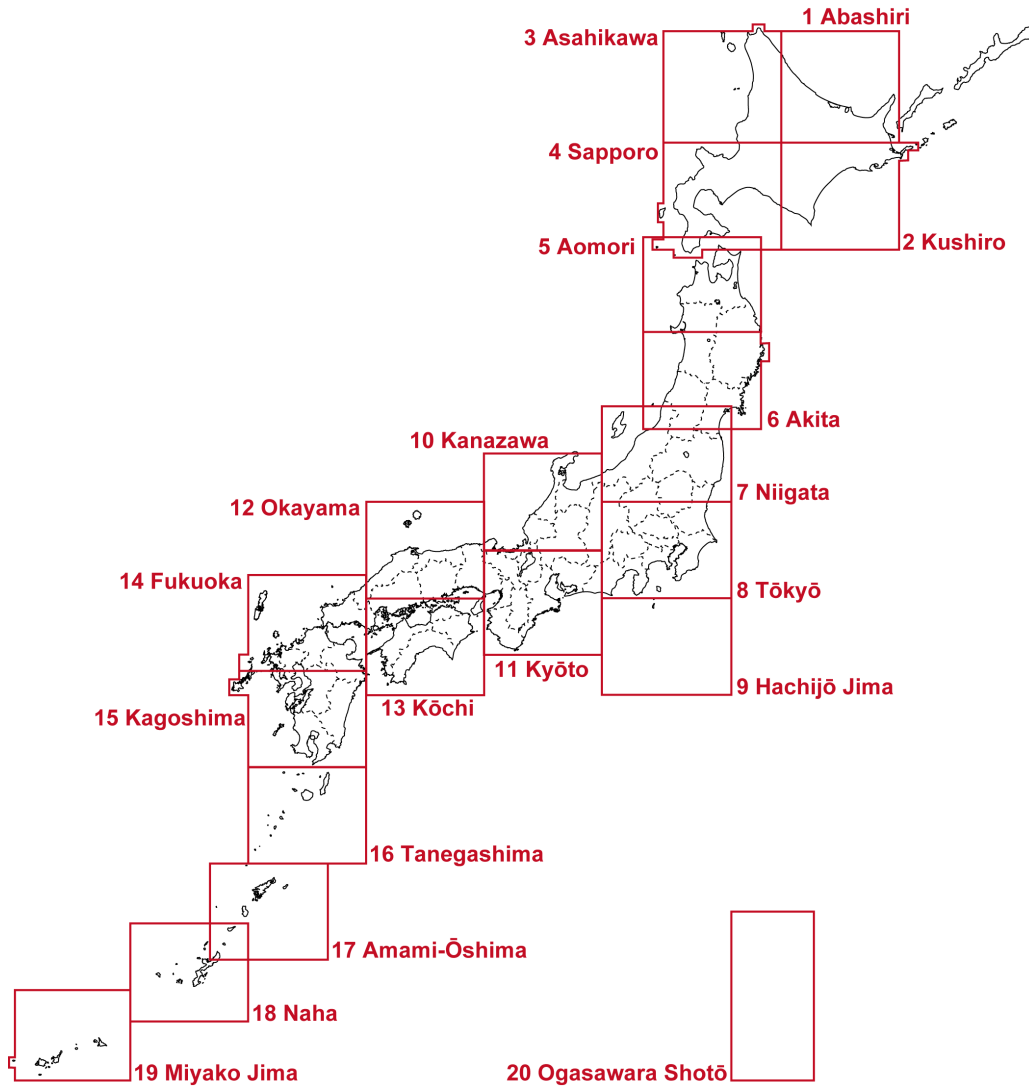


Fig. 3 Japanese Islands with quadrangular areas of 1:500,000.

2. Notable descriptions of radiolarian research in GSJ publications

2.1 Geological Maps of the Quadrangle Series (1:50,000)

The Geological Maps of the Quadrangle Series (1:50,000) have been published since the 1950s (Fig. 1). Between 1951 and 2019, over 700 geological maps in the series have been published. Among them, 252 geological maps contain descriptions of radiolaria (Table 1).

Until the early 1970s, radiolaria were not an important index fossil because they could not be extracted from hard rocks. Pessagno and Newport (1972) discovered a radiolarian extraction method by using hydrofluoric acid (HF). Since their discovery, radiolarian studies have rapidly progressed (e.g. Ichikawa, 1982; Yao *et al.*, 2001; O'Dogherty *et al.*, 2009; Danelian *et al.*, 2017). High-resolution biochronology based on microfossils (radiolaria and conodont) prompted the overturn of

previously believed scenarios for geologic history of the Japanese Islands (e.g. Sakai *et al.*, 1982; Nakaseko *et al.*, 1983; Ichikawa *et al.*, 1985; Ichikawa, 1990; Isozaki and Maruyama, 1991; Yao and Mizutani, 1993; Isozaki *et al.*, 2010; Agematsu-Watanabe and Kamata, 2018). This research progress and the revision of the geologic history are commonly referred to as the Radiolarian Revolution (e.g. Ishigaki and Yao, 1982; Nakaseko, 1984; Sato, 1989; Suzuki and Aita, 2011; Matsuoka and Ito, 2017). Likewise, the GSJ Publications related to radiolarian research also increased in the early 1980s (Fig. 4). Among the geological maps from 1981 to 2019 that include radiolarian descriptions, approximately half present an occurrence list and/or an image of radiolarians in addition to the text.

Meanwhile, even before the 1970s, radiolarian had been described in some geologic maps. In the 1950s, S. Igi had shown an occurrence species list of radiolarians (identified by K. Ichikawa) from chert of the Hidaka Group in the

Table 1 Bibliographic list from the Geological Maps of the Quadrangle Series (1:50,000) that include radiolarian descriptions.
 +: Appearance from its district. -: Appearance from adjacent district(s).

No.	District		Author(s)	Year	Text	Occurrence list	Illustrations			Radiolarian age									
	English	Japanese					SEM	Etched surface	Thin section	Silurian	Devonian	Carboniferous	Permian	Triassic	Jurassic	Cretaceous	Paleogene	Neogene	Quaternary
1 Abashiri																			
9	Kamiokoppe	上興部	Hasegawa K. <i>et al.</i>	1969	+														
10	Okoppe	興部	Hasegawa K. and Uozumi	1975	+														
14	Takinoue	滝上	Matsunami	2002	+							+							
15	Kamishokotsu	上渚滑	Matsunami <i>et al.</i>	2002	-							-							
23	Maruseppu-Hokubu	丸瀬布北部	Yahata <i>et al.</i>	1988	+							+							
24	Engaru	遠軽	Tajika and Yahata	1991	+					+		+							
26, 27	Abashiri	網走	Kawakami G. <i>et al.</i>	2018	-							-							
47	Kitami	北見	Ishida and Sawamura	1968	+														
58	Honki	本岐	Yamaguchi and Sawamura	1965	+														
2 Kushiro																			
釧路																			
2	Tokachigawajoryū	十勝川上流	Sakō and Hasegawa	1957	+														
5	Rikubetsu	陸別	Mitani <i>et al.</i>	1960	+														
19	Ashorobuto	足寄太	Mitani <i>et al.</i>	1958	+														
32	Honbetsu	本別	Mitani <i>et al.</i>	1959	+														
33	Kamicharo	上茶路	Sato S. <i>et al.</i>	1961	+														
53	Nukanai	糠内	Yamaguchi and Satoh	1989	+														
56	Idonnappudake	イドンナツブ岳	Suzuki M. <i>et al.</i>	1961a	+														
59	Chūruī	忠類	Yamaguchi <i>et al.</i>	2003	+														
65	Mitsuishi	三石	Wada <i>et al.</i>	1992	-														
66	Nishicha	西舎	Sakai and Kanie	1986	+									+					
69	Urakawa	浦河	Kanie and Sakai	2002	+									+					
72	Erimo-Misaki	襟裳岬	Igi and Kakimi	1956	+														
3 Asahikawa																			
旭川																			
16	Kamisamufutsu	上猿払	Tanaka K.	1960	+														
20	Pimneshiri	敏音知	Igi	1959	+														
25	Hatsuura	初浦	Hata	1961	+														

Table 1 Continued.

No.	District		Author(s)	Year	Text	Occurrence list	Illustrations			Radiolarian age							
	English	Japanese					SEM	Etched surface	Thin section	Silurian	Devonian	Carboniferous	Permian	Triassic	Jurassic	Cretaceous	Paleogene
35	Soeushinai	添牛内	Hashimoto <i>et al.</i>	1965	+												
39	Horokanai	幌加内	Igi <i>et al.</i>	1958	+	+											
40	Kenbuchi	剣淵	Matsushita <i>et al.</i>	1977	+												
41	Rumoi	留萌	Tsushima and Yamaguchi	1954	+												
42	Ebshima	恵比島	Watanabe and Yaoshida	1995	+												
44	Pippu	比布	Suzuki J.	1957	+												
48	Fukagawa	深川	Suzuki J.	1953	+												
49	Asahikawa	旭川	Suzuki J.	1955	+												
54	Biei	美瑛	Suzuki M. <i>et al.</i>	1961b	+												
4	Sapporo		札幌														
7	Shimofurano	下富良野	Hashimoto	1955	+												
15	Ikushumbetsu-Dake	幾春別岳	Yoshida T. and Kambe	1955	+												
16	Yamabe	山部	Hashimoto	1953	+												
23	Yūbari	夕張	Sasa <i>et al.</i>	1964	+												
24	Ōyubari	大夕張	Nagao <i>et al.</i>	1954	+												
25	Ishikarikanayama	石狩金山	Osanaï <i>et al.</i>	1958	+												
32	Ōiwake	追分	Matsuno and Hata	1960	+												
33	Momijiyama	紅葉山	Takahashi Koh. <i>et al.</i>	2002	+												
34	Hidaka	日高	Takahashi Koh. and Suzuki	1986	+												
35, 46	Harauta and Karibayama	原歌及ひ狩場山	Yamagishi and Kurosawa	1987	-												
41	Tarumaizan	樽前山	Doi	1957	+												
45	Iwachishi	岩知志	Takahashi Koh. and Suzuki	1978	+												
47	Ohbirayama	大平山	Kurosawa <i>et al.</i>	1993	+	+											
52	Shiraoi	白老	Doi	1953	+												
55	Tomikawa	富川	Imai and Sumi	1957	+												
56	Biu	比宇	Yoshida T. <i>et al.</i>	1959	+												

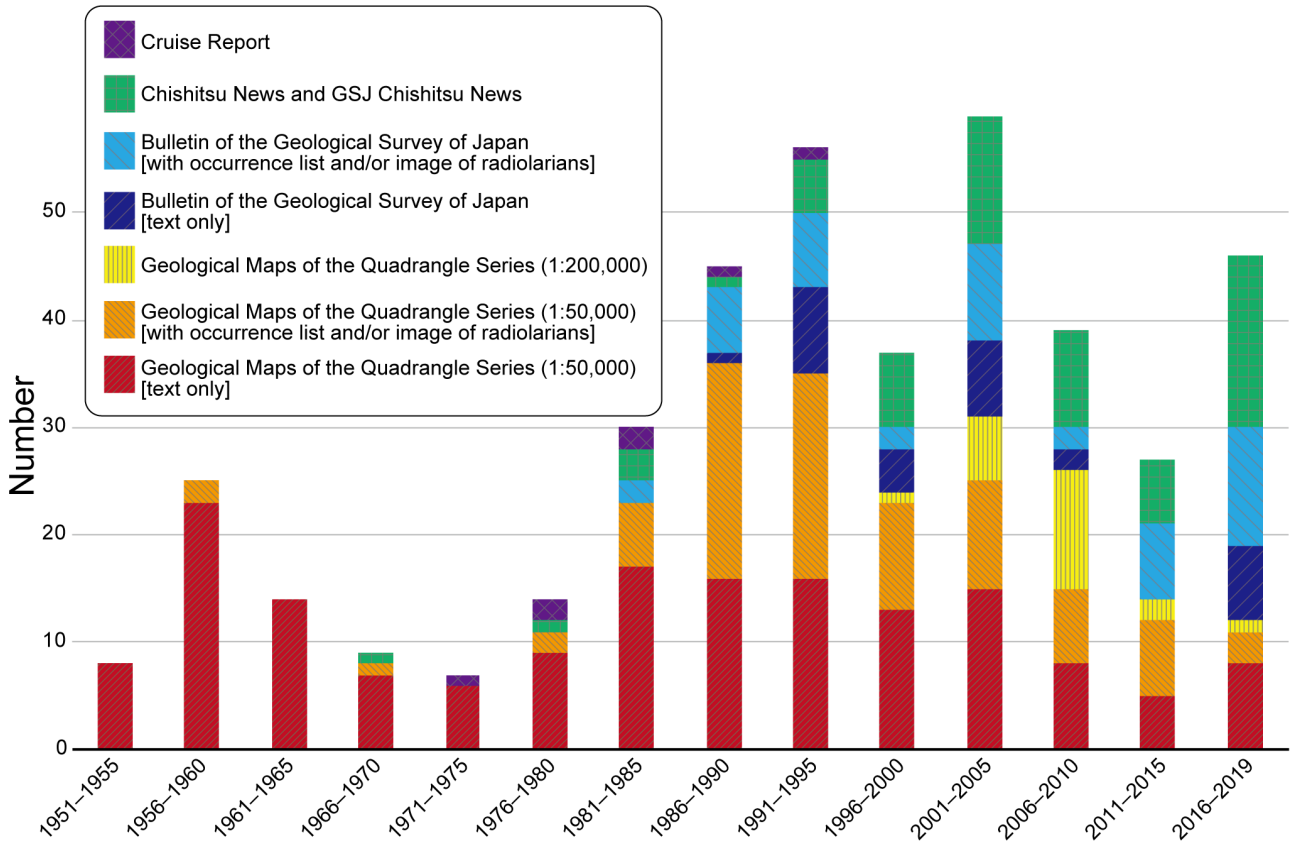


Fig. 4 Quinquennial number of GSJ publications that contain the term RADIOLARIA in Japanese (=“放散虫”).

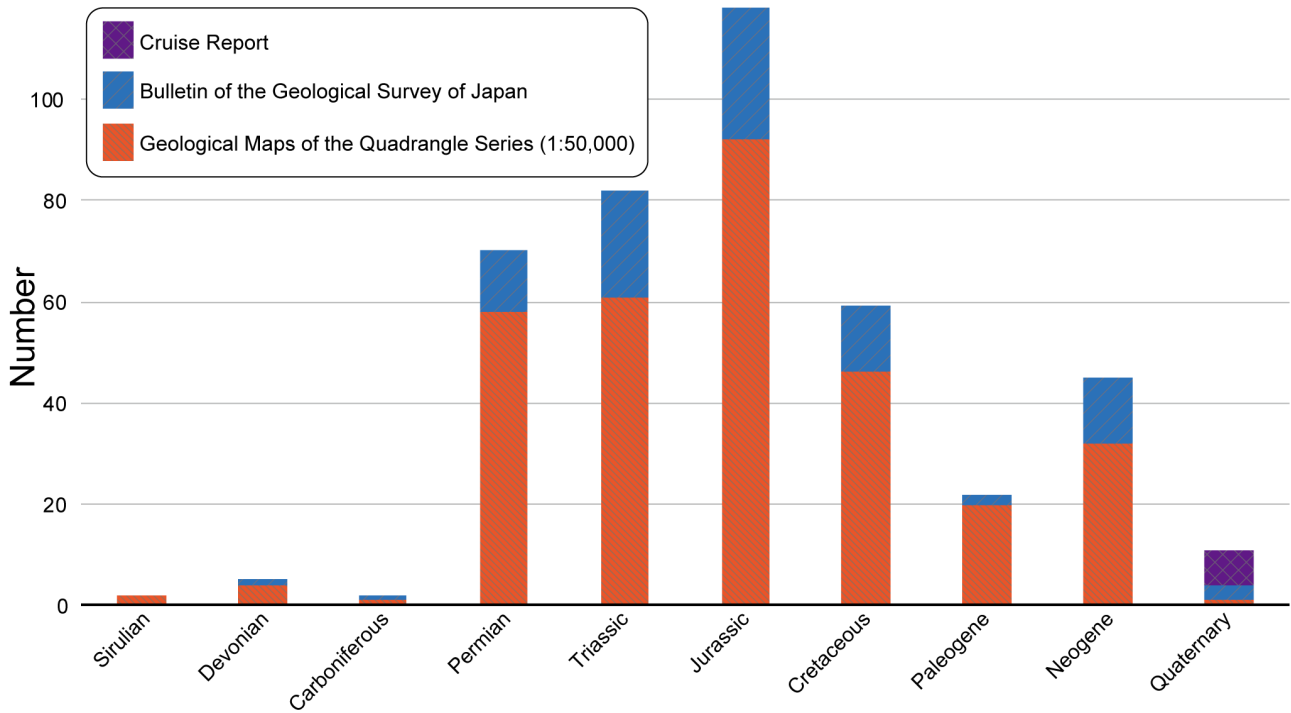


Fig. 5 Age distribution of radiolarian descriptions in GSJ publications.

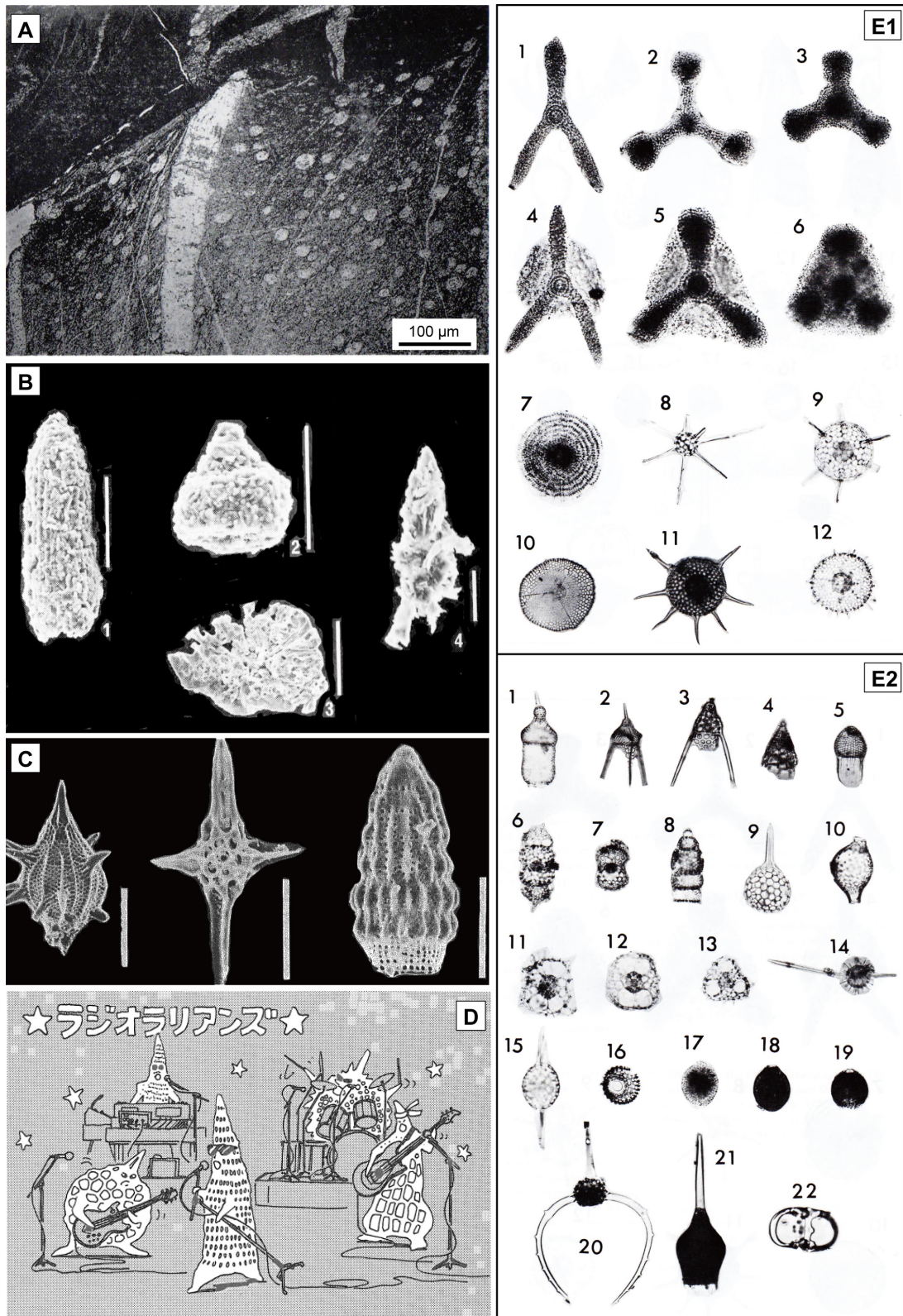


Fig. 6 Notable figures of radiolarian research by GSI. A: First photograph of radiolarians within phyllite (thin section) in the Geological Maps of the Quadrangle Series (1:50,000) (reprints from Yamada, 1966). B: First SEM images of radiolarians in Geological Maps of the Quadrangle Series (1:50,000) (reprints from Sakamoto *et al.*, 1984). C: First SEM images of radiolarians in GSI publications (Chishitsu News) (reprints from Sakai *et al.*, 1982). D: Illustrations of cartoon radiolarians shown in Chishitsu News (reprints from Wakita and Kawamura, 1985). E: Possible first images of radiolarian individuals in GSI publications (Cruise Report) (reprints from Arita and Mizuno, 1977).

Table 2 Bibliographic list from the Geological Maps of the Quadrangle Series (1:200,000) that include radiolarian descriptions.

District		Author(s)	Year
English	Japanese		
Ichinoseki	一関	Takeuchi K. <i>et al.</i>	2005
Ise	伊勢	Nishioka <i>et al.</i>	2010
Ishigaki Jima	石垣島	Nakae <i>et al.</i>	2009
Kaimon Dake and a part of Kuro Shima	開聞岳及び黒島の一部	Kawanabe <i>et al.</i>	2004
Kōfu	甲府	Ozaki <i>et al.</i>	2002
Kubokawa	窪川	Hara <i>et al.</i>	2006
Mito (2nd Edition)	水戸(第2版)	Yoshioka <i>et al.</i>	2001
Nakatsu	中津	Ishizuka <i>et al.</i>	2009
Niigata (2nd Edition)	新潟(第2版)	Takahashi Yut. <i>et al.</i>	2010
Ōita (2nd Edition)	大分(第2版)	Hoshizumi <i>et al.</i>	2015
Okayama and Marugame	岡山及丸亀	Matsuura <i>et al.</i>	2002
Shirakawa	白河	Kubo <i>et al.</i>	2007
Shizuoka and Omae Zaki (2nd Edition)	静岡及び御前崎(第2版)	Sugiyama <i>et al.</i>	2010
Toyohashi and Irigo Misaki	豊橋及び伊良湖岬	Makimoto <i>et al.</i>	2004
Urakawa	浦河	Sakai <i>et al.</i>	2000
Wajima (2nd Edition)	輪島(第2版)	Ozaki <i>et al.</i>	2019
Yaku Shima	屋久島	Saito <i>et al.</i>	2007b
Yamaguchi and Mishima	山口及び見島	Matsuura <i>et al.</i>	2007
Yatsushiro and a part of Nomo Zaki	八代及び野母崎の一部	Saito <i>et al.</i>	2010
Yokosuka (2nd Edition)	横須賀(第2版)	Takeuchi K. <i>et al.</i>	2015
Yoron Jima and Naha	与論島及び那覇	Nakae <i>et al.</i>	2010

2.4 Chishitsu News and GSJ Chishitsu News

Chishitsu News and GSJ Chishitsu News are monthly newsletters published by GSJ. Chishitsu News was published from 1953 to 2011. From 2011 and onward, GSJ Chishitsu News has been published as a successor to Chishitsu News.

In total, 14 and 21 articles containing radiolarian descriptions were published in Chishitsu News and GSJ Chishitsu News, respectively (Tables 4, 5). Fukuda and Natori (1977) showed a transmitted photomicrograph of Neogene radiolaria reprinted from Nakaseko and Sugano (1973). This was possibly the first isolated radiolarian images presented in the GSJ publications. Sakai *et al.* (1982) introduced a micropaleontological study on conodont and radiolaria and described their significance in Chishitsu News at the dawn of the Radiolarian Revolution. They also showed SEM images (Fig. 6C), which were the first SEM images published in the GSJ publications. Wakita and Kawamura (1985) wrote an essay about radiolarians, which included some SEM images, thin section and cartoons (Fig. 6D).

Since 1997, the Geological Museum owned by GSJ has displayed radiolarian exhibits, such as panels and models (Toshimitsu and Saito, 1997). The Geological Museum also made a poster showing radiolarians with

reconstructed oceanic plate stratigraphy in Jurassic accretionary complexes (Fig. 7). Special exhibitions related to radiolarians have often been displayed in the museum (e.g. Shibahara *et al.*, 2012; Ito *et al.*, 2017).

2.5 Cruise Report

GSJ had published the Cruise Report from 1972 to 1997. Twenty-four issues of the Cruise Report were published during this time period. Among them, seven articles contain descriptions of radiolarians (Table 6).

Arita and Mizuno (1977) showed photomicrographs of living radiolarians from the central–eastern part of the Central Pacific Basin (Fig. 6E). These were possibly one of the first isolated radiolarian images in the GSJ publications like Fukuda and Natori (1977).

Acknowledgments: Dr. UCHINO Takayuki (Geological Survey of Japan, AIST) has carefully reviewed the manuscript and has provided constructive comments. Dr. Toshimitsu Seiichi (Geological Survey of Japan, AIST) provided a reprinted poster of “Reconstructed Oceanic Plate Stratigraphy in Jurassic accretionary complexes and radiolarian fossils in Japan” made by the Geological Museum of GSJ.

Table 3 Bibliographic list from the Bulletin of the Geological Survey of Japan that include radiolarian descriptions.

Author(s)	Bibliography				Text	Occurrence list	Images			Radiolarian age							
	Year	Vol.	No.	Pages			SEM	Transmitted	Thin section	Devonian	Carboniferous	Permian	Triassic	Jurassic	Cretaceous	Paleogene	Neogene
Hara and Hara	2019	70	1/2	117–123	+		+							+			
Hara <i>et al.</i>	2012	63	11/12	301–308	+		+		+					+			
Hattori	1993	44	7	455–469	+												
Hori N.	2004a	55	9/10	271–285	+						+	+	+				
	2004b	55	9/10	287–301	+	+	+				+						
	2004c	55	9/10	303–334	+	+	+					+					
	2004d	55	9/10	335–388	+	+	+						+				
	2005	56	1/2	37–83	+	+	+				+	+	+				
Hori N. <i>et al.</i>	2002	53	9/10	689–724	+	+	+						+				
Hori R. S.	1993	44	9	555–570	+		+						+				
Imoto and Saijyo	1993	44	9	547–554	+						+	+					
Ishiga and Yamakita	1993	44	7	419–423	+												
Ishiga <i>et al.</i>	1993	44	12	721–726	+												
Ito	2019a	70	1/2	225–247	+	+	+	+	+			+	+	+			
Kakuwa	1993	44	9	533–546	+												
Kametaka <i>et al.</i>	2005	56	7/8	237–243	+		+		+			+					
Kamikuri	2019a	70	1/2	137–161	+	+		+									+
	2019b	70	1/2	163–194	+	+		+									+
Kashiwagi and Kurimoto	2003	54	7/8	279–293	+	+	+					+	+				
Kimura	1997	48	6	313–337	+												
Kimura and Nakae	1993	44	12	727–743	+	+	+		+				+				
Kojima <i>et al.</i>	1994	45	2	63–97	+	+	+					+	+	+			
Kojima and Saito	2000	51	4	143–165	+		+					+	+				
Kurimoto	1987	38	2	69–80	+	+	+					+	+				
	1989	40	2	55–63	+	+	+					+	+				
	1994a	45	5	235–255	+	+	+	+						+			
Kurimoto and Kuwahara	1991	42	2	69–73	+	+	+					+	+				
Kurimoto <i>et al.</i>	2015	66	3/4	41–79	+	+	+							+			
Matsuzaki and Itaki	2019	70	1/2	195–209	+	+		+									+
Mizutani	2019	70	1/2	261–265	+												
Motoyama	2019	70	1/2	125–136	+												+
Motoyama and Itaki	2019	70	1/2	1–4	+												+
Motoyama and Maruyama	2019	46	7	333–374	+	+	+	+									+
Motoyama <i>et al.</i>	2010	61	3/4	87–103	+	+		+							+	+	+
Musashino	1993	44	12	699–705	+												
Muto <i>et al.</i>	2019	70	1/2	43–89	+							+					
Nakae	1993	44	7	471–481	+												
	2000	51	4	113–128	+	+	+		+				+				
	2001	52	6/7	245–252	+	+	+				+						
	2002	53	1	51–59	+	+	+					+	+				
	2006	57	1/2	29–50	+				+								
	2011	62	11/12	441–453	+	+	+				+						
	2012	63	9/10	269–281	+	+					+						
	2013a	64	3/4	85–112	+	+	+					+	+				
	2013b	64	5/6	151–190	+	+	+					+	+				
2016	67	3	81–100	+	+	+		+				+					

Table 3 Continued.

Author(s)	Bibliography				Text	Occurrence list	Images			Radiolarian age							
	Year	Vol.	No.	Pages			SEM	Transmitted	Thin section	Devonian	Carboniferous	Permian	Triassic	Jurassic	Cretaceous	Paleogene	Neogene
Nakae and Kurihara	2017	68	2	57–86	+	+	+	+						+			
Nakato <i>et al.</i>	2005	56	5/6	225–236	+												+
Noda and Kurihara	2016	67	4	119–131	+	+	+							+			
Saito	1993	44	9	571–596	+	+	+				+	+	+				
Suto <i>et al.</i>	2005	56	11/12	375–409	+												+
Sugiyama K. and Saito	1994	45	7	383–404	+	+	+	+								+	
Takahashi M. <i>et al.</i>	1999	50	3	225–243	+												+
Takemura	2019	70	1/2	267–272	+												
Takeuchi M. and Takizawa	1991	42	9	439–472	+												
Teraoka and Kurimoto	1986	37	8	417–453	+	+	+	+						+			
Tominaga <i>et al.</i>	2019	70	3	299–314	+						+		+	+			
Tuzino	2010	61	3/4	125–136	+												
Uchino	2010	61	9/10	365–381	+												
	2017a	68	2	23–24	+												
Uchino and Hori	2011	62	3/4	191–196	+		+		+			+					
Uchino and Ishida	2017	68	2	25–39	+	+	+						+				
Uchino and Kurihara	2019	70	1/2	109–115	+		+		+	+							
Wakita	1983	34	7	329–342	+	+	+					+	+	+			
	1988a	39	6	367–421	+	+	+		+			+	+	+			
	1988b	39	11	675–757	+							+	+	+			
Wakita and Isomi	1986	37	6	325–333	+		+					+	+				
Wakita and Okamura	1982	33	4	161–185	+		+					+	+	+			
Yanagisawa	1999	50	3	167–213	+												+
	2003a	54	1/2	1–13	+												+
Yanagisawa <i>et al.</i>	1989	40	8	405–467	+	+		+									+
	2003a	54	1/2	29–47	+												+
	2003b	54	11/12	351–364	+												+
Yao	2019	70	1/2	246–260	+												
Yoshii <i>et al.</i>	1997	48	10	567–584	+												

Table 4 Bibliographic list from Chishitsu News that include radiolarian descriptions.

Author(s)	Bibliography			Contents	Images	Age
	Year	No.	Pages			
Endo and Sarashina	2007	632	41–45	Structure and function of proteins		
Fukuda and Natori	1977	273	32–43	Report on international congress about the Neogene in the Pacific	Transmitted photomicrograph	Neogene
Hara <i>et al.</i>	2005	611	49–59	Geology of Lao		
Kanie	2007	633	22–30	Cenozoic stratigraphy of the Ryukyu Arc		
Kanie	1998	532	59–61	Mollusks research in museum		
Kano K. <i>et al.</i>	2003	584	48–49	Outline of the Geological map of Japan (1:2,000,000)		
Kashiwagi <i>et al.</i>	2004	604	15–22	Geology of Mongolia	SEM images	Devonian?
	2005	605	55–60	Geology of Mongolia		
Katada <i>et al.</i>	1970	186	48–51	Thin section of limestone and chert	Thin section	
Kishimoto	1991	437	41–55	Mineral resources in China		
Kurimoto	1994b	482	21–30	Radiolarian biostratigraphy and the Geological maps of Quadrangle Series, 1:50,000	Sketches; SEM images	Mainly Permian to Jurassic
Matsuura <i>et al.</i>	1996	498	22–24	Outline of the Geological map of the Hirone district (Quadrangle Series, 1:50,000)		
Nakajima	1986	387	6–15	Geology around Himalayan		
Nakano <i>et al.</i>	2005	612	53–57	Outline of the Geological map of the Minakuchi district (Quadrangle Series, 1:50,000)		
Nohara <i>et al.</i>	1983	343	9–21	Research on deep sea mineral resource	Transmitted photomicrograph	Quaternary
Saito	1997	514	14–22	Research progress of the Jurassic accretionary complex of Japan		
Saito and Nishioka	1995	487	63–66	Outreach with using geological maps		
Saito and Ozaki	2000	548	59–61	Outreach with using geological maps		
Saito <i>et al.</i>	1997	514	7–13	Exhibition of the Geological Museum		
	2006	619	56–60	Outline of the Geological map of the Tomochi district (Quadrangle Series, 1:50,000)		
	2008	647	52–60	Outline of the Geological map of the Yakushima (1:200,000)	SEM images	Jurassic
Sakai <i>et al.</i>	1982	337	166–167	Research progress of radiolaria and conodont	SEM images	Jurassic
Sato T.	1991a	438	13–25	Research progress of Paleozoic–Mesozoic in Japan		
	1991b	440	19–33	Research progress of Paleozoic–Mesozoic in Japan		
Shinbo	2006	624	42–47	Observation of foraminifera in beach sands		
Takechi	2010	666	48–52	Fossil in Okayama Prefecture	Thin section	

Table 4 Continued.

Author(s)	Bibliography			Contents	Images	Age
	Year	No.	Pages			
Tanaka Y.	2007	634	29–34	Biogenic particles including radiolarian one		
Takahashi Koz.	2002	576	37–43	Diatom and radiolarian (including Phaeodarian)	SEM images	Recent
Takahashi Yut.	1996	506	7–14	Geology of the Iide Mountains		
	2005	607	57–62	Outline of the Geological map of the Suhara district (Quadrangle Series, 1:50,000)		
Teraoka	2004	599	40–48	Geology of Shimanto accretionary complex		
Tokuhashi	2008	645	26–52	Report on excursion of 17th International Sedimentology Congress 2006, Fukuoka		
Toshimitsu and Saito	1997	514	frontispiece	Exhibition of the Geological Museum	Hand-size model	Mainly Mesozoic
Wakita	2001	567	52–66	Geology of Indonesia		
	2002a	574	53–67	Geology of Indonesia		
	2002b	576	44–59	Geology of Indonesia		
Wakita and Kawamura	1985	376	60–66	Research progress of radiolaria	SEM images; Thin section; Cartoon	Devonian–recent
Yamada	2009	660	32–47	Geological map compiled by T. Harada		
	2011	679	8–22	Geological map compiled by T. Kochibe and others		
Yoshida F.	2003	592	61–63	Photographs of Shirasaki coast		

Table 5 Bibliographic list from GSJ Chishitsu News that include radiolarian descriptions.

Author(s)	Bibliography				Contents	Images	Age
	Year	Vol.	No.	Pages			
Hara and Ito	2018	7	11	259–261	International training course in the Kanto Mountains including field excursion and observation on radiolarian fossil	SEM images	Permian
Itaki	2019	8	5	125–127	Artificial intelligence technology for accurate identification and sampling of radiolarians	Transmitted photomicrograph	Quaternary
Ito	2017a	6	5	166–174	Author's radiolarian research in China		
	2017b	6	5	175–178	Chinese signage of words of sedimentology including "radiolarian ooze"		
	2017c	6	11	373–376	Chinese signage of words of paleontology including "radiolarian"		
	2017d	6	11	377–380	Chinese local name based on scientific name with examples of Permian radiolarians		
	2019b	8	7	175–180	Report on the 5th International Palaeontological Congress, with observation of samples of Deflandre (1952)	Transmitted photomicrograph	Carboniferous
Kano K.	2013	2	8	235–238	Outline of the Geological map of the Aniai district, 2nd Edition (Quadrangle Series, 1:50,000)		
Kato	2012	1	10	293–309	Excursion in the Chichibu area including "radiolarian slate" noted by K. Hosaka		
Kawabata	2016	5	8	263–265	Introduction of new staffs of the Geological Survey of Japan in 2016		
Nakashima <i>et al.</i>	2015	4	8	230–234	Lecture of Cenozoic stratigraphy in Japan		
Ozaki	2019	8	2	31–40	Outline of the Geological map of the Minobu district (Quadrangle Series, 1:50,000)		
Takahashi M.	2017	6	5	149–157	Discussion on tectonic boundary between Northeast and Southwest Japan		
Takahashi Yut. <i>et al.</i>	2018	7	11	303–308	International training course in the Abukuma Mountains		
Toshimitsu <i>et al.</i>	2019	8	12	322–335	Chronological timetable of the Geological Museum		
Tuzino <i>et al.</i>	2019	8	10	261–272	Exhibition of rocks (inc. radiolarian chert) in the Geological Museum		
Uchino	2015	4	3	69–74	Origin of "Shiraishi" in the "Shikinen Sengu ceremony" at Ise Jingu		
	2017b	6	9	283–288	Vegetation on chert in the Toba District		
	2018	7	4	91–101	Outline of the Geological map of the Toba district (Quadrangle Series, 1:50,000)		
Uchino and Kawamura	2014	3	11	329–333	Outline of the Geological map of the Hayachine San district (Quadrangle Series, 1:50,000)		
Utsunomiya	2018	7	9	223–226	Report on the 16th International Nannoplankton Association Meeting		

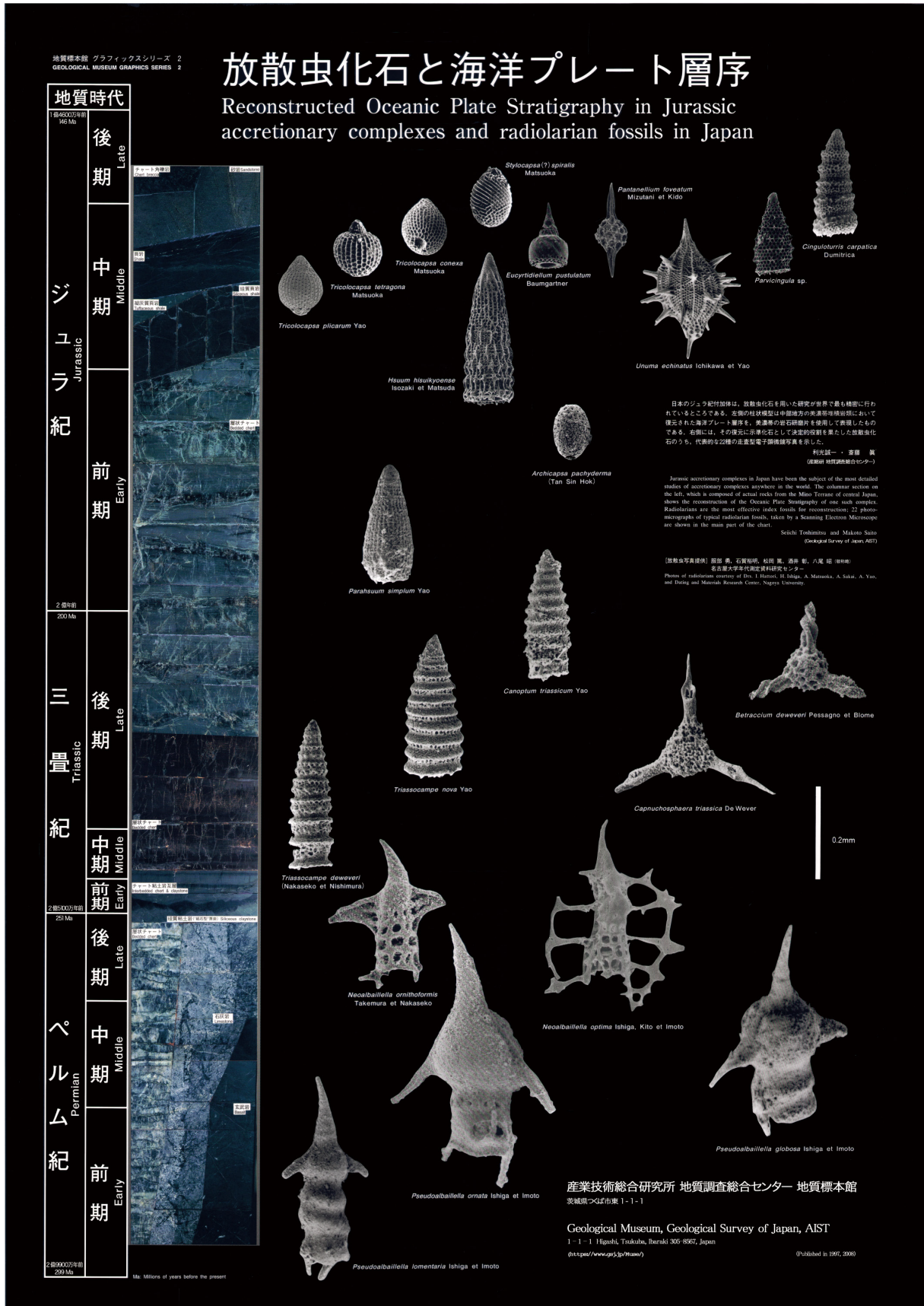


Fig. 7 Reprinted poster of “Reconstructed Oceanic Plate Stratigraphy in Jurassic accretionary complexes and radiolarian fossils in Japan” made by the Geological Museum of GSJ.

Table 6 Bibliographic list from the Cruise Report that include radiolarian descriptions.

Author(s)	Bibliography			Ocean area	Occurrence list	Images	Age
	Year	No.	Pages				
Arita and Mizuno	1977	8	301–308	Central Pacific		Transmitted photomicrograph	Quaternary
Hasegawa S. <i>et al.</i>	1976	7	80–85	Southern Kurile Trench and Slope			Quaternary
Inoue <i>et al.</i>	1972	1	20–33	Northwest Pacific			Quaternary
Nishimura	1984	20	67–89	Magellan Trough		Smear slide	Quaternary
	1986	21	56–83	Central Pacific		Transmitted photomicrograph; SEM image	Quaternary
Nishimura and Ikehara	1992	22	85–96	Central Pacific	+		Quaternary
Takayangi <i>et al.</i>	1982	18	301–308	Wake-Tahiti Transect in the Central Pacific	+		Quaternary

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Appendix: How to access and download the publications mentioned in this article

A1. Geological Map of the Quadrangle Series (1:50,000 and 1:200,000)

Access the web page (<https://gbank.gsj.jp/datastore/download.php?lang=en>) and do as follows. The numbers correspond to the numbers in the circles in Fig. A1.

- 1: Enter or copy-paste a district name.
- 2: Click the “search” button.
- 3: Select the check box that you need.
- 4: Click the “download” button.

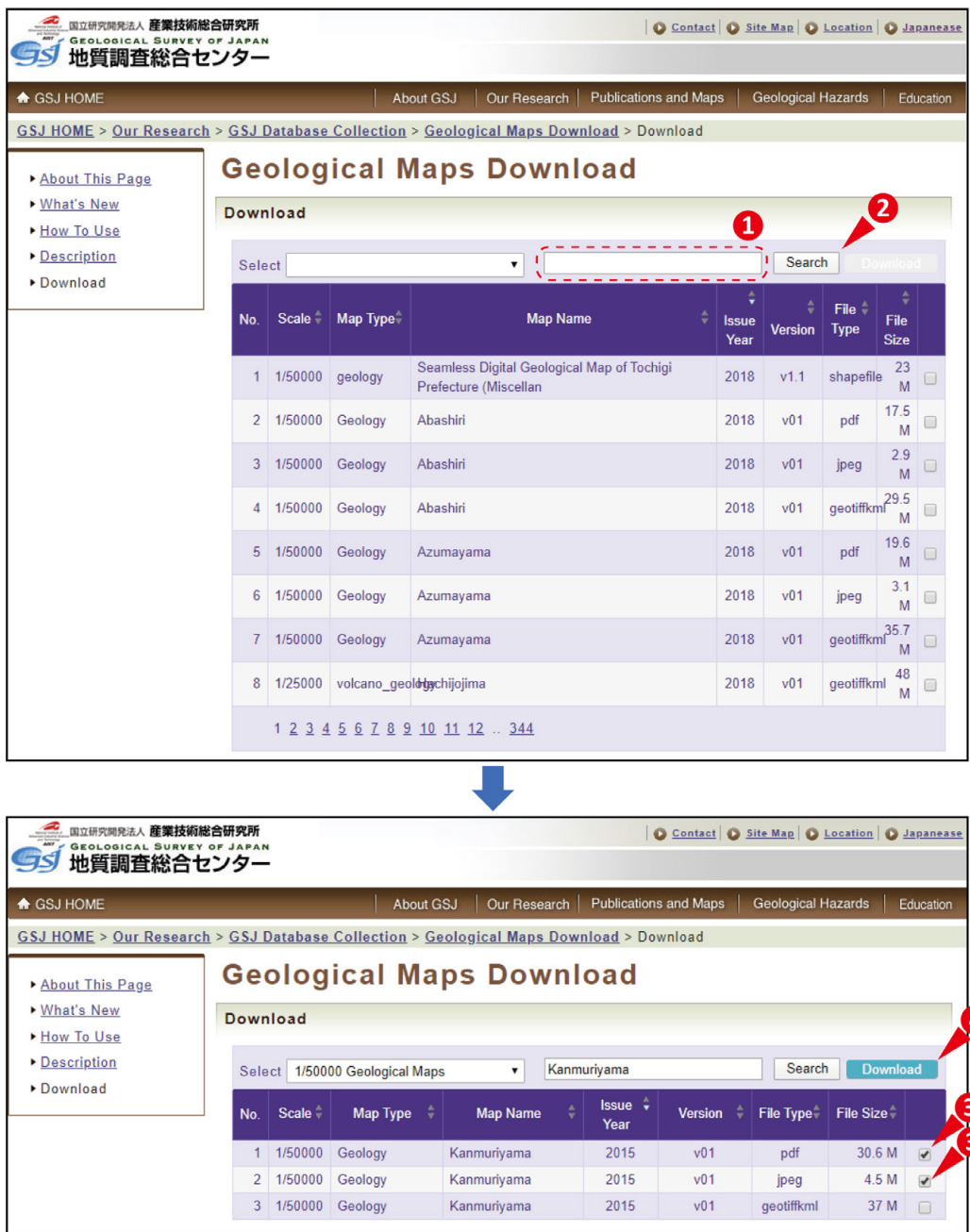


Fig. A1 Image captures of the webpage for search and download of geological maps from GSJ.

A2. Bulletin of the Geological Survey of Japan (1950 to 2011, vol. 52, no. 2/3)

Access the web page (<https://www.gsj.jp/en/publications/bull-gsj/index.html>) and download the article of interest.

A3. Bulletin of the Geological Survey of Japan (2011, vol. 52, no. 4/5 to present)

Access the web page (<https://www.gsj.jp/en/publications/bulletin/index.html>) and download the article of interest.

A4. Chishitsu News [Japanese only]

Access the web page (<https://www.gsj.jp/publications/pub/chishitsunews/news-contents.html>) and download the article of interest.

A5. GSJ Chishitsu News [Japanese only]

Access the web page (<https://www.gsj.jp/publications/gcn/gcn.html>) and download the article of interest.

A6. Cruise Report

Access the web page (<https://www.gsj.jp/en/publications/cruise-rep/index.html>) and download the article of interest.

地質調査総合センターにおける放散虫研究の歴史及び
1950年～2019年(昭和25年～令和元年)の関連出版物目録

伊藤 剛・中江 訓・板木 拓也

要 旨

1882年(明治15年)に設立された地質調査所(現産業技術総合研究所地質調査総合センター)は、2017年に創立135周年を迎えた。その歴史の中で、地質図、論文、ニュース誌など、数多くの出版物を刊行してきた。本論では、これらの出版物の中で放散虫に関係するものを纏めた。1950年(昭和25年)から2019年(令和元年)の間の出版物の中で、「放散虫」という単語は、5万分の1地質図幅では252編、20万分の1地質図幅では21編、地質調査所月間報告及び地質調査研究報告では75編、地質ニュースでは14編、GSJ地質ニュースでは21編、Cruise Reportでは7編の論文・記事で記述されている。放散虫研究にかかわる出版物の数は1980年代に増加しており、これはいわゆる放散虫革命と同時期である。

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- 702 新規富士火山降下火砕物の層序、平均最大粒径、最小体積及び化学組成 山元 孝広・中野 俊・石塚 吉浩・高田 亮
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